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Critical review of fluorescence and absorbance measurements as surrogates for the molecular weight and aromaticity of dissolved organic matter†

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Dissolved organic matter (DOM) is ubiquitous in aquatic environments and challenging to characterize due to its heterogeneity. Optical measurements (*i.e.*, absorbance and fluorescence spectroscopy) are popular characterization tools, because they are non-destructive, require small sample volumes, and are relatively inexpensive and more accessible compared to other techniques (*e.g.*, high resolution mass spectrometry). To make inferences about DOM chemistry, optical surrogates have been derived from absorbance and fluorescence spectra to describe differences in spectral shape (*e.g.*, E2:E3 ratio, spectral slope, fluorescence indices) or quantify carbon-normalized optical responses (*e.g.*, specific absorbance (SUVA) or specific fluorescence intensity (SFI)). The most common interpretations relate these optical surrogates to DOM molecular weight or aromaticity. This critical review traces the genesis of each of these interpretations and, to the extent possible, discusses additional lines of evidence that have been developed since their inception using datasets comparing diverse DOM sources or strategic endmembers. This review draws several conclusions. More caution is needed to avoid presenting surrogates as specific to either molecular weight or aromaticity, as these physicochemical characteristics are often correlated or interdependent. Many surrogates are proposed using narrow contexts, such as fractionation of a limited number of samples or dependence on isolates. Further study is needed to determine if interpretations are generalizable to whole-waters. Lastly, there is a broad opportunity to identify why endmembers with low abundance of aromatic carbon (*e.g.*, effluent organic matter, Antarctic lakes) often do not follow systematic trends with molecular weight or aromaticity as observed in endmembers from terrestrial environments with higher plant inputs.

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Environmental significance

Dissolved organic matter (DOM) is ubiquitous in aquatic environments and plays important roles in environmental biogeochemical processes and water treatment operations. A common method to detect and characterize DOM is through readily accessible optical measurements. This critical review discusses the foundations and more recent insights into how optical measurements are used as surrogates for DOM chemistry.

1 Introduction

Ubiquitous in aquatic environments, dissolved organic matter (DOM) is a heterogeneous mixture of organic compounds composed primarily of carbon, nitrogen, oxygen, and hydrogen (with small contributions from phosphorous and sulfur).[1,2] DOM optical properties impact a multitude of processes in natural and engineered systems. For example, in lakes and rivers, DOM absorbs light, which decreases the light penetration in the water column and simultaneously produces transient oxidants.[3–6] In engineered systems, DOM can both inhibit treatment through light screening or enhance it through reactive intermediate formation.[7,8] Across contexts, DOM impacts biogeochemical processes through the surface reactivity of metal (nano)particles and other mineral surfaces.[9] By characterizing either DOM quantity or quality, optical measurements are indirect surrogates for DOM reactivity, such as biodegradability,[10] contaminant sorption,[11,12] water treatment efficiency,[13,14] and disinfection byproduct formation.[15] Given these

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† Electronic supplementary information (ESI) available: Contains search terms for the citation analysis in Fig. 1, supplementary graphs, derivations related to inner filter effects and the humification index, and the methodology for estimating specific fluorescence intensity from PARAFAC components. A spreadsheet is also provided that compiles all of the data presented in the figures. See DOI: https://doi.org/10.1039/d4em00183d

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multiple roles, a significant research effort aims to advance the use of optical surrogates to characterize DOM physicochemical properties and the impact of natural and engineered processes on these properties.

Optical measurements (*i.e.*, absorbance and fluorescence spectroscopy) are among the most frequent and oldest techniques to study DOM. Two chapters by Berzelius (1806) are commonly attributed as first reports of DOM color,[16] and its inquiry is as old as the debate over the color of water. Predating the discovery of Raman scattering, Bancroft (1919)[17] outlines the contemporary debates about color and the challenge of reconciling observations of blue water bodies with yellow residues upon evaporation. Concomitantly in Saville (1917),[18] the drinking water industry was discussing the implications of color on water treatment. To the best of our knowledge, the first recognition of fluorescent DOM was Dienert (1908),[19] who observed DOM fluorescence as a source of error during a fluorescein tracer study prior to deliberately investigating different source waters in a 1910 study.[20]

Illustrated in Fig. 1c and d, current DOM research using optical measurements leverages spectral parameters (*e.g.*, specific ultraviolet absorbance (SUVA) and fluorescence index (FI)) to serve as *surrogates* for DOM physicochemical properties like molecular weight (or size) and aromaticity. The number of publications reporting optical surrogates has continued to grow (Fig. 1a and b). The draw of optical measurements is their ease of use compared to other methods. For most applications, whole-water samples can be characterized directly with as little as ~4 mL of sample. The limited sample volume and ease of measurement permits high coverage across spatiotemporal scales, whether in natural systems or engineering applications. Recent studies have also taken advantage of the short analysis time to characterize spectra across multiple chemical (*e.g.*, pH,[21] borohydride reduction,[22] photooxidation[23]) or fractionated sample (*e.g.*, size exclusion,[24] solid phase extraction[25]) dimensions.

Despite their success and frequent use, optical surrogates are often not paired with independent measures of molecular weight and aromaticity, such as size exclusion chormatography (SEC) or nuclear magnetic resonance (NMR), respectively. This decision is understandable due to instrumental, sample, or cost limitations, but it leads to applying generalized interpretations in contexts different from the original studies in which physicochemical relationships were proposed. As a result, trends

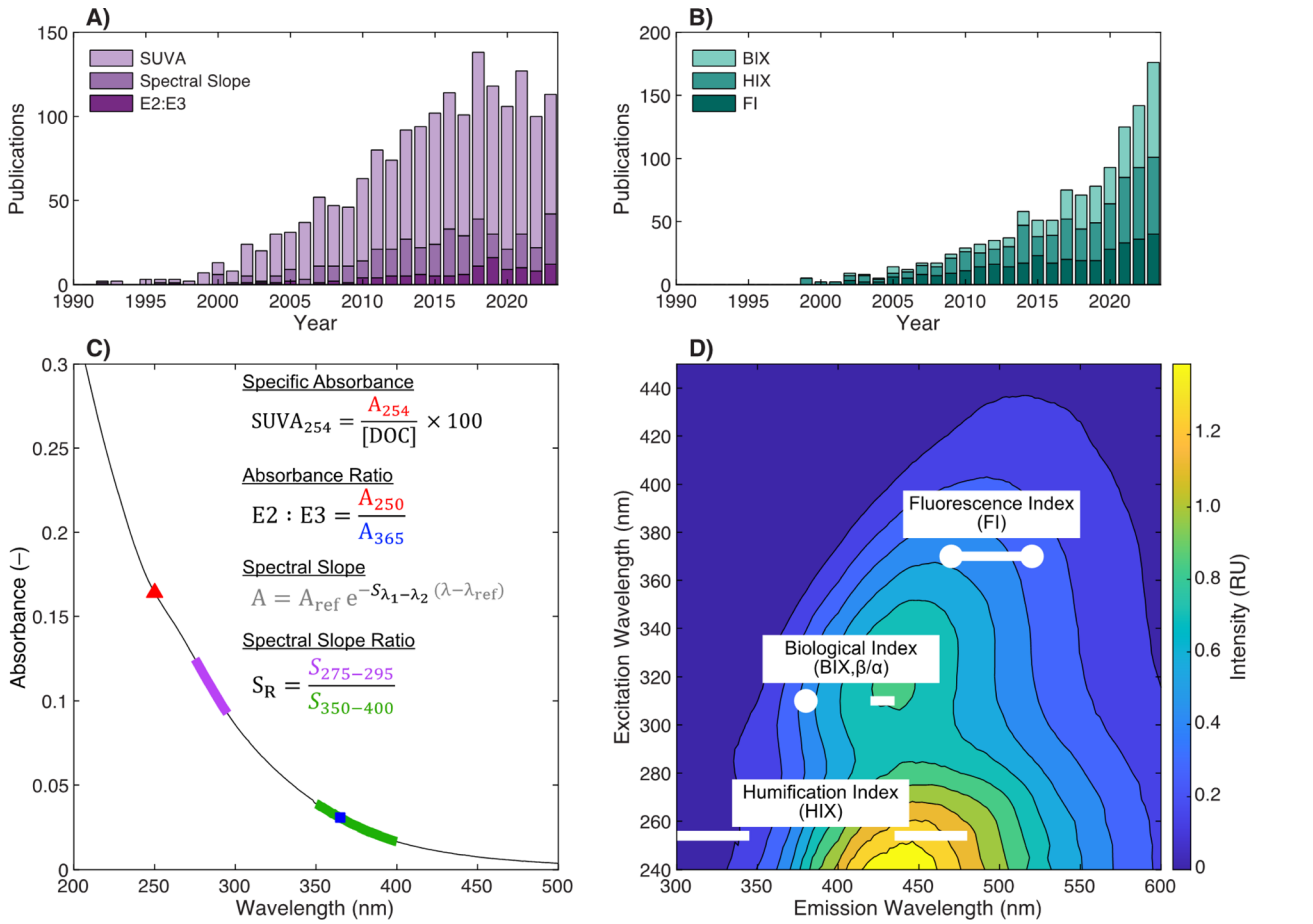

Fig. 1 Overview of optical surrogates for characterizing dissolved organic matter (DOM). Annual number of publications referencing (A) absorbance- or (B) fluorescence-based surrogates in the indexed abstract. Details about search terms are in ESI Text 1.† Depictions of (C) absorbance- and (D) fluorescence-based surrogates.

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observed in one context are applied to explain results in another. Within DOM studies, changes in optical surrogates have been leveraged to explain qualitative changes in physicochemical characteristics, such as its photochemical reactivity[26] or selective removal by coagulation.[27] A quantitative example is applying regression models developed in other studies to report quantitative values for aromaticity and molecular weight in a new data set using only absorbance measurements.[28] More broadly, interpretations derived from DOM samples in natural systems have been applied to different contexts without independent verification. Examples include treatment of highly specific feedstocks or anthropogenic waste streams[29,30] and production of microplastic-derived organic matter.[31,32] The lack of independent measures of DOM molecular weight and aromaticity in diverse contexts creates considerable uncertainty about interpretations derived from optical measurements, limiting progress in the field.

This critical review examines the foundation for optical surrogates commonly used to assess DOM molecular weight and aromaticity. For each surrogate, we review and scrutinize (1) the earliest known studies defining the surrogate and subsequent variations in definition that may be points of ambiguity in current literature, (2) the earliest known studies relating optical surrogates to aromaticity and/or molecular weight, (3) the experimental context for original studies (*e.g.*, soil *vs.* water) that may constrain current interpretations, and (4) the continued inquiry into direct lines of independent evidence (*e.g.*, NMR, membrane fractionation, SEC, *etc.*) for each surrogate. Although we focus the scope of this review on aquatic environments, many surrogates originated from soil science. We expect that this information will be useful to scientists and engineers studying DOM in aquatic systems and may serve as a framework for other environments, such as atmospheric aerosols.[33,34]

With respect to reviewing more recent, continued inquiry into optical surrogates, we focus on studies that (1) included a diverse range of source materials, (2) contrasted allochthonous and autochthonous endmembers, and (3) chemically characterized samples by multiple methods. Five papers are a consistent thread throughout the article due to data availability and breadth of organic matter sources.

First, Kellerman *et al.* (2018)[35] presents optical surrogates paired with Fourier-transform ion cyclotron resonance mass spectrometry (FT-ICR MS) data from 37 isolates collected from diverse aquatic environments, representing arguably one of the strongest available datasets for this inquiry. Their samples were collected from aquatic systems and isolated using either reverse osmosis (RO) or XAD resin to produce natural organic matter (NOM), hydrophobic organic acid (HPOA), and fulvic acid (FA) fractions. Compared to the scope of this review article, all the optical surrogates discussed were published in the original paper with one exception. The original paper did not publish peak intensities or full excitation-emission matrices (EEMs). A fluorescence intensity was estimated by reconstructing intensities from decomposed parallel factor analysis (PARAFAC) components (ESI Text 4†). The measure of aromatic carbon in this paper was the relative abundance of formulae classified as condensed or polycyclic aromatic formulae as described in Section 2.3.2.

Two papers by Maizel and Remucal also characterized aromatic carbon using FT-ICR MS. Maizel and Remucal (2017)[36] compared two endmember isolates, Suwannee River fulvic acid (SRFA) with Pony Lake fulvic acid (PLFA), and provided insight into molecular weight trends using ultrafiltration (UF) size fractionation. Another paper, Maizel and Remucal (2017),[37] collected samples from seven different lakes in northern Wisconsin of diverse trophic status. Both papers interpreted FT-ICR MS data by calculating the double bond equivalents (DBE).[38]

The fourth paper is McKay *et al.* (2018)[39] presenting optical surrogates from both aquatic and soil isolates, predominantly from the International Humic Substances Society (IHSS). These samples were paired with aromaticity data (^{13}C NMR) from other primary sources.[40–43] The last study, Mostafa *et al.* (2014),[38] used UF to contrast two endmember samples: Suwannee River natural organic matter (SRNOM) and a secondary treated, wastewater effluent (EfOM).

Although measures of aromatic carbon derived from ^{13}C NMR and FT-ICR MS are not directly comparable, both have become widely used methods to examine relationships between optical surrogates and aromatic carbon across diverse sources. Readers are referred to original sources for more details about study-specific instrumentation and methods. Lastly, many studies use IHSS isolates, and these materials have been isolated in different batches, each with a unique reference number. Readers are referred to original sources to determine if isolates presented across studies originated from the same batch.

Across the figures which synthesize data from multiple studies, several conventions are applied. In figures where correlations are calculated for literature data, the Spearman rank correlation coefficient (ρ_S) and associated p value (p_S) are presented. This approach does not assume linearity between variables. If the original paper fit a non-linear model, these models are shown with the annotated equation (*e.g.*, Fig. 4b). Least-squares linear regressions are shown selectively to highlight trends within a dataset. Regressions are not shown if rank correlations were not statistically significant (*e.g.*, Fig. 4c, PLFA), or if generalized regressions would be suspect due to data clustering (*e.g.*, Fig. 4c, Lakes). To include data from three literature sources,[38,39,44] some surrogates were not calculated in the original study and were later calculated from spectra in the Korak and McKay (2024)[45] meta-analysis; the original study that generated the data is attributed in the text and figures. Lastly, the conventional "*et al.*" is intentionally not printed in figure annotations due to space constraints; reference numbers are noted in the captions.

This review has three main Sections (2–4) followed by conclusions. Section 2 (background) presents some of the fundamental principles of absorbance and fluorescence spectroscopy, because some surrogates directly stem from these equations. Sections 3 and 4 cover absorbance and fluorescence surrogates, respectively. Within these sections, subsections focus on individual optical surrogates, detailing their genesis

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and exploring continued inquiry. These sections could be read in any order according to reader interest.

2 Background

In this section, terminology follows recommendations by the International Union of Pure and Applied Chemistry (IUPAC)[46] unless otherwise specified. Generally, fundamental equations are presented for single compounds. DOM is a heterogeneous mixture, and the terms DOM and dissolved organic carbon (DOC) are often used interchangeably; distinctions are made when the context is specific to DOC concentration ([DOC]) measured on a carbon basis. Since not all DOM is optically active, compositional surrogates would be interpreted as "apparent" values for the mixture.

2.1 Absorbance

Absorbance is the process by which a molecule absorbs light energy (*i.e.*, photons). The energy required to promote an electron is determined by the energy difference between the ground and excited states, which is a function of the type of molecular orbital involved in the transition (n *vs.* π) and the presence of electron delocalization or conjugation.[47,48] In DOM, absorbance in the ultraviolet-visible wavelength range (200–700 nm) primarily promotes π bond electrons associated with aromatic chromophores.[49] The conjugation of the aromatic ring can be extended through the addition of electron withdrawing groups, like carbonyls, or electron donating groups, like hydroxy and alkoxy groups. Extended conjugation increases the absorbance maximum wavelength (lower energy transition) relative to benzene (Fig. 2a). Furthermore, extending the conjugation *via* fusion of two benzene rings, as in naphthalene, also results in lower energy transitions.[51] The promotion of electrons associated with double bonds at 254 nm is why higher carbon-normalized absorbance is associated with higher aromaticity (Fig. 2b).[50] In the visible wavelength range, chromophore identity is less clear but could originate from highly conjugated aromatics, charge-transfer interactions between aromatic moieties, or a combination of these two.[49,52,53]

Quantitatively, absorbance is defined as the ratio of incident (P_λ^0) to radiant ($P_\lambda(\ell)$) spectral power at a specific wavelength (λ) (eqn (1)). The Bouguer–Beer–Lambert law relates absorbance to concentration (c) and pathlength (ℓ) using a proportionality constant (ε or κ). Depending on the logarithm convention, calculations using a base 10 logarithm pairs the terms absorbance ($A(\lambda)$) and molar decadic absorption coefficient (ε) following eqn (1). Calculations using the natural logarithm pair the terms Napierian absorbance ($A_e(\lambda)$) and molar Napierian absorption coefficient (κ) following eqn (2). Formal derivations are summarized elsewhere.[54,55] The molar absorption coefficients κ and ε are related through eqn (3).

$$A(\lambda) = \log_{10}\frac{P_\lambda^0}{P_\lambda(\ell)} = \varepsilon c\ell \quad (1)$$

$$A_e(\lambda) = \ln\frac{P_\lambda^0}{P_\lambda(\ell)} = \kappa c\ell \quad (2)$$

$$\kappa = 2.303\varepsilon \quad (3)$$

Differentiating these conventions is important for calculating DOM optical surrogates.[56] For example, decadic absorbance ($A(\lambda)$) is commonly used to calculate SUVA, whereas spectral slope calculations fit regressions to linear Napierian absorption coefficients ($\alpha = \kappa c$). Note, the DOM community commonly uses the acronym a for the linear Napierian absorption coefficient,[57] which is inconsistent with current IUPAC conventions.[46] Decadic absorption coefficients (ε) are commonly reported for freshwaters, whereas the marine community typically reports Napierian absorption coefficients.

DOM is generally assumed to follow the Bouguer–Beer–Lambert law across environmentally relevant concentrations.[58,59] For dilution series, non-zero intercepts indicate the contribution of non-chromophoric carbon. The effects of the cuvette or solvent are eliminated by pairing measurements with a reference cell through either a double beam configuration or

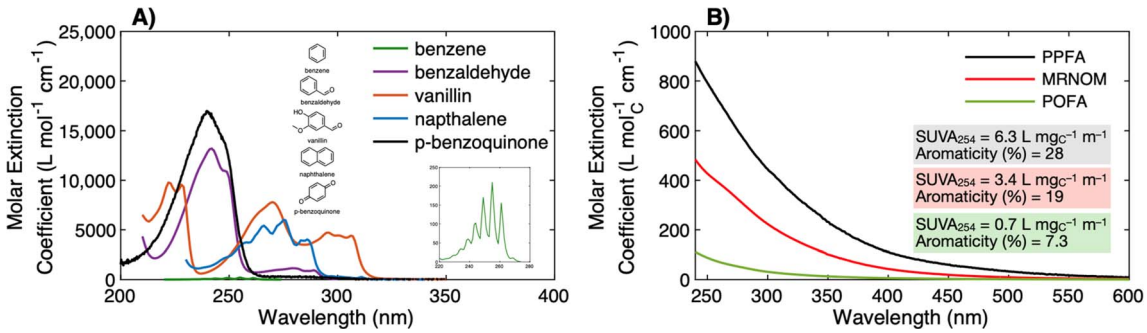

Fig. 2 (A) Molar extinction (ε, M^{-1} cm^{-1}) spectra for model aromatic chromophores demonstrating the impacts of electron withdrawing groups (benzaldehyde, –CHO), donating groups (vanillin; –OH, $-OCH_3$), and extended π conjugation on the energy of electronic transitions. The inset shows the spectrum of benzene. (B) Molar extinction spectra (ε, M_C^{-1} cm^{-1}) for DOM isolates from diverse sources with wide variations in aromaticity and specific ultraviolet absorbance at 254 nm ($SUVA_{254}$). Spectra are from McKay *et al.* (2018)[39] and paired with ^{13}C NMR data from other sources.[40,50] Samples include Pahokee Peat Fulvic Acid (PPFA), Mississippi River Natural Organic Matter (MRNOM), and Pacific Ocean Fulvic Acid (POFA).

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subtracting a blank spectrum. To measure absorbance at ultraviolet wavelengths, a cuvette material with high transmittance (*e.g.*, quartz) is necessary.

The Bouguer–Beer–Lambert law only describes absorbance – not attenuance due to light scattering by suspended particles. Although also measured on a spectrophotometer, attenuance is a function of both absorbance and light scattering. Isolating the absorbance phenomenon requires sample filtration. For absorbance measurements alone (not fluorescence), regulatory methods (*e.g.*, Standard Method 5310 or USEPA 415.3)[60,61] often define the dissolved fraction as <0.45 μm using a range of organic-based materials (*e.g.*, nylon, polyethersulfone). These filters have low potential to adsorb DOM or leach material that interferes at 254 nm after sufficient rinsing.[62] Across the research community, selection of filter material and nominal pore size (*e.g.*, 0.2–0.7 μm) is highly variable. This distinction between absorbance and attenuance is particularly important for online sensor data.

2.2 Fluorescence

Generated by light absorbance, singlet excited states can return to the ground state (called S_0) through several different pathways, one of which is fluorescence. Initially, excited molecules undergo relaxation to the lowest vibrational level of the first singlet excited state (called S_1),[63] and fluorescence occurs from this state when the excited molecule emits a photon with an energy ($\propto$ wavelength^{-1}) corresponding to the energy gap between S_1 and S_0. Fluorescence always occurs at emission wavelengths longer than the excitation/absorbance wavelength due to the energy lost during relaxation of the singlet excited state *via* vibrations and solvent reorientation, which is called the Stokes Shift.[54,64] In addition to fluorescence, relaxation from S_1 to S_0 can occur through non-radiative pathways such as internal conversion (IC) and intersystem crossing (ISC) to a triplet state (*e.g.*, T_1). From the triplet state, relaxation can occur through radiative (*i.e.*, phosphorescence) or non-radiative IC processes. The fluorescence quantum yield (Φ_f) is the ratio of the fluorescence rate constant (k_f) relative to the sum of the rate constants for radiative and nonradiative (k_{nr}) decay pathways (eqn (4)).[54,65]

$$\Phi_f = \frac{k_f}{k_f + k_{nr}} \tag{4}$$

An analogous quantum yield can be defined for relating ISC to a triplet state to other pathways.[46,65] For DOM, fluorescence quantum yields[39,66,67] are typically ~1%, and ISC quantum yields[68] are ~5%, suggesting that most photons absorbed by DOM are lost through non-fluorescence pathways. Although DOM fluorescence studies have taken advantage of both time-resolved and steady-state methods,[69–72] we focus here on steady-state methods used to calculate optical surrogates.

Benchtop spectrofluorometers use narrow slits to focus semi-monochromatic light on a narrow cross section of the cuvette. Emitted light is measured from a small interrogation zone perpendicular to the incident light (Fig. 3). Following

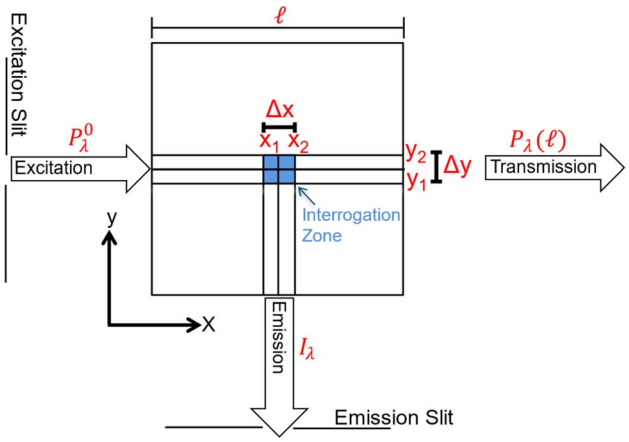

Fig. 3 Schematic of the cross-section (top view) of a square cuvette with annotated dimensions used in fundamental equation derivations.

derivations published elsewhere,[73,74] fluorescence intensity (I_f) is proportional to Φ_f and the power of light absorbed ($P_\lambda(x_1) - P_\lambda(x_2)$) across the interrogation zone (eqn (5)). Applying the Bouguer–Beer–Lambert law across the interrogation zone, I_λ is not proportional to concentration. Note, this source of non-linearity is different from inner filter effects (*vide infra*).

$$I_\lambda = \Phi_f(P_\lambda(x_1) - P_\lambda(x_2)) = \Phi_f P_\lambda(x_1)(1 - e^{-\kappa c \Delta x}) \tag{5}$$

However, in practice, a linear relationship between I_λ and c is commonly observed and is an underlying assumption for many intrinsic fluorescence surrogates (*e.g.*, ratio of fluorescence intensities or [DOC]-normalized fluorescence intensity).[75–77] A linear relationship between fluorescence intensity and concentration is supported mathematically by applying a power series expansion (eqn (6)) and assuming absorbance across the interrogation zone (Δx) is small. This approximation simplifies eqn (5) to (7). In practice, regressions between [DOC] and fluorescence intensity often have a non-zero intercept due to non-fluorescent DOM.[75,76]

$$e^{-\kappa c \Delta x} \approx 1 - \frac{\kappa c \Delta x}{1!} + \frac{(\kappa c \Delta x)^2}{2!} - \frac{(\kappa c \Delta x)^3}{3!} + \frac{(\kappa c \Delta x)^4}{4!} - \cdots \tag{6}$$

$$I_\lambda \approx \Phi_f P_\lambda(x_1)\left(1 - \left(1 - \frac{\kappa c \Delta x}{1!}\right)\right) \approx \Phi_f P_\lambda(x_1) \kappa c \Delta x \tag{7}$$

The practical application of eqn (7) requires fluorescence measurements that are free of instrumental or other optical artifacts. Like absorbance, filtration prevents light attenuation due to suspended particles. For fluorescence, glass fiber filters (GF/F) are commonly used with a nominal pore size 0.7 μm. Although GF/F filters have a lower potential to leach fluorescent material, they still need to be muffled and thoroughly rinsed to remove any binding material.[78] Notably, the common choice of GF/F filters for fluorescence conflicts with the 0.45 μm cut-off specified in USEPA method 415.3 for absorbance.[60] Within the DOM field, Murphy *et al.* (2010)[79] outlines the broadly accepted

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fluorescence correction procedures.[58] These corrections include instrument-specific correction factors, blank subtraction, scatter masking, intensity normalization, and inner filter corrections. Software packages that perform these corrections are available in MATLAB,[80,81] R,[82] and proprietary software (*e.g.*, Horiba's Aqualog). Instrument-specific correction factors are unique to each spectrofluorometer and are often provided by manufacturers to account for wavelength-specific efficiencies of light transmission.[83,84] Blank subtraction, like absorbance reference cells, isolates the sample fluorescence independent of the solvent and cuvette. Blank subtraction is partially effective to remove Rayleigh and Raman scattering, but most analyses excise and interpolate scatter.[80,81,85,86] Finally, signal normalization scales the intensities by either the integrated area of the Raman peak for deionized water (Raman Units; RU) or the emission from a model fluorophore like quinine sulfate (Quinine Sulfate Units; QSU).[87,88]

Inner filter corrections account for light absorbance to and from the interrogation zone (Fig. 3). Primary inner filtering is the loss of light between the incident cuvette edge and the interrogation zone. Secondary inner filtering is the loss of emitted light between the interrogation zone and the cuvette edge perpendicular to the incident light. Inner filter corrections apply broadly across fluorescence spectroscopy, and several studies have proposed or derived correction procedures over the past 60 years using absorbance-based approaches,[59,89–95] controlled dilution approaches,[59,96] and cell shift methods.[89,97] The latter two approaches are less common in the DOM community; readers are referred to the cited references for more details. The absorbance-based approach is the most common correction following eqn (8). The correction factor is a function of the sum of (decadic) absorbance values at the excitation ($A(\lambda_{ex})$) and emission wavelengths ($A(\lambda_{em})$), assuming the same pathlength for fluorescence and absorbance spectroscopy. If absorbance and fluorescence measurements use different pathlengths, absorbance must be normalized to the fluorescence pathlength before applying eqn (8). Kothawala *et al.* (2013) reported that eqn (8) performed sufficiently up to an absorbance sum of 1.5.[59]

$$I_{\lambda,\mathrm{corr}} = I_{\lambda} \times 10^{\frac{A(\lambda_{ex})+A(\lambda_{em})}{2}} \quad (8)$$

To illustrate the magnitude of corrections, consider a scenario where $A(\lambda_{ex})$ is 0.1 and $A(\lambda_{em})$ is 0.05 making the absorbance sum 0.15. The observed fluorescence intensity (I_{λ}) would be corrected by multiplying it by a factor of 1.19. If uncorrected, the observed fluorescence would be 19% too low due to inner filtering (Fig. S1†). Some studies cite a DOC concentration threshold to justify the need for inner filter corrections (or lack thereof). This approach is strongly discouraged, because inner filter effects are an absorbance-based phenomenon. A DOC concentration criterium would assume all DOM has the same absorbance per unit carbon.[59]

Although broadly used, the derivation of eqn (8) includes some simplifying assumptions that may not be appropriate in all cases. Similar to the linearization of eqn (5), inner filter corrections also use a power-series expansion (eqn (6)), which assumes the quantity $\kappa c \Delta x$ is small, to linearize an exponential term. The second assumption is that the interrogation zone is in the middle of cuvette (pathlength $= \ell/2$). The application of eqn (8) on benchtop spectrofluorometers is generally appropriate. Kubista *et al.* (1994) showed that quantifying the interrogation zone size is not important for most practical applications, and the same empirical correction equation could be used for bandpasses between 0.5 and 15 nm.[91] However, adaptation of benchtop methods to other applications, such as field sensors, may need to reevaluate these assumptions. Full derivations are provided elsewhere.[90,92,98]

Spectrofluorometers from many manufacturers have been used by the DOM research community, including Aminco-Bowman, PerkinElmer, Varian, and Horiba. Small but meaningful differences in fluorescence spectra have been documented and would be expected between different spectrofluorometers given differences in hardware.[79] Even within Horiba instruments, substantive instrument bias has been reported between the Aqualog and Fluoromax (*e.g.*, F3 and F4), which, in part, results from differences between excitation gratings in the Fluoromax-4 (plane ruled), that passes more stray light, compared to the Aqualog (concave holographic).[99] Past research has shown that apparent quantum yields at excitation wavelengths less than 350 nm are systematically larger on the Fluoromax-4 compared to the Aqualog.[39] Unfortunately, the impact of these instrument biases on fluorescence-derived optical surrogates are not well-constrained in the DOM research community. In contrast, fluorescence-based surrogates that rely on intensity ratios at the same excitation wavelength may be less impacted.

2.3 Estimation of DOM molecular weight and aromaticity

2.3.1 Molecular weight. There are several methods to characterize DOM molecular weight, including ultrafiltration (UF),[100] size exclusion chromatography (SEC, in the absence[101–105] or presence of high-pressure pumps[106]), field-flow fractionation (FFF),[107] and diffusion ordered spectroscopy.[108] Other methods include small-angle X-ray scattering,[109] vapor pressure osmometry,[110] viscosimetric analysis,[111] and cryoscopic techniques.[112] The most common separation techniques paired with optical surrogates are UF fractionation, gel permeation chromatography (GPC), and SEC. In these techniques, separation is based on molecular size, but data is usually communicated as DOM molecular weight (*e.g.*, Daltons (Da)) by calibrating elution volume or membrane cutoffs using polymer standards (*e.g.*, polyethylene glycol, polystyrene sulfonate). The assumption is that the molecular size of DOM is well-represented by the polymer calibration, which has well-documented limitations.[113,114] With this understanding, it is important to recognize that size-based separation techniques are not direct measurements of molecular weight, even though data are presented in units of molecular weight herein (*e.g.*, kDa).

UF fractionated samples yield categorical molecular weight classifications (*e.g.*, <1 kDa, 1–3 kDa, >3 kDa). Each filter has a nominal molecular weight cutoff, but two filters with the same

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nominal cut-off may perform differently due to membrane surface chemistry, DOM concentration, and aqueous ionic composition.[115,116] GPC, without the use of a high-pressure pump, was commonly used in many early papers by packing columns with Sephadex resins of different pore sizes.[103,104] To decrease sample volumes and increase resolution, high-pressure SEC uses macroporous resins with online detectors.[106,117,118] In both GPC and SEC, smaller molecules are retained in the column and elute after longer times compared to larger molecules. The detector signal can be numerically integrated across the SEC chromatogram to calculate either a number-average (M_n) and weight-average (M_w) molecular weight. Polydispersity (M_w/M_n) is generally greater than unity for DOM, implying that M_w is greater than M_n.

2.3.2 Aromaticity. Characterization of aromatic carbon in humic substances has transitioned from wet chemistry methods, like permanganate oxidation,[119,120] to instrumental techniques like Fourier transform infrared spectroscopy (FTIR),[121–123] pyrolysis gas chromatography-mass spectrometry,[124,125] ^{13}C or proton (^{1}H) NMR,[126–130] and most recently FT-ICR MS.[129,131–137] Studies comparing DOM aromaticity to optical surrogates have mainly used quantitative ^{13}C NMR on solid isolate materials.[138] Although it achieves higher sensitivity than liquid-state NMR, existing solid-state ^{13}C NMR data are predominately from HPOA and fulvic acid isolates and, importantly, do not represent hydrophilic or transphilic fractions. Although there are several methods for ^{13}C NMR,[1,40,42,50,139] integrating the 110–160 ppm chemical shift region relative to the total area informs the relative abundance of aromatic carbon (aromaticity (%)).

For FT-ICR MS, ions are detected, and formulae are assigned based on the accurate mass using automated programming algorithms and a set of rules.[133] FT-ICR MS has both instrumental limitations, where ion detection may be biased to low molecular weights,[135] and data analysis limitations, where unambiguous formula assignment is often limited to mass-to-charge ratios (m/z) between 150–1000.[132,133,140] Using the assigned formulae, metrics can be calculated, such as modified aromaticity index (AI_{mod})[132] and double bond equivalents (DBE).[141] These metrics are used to group assigned molecular formulae into chemical characterization categories that suggest aromatic or condensed aromatic moieties in the DOM samples. Mass spectral peak intensity-weighted averages, based on all assigned formulae in a sample, are also reported. For example, Kellerman *et al.* (2018)[35] used AI_{mod}[132] based on FT-ICR MS data to calculate relative abundance (%) of two classes: condensed or polycyclic aromatic formulae ($AI_{mod} > 0.66$) and polyphenolic formulae ($0.66 \geq AI_{mod} > 0.5$) using the formula bounds of $C_{1-45}H_{1-92}N_{0-4}O_{1-25}S_{0-2}$.

Estimates of aromaticity are not directly comparable when derived using different instrumental (*e.g.*, NMR or FT-ICR MS) or calculation methods (*e.g.*, AI_{mod} boundary conditions). NMR probes all ^{13}C carbons, albeit at different shifts depending on chemical environment. For example, considering two structural isomers, cyclohexane and hexene (each having one degree of unsaturation), hexene would have a ^{13}C resonance at about 120 ppm (from the sp^2 hybridized carbon), whereas cyclohexane would have only aliphatic carbon resonances <60 ppm.[141] In contrast to NMR, FT-ICR MS only detects molecules that can be ionized (in either positive or negative ion mode) and characterized as having aromatic nature, subject to meeting signal-to-noise thresholds.[132,142] Past work suggests that UV chromophoric DOM is poorly ionized by negative mode electrospray ionization[143] and is therefore poorly represented in the majority of studies characterizing the chemical nature of DOM with ultra-high resolution mass spectrometry techniques. In contrast, a study by Laszakovits *et al.* (2020) demonstrated that a higher percent of assigned formulae were characterized as aromatic and condensed aromatic using laser desorption ionization (in both positive and negative mode) compared to electrospray ionization.[144] For future inquiries relating optical surrogates to composition, there is an opportunity to explore multiple ionization techniques.

In this review, FT-ICR MS is considered as a technique to assess the abundance of aromatic carbon, despite its limitations, due to its growing popularity in DOM research.[145] FT-ICR MS data has also been analyzed to calculate molecular weight distributions of DOM[146,147] but will not be used as a comparison measure in this review. Due to incomplete DOM recovery by SPE[148–151] and limitations of the analytical mass range[132,133,140] and ionization efficiency,[131,144,152] values derived from FT-ICR MS data are far from comparable to other methods that characterize molecular weight. This review focuses on UF and SEC to characterize molecular weight.

Overall, all characterization methods for molecular weight and aromatic carbon are subject to sampling, analytical, and methodological constraints. Readers are referred to primary sources for details about sample preparation, instrument biases, method background and limitations, and other challenges with DOM chemical characterization.

3 Absorbance surrogates

Absorbance-based optical surrogates are listed in Table 1 along with the equations and primary sources. The following subsections explore each surrogate.

3.1 Specific absorbance

3.1.1 Definition and genesis. Related to κ and ε in the Bouguer–Beer–Lambert Law (eqn (1)), specific absorbance normalizes absorbance to the DOM concentration and pathlength. With different approaches for measuring both absorbance (decadic *vs.* Napierian) and concentration (DOM mass or carbon mass), specific absorbance can be reported using a variety of different conventions. As advocated for in Chin *et al.* (1994),[154] a carbon basis (either $mg_C\ L^{-1}$ or M_C) is preferred, because a DOM mass basis is biased by residual ash. When normalizing to carbon concentration, specific absorbance is typically reported on either a mass basis ($L\ mg_C^{-1}\ m^{-1}$) or a molar basis as the molar absorption coefficient ε, ($M_C^{-1}\ cm^{-1}$). Lastly, specific absorbance is wavelength dependent, reported at either selected wavelengths (*e.g.*, 254, 280 nm *etc.*) or as a normalized spectrum (Fig. 2b). When wavelengths

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Table 1 Definitions of optical surrogates based on absorbance spectroscopy

Metric	Examples	Calculation	Units	Notes	References
Specific absorbance	$SUVA_{254}$ $SUVA_{280}$ $SUVA_{300}$	$SUVA_\lambda = \frac{A_\lambda}{[DOC]} \times 100$	L mg_C^{-1} m^{-1}	λ: wavelength of incident light A_λ: decadic absorbance at λ normalized to a pathlength of 1 cm [DOC]: dissolved organic carbon concentration	Traina *et al.* (1990)[153] Chin *et al.* (1994)[154] Peuravuori and Pihlaja (1997)[155] Weishaar *et al.* (2003)[50]
Absorbance ratio	E2 : E3	$E2:E3 = \frac{A_{250nm}}{A_{365nm}}$	—	A_{250nm}: decadic absorbance at 250 nm A_{365nm}: decadic absorbance at 365 nm	De Haan (1972)[156] Peuravuori and Pihlaja (1997)[155]
	E4 : E6	$E4:E6 = \frac{A_{465nm}}{A_{665nm}}$	—	A_{465nm}: decadic absorbance at 465 nm A_{665nm}: decadic absorbance at 665 nm	Chen *et al.* (1977)[157]
Spectral slope (non-linear)	$S_{300-600}$ $S_{300-650}$ $S_{300-700}$	$A_{e,\lambda} = A_{e,\lambda_{ref}} e^{-S_{\lambda_1-\lambda_2}(\lambda-\lambda_{ref})}$	nm^{-1}	$S_{\lambda_1-\lambda_2}$: spectral slope between λ_1 and λ_2 λ_{ref}: reference wavelength, typically 350 nm	Stedmon *et al.* (2000)[158] Kowalczuk *et al.* (2005)[159] Twardowski *et al.* (2004)[160] Helms *et al.* (2008)[57]
Spectral slope (linear)	$S_{275-295}$ $S_{350-400}$	$\log_{10}\left(\frac{2.303 \cdot A_{e,\lambda}}{\frac{0.01\ m}{\ell\ /\ cm}}\right) = -S_{\lambda_1-\lambda_2} \cdot \lambda + b$	nm^{-1}	b is the intercept (typically not reported)	Helms *et al.* (2008, 2013)[57,161] Fichot and Benner (2012)[162]
Spectral slope ratio	S_R	$S_R = \frac{S_{275-295}}{S_{350-400}}$	—		Helms *et al.* (2008)[57]

are in the ultraviolet range, the surrogate is termed specific ultraviolet absorbance (SUVA).[15] The convention in Table 1 denotes the wavelength using the subscript (*e.g.*, $SUVA_{254}$, $SUVA_{280}$).

Since specific absorbance is an application of the Bouguer–Beer–Lambert law, the genesis within DOM research dates to the earliest inquiries. Juday and Birge (1933)[3] explored specific absorbance indirectly by plotting color (in platinum-cobalt units) *versus* DOC concentration, showing a correlation for >500 lakes in Wisconsin (USA). Later, James (1941)[58] diluted lake samples, calculated a "molecular absorption coefficient", and showed agreement with the Bouguer–Beer–Lambert law for up to a 20-fold dilution across 5 visible wavelengths (407.9 to 700 nm). Over the following decades, most studies commonly reported color per unit carbon for individual samples or the slope of a linear regression between color and DOC concentration for multiple samples.[2,105,122,163–166] For example, Packham (1964) reported that the specific absorbance for humic acid (HA) isolates was greater than paired the fulvic acids for 7 samples at 300 and 450 nm.[165]

Specific absorbance has been reported at wavelengths spanning the ultraviolet[50,154,167,168] and visible regions,[121,169–171] with early studies mostly reporting visible wavelengths and more recent studies focusing on ultraviolet wavelengths. Ghassemi and Christman (1968)[169] fractionated samples with Sephadex and noted higher absorbance per unit carbon at 350 nm in the larger molecular weight fractions. Despite a 1953 study which advocated for monitoring at 275 nm,[172] 254 nm gained popularity in the 1960s and 1970s due to this wavelength aligning with an emission maxima of low-pressure mercury lamps.[173–176] Focus shifted from using absorbance as a surrogate for DOC concentration to using specific absorbance as an intrinsic surrogate of DOM composition. Studies demonstrated relationships with coagulation efficiency[14] and disinfection byproduct formation using both the color-to-[DOC] ratio in 1983 (ref. 177) and then $SUVA_{254}$ in 1985.[15] Although independent measures of molecular weight or aromaticity were not the focus of these early studies, the context for measuring UV absorbance was consistently framed as probing the $\pi \rightarrow \pi^*$ transitions occurring in O- and N-substituted aromatic compounds such as phenols.[15,173,174,178,179] However, some early studies also correlated specific absorbance to molecular weight[121,180] or qualitatively described the UV absorbance as dependent on both chemical characteristics[181] (and not mutually exclusive).

3.1.2 Continued inquiry into aromaticity and molecular weight

3.1.2.1 Aromaticity. Like many optical surrogates, early explorations of aromaticity are rooted in soil humic substances. Traina *et al.* (1990)[153] correlated ^{13}C aromaticity with $SUVA_{272}$ for humic substances extracted from soil, showing good agreement with data for humic substances extracted from marine sediments and terrestrial soils from Gauthier *et al.* (1987).[179] For the combined dataset, the linear regression was strong ($n = 12$, $r = 0.937$),[153] but as a note of caution, absorbance was normalized to the material mass, not carbon mass, which may underestimate specific absorbance due to residual water and ash.

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Transitioning to aquatic DOM, Chin *et al.* (1994)[154] studied the relationship between $SUVA_{280}$ and both ^{13}C aromaticity and molecular weight for aquatic fulvic acids, presenting a strong positive correlation between $SUVA_{280}$ and ^{13}C aromaticity spanning autochthonous and allochthonous sources (Fig. 4a). Similarly, $SUVA_{300}$ and ^{13}C aromaticity were positively correlated in McKnight *et al.* (1997)[182] for fulvic acids from diverse origins, including the Suwannee River, two Antarctic lakes (Pony Lake and Lake Fryxell), and an alpine watershed in Colorado (USA). Notably, the coefficient of determination (R^2) for the regression with ^{13}C aromaticity was stronger for $SUVA_{300}$ ($R^2 = 0.76$) than $SUVA_{450}$ ($R^2 = 0.43$). Croue *et al.* (2000)[125] fractionated four waters using RO and XAD resins, also affirming a positive correlation between $SUVA_{254}$ and ^{13}C aromaticity ($R^2 = 0.72$, $n = 27$). Currently, the highest cited paper relating $SUVA_{254}$ to ^{13}C aromaticity is Weishaar *et al.* (2003),[50] which included fulvic acid-dominated isolates from diverse aquatic origins ranging from the Pacific Ocean to the Florida Everglades (^{13}C aromaticity (%) $= 6.52 \times SUVA_{254}$, $n = 13$, $R^2 = 0.97$) (Fig. 4a). Pairing the specific absorbance data from McKay *et al.* (2018)[39] with ^{13}C data from other studies,[40–43] there is continuity in the correlation for the sample set expanded to include an aquatic humic acid and three soil isolates, which is not be the case for some other surrogates (*vide infra*).

More recently, molecular formulae assigned using FT-ICR MS data were analyzed to characterize differences in aromatic carbon. For DOM samples from the Florida Everglades, molecular formulae with lower hydrogen-to-carbon (H/C) ratios (higher DBE) correlated positively with $SUVA_{254}$.[183] In Maizel and Remucal (2017),[36] increases in DBE were associated with increases in $SUVA_{254}$ between two clusters of lakes in Wisconsin of diverse trophic status (Fig. 4c). Encompassing the broadest set of NOM, HPOA, and fulvic acid isolates ($SUVA_{254}$ 0.6–4.9 L mg_C^{-1} m^{-1}), Kellerman *et al.* (2018) reported an exponential relationship between $SUVA_{254}$ and the relative abundance of condensed aromatic and polyphenolic formulae (Fig. 4b).[35]

3.1.2.2 Molecular weight. The contemporary understanding for DOM is that SUVA increases as molecular weight increases, although some early studies proposed the opposite.[101,184] Utilizing cultures from soil-derived microorganisms (actinomycetes), Ewald *et al.* (1988)[180] fractionated DOM using XAD-2 resin[185] and UF (500 and 100 000 Da) followed by pH adjustment to 2 and 13. For the UF fractions, molar absorption coefficients at 370 nm (ε_{370}) increased proportionally to the log of the molecular weight with higher ε_{370} values at pH 13 than pH 2.[180] Contrasting a Nordic lake with rivers in southeastern USA, Alberts and colleagues[186,187] fractionated samples by UF, characterized the fractions by SEC (DOC and UV_{254} detectors), and confirmed increasing $SUVA_{254}$ with increasing molecular weight. More recently, Mostafa *et al.* (2014)[38] and Maizel and Remucal (2017)[36] fractionated endmember aquatic DOM samples with UF, also affirming increased $SUVA_{254}$ with

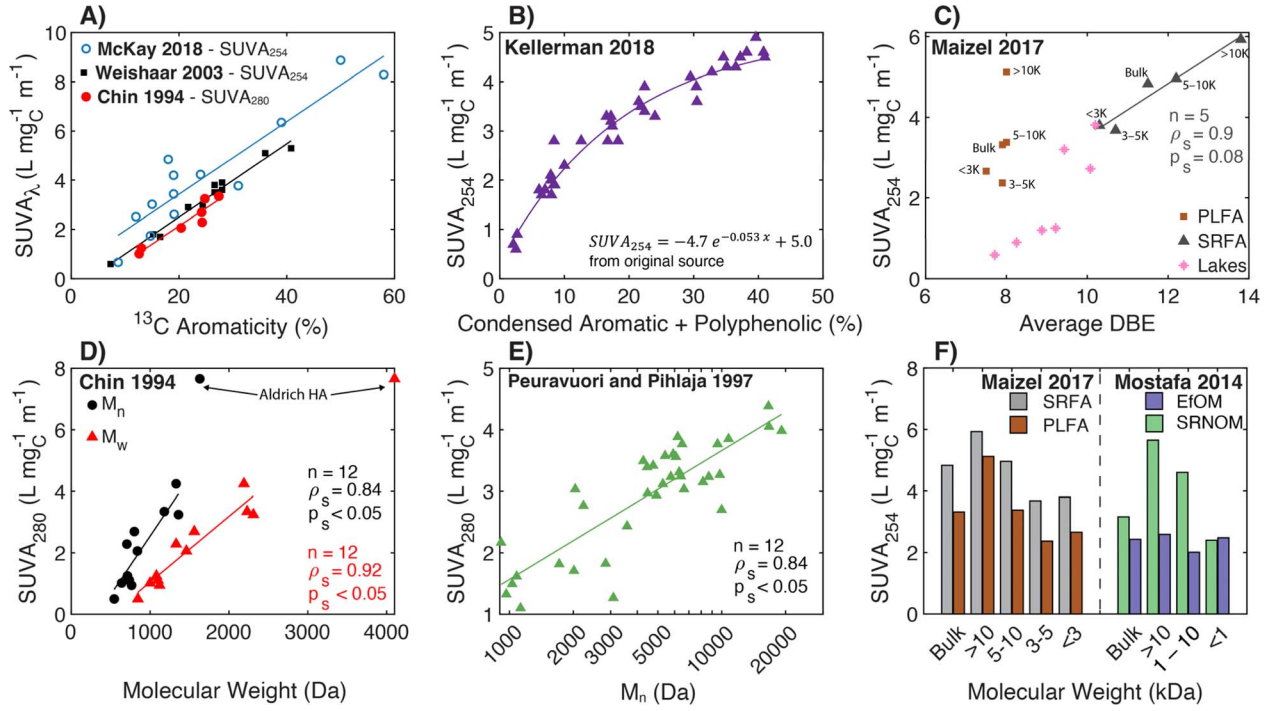

Fig. 4 Relationship between specific ultraviolet absorbance (SUVA) and either molecular weight or different measures of aromatic carbon. (A) $SUVA_{254}$ and $SUVA_{280}$ *versus* aromaticity (%) per sample calculated from ^{13}C NMR data.[39–43,50,154] (B) $SUVA_{254}$ *versus* the relative abundance of condensed aromatic and polyphenolic formulae (%) per sample calculated based on AI_{mod}[132] from FT-ICR MS data.[35] (C) $SUVA_{254}$ *versus* average double bond equivalents (DBE) calculated based on FT-ICR MS data for size-fractionated fulvic acid isolates (PLFA and SRFA)[36] and seven different lakes in northern Wisconsin of diverse trophic status.[37] (D) $SUVA_{280}$ *versus* number-average (M_n) and weight-average (M_w) molecular weight determined by SEC with UV detection.[154] (E) Number-average molecular weight *versus* $SUVA_{280}$ for SEC with UV detection.[155] (F) $SUVA_{254}$ *versus* UF molecular weight fraction for three isolates (SRFA, PLFA, and SRNOM) and effluent organic matter (EfOM).[36,38]

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increasing nominal filter cutoffs for most isolates (Fig. 4f). Notably, UF fractionated EfOM did not show a dependence of $SUVA_{254}$ on molecular weight.

Use of specific absorbance to interpret aromaticity and molecular weight as independent characteristics is limited, because they correlate with one another. Both Chin *et al.* (1994)[154] and Peuravuori and Pihlaja (1997)[155] reported a positive correlation between $SUVA_{280}$ and both ^{13}C aromaticity and number-average molecular weight (M_n) (Fig. 4a, d and e). Maizel and Remucal (2017)[36] fractionated SRFA with UF and used FT-ICR MS to show average DBE increased with increasing molecular weight, both correlating with $SUVA_{254}$ (Fig. 4C). However, size-fractioned PLFA did not exhibit the same internally consistent relationship between DBE and $SUVA_{254}$. Despite these correlated relationships, one cannot assume that high molecular weight DOM will also have high $SUVA_{254}$; large biopolymers, like polysaccharides, that are abundant in EfOM often have low molar absorption coefficients.[117,118]

Taken as a whole, both historical and more recent literature support that SUVA generally increases with both molecular weight and aromaticity across samples from diverse geographic sources, but microbial endmembers may deviate from this generalization. In addition, asserting that SUVA is unique to aromaticity and not molecular weight is inconsistent with the frequent correlation of these characteristics reported in the primary literature.

3.2 E2:E3

3.2.1 Definition and genesis. E2:E3 is defined as the ratio of absorbance at 250 to 365 nm (Table 1). Although variations have been reported using absorbance at 254 nm in the numerator, there are noticeable differences (>5%) between conventions (A_{254}/A_{365} *vs.* A_{250}/A_{365}).[45] The use of A_{254} is not consistent with the original definition of E2:E3, and A_{250}/A_{365} is recommended.

The genesis of E2:E3 appears in De Haan (1972),[156] although several earlier papers used absorbance ratios at different excitation wavelengths (see the Spectral Slope Section 3.4). De Haan (1972) concentrated samples from "shallow peaty lakes" in the Netherlands using pH adjustment (pH 7.0), decalcification (ion exchange), and freeze drying. The freeze-dried samples were reconstituted and fractionated by Sephadex size exclusion resin (G-25) measuring the absorbance at 250 nm semi-continuously. Sephadex G-25 produced three chromatographically resolved fractions (I, II, and III). Fig. 5d illustrates the correlation between the whole-water E2:E3 (pH 7.0) and the percentage of low molecular weight material (fraction III) from the freeze-dried, decalcified sample. Compared to the lower molecular weight fraction, correlations were weaker between E2:E3 and the percentage of either medium (II) or high (I) molecular weight DOM.[46]

Two later papers by De Haan and co-workers[77,189] are often cited to support the inference of DOM molecular weight from

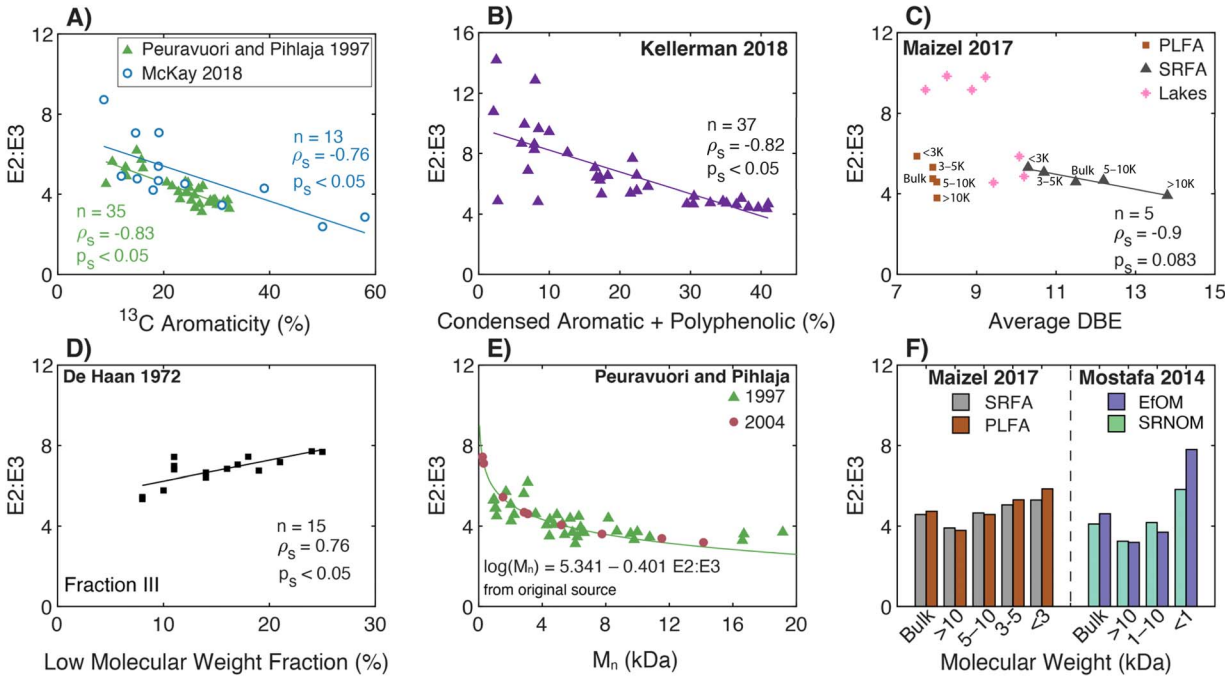

Fig. 5 Relationship between E2:E3 (A_{250}/A_{365}) and either molecular weight or different measures of aromatic carbon. (A) E2:E3 *versus* aromaticity (%) per sample calculated from ^{13}C NMR data.[39–43,155] (B) E2:E3 *versus* the relative abundance of condensed aromatic and polyphenolic formulae (%) per sample calculated based on AI_{mod}[132] from FT-ICR MS data.[35] (C) E2:E3 *versus* average double bond equivalents (DBE) calculated based on FT-ICR MS data for UF fractionated isolates (PLFA and SRFA)[36] and seven different lakes in northern Wisconsin of diverse trophic status.[37] (D) E2:E3 of the whole-water *versus* the percentage of low molecular weight DOM (longest retention time) for a lake.[156]. (E) E2:E3 *versus* number-average molecular weight (M_n) for DOM from Nordic surface waters (humic substances and ultrafiltered samples).[155,188] (F) E2:E3 *versus* nominal molecular weight fractions for endmembers following UF fractionation.[36,38]

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E2:E3 measurements. First, De Haan *et al.* (1983)[189] measured E2:E3 for a ditch water sample derived from peat soil drainage at pH values between 2.0 and 10.5. Compared to pH 2, the relative change in absorbance ($[A_{pHX} - A_{pH2}]/A_{pH2} \times 100$) increased by over 30% from pH 2 to 10.5 at 365 nm, whereas the increase was only about 5% at 250 nm. Therefore, E2:E3 decreased with increasing pH. Furthermore, the ditch water DOM was fractionated using dialysis and UF to assess the molecular weight distribution of chromophoric DOM as measured by absorbance spectroscopy. Increasing the pH decreased the relative fraction of chromophoric DOM that permeated through a 25 Å cutoff dialysis bag. The relative fraction was positively correlated to the peak area of the highest molecular weight fraction separated by Sephadex G-25, corroborating the dialysis measurements. In addition, for ultrafilters of a given pore size (between 50 and 150 Å), a larger fraction of chromophoric DOM was retained as pH increased from 3 to 7, also consistent with the dialysis results. Based on these observations and prior work,[156] De Haan *et al.* (1983)[189] concluded that the, "decrease of E250/E365 with increasing pH indicated increasing molecular weight and size of the fulvic acid with increasing pH" (page 71, paragraph two). The second study, De Haan and De Boer (1987),[77] reported that E2:E3 increased with smaller membrane cutoffs for 38 UF fractionations from a single lake (Fig. 5b).

To summarize, the genesis of E2:E3 as a surrogate for DOM molecular weight is based on individual DOM samples subjected to molecular weight fractionation or pH titration. This surrogate was not developed to compare DOM from multiple geographic sites with independent molecular weight distributions. In practice, E2:E3 is often interpreted with specificity for molecular weight. However, another early study by De Haan (1983) complicates this interpretation, because curie point pyrolysis/mass spectrometry showed that increased molecular weight coincides with increased aromatic compounds (*i.e.*, phenols) and decreased E2:E3 ratios.[190]

3.2.2 Continued inquiry into aromaticity and molecular weight

3.2.2.1 Molecular weight and aromaticity. Since the original work of De Haan and colleagues, Peuravuori and Pihlaja (1997, 2004)[155,188] are cornerstones for generalized interpretations of E2:E3 for DOM molecular weight and aromaticity. Peuravuori and Pihlaja (1997)[155] evaluated the relationships between molecular weight (by SEC) and aromaticity (by ^{13}C NMR) with both $SUVA_{280}$ (L mol_C^{-1} cm^{-1}) and E2:E3 for lake samples ($n = 2$), river water ($n = 1$), and many ($n > 60$) isolates derived from these three whole-water samples (see Section 3.1 for $SUVA_{280}$ discussion). Humic acid, fulvic acid, and hydrophobic neutral fractions were isolated using an XAD-8 procedure.[191] In addition, each whole-water was ultrafiltered using membranes with various nominal molecular weight cutoffs (NMW, in Da) into three fractions (I, NMW>100 000; II, NMW 10,000–100,000; III, NMW 1000–10 000). Each UF fraction was further subjected to three further treatments: freeze-drying, cation exchange followed by freeze-drying again, followed by XAD-8 isolation of humic- and fulvic acid fractions from the UF fractions. Finally, selected samples were ultrafiltered to obtain an NMW <1000 fraction (IV) followed by XAD-8 isolation.

Peuravuori and Pihlaja (1997)[155] presented single parameter regressions using E2:E3 as a predictor of either aromaticity using XAD isolates (eqn (9), linear) or M_n using UF fractionated samples (eqn (10), semilog).

$$\text{Aromaticity (\%)} = 52.059 - 6.780 \times \text{E2:E3},\ n = 39,\ R^2 = 0.78 \quad (9)$$

$$\log M_n = 5.341 - 0.401 \times \text{E2:E3},\ n = 40,\ R^2 = 0.70. \quad (10)$$

The regressions used a subset ($n = 39$ or 40) of the total sample set ($n > 60$), because the authors excluded UF samples that were further treated by freeze drying alone or freeze drying after calcium removal by ion exchange. Thus, the sample context at the core of these regressions is limited to hydrophobic acids and hydrophobic neutrals.

Peuravuori and Pihlaja (2004)[188] used preparative scale SEC to separate DOM into eight molecular weight fractions. The sample context was Lake Savojärvi (southwestern Finland), a marsh containing ~20 mg_C L^{-1}. Fig. 5e shows that there is an exponential relationship between E2:E3 and M_n for the 2004 dataset, agreeing with the general trend from the 1997 study.[155] Unfortunately, aromaticity was not reported in this later study. In both studies, E2:E3 values have a lower limit of ~3, even for the highest molecular weight fractions.[155,188]

More recently, two studies used UF to fractionate endmember samples (*i.e.*, PLFA and SRFA).[36,39] For both microbial and terrestrial endmembers, there were internally consistent trends for each sample; E2:E3 increased as molecular weight decreased (Fig. 5f).

3.2.2.2 Aromaticity. Using FT-ICR MS, Kellerman *et al.* (2018) applied nonmetric multidimensional scaling to show that DOM molecular formulae associated with aliphatic and less aromatic chemical species typically had higher E2:E3 values.[35] Fig. 5b shows an inverse correlation between E2:E3 and the relative abundance of condensed aromatic and polyphenolic formulae ($\rho_s = -0.83$, $p_s < 0.05$). Although statistically significant, the data are more scattered as aromaticity decreased. For example, isolates with condensed aromatic and polyphenolic formulae > 30% consistently have low E2:E3 values (~5), while samples with the lowest formula abundance (<10%) have E2:E3 values that span the dataset range.[35]

Also using FT-ICR MS, Maizel and Remucal (2017)[36] compared DOM molecular formulae to E2:E3 for lakes of diverse trophic status. Of formulae assigned across all samples, 97% of those formulae that correlated positively with E2:E3 had a greater relative intensity in oligotrophic lakes. These formulae were also more oxidized and aliphatic than the 452 formulae that correlated negatively with E2:E3,[36] consistent with Kellerman *et al.* (2018).[35] Fig. 5c shows that size-fractioned SRFA exhibited an inverse correlation between E2:E3 and average DBE with borderline statistical significance. Lake samples clustered into 2 groups where samples with higher E2:E3 had lower average DBE. However, the size-fractionated microbial endmember, PLFA, showed no systematic trend between these same variables. Lastly, Fig. 5a overlays the optical data from

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McKay *et al.* (2018)[39] to show both general agreement with Peuravuori and Pihlaja (1997)[155] and continuity of the correlation with three soil isolates having ^{13}C aromaticity >35%. However, there was no statistically significant correlation across a subset including only the aquatic HPOA, NOM, and fulvic acid isolates ($p_s = 0.3$, aromaticity <25%).

Collectively, this review highlights that the foundation for relating E2:E3 to molecular weight hinges on a few samples that have been fractionated or isolated to reveal trends that are internally consistent. Particularly for molecular weight, without a breadth of samples across different spatiotemporal scales, E2:E3 may be better suited for examining changes across natural or engineered treatment gradients for a single sample or samples related closely in origin. In future assessments of broader sample contexts, it is imperative to keep pH constant between samples when measuring E2:E3, as the pH dependence of E2:E3 served as one of the earliest citations[189] to the dependence of this surrogate on molecular size. Even though the effects of pH were evaluated in the E2:E3 genesis papers, pH may impact both absorbance- and fluorescence-based surrogates more broadly,[21,192–197] and there is a general need to standardize pH adjustment practices across the DOM research field.

3.3 E4:E6

3.3.1 Definition and genesis. E4:E6 is most commonly defined as the ratio of absorbance at 465 to 665 nm. Early work often attributed the genesis to two papers, Scheffer (1954)[198] and Welte (1955),[199] focused on soil humic acid isolates. Different conventions have been used, such as 472/664,[199] 400/600,[200] 472/666,[201] and 465/675.[202] Use of E4:E6 is strongly rooted in characterizing soil isolates.[121,181,200,203–205] Use is less common for aquatic DOM, in part because the denominator (A_{665}) approaches the quantification limit of most spectrophotometers, unless the DOC concentration or cuvette pathlength are large.

For aromaticity, Kononova (1966)[206] postulated that E4:E6 is a surrogate for the degree of condensation of "aromatic nets of carbon atoms" in soil humic acids. The position recognized that E4:E6 was higher for fulvic acids than humic acids. Within humic acids, E4:E6 decreased from podzolic soils to chernozems, which is associated with a simultaneous decrease in aliphatic side chains.[206] This logic was later refuted in Chen *et al.* (1977)[157] but applied by Ghosh and Schnitzer (1979)[207] to describe macromolecular characteristics with changing pH and ionic strength.

Two studies investigated the relationship between E4:E6 and molecular weight using Sephadex gels. Tan and Giddens (1972)[200] characterized poultry litter, sewage sludge, and loamy sand, and Chen *et al.* (1977)[157] characterized soils from a range of classifications. Both studies reported lower E4:E6 ratios for humic acids than fulvic acids and an inverse correlation between molecular weight and E4:E6.[157,200] In Tan and Giddens (1972),[200] E4:E6 correlated positively with the Sephadex G-50 partition coefficients for fulvic acids, but a similar relationship was not observed for humic acids.[200] For size-fractionated soil isolates, Chen *et al.* (1977) reported significant correlations between E4:E6 and several physicochemical properties, including reduced viscosity ($r = -0.95$ a measure of molecular size), % carbon ($r = -0.73$), % oxygen ($r = 0.82$), total acidity ($r = 0.62$), and carboxyl content ($r = 0.66$).[157] The same study explored the relationship between E4:E6 and M_w as a function of pH, similar to the approach used by De Haan *et al.* (1983)[189] between E2:E3 and M_n for aquatic DOM. However, the pH-dependence of E4:E6 (soil isolates) and E2:E3 (aquatic isolates) was opposite. For soils, E4:E6 increased with increasing pH (from 2 to 7) with concomitant increases in M_w,[157] whereas E2:E3 decreased with increasing pH for aquatic DOM even though M_n increased.[189] In the end, Chen *et al.* (1977) argued that E4:E6 is a better indicator of molecular weight than of the abundance of condensed aromatic rings.[157]

3.3.2 Continued inquiry into aromaticity and molecular weight. Given the difficulty of measuring E4:E6 at low DOC concentrations, there has been less focus on this surrogate compared to E2:E3 for aquatic DOM. In Summers *et al.* (1987),[208] E4:E6 correlated inversely with molecular weight for native and size-fractionated isolates, including Suwannee River samples and a commercial sample (Aldrich Humic Acid). In Chin *et al.* (1994), E4:E6 was only weakly correlated to molecular weight for aquatic fulvic acids, whereas aromaticity was more strongly correlated to the molar absorption coefficient at 280 nm.[154] We do not recommend E4:E6 in future studies given the forementioned difficulties in measuring A_{665}, the lack of studies relating molecular-level information to this parameter, and the ongoing debate about whether soil humic substances are an artifact of the extraction process.[209–211]

3.4 Spectral slopes and the spectral slope ratio

3.4.1 Definition and genesis. Spectral slope calculations assume that, within a defined wavelength range, absorbance decays exponentially (Fig. 1c).[212–214] If exponential decay is assumed, the spectral slope ($S_{\lambda_1-\lambda_2}$) is the decay constant of a single exponential model for absorbance *versus* wavelength (Table 1), where the subscripts λ_1 and λ_2 indicate the wavelengths between which the slope is fit. In practice, there are two calculation approaches: fitting a non-linear exponential function[57,158–160] or log transforming the data to fit a linear regression.[57,162,213,215]

Interestingly, the earliest references to use the term "spectral slope" describe what today would be called an absorbance ratio. For example, Packham (1964) reports that the absorbance ratio of 300 to 450 nm is greater for fulvic acids than paired humic acids across 7 samples.[165] Kalle (1966) compared spectral steepness for marine DOM using an absorbance ratio of 420 to 665 nm.[216] Similarly, Brown (1977) defined spectral slope as the ratio of 280 to 310 nm.[217] Thus, absorbance ratios like E2:E3 are historically a spectral slope.

One of the first studies to use >2 wavelengths was Kopelevich and Burkenov (1977), in which log-transformed data was fit using a linear regression between 390 and 490 nm.[215] In early studies, S was leveraged as both a characterization tool, focusing on variability between samples, and as a modeling

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tool, assuming little variability between samples. Often cited in later key studies,[57,160,218] Zepp and Schlotzhauer (1981)[67] characterized spectral slopes between 300 and 500 nm noting similar values (range 0.0128–0.0175 nm^{-1}) for soil fulvic acids and DOM from freshwater and marine environments (linearization not specified). Bricaud *et al.* (1981) advocated that the small variability between samples would enable an average value of S to be used to extrapolate absorbance spectra from the UV range (375 nm) to the visible range (440 nm).[213] Blough and coworkers used spectral slopes, calculated from log-transformed data, to characterize differences in water samples from the Orinoco River outflow,[219] Gulf of Mexico,[220] and Amazon River.[220] With linearization, however, absorbance near or below detection disproportionately impacts fitted model parameters. To minimize impacts, Blough, Green, and coworkers[219,220] defined a detection limit (absorption coefficient = 0.1 m^{-1}) and fit linear regressions for sample-specific wavelength ranges where absorbance exceeded the limit.

These examples highlight the challenges in comparing spectral slopes between studies that arise due to differences in wavelength range and regression technique. Twardowski *et al.* (2004) recognized and addressed this issue comprehensively, providing a literature review (see Table 1 of Twardowski *et al.* (2004)).[160] Although this paper ultimately recommended a hyperbolic model,[160] the DOM field has largely continued to use a single exponential model with closer attention paid to wavelength ranges and regression techniques. Related to spectral slope, Helms *et al.* (2008)[57] defined the spectral slope ratio (S_R) as the ratio of $S_{275–295}$ divided by $S_{350–400}$.[221] Our recent meta-analysis of >700 paired optical surrogates identified that E2:E3 and spectral slope (especially $S_{300–600}$), but not S_R, are equally as good at describing spectral tailing.[222]

3.4.2 Continued inquiry into aromaticity and molecular weight

3.4.2.1 Molecular weight. Although Helms *et al.* (2008)[57] is the most cited reference for linking spectral slopes to DOM molecular weight, there are earlier investigations. Hayase and Tsubota (1985) extracted humic- and fulvic acid isolates from sediment, followed by molecular weight fractionation using UF. For the fulvic acid, S decreased with increasing molecular weight from the <10 to >300 kDa fractions.[218] Yacobi *et al.* (2003)[171] fractionated samples by UF into <10, 10–50, and >50 kDa nominal molecular weights for 6 rivers in Georgia (USA). For every sample, $S_{300–450}$ was greater in the smallest fraction compared to the largest. However, the medium fraction did not always support a monotonic trend with molecular weight for individual samples.

In their formative study, Helms *et al.* (2008)[57] sampled diverse aquatic systems (*i.e.*, marsh, coast, open ocean) and compared spectral slopes derived from two regression methods and three wavelength ranges. $S_{275–295}$ and $S_{350–400}$ were fit using linear regressions of log-transformed absorbance spectra and

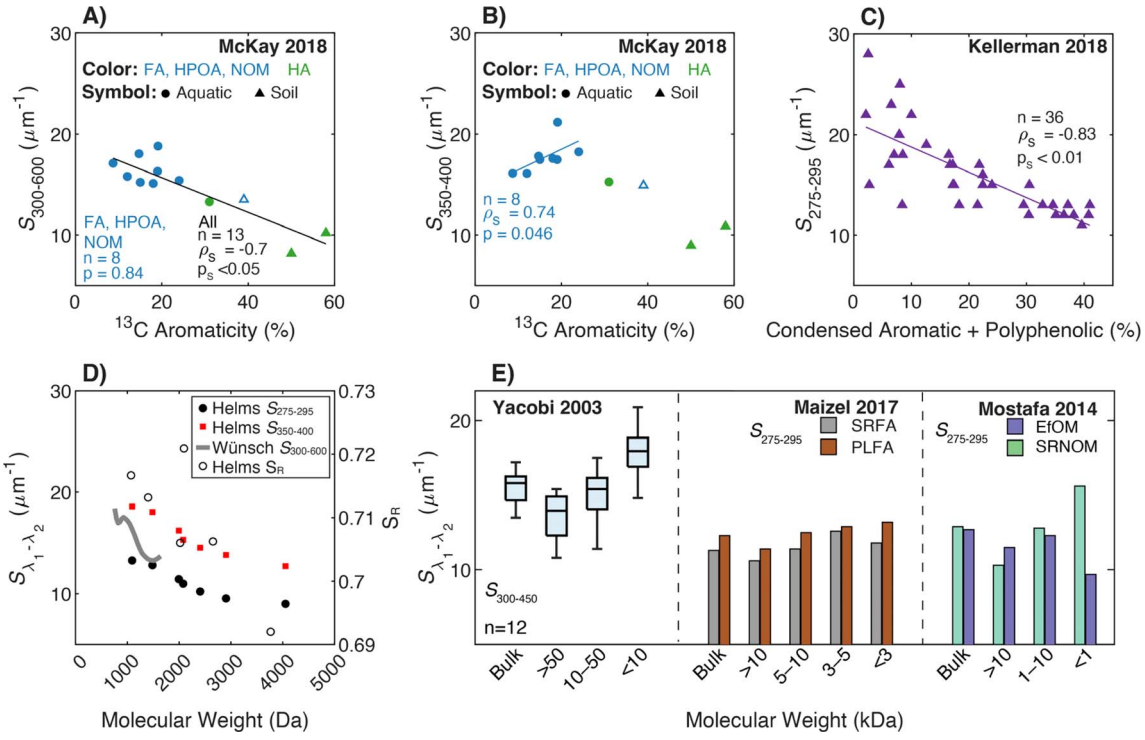

Fig. 6 Relationship between spectral slope (S), spectral slope ratio (S_R), and either molecular weight or different measures of aromatic carbon. (A) $S_{300–600}$ and (B) $S_{350–400}$ *versus* aromaticity (%) per sample calculated from ^{13}C NMR data[39–43] with isolate type indicated by the marker and color. (C) $S_{275–295}$ *versus* the relative abundance of condensed aromatic and polyphenolic formulae (%) per sample calculated based on AI_{mod}[132] from FT-ICR MS data.[35] (D) $S_{\lambda_1-\lambda_2}$ and S_R *versus* molecular weight determined by SEC.[57,223] Wavelength ranges are defined in the legend. Spectral slope, S_R, and M_n values from Helms *et al.* (2008) were extracted using Web Plot Digitizer (https://apps.automeris.io/wpd/). (E) $S_{\lambda_1-\lambda_2}$ *versus* nominal molecular weight by UF fractionation.[36,38,171]

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showed high variability between samples from different environments. Samples were also fractionated by UF (1 kDa) calculating the distribution of DOC concentrations, or integrated absorbance, in the permeate (labeled low molecular weight, LMW) compared to the retentate (labeled high molecular weight, HMW). In addition to whole-water samples, Helms *et al.* (2008) also size-fractionated SRNOM by Superdex-30 characterizing each fraction by SEC with UV detection.[57]

For size-fractionated SRNOM, Helms *et al.* (2008) showed that $S_{275–295}$ and $S_{350–400}$ both decreased with increasing molecular weight, indicating increased absorbance tailing (Fig. 6d). However, S_R had no statistically significant correlation with molecular weight ($p = 0.3$) (Fig. 6d). The lack of correlation for size-fractionated SRNOM contrasted with the whole-water samples (Fig. 4 in Helms *et al.* (2008)), where S_R correlated positively with LMW : HMW ratio across diverse environmental contexts.[57] It is worth noting that the relationship between S_R and the LMW : HMW ratio has steeper slope and less scatter in the data (higher R^2) for integrated absorbance compared to DOC concentration, but the number of samples was different.[57] The difference in slope highlights the limitation of UV absorbance to track DOC concentration because not all DOM is chromophoric.

More recently, endmember analysis revealed conflicting trends. Using coarse fractions, $S_{300–600}$ increased with decreasing molecular weight in Mostafa *et al.* (2014).[38] However, with more resolution, SRFA defied this trend in Maizel and Remucal (2017),[36] with the <3 kDa fraction having a lower S than the 3–5 kDa fraction (Fig. 6e). Several studies examined spectral slope at even finer size resolutions than UF. Guégen and Cuss (2011) coupled asymmetric field-flow fractionation with a diode array detector to demonstrate that $S_{275–295}$, $S_{350–400}$, and S_R decreased with increasing molecular weight.[224] Using SEC, Wünsch *et al.* (2018) affirmed the same trend for $S_{300–600}$ (Fig. 6d).[223] It is worth noting, however, that in all SEC studies, the dependence of spectral slope on molecular weight was most sensitive at intermediate weights (1500–2500 Da), and trends were not monotonic in some studies.

3.4.2.2 Aromaticity. By plotting FT-ICR MS data from Kellerman *et al.* (2018), there was an inverse correlation between $S_{275–295}$ and the relative abundance of condensed aromatic and polyphenolic formulae (Fig. 6c), but like E2:E3, data are highly scattered for samples with lower relative abundance of aromatic formulae. Endmember samples in Maizel and Remucal (2017) exhibited similar trends to E2:E3; there was a negative correlation between $S_{275–295}$ and DBE for SRFA but not PLFA.[36]

Optical data from McKay *et al.* (2018) showed different relationships depending on the data subset (*i.e.*, with or without soils or humic acids). Including all isolates, there was a statistically significant, inverse correlation between ^{13}C aromaticity and both $S_{300–600}$ (Fig. 6a) and $S_{275–295}$ (data not shown), but there was no correlation between ^{13}C aromaticity and either $S_{350–400}$ (Fig. 6b, $p_s = 0.12$) or S_R ($p_s = 0.09$, data not shown). Constrained to only fulvic acid-dominated aquatic isolates (excluding SRHA), $S_{350–400}$ had the lowest p_s value (0.046) but the correlation was positive ($\rho_S = 0.74$) contradicting the broad trends observed for $S_{275–295}$ and $S_{300–600}$ (Fig. 6b).

Taken as a whole, spectral slope appears to vary with both molecular weight and descriptors of aromatic carbon content across wide samples gradients (*e.g.*, endmember comparison) or sample fractionation. The conventional interpretation relating slope to molecular weight shows the strongest evidence for single-sample fractionations with noted contradictions. Between samples of different origin, there is overlap in surrogate values between molecular weight classes (Fig. 6e). Lastly, comparing Fig. 5 (E2:E3) and Fig. 6 ($S_{275–295}$ or $S_{300–600}$), there is no indication that one surrogate is more robust than the other.[45]

Considering absorbance surrogates for calculations of aromatic nature, the Kellerman *et al.* (2018)[35] data suggest that $SUVA_{254}$ is not a redundant surrogate to measures of spectral slope, such as E2:E3 and S. There is less scatter in the $SUVA_{254}$ data, and low $SUVA_{254}$ values are specific to samples with a low (<15%) abundance of condensed aromatic and polyphenolic formulae (Fig. 4d, 1–2 L mg_C^{-1} m^{-1}). However, a wide range of E2:E3 (Fig. 5d) and spectral slope values (Fig. 6c) are observed at low relative abundances of the same formulae, showing too much scatter in the data to be informative.

4 Fluorescence surrogates

Fluorescence-based optical surrogates are listed in Table 2 along with the equations and primary sources. The following subsections explore each surrogate. This review focuses on indices that are not derived from PARAFAC models, because PARAFAC components are unique to individual studies. Although relationships between PARAFAC components and DOM physicochemical characteristics are not investigated, PARAFAC data is used to reconstruct fluorescence intensities from one source in Section 4.2.

4.1 Apparent quantum yield

4.1.1 Definition and genesis. The use of quantum yields to characterize DOM predates other fluorescence surrogates. As introduced in Section 2.2, fluorescence quantum yield (Φ_f) characterizes the proportion of absorbed photons that are emitted as fluorescence. Unlike other fluorescence surrogates (*e.g.*, HIX, FI, BIX, *etc.*), quantum yields are rooted in fundamental principles, used broadly across photophysical inquiries, and not limited to DOM characterization.[54,65,225] A unique consideration is that DOM is a heterogeneous mixture and not a single fluorophore. Therefore, the term apparent quantum yield (AQY) is used.

In practice, AQY compares the fluorescence of an unknown sample to a standard with a known quantum yield by calculating the ratio of the integrated fluorescence intensity (across all emission wavelengths) relative to the absorbance at the same excitation wavelength (Table 2). The same measurements are performed on a standard with a known quantum yield. Quinine sulfate in 0.1 N H_2SO_4 is well-characterized ($\Phi_f = 0.51$)[237] and a commonly used reference for humic substances.[39,66,220,234,238] However, salicylic acid ($\Phi_f = 0.35$) has been preferred for SEC studies over quinine sulfate in 0.1 N H_2SO_4 due to

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Table 2 Definitions of optical surrogates based on fluorescence spectroscopy[a]

Metric	Abbreviation	Example calculation	Notes	Ref.
Quantum yield	Φ_f	$\frac{\phi_{unknown}}{\phi_{ref}} = \frac{{n_{unk}}^2}{{n_{ref}}^2} \frac{\int_0^\infty I_{unk}(\lambda_{ex},\lambda_{em})d\lambda_{em}}{1-10^{A_{unk}(\lambda_{ex})}} \frac{1-10^{A_{ref}(\lambda_{ex})}}{\int_0^\infty I_{ref}(\lambda_{ex},\lambda_{em})d\lambda_{em}}$	unk: unknown ref: reference	Birks (1970)[225] Würth *et al.* (2013)[226]
Humification index	HIX	$HIX_{em,1999} = \frac{\sum_{\lambda_{em,3}}^{\lambda_{em,4}} I(254\ nm,\lambda_{em})}{\sum_{\lambda_{em,1}}^{\lambda_{em,2}} I(254\ nm,\lambda_{em})}$ $HIX_{em,2002} = \frac{\sum_{\lambda_{em,3}}^{\lambda_{em,4}} I(254\ nm,\lambda_{em})}{\sum_{\lambda_{em,1}}^{\lambda_{em,2}} I(254\ nm,\lambda_{em}) + \sum_{\lambda_{em,3}}^{\lambda_{em,4}} I(254\ nm,\lambda_{em})}$	$\lambda_{em,1} = 300$ nm $\lambda_{em,2} = 345$ nm $\lambda_{em,3} = 435$ nm $\lambda_{em,4} = 480$ nm	Zsolnay *et al.* (1999)[227] Ohno (2002)[228]
Biological/freshness index	β/α	$\beta/\alpha = \frac{I(310\ nm,380\ nm)}{\max(I(310\ nm,420-435\ nm))}$		Parlanti *et al.* (2000)[229] Huguet *et al.* (2009)[230] Wilson and Xenopoulos (2009)[231]
	BIX	$BIX = \frac{I(310\ nm,380\ nm)}{I(310\ nm,430\ nm)}$		
Fluorescence index	FI	$FI = \frac{I(370\ nm,470\ nm)}{I(370\ nm,520\ nm)}$		McKnight *et al.* (2001)[232] Cory and McKnight (2010)[84]
Maximum emission wavelength	$\lambda_{em,max}$ at $\lambda_{ex} = 310$ nm	$\lambda_{max}(310) = \max(I(310\ nm, 310-590\ nm))$		Parlanti *et al.* (2000)[229] McKnight *et al.* (2001)[232] Wilson and Xenopoulos (2009)[231] Cory and McKnight (2010)[84]
	$\lambda_{em,max}$ at $\lambda_{ex} = 370$ nm	$\lambda_{max}(370) = \max(I(370\ nm, 370--710\ nm))$		
Specific fluorescence intensity	SFI A	$SFI\ A = \frac{I(260\ nm,426\ nm)}{[DOC]}$		Coble (1996)[233] Alberts and Takács (2004)[234] Jaffé *et al.* (2004)[235] Hudson *et al.* (2007)[236] Korak *et al.* (2014)[75]
	SFI B	$SFI\ B = \frac{I(280\ nm,310\ nm)}{[DOC]}$		
	SFI C	$SFI\ C = \frac{I(320\ nm,\ 440\ nm)}{[DOC]}$		
	SFI T	$SFI\ T = \frac{I(280\ nm,338\ nm)}{DOC}$		

[a] λ_{ex} is the excitation wavelength in nm. λ_{em} is the emission wavelength in nm. $A_i(\lambda_{ex})$ is decadic absorbance of sample i at λ_{ex}. $I(\lambda_{ex},\lambda_{em})$ is the corrected fluorescence intensity at λ_{ex} and λ_{em} in normalized units. n_i is the refractive index of sample i at the average emission wavelength. [DOC] is the dissolved organic carbon concentration.

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incompatibility of acidic solutions with the stationary phase of chromatography columns.[66,239]

The earliest studies reported ratios of fluorescence to absorbance (or color) rather than the formally defined AQY. For example, Kalle (1949) followed by Duursma (1965) and Christman and Ghassemi (1966) investigated linear relationships between fluorescence intensity and absorbance,[164,240,241] the slope of which is related to AQY. Levesque (1972) and Hall and Lee (1974) fractionated samples with Sephadex and found the ratio of fluorescence to absorbance increased with decreasing molecular weight for a fulvic acid, extracts from leaves, sediment, and a lake sample.[105,242] Zepp and Schlotzhauer (1981) formally calculated AQY for a range of aquatic and soil sources and noted that, despite breadth in origin, the range was remarkedly small (0.001–0.004), except for a sample from the Florida Everglades (0.012).[67] AQY for most aquatic isolates is <0.02,[39,66,220,238] with higher values (0.02–0.1) reported for specific contexts, such as marine and estuary environments,[66,220,243] EfOM,[38,244] pyrogenic organic carbon,[245,246] and samples treated by physicochemical processes.[247,248] In Alberts and Takács (2004),[234] AQY was higher for fulvic acids than humic acids for 13 IHSS isolates. Overall, AQY values for DOM are small compared to pure, model organic compounds common in broader fluorescence inquiries such as tyrosine ($\Phi_f = 0.21$),[249] quinine sulfate ($\Phi_f = 0.51$),[237] and fluorescein ($\Phi_f = 0.79$).[250] However, model compounds leveraged as DOM building blocks, such as gallic-, syringic-, and vanillic acids, typically have quantum yields of similar magnitude to DOM.[66]

4.1.2 Continued inquiry into aromaticity and molecular weight

4.1.2.1 Molecular weight. For both aquatic samples and plant leachates, Stewart and Wetzel (1980)[251] fractionated samples using Sephadex and dialysis, where the fluorescence-to-absorbance ratio (Fluor/Abs) increased with decreasing molecular weight following internally consistent trends within, but not across, samples (Fig. 7e). For example, the Fluor/Abs ratio for the <3.5 kDa fraction of leaf litter leachate was similar to the >3.5 kDa fraction from Lawrence Lake (Fig. 7e). Similarly, De Haan and De Boer (1987) fractionated lake samples using UF and presented a log-linear correlation between Fluor/Abs and molecular size (Fig. 7b).[77] While this trend was strong, the concentration-weighted Fluor/Abs values from each fraction did not reconcile with the whole-water value, suggesting non-conservative mixing of optically active moieties.[77] A note of caution, some early studies used different fluorescence excitation and absorbance wavelengths,[77,116,251] whereas fundamental theory and current practice utilize the same wavelength.

Ewald *et al.* (1988) formally calculated the AQY for a fulvic acid that was isolated from a microbial culture inoculated with a soil actinomycete and fractionated by UF. For this narrow context, AQY decreased with increasing log-molecular weight following a pH-dependent, linear relationship (Fig. 7a).[180] More representative of aquatic systems, Belin *et al.* (1993) fractionated two waters comparing UF and XAD techniques (Fig. 7a and d). For UF fractionation, AQY decreased as molecular weight increased following a sample-specific, log-linear relationship.[252] Similar sample-specific trends of increasing AQY with decreasing molecular weight have been demonstrated for a broader range of DOM samples, including isolates (*e.g.*, terrestrial and marine)[70,186,253] and EfOM (Fig. 7d),[38] but it has also been contradicted in deep oligotrophic waters.[116]

Although many studies have presented SEC chromatograms using absorbance and fluorescence detectors,[118,254,255] Hanson *et al.* (2022)[239] was the first to couple the responses from both detectors and calculate AQY as a function of elution volume. This coupling also required development of instrument correction factors for SEC. Within a single sample, Hanson *et al.* (2022) offers additional lines of evidence that AQY increases as molecular weight decreases for a surface water (Colorado, USA), two humic substance isolates, and ozonated samples. Comparing SEC chromatograms of AQY and DOC concentration, the DOM fraction with the highest AQY was associated with a relatively small fraction of the total organic carbon.[239]

4.1.2.2 Aromaticity. Compared to size fractionation, fewer studies have assessed AQY as a function of aromatic carbon content. As an indirect approach, XAD methods isolate fractions with more (HPOA) or less (transphilic acid, TPIA) aromatic carbon.[139,256] In Belin *et al.* (1993), the humic acid isolate had lower AQY compared to the paired fulvic acid,[252] which can be inferred as AQY decreasing with increasing abundance of aromatic carbon.[1] Applying a conservative mass balance to fluorescence intensity for one lake sample, Belin *et al.* (1993)[252] concluded that roughly 60% of the whole-water fluorescence intensity was attributed to the fulvic acid fraction, followed by hydrophilic acids (35%) then humic acids (5%). Baker *et al.* (2008)[257] applied DAX-8 resin fractionation to 25 surface waters demonstrating that the Fluor/Abs ratio (Peak C to absorbance at 340 nm) was proportional to the fraction of whole-water DOC recovered in as hydrophilic DOM. There are mixed results about the impact of solid phase extraction on AQY. Although some studies show decreasing spectral slope after isolation with C18 cartridges,[70,220] others report mixed outcomes for how solid phase extraction impacts fluorescence AQY.[220,243,258]

Although the aforementioned studies provide indirect evidence relating AQY and aromaticity, optical data from McKay *et al.* (2018)[39] paired with ^{13}C aromaticity data[40–43] reveal two main trends (Fig. 7f). Across most aquatic isolates, AQY decreased with increasing aromaticity with a continuous, linear relationship across fulvic-dominated, aquatic isolates (NOM, HPOA, and FA) and SRHA. As a subset, humic acids (1 aquatic and 2 soil) showed no trend. Two fulvic-dominated isolates were outliers. Yukon HPOA (YHPOA), a winter baseflow HPOA isolate from the Yukon River,[42] had a higher AQY than all other isolates. Similarly, Pahokee Peat fulvic acid (PPFA) from the Florida Everglades exhibited an AQY similar to aquatic isolates, despite this sample having a higher aromaticity based on ^{13}C NMR.

Across diverse aquatic environments, there is an inverse relationship between AQY and peak emission wavelength. This trend could be predicted from first-principles relating the energy gap between S_1 and S_0 to non-radiative decay rates.[52,225] In Ewald *et al.* (1988) and Belin *et al.* (1993), emission wavelengths of maximum fluorescence intensity shifted 30–50 nm

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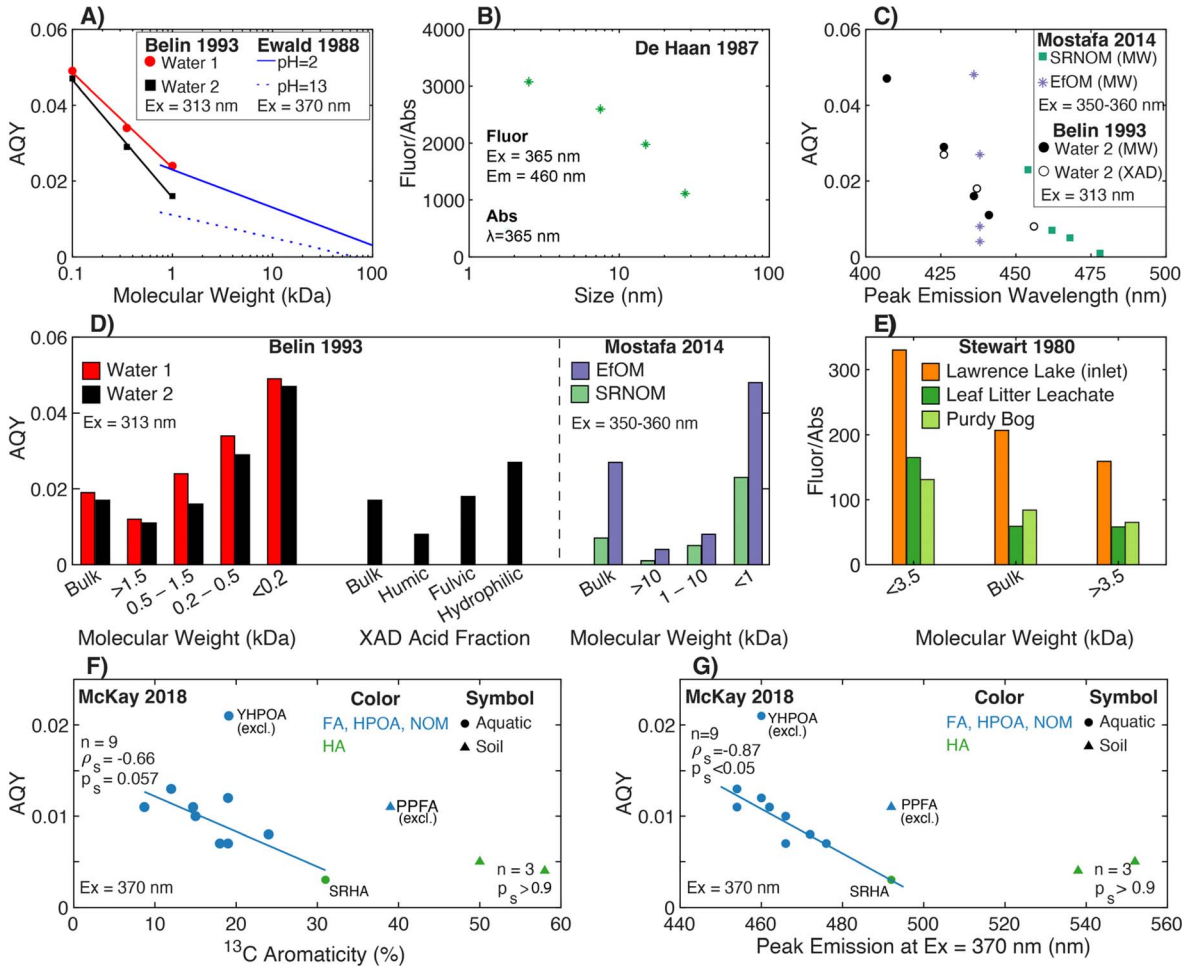

Fig. 7 Relationship between fluorescence efficiency and either molecular weight, isolate fraction, peak emission wavelength, or different measures of aromatic carbon. Fluorescence efficiency is quantified through either fluorescence AQY (A, C, D, F, and G) or as the ratio of fluorescence-to-absorbance at discrete wavelengths (B and E). (A) AQY *versus* molecular weight for UF fractions[252] plotted at the mid-point of each bounded molecular weight fraction. The unbounded fraction (>1.5 kDa) is not plotted. Regression lines are from Ewald *et al.* (1988).[180] (B) Fluorescence-to-absorbance ratio (Fluor/Abs) *versus* molecular size for Lake Tjeukemeer DOM. Molecular size data are plotted at the midpoint for each bounded size fraction.[77] (C) AQY *versus* peak emission wavelength for one water fractionated by both UF and XAD techniques[252] and two endmembers fractionated by UF.[38] (D) AQY *versus* either UF fraction or XAD fraction.[38,252] (E) Fluor/Abs *versus* UF fraction.[251] Fluorescence efficiency is reported using fluorescence at λ_{ex} 360 nm and λ_{em} 460 nm relative to absorbance at 250 nm. AQY[39] *versus* (F) aromaticity (%) per sample calculated from ^{13}C NMR data[40–43] or (G) peak emission wavelength.[39] Yukon HPOA (YHPOA) and Pahokee Peat Fulvic Acid (PPFA) isolates were excluded from the linear regression.

between fractions, and maximum wavelengths increased as AQY decreased.[180,252] In the latter study, there was good agreement between UF and XAD fractionation techniques for a single sample (Fig. 7c).[252] This trend also held for UF-fractionated SRNOM but not EfOM, which showed no change in peak emission wavelength (Fig. 7c).[38] Fig. 7g shows that the trend generally held across the aquatic isolates in McKay *et al.* (2018).[39] Similar to relationships with aromaticity (Fig. 7f), however, humic acids, YHPOA, and PPFA were exceptions.

Broadly, systematic relationships between AQY and molecular weight appear to hold for fractionated samples; however, variations in AQY should not be used to infer differences in molecular weight across samples. An inverse relationship between AQY and both aromaticity and peak emission wavelength holds across a range of isolates; however, future work is needed to investigate the outliers (*e.g.*, YHPOA, PPFA, and EfOM). Lastly, the increased resolution gained by online measurements of AQY using SEC highlights a future opportunity for establishing relationships between AQY and DOM molecular size.

4.2 Specific fluorescence intensity (SFI)/relative fluorescence intensity

4.2.1 Definition and genesis. Specific fluorescence intensity (SFI) is older than more common fluorescence surrogates, although its name and definition have varied. Normalizing fluorescence intensity to DOC concentration is analogous to SUVA, as an intrinsic surrogate independent of the concentration. Three-dimensional EEMs, two-dimensional spectra, individual intensities at discrete wavelength pairs, and integrated

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EEMs can all be carbon-normalized. Compared to other surrogates, SFI is not calculated at consistent wavelengths across studies. Users have discretion over what fluorescence data to normalize, which is analogous to $SUVA_{254}$, $SUVA_{272}$, and $SUVA_{280}$ described in Section 3.1. The most common wavelength pairs align with the peak names defined by Coble (1996)[233] and summarized in Table 2. Over time, this surrogate has been called relative fluorescence intensity,[187,234,259,260] specific fluorescence,[75,261,262] carbon-specific fluorescence,[235,263] or carbon-normalized intensity.[166,245,264,265] For the remainder of this review, we will use the term SFI, leveraging the parallel to specific absorbance, regardless of the term used in individual studies. Despite its legacy, specific intensities have not been traditionally included in reviews of other fluorescence surrogates.[266,267]

There are different approaches to calculate SFI with different unit conventions. For example, samples may be prepared at the same DOC concentration (*e.g.*, 1, 10 or 100 $mg_C\ L^{-1}$) before measuring fluorescence,[259,268–271] or samples may be analyzed as-is before dividing fluorescence intensities by the DOC concentration.[235,261] The two approaches should yield equivalent information, assuming that (1) sample chemistry follows fundamental equations (Bouguer–Beer–Lambert law (eqn (1)) and linearized fluorescence (eqn (7))) and (2) best practices are followed for spectral corrections and instrument limitations (*e.g.*, detector linearity). The units depend on the units of both fluorescence intensity (*e.g.*, arbitrary units, counts per second, Raman units, or quinine sulfate equivalents) and DOC concentration (*e.g.*, $mg_C\ L^{-1}$ or $mmol_C\ L^{-1}$), leading to units analogous to SUVA (*e.g.*, RU L mg_C^{-1}, RU L g_C^{-1} or QSU L $mmol_C^{-1}$).[75,235,261,272] In some cases, captions or methods describe the normalization process, not the intensity units explicitly.[273–275]

There is an opportunity for the research community to be more explicit and consistent in unit conventions to facilitate inter-study comparisons and reduce ambiguity. It is important to note that diluting to a common DOC concentration does not negate the need for inner filter corrections as each sample has a unique absorbance spectrum. Care is needed when interpreting studies without spectral corrections.

SFI is the slope relating fluorescence intensity to DOC concentration.[75] As early as 1968, Ghassemi and Christman compared different source waters using SFI.[169] In addition to highlighting compositional differences, research has also leveraged contexts with little variability to predict DOC concentration from fluorescence.[76,276] Across samples, stronger correlations between fluorescence intensity and DOC concentration indicate more homogeneity in DOM composition. In contexts with greater spatial and temporal variability, (*e.g.*, diverse surface waters, variable land use, and changing hydrological conditions), relationships between DOC concentration and fluorescence exhibit site-specific slopes with lower correlation coefficients[76,272,276] compared to narrower contexts.[76,272,276] For example, consistent slopes are observed within narrow contexts, such as short-term sampling campaigns across an estuary salinity gradient dominated by conservative dilution.[71,263] Therefore, using fluorescence as a proxy for DOC concentration depends on an underlying assumption of constant SFI.

4.2.2 Continued inquiry into aromaticity and molecular weight

4.2.2.1 Aromaticity. SFI was not introduced as a quantitative surrogate for a specific compositional characteristic, but differences were often interpreted through the lens of aromatic carbon. Using soil isolates prepared at comparable concentrations, Mukherjee and colleagues (1964, 1971)[270,277] attributed fluorescence to aromatic moieties with electron donating groups or conjugated, unsaturated compounds with resonance. Datta *et al.* (1971)[270] reported higher SFI for fulvic acids than humic acids and attributed differences to aromaticity. Later, Senesi, Miano, and colleagues measured excitation, emission, and synchronous scans for soil and aquatic isolates prepared at a constant concentration. These studies proposed SFI as a tool to distinguish humic acids from fulvic acids, noting that fulvic acids consistently exhibited higher SFI and blue-shifted emission maxima.[268,275,278,279] Focused on soils, Senesi *et al.* (1991)[268] attributed these differences to fulvic acids having less aromatic condensation, less conjugation between chromophores, and higher abundance of electron-donating groups. Although independent measures were not presented, the study advocated for complementary techniques like NMR in future work.

There are some contradictions in literature relating SFI and aromaticity between intra- and inter-sample comparisons. Croue *et al.* (2000) fractionated SRNOM with XAD resin, comparing SFI at excitation 320 nm between 9 fractions. In general, SFI increased as relative hydrophilicity of the fraction increased. For example, the hydrophilic base fraction accounted for 0.04% of the total carbon but was about four times more fluorescent.[125] This study would suggest that SFI increases as aromaticity decreases. However, in McKnight *et al.* (2001),[232] SFI was calculated at excitation 370 nm with paired ^{13}C aromaticity measurements ($n = 6$), and the median SFI was higher for terrestrial endmembers with higher aromaticity than microbial endmembers (Fig. 8a). Since the highest SFI in the whole dataset ($n = 11$) was a microbial endmember, fluorescence index (FI) was justified as a better measure to differentiate sources.[232] Alberts and Takács (2004)[234] paired fluorescence measurements with IHSS-reported ^{13}C NMR[40,41] and carboxyl and phenol content[280] for 13 IHSS isolates. SFI near Peak C (λ_{ex} 330 nm, λ_{em} 440 nm) did not correlate with the relative abundance of functional groups determined by NMR. Using titration, however, SFI correlated positively with carboxyl content and inversely with phenolic content for the fulvic acids. The same correlations were not significant for humic acids, where SFI only correlated with $SUVA_{254}$.[234]

The optical data in McKay *et al.* (2018) illustrates a discontinuity between humic- and fulvic acid isolates. Fig. 8b shows two distinct groups in the data. Isolates from aquatic fulvic acids, HPOA and NOM formed one group where Peak C SFI (λ_{ex} 320 nm, λ_{em} 440 nm) and ^{13}C aromaticity correlated ($\rho_S = 0.93$, $n = 9$). Humic acids formed a second group where Peak C SFI was relatively insensitive to ^{13}C aromaticity. Detailed in ESI Text 4,† Peak C SFI was estimated by reconstructing intensities from PARAFAC data presented in Kellerman *et al.*

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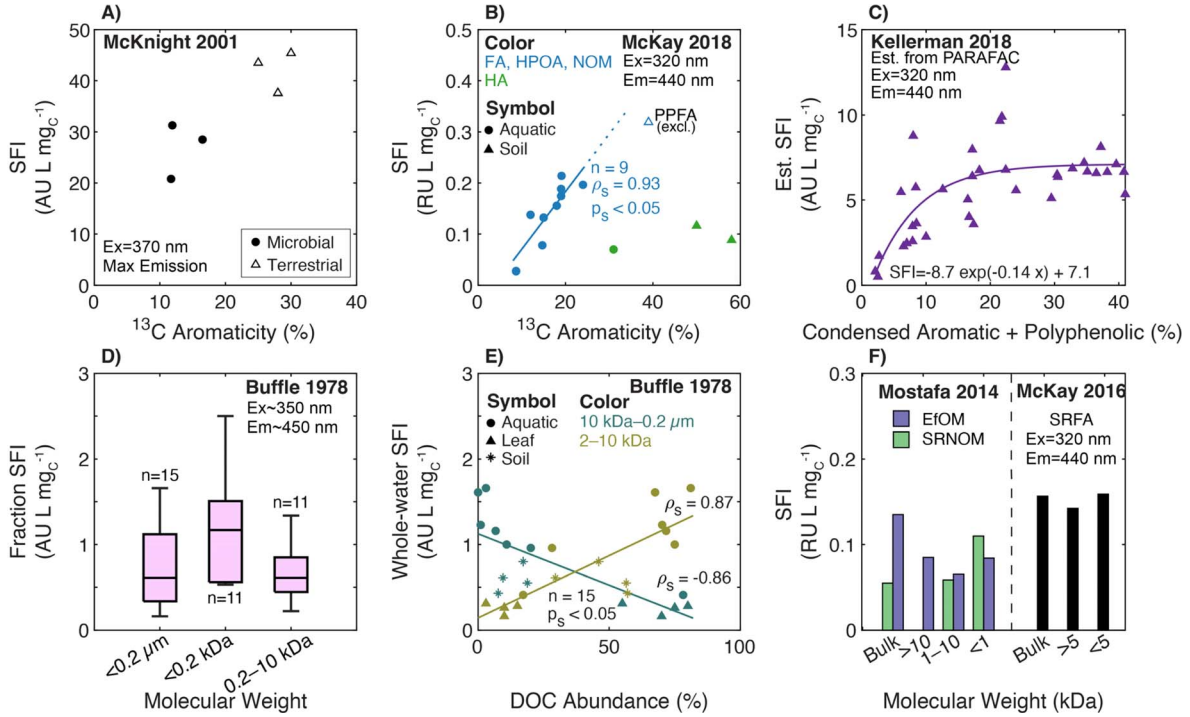

Fig. 8 Relationship between specific fluorescence intensity (SFI) and either molecular weight or different measures of aromatic carbon. (A) SFI *versus* aromaticity (%) per sample calculated from ^{13}C NMR data for fulvic acids.[232] (B) SFI *versus* aromaticity (%) per sample calculated from ^{13}C NMR data pairing optical[39] and NMR data.[40–43] (C) Estimated (Est.) SFI (calculated from PARAFAC components, see ESI Text 4†) *versus* the relative abundance of condensed aromatic and polyphenolic formulae (%) per sample calculated based on AI_{mod}[132] from FT-ICR MS data.[35] (D) SFI of UF molecular weight fractions from lakes, soils, and leaf leachates.[100] (E) Whole-water (<0.2 μm) SFI *versus* the relative abundance of whole-water DOC (%) in two UF fractions.[100] (F) FI *versus* UF molecular weight fractions.[38,44]

(2018),[35] because peak intensities were not reported. Fig. 8c suggests that Peak C SFI increases as relative abundance increases only at low abundances of aromatic and polyphenolic formulae. When abundance exceeds about 25%, the estimated SFI is comparatively insensitive to compositional changes measured by FT-ICR MS.

4.2.2.2 Molecular weight. Using Sephadex separation, Ghassemi and Christman (1968) characterized a river sample (British Columbia, Canada) by GPC and reported chromatograms of DOC concentration, absorbance (350 nm), and fluorescence intensity (λ_{ex} 360 nm, λ_{em} 415 nm). The elution volume at which the maximum signal occurred was lower for both DOC concentration and absorbance compared to fluorescence, suggesting SFI increased in smaller molecular weight fractions.[169] At similar fluorescence wavelengths, several other studies have reported higher SFI associated with smaller molecular weight for lakes[105] and extracts from soil, leaves, and sediment.[105,281]

Using UF, Buffle *et al.* (1978)[100] fractionated eight natural waters, five soil extracts, and four leaf extracts into three molecular weight fractions (0.2 μm–10 kDa, 0.21–10 kDa, and <0.21 kDa). Compared to the filtered whole-waters (<0.2 μm), median SFI values were higher in the smaller weight fractions (Fig. 8d). Applying a mass balance to DOC concentration, Fig. 8e shows that the whole-water SFI increased as (1) the fraction of whole-water DOC recovered in the middle fraction (0.21–10 kDa) increased ($\rho_S = 0.87$, $n = 15$) and (2) the DOC recovered in the large fraction (0.2 μm–10 kDa) decreased ($\rho_S = 0.87$, $n = 15$). There was no statistically significant correlation between SFI and the fraction of whole-water DOC recovered in the small fraction (<0.2 kDa) ($p_s = 0.2$).

Multiple studies have presented additional evidence supporting greater SFI in smaller molecular weight fractions, but a pattern of inconsistencies with the smallest weight fractions are noted.[77,282,283] For example, in De Haan and De Boer (1987), SFI (λ_{ex} 365 nm, λ_{em} 470 nm) was highest in the intermediate fractions (5–35 nm) and smaller in both remaining (<5 nm and >35 nm) fractions.[77] For a membrane bioreactor, Xiao *et al.* (2018)[284] noted a discontinuity for the <0.5 kDa fraction with lower than expected SFI and relatively little change in SFI at molecular weights >10 kDa. For a sediment sample, Hayase and Tsubota (1985) fractionated humic- and fulvic acid isolates into five molecular weight fractions ranging from <10 kDa to >300 kDa, and SFI increased with decreasing molecular weight. Interestingly, SUVA at the excitation wavelengths of maximum fluorescence intensity (320 and 480 nm for fulvic- and humic acid isolates, respectively) had opposite correlations with SFI. For humic acids, $SUVA_{480}$ correlated positively with SFI, but $SUVA_{320}$ correlated negatively with SFI for fulvic acids. This difference was attributed to $SUVA_{480}$ increasing with decreasing humic acid molecular weight fraction,[218] which contradicts the trend most often observed (Section 3.1.2). In Fig. 8f, two IHSS isolates followed the trend of increasing SFI as molecular weight fraction decreased. The

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wastewater endmember (EfOM) deviated with similar a SFI in the largest and smallest fractions and an uncharacteristically high SFI in the bulk water, warranting further investigation.

Between 1999 and 2004, several studies by Alberts, Takács, and colleagues used UF and/or SEC to characterize DOM in aquatic environments using SFI near Peak C.[186,187,285,286] Some samples, such as Suwannee River and estuarine UF fractions, matched expectations showing increasing SFI with decreasing molecular weight.[187,285,286] However, some waters had higher SFI in the middle fraction (10–50 kDa) compared to the small fraction (<10 kDa), and differences between the Peak A and C regions.[187] Although some studies do not appear to correct for inner filter effects, the lower SFI in the small fraction cannot be explained by inner filtering. If absorption coefficients typically increase with molecular weight (Section 3.1), then SFI would be suppressed more in larger fractions. More likely, inconsistencies can be attributed to the differences in membrane cut-off relative to the expected weight distribution in DOM. A 10 kDa membrane cut-off may offer appropriate resolution to fractionate soil isolates, but most carbon by mass is associated with <10 kDa fractions in aquatic systems.[1,100] By using SEC to fractionate two samples, the increased resolution in Alberts *et al.* (2002) supports increasing SFI with decreasing molecular weight for wavelengths near Peaks A and C.[187]

Molecular weight resolution is further improved with online fluorescence and DOC concentration detectors for SEC analysis. Her *et al.* (2003, 2004)[262,269] characterized contrasting samples, including groundwater, hypereutrophic surface water, secondary wastewater effluent, and algal organic matter. $SUVA_{254}$ decreased with decreasing molecular weight, but SFI near Peak C (λ_{ex} 337 nm, λ_{em} 423 nm) either remained relatively constant or increased. In samples from microbiologically active systems, the large molecular weight peak, commonly associated with biopolymers,[117] exhibited higher SFI near Peak T (λ_{ex} 278 nm, λ_{em} 353 nm) compared to Peak C.[118,269] Allpike *et al.* (2005) presented additional evidence that SFI near Peak T increased as both $SUVA_{254}$ and molecular weight decreased.[254]

Overall, there appears to be inconsistency in the logic relating the SFI near Peak C to aromaticity and molecular weight. Within aquatic systems, Fig. 8b and c shows increasing SFI with increasing aromatic carbon for a breadth of aquatic isolates. However, multiple lines of evidence support increasing SFI with decreasing molecular weight (Fig. 8d–f). Yet, the conventional interpretation is that more aromatic DOM is associated with larger molecular weights. This inconsistency could be explained by the increases in AQY as molecular weight decreases (*e.g.*, lower molecular weight DOM is less prone to deactivation of singlet excited states through intramolecular energy transfer).[49,52,53] Even though fewer photons are absorbed per unit carbon (*i.e.*, lower SUVA), the AQY increases to yield a higher SFI.

4.3 Humification index (HIX)

4.3.1 Definition and genesis. Although originally defined following multiple approaches (Table 3), Humification Index (HIX) had a unified goal of describing the relative abundance of red-shifted (*i.e.*, higher emission wavelength) fluorescence as an indicator of increased condensation and aromaticity.[260] The original studies focused on soil extracts, the impact of land use, and the relationship between soil organic matter and DOM in hydraulically connected waters.[227,259,260]

In the same year (1999), both Kalbitz *et al.*[260] and Zsolnay *et al.*[227] proposed independent definitions of HIX. Kalbitz *et al.* (1999) calculated peak ratios from synchronous fluorescence scans ($\Delta\lambda$ = 18 nm) proposing 3 different definitions ($HIX_{syn,400/360}$, $HIX_{syn,470/360}$, and $HIX_{syn,470/400}$). Synchronous scans are a two-dimensional spectrum that steps through a range of values for λ_{ex} and λ_{em}, while maintaining a constant offset ($\Delta\lambda = \lambda_{em} - \lambda_{ex}$), effectively measuring a diagonal cross-section of a three-dimensional EEM. $HIX_{syn,400/360}$ and $HIX_{syn,470/360}$ had a high correlation coefficient (r) between each other (r = 0.91), and comparisons with other characterization techniques used $HIX_{syn,470/360}$.[260] In a subsequent paper, the list expanded defining a new ratio at lower wavelengths for whole-waters ($HIX_{syn,390/355}$) compared to fulvic acids ($HIX_{syn,470/360}$); HIX values between whole-waters and paired isolates were strongly correlated ($\rho > 0.88$).[259] In contrast, Zsolnay *et al.* (1999) defined $HIX_{em,1999}$ by integrating and then dividing areas under two regions of an emission scan (Table 3).[227] The emission scan approach has since gained in popularity, in part due to the ease of collecting EEMs and availability of fluorometers with charge-coupled device detectors. Kalbitz *et al.* (2003) noted that $HIX_{em,1999}$ yielded stronger correlations with other

Table 3 Chronology and other calculation variations for HIX

Study	HIX definition
Kalbitz *et al.* (1999)[260]	$HIX_{syn,400/360} = \frac{I(382\text{ nm}, 400\text{ nm})}{I(342\text{ nm}, 360\text{ nm})}$
	$HIX_{syn,470/360} = \frac{I(452\text{ nm}, 470\text{ nm})}{I(342\text{ nm}, 360\text{ nm})}$
	$HIX_{syn,470/400} = \frac{I(452\text{ nm}, 470\text{ nm})}{I(382\text{ nm}, 400\text{ nm})}$
Zsolnay *et al.* (1999)[227]	$HIX_{em,1999} = \frac{\sum_{\lambda_{em,3}}^{\lambda_{em,4}} I(254\text{ nm}, \lambda_{em})}{\sum_{\lambda_{em,1}}^{\lambda_{em,2}} I(254\text{ nm}, \lambda_{em})}$
	Where
	$\lambda_{em,1}$ = 300 nm
	$\lambda_{em,2}$ = 345 nm
	$\lambda_{em,3}$ = 435 nm
	$\lambda_{em,4}$ = 480 nm
Kalbitz *et al.* (2000)[259]	$HIX_{390/355} = \frac{I(372\text{ nm}, 390\text{ nm})}{I(337\text{ nm}, 355\text{ nm})}$
Ohno (2002)[228]	$HIX_{em,2002} = \frac{\sum_{\lambda_{em,3}}^{\lambda_{em,4}} I(254\text{ nm}, \lambda_{em})}{\sum_{\lambda_{em,1}}^{\lambda_{em,2}} I(254\text{ nm}, \lambda_{em}) + \sum_{\lambda_{em,3}}^{\lambda_{em,4}} I(254\text{ nm}, \lambda_{em})}$
	Where
	$\lambda_{em,1}$ = 300 nm
	$\lambda_{em,2}$ = 345 nm
	$\lambda_{em,3}$ = 435 nm
	$\lambda_{em,4}$ = 480 nm

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physicochemical properties than the synchronous scan approaches.[287] Another definition ($HIX_{em,2002}$) was proposed by Ohno (2002) using the sum of both integrated areas in the denominator to mitigate inflating uncertainty when blue-shifted (*i.e.*, lower emission wavelength) intensities are low.[228] In a published comment, Zsolnay (2002) concurred that the revised approach has statistical advantages but also argued that it may compress the scale and decrease the sensitivity to differentiate samples with high HIX values.[288]

Both definitions are related to one another following eqn (11) with the derivation in ESI Text 3.† Regardless of the definition, the intent of the surrogate is similar; however, a meta-analysis of >700 samples showed that neither definition was strongly correlated with peak emission wavelength.[45]

$$HIX_{em,2002} = \frac{HIX_{em,1999}}{(1 + HIX_{em,1999})} \quad (11)$$

A challenge with some early studies is the lack of specificity regarding inner filter corrections. Some spectra were measured at constant DOC concentration but did not explicitly state if spectra were corrected for inner filter effects.[259,260] To minimize bias, Zsolnay *et al.* (1999)[227] diluted samples and applied primary inner filter corrections, but the correction approach is inconsistent with common practice today.[59,79] Independent of fluorescence characteristics, a sample with a higher spectral slope (or E2:E3) would introduce bias that increases HIX values. Systematic bias is introduced, because blue-shifted fluorescence would be suppressed due to secondary inner filtering. Since DOM absorbance spectra typically follow exponential decay, eqn (8) shows that the correction factor increases as wavelength decreases. If studies did not correct for inner filter effects, the reported correlations between HIX and chemical characteristics are biased to some extent, and it may not be possible to attribute relationships uniquely to fluorescence characteristics.

In the Kalbitz *et al.* (1999 and 2000) papers, the primary objective evaluated HIX_{syn} as a surrogate to either (1) characterize soil organic matter extracts or (2) understand the impact of land use on whole-water optical characteristics in hydraulically connected systems. Using fulvic acid isolates from water extracted soil, groundwater, and surface water, $HIX_{syn,470/360}$ positively correlated with $SUVA_{285}$ ($r = 0.85$) and inversely correlated with the carbon-to-nitrogen ratio (C/N) ($r = -0.63$).[260] The relative abundance of aromatic carbon (C=C stretch at 1620 cm^{-1}) and carboxyl groups (C=O stretch at 1725 cm^{-1}) was estimated using FTIR.[260] $HIX_{syn,470/360}$ was positively correlated to the relative intensity of aromatic carbon ($r = 0.71$) and inversely correlated to the ratio of aromatic to carboxyl intensities (1725 cm^{-1}/1620 cm^{-1}, $r = -0.69$).[260] These results supported $HIX_{syn,470/360}$ as a surrogate for aromaticity based on either direct (FTIR) or indirect ($SUVA_{285}$) measurements.[260] In Kalbitz *et al.* (2000),[259] HIX_{syn} for paired whole-water and fulvic acid isolates were more strongly correlated ($r = 0.89$) than $SUVA_{285}$ ($r = 0.62$). Additionally, $HIX_{syn,390/355}$ values for whole-waters positively correlated with aromatic carbon by FTIR ($r = 0.66$) and inversely correlated with the C/N ratio ($r = -0.48$).[259] Collectively, chemical composition correlated more strongly with HIX_{syn} than $SUVA_{285}$, rationalizing an advantage of fluorescence to infer differences in aromaticity without isolating fulvic acids.[259] One disadvantage of these studies is that FTIR is a semi-quantitative assessment of aromaticity, whereas SUVA has strong support from ^{13}C NMR measurements.[2]

Defining the emission spectra approach, Zsolnay *et al.* (1999)[227] assessed how fluorescence could differentiate DOM leached for different soil treatments. Fumigation with chloroform released DOM with blue-shifted fluorescence and a lower SFI relative to dried soil. The interpretation linking increased $HIX_{em,1999}$ to more condensed, polyaromatic carbon was made by inference from previous studies,[123,180,252,289,290] not direct evidence. Two of the previous studies,[180,252] also described in Section 4.1, reported shifts in peak emission wavelength for both UF and XAD fractionation. Importantly, Belin *et al.* (1993) cites previous work[256,291] for how aromaticity (primary) and molecular weight (secondary) are confounding factors in XAD fractionation. Also referenced, Kumke *et al.* (1995) reported a positive correlation between the wavelength of maximum fluorescence intensity and aromaticity for polycyclic aromatic hydrocarbons (PAHs), not humic substances.[289] Overall, the definition of $HIX_{em,1999}$ was not originally paired with a secondary, quantitative measure for aromaticity.

With soil microcosm experiments, Kalbitz *et al.* (2003) used ^{1}H NMR to demonstrate that $HIX_{em,1999}$ positively correlated with percent aromatic protons and negatively correlated with the proportion of oxygen-containing functional groups (Fig. 9a and b).[287] NMR data correlated more strongly with $HIX_{em,1999}$ than $HIX_{syn,460/345}$, strengthening the utility of the emission-based approach.[287]

4.3.2 Continued inquiry into aromaticity and molecular weight

4.3.2.1 Aromaticity. Since this initial cohort of studies, data from Kellerman *et al.* (2018)[35] presents the strongest evidence for $HIX_{em,1999}$ as a surrogate for aromatic carbon. Fig. 9c shows that $HIX_{em,1999}$ increased linearly with the relative abundance of condensed aromatic and polyphenol formulae. The positive intercept implies a positive $HIX_{em,1999}$ value is possible with no contributions from these aromatic formulae. However, electrospray ionization does not ionize all aromatic moieties in DOM,[143] and excitation without aromatic moieties is unlikely.

Fig. 9d and f compares HIX calculation methods using data from McKay *et al.* (2018).[39] For HPOA and fulvic acid isolates from aquatic systems, there is a positive correlation between HIX and aromaticity by ^{13}C NMR for both HIX calculation methods. Extrapolating $HIX_{em,1999}$ tracks towards PPFA, the only soil-derived fulvic acid in the dataset, which is not the case for $HIX_{em,2002}$. Similar to AQY (Fig. 7f), $HIX_{em,2002}$ is relatively insensitive to ^{13}C aromaticity for humic acids. Fig. 9f illustrates the sensitivity loss for $HIX_{em,2002}$ to differentiate isolates as described by Zsolnay (2002).[288] With the exception of marine isolates and the Antarctic endmember, $HIX_{em,2002}$ for all other isolates was between 0.92 and 0.98 and relatively insensitive to changes in ^{13}C aromaticity.

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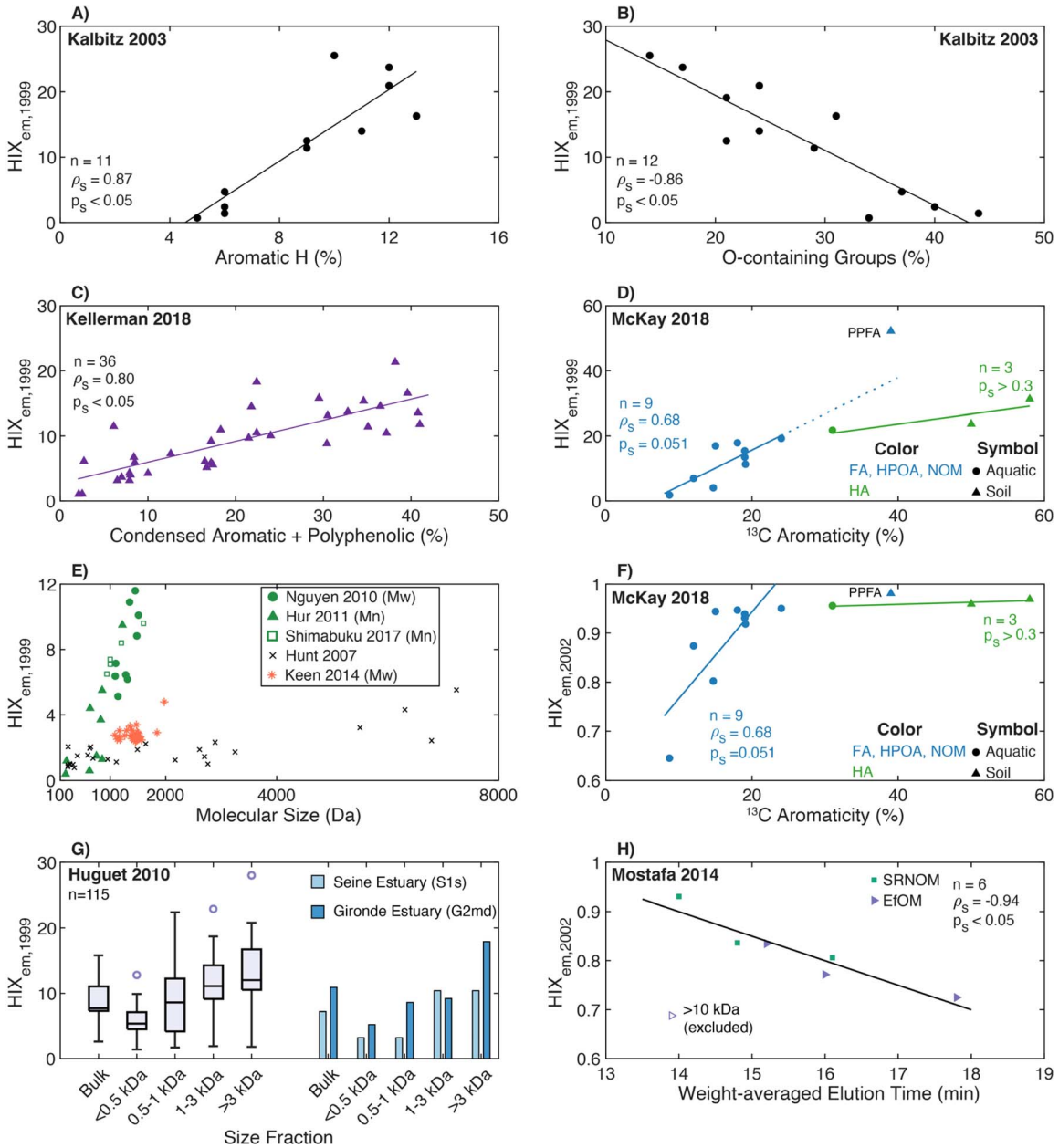

Fig. 9 Relationship between humification index (HIX) and either molecular weight or different measures of aromatic carbon. $HIX_{em,1999}$ *versus* (A) aromatic proton abundance and (B) oxygen-containing group abundance.[287] (C) $HIX_{em,1999}$ *versus* the relative abundance of condensed aromatic and polyphenolic formulae (%) per sample calculated by AI_{mod}[132] from FT-ICR MS data.[35] (D) $HIX_{em,1999}$ *versus* % aromaticity per sample calculated from ^{13}C NMR pairing optical[39] and NMR data.[40–43] (E) $HIX_{em,1999}$ *versus* molecular weight determined by SEC and UV detection at 254 nm.[292–296] Legend indicates calculation method (M_w *vs.* M_n) when available. (F) $HIX_{em,2002}$ *versus* % aromaticity per sample pairing optical[39] and ^{13}C NMR data.[40–43] (G) $HIX_{em,1999}$ *versus* UF fraction including a boxplot with all data and bar chart with two samples.[297] (H) $HIX_{em,2002}$ *vs.* weight-averaged elution time from SEC for samples fractionated by UF for a DOM isolate (SRNOM) and wastewater effluent (EfOM).[38] HIX values were not reported in the initial study but calculated from corrected EEMs. Applying a quality control approach,[45] $HIX_{em,2002}$ is not reported for SRNOM in the >10 kDa fraction due to spectral noise. For EfOM, the >10 kDa sample was excluded from the regression.

Several other papers demonstrate quantitative relationships between HIX and measures of aromatic carbon for narrower contexts. For example, Zhou *et al.* (2021) incubated 4 diverse DOM sources (*i.e.*, algal-, macrophyte-, sewage-, and soil-derived). Not only did $HIX_{em,1999}$ increase, but principal component analysis (PCA) identified strong associations between $HIX_{em,1999}$ and two FT-ICR MS measures of aromatic carbon (*i.e.*, polyphenolic abundance and AI_{mod}).[298] Using a HIX_{syn} approach, Hur *et al.* (2021) extracted humic acids from soils and sediments and showed that HIX_{syn} correlated positively with aromatic carbon and negatively with o-alkyl carbon.[299] In Kellerman *et al.* (2015) for Nordic lakes, $HIX_{em,2002}$ was strongly associated with aromatic, oxygen-rich compounds measured by FT-ICR MS.[300] Finally, Wagner *et al.*

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(2015) reported an association between $HIX_{em,2002}$ and less condensed aromatic structures, such as polyphenols, in the Florida Everglades.[183]

4.3.2.2 Molecular weight. Most studies have focused on demonstrating within-sample relationships using UF for narrow sample contexts. Huguet *et al.* (2010)[297] presented one of the largest datasets of UF fractionated samples with a narrow geographic context of the Seine and Gironde estuaries (France). Aggregating the entire data set, $HIX_{em,1999}$ generally increased as molecular weight increased (Fig. 9g). However, there is substantial overlap between fractions, where molecular weight could not be uniquely inferred from $HIX_{em,1999}$ alone. Contrasting endmembers in Fig. 9h, $HIX_{em,2002}$ was inversely correlated with weight-averaged elution time from SEC (UV_{254} detector) for <10 kDa fractions.[38] However, the largest fraction (>10 kDa) of EfOM was an outlier. A similar disconnect was reported for leaf litter extracts, where $HIX_{em,1999}$ decreased as molecular weight increased from 10 to 100 kDa.[301]

In other SEC studies using UV detection, Fig. 9e shows mixed trends. Quantitative comparisons between studies are challenging due to different SEC methods, calibration standards, and reporting conventions (M_w *vs.* M_n). Hur (2011) offers the most diverse dataset including EfOM, algal DOM, terrestrial DOM, and plant-derived DOM; $HIX_{em,1999}$ increased as M_n increased.[296] Similar positive correlations were reported for stormwater runoff,[292] sediments from two locations,[302] and a single source water subjected to UF fractionation and coagulation.[293] However, $HIX_{em,1999}$ was insensitive to molecular weight across different wastewaters[295] and leachates from plant biomass and manure residues.[294] Across a water treatment process (*e.g.*, nanofiltration, granular activated carbon) in Köhler *et al.* (2016),[303] $HIX_{em,2002}$ was positively correlated with the average molecular weight of the humic substance fraction, according to the Huber *et al.* (2011)[117] SEC method. Lee *et al.* (2019) coupled a fluorescence detector with SEC to calculate $HIX_{em,1999}$ from emission spectra taken at one minute intervals. Testing four IHSS isolates, elution profiles supported increasing $HIX_{em,1999}$ with increasing apparent molecular weight.[304]

As a whole, these studies reveal opportunities for further inquiry and convergence in the DOM community. Relationships between HIX and either aromaticity or molecular weight appear to have contextual limitations. Correlations for fulvic-acid dominated isolates (NOM, HPOA, or FA) may not be broadly applicable to whole-waters, EfOM, or humic acid isolates, of which the latter two both showed lower sensitivity to aromaticity. Within-sample trends appear to be stronger than comparisons between samples of different origins, urging caution against using differences in HIX to infer differences in molecular weight between sources. Between XAD fractionation (where non-humic substances are not recovered) and the reliance on UV detection for SEC (which does not detect non-chromophoric DOM), the aromaticity- and SEC-derived relationships in Fig. 9 are limited. SEC studies that combine DOC concentration and UV absorbance detection can help resolve the molecular weight fractions of DOM that are non-chromophoric.

4.4 Biological index, Freshness Index, or β/α

4.4.1 Definition and genesis. Using various names, Biological Index (BIX), Freshness Index, and the ratio of beta-to-alpha peaks (β/α) all recognize that fluorescence emission spectra often shift to lower emission wavelengths (blue shift) with increased biological activity. Parlanti *et al.* (2000)[229] described how fluorescence changes with eutrophication and the decomposition of algal material (exudates) using three fluorescence peaks and DOM samples from marine and freshwaters. Biological activity increased the fluorescence intensity of the β peak (λ_{ex} 310–320 nm, λ_{em} 380–420 nm, similar to Peak M in Coble (1996)[233]) and the γ peak (λ_{ex} 270–280, λ_{em} 300–320 nm, similar to Peak B), both of which were attributed to autochthonous DOM. In macro-algae degradation experiments, intensity of the allochthonous α peak (λ_{ex} 330–350, λ_{em} 420–480 nm, similar to Peak C) increased as the β peak decreased.[229] Therefore, a peak ratio (β/α) was proposed as a surrogate for "fresh" DOM from autochthonous sources relative to more "humified" material, which was described as degraded products that has been transformed through, "successive condensation reactions or structural arrangements".[229] Parlanti *et al.* (2000)[229] specified wavelength ranges rather than discrete wavelengths for each peak, creating ambiguity in the calculation method.

Huguet *et al.* (2009)[230] introduced the acronym BIX but interestingly did not define it as Biological Index, as commonly referred to today. The acronym was introduced as the "index of recent autochthonous contribution". The earliest known paper that attributed the name Biological Index to BIX was Carstea *et al.* (2009).[305] Reducing ambiguity, Huguet *et al.* (2009)[230] explicitly defined a calculation method for BIX (Table 2) as the ratio of two fluorescence intensities (380 to 430 nm) measured at 310 nm excitation. Huguet *et al.* (2009) also defined an interpretation scale, where BIX values between 0.6–0.7 have low autochthonous contributions and BIX value >1 have high contributions. That paper concluded that HIX and BIX are complementary indices and could capture different characteristics of DOM composition. Across an estuarine-marine gradient, HIX responded to changes in the DOM fluorescence at lower salinity ($S < 25$), where BIX was comparatively insensitive, yet both surrogates trended in opposite directions at higher salinity.

Wilson and Xenopoulos (2009)[231] applied the same concept using a different calculation method. Continuing to use the term β/α coined by Parlanti *et al.* (2000), Wilson and Xenopoulos (2009) modified the denominator to calculate the maximum fluorescence intensity between λ_{em} 420–435 nm (Table 2). They showed that β/α increased as both the percentage of continuous cropland and total dissolved nitrogen increased.[231] A meta-analysis by Korak and McKay (2024)[45] compared the fixed and variable λ_{em} calculation methods across >700 samples and showed good agreement at low BIX values with small, systematic deviations at higher values.

Of these three early papers, none of them used the term "Freshness Index", which is also commonly used today. To the best of our knowledge, the earliest reference for this term is the highly cited review article by Fellman *et al.* (2010).[267]

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4.4.2 Continued inquiry into aromaticity and molecular weight

4.4.2.1 Molecular weight. Of the three genesis papers,[229–231] none paired the optical data with independent measures of aromaticity or molecular weight. Huguet *et al.* (2010)[297] was the first to fractionate samples with UF. For the two estuarine gradients (*i.e.*, Seine and Gironde rivers), sample-specific trends generally held where BIX increased as molecular weight decreased (Fig. 10a). Aggregating the entire dataset, boxplots show that differences in BIX were most pronounced in the smallest molecular weight fractions (*i.e.*, <0.5 kDa and 0.5–1 kDa) with no difference between the 1–3 and >3 kDa fractions (Fig. 10a). Since previous work[1,297] has shown more than 50% of DsOM in freshwaters, quantified as DOC concentration, is typically associated with a >1 kDa fraction, we caution using BIX to infer differences in molecular weight between whole-water samples. Since the Huguet *et al.* (2010)[297] study, there is a noticeable gap in the literature pairing BIX with independent characterization techniques, recognizing an opportunity for future work to prioritize diverse DOM sources.

Following one source water across an engineering treatment gradient, Köhler *et al.* (2016) presented an inverse, linear relationship between BIX and the relative abundance of the humic substance fraction,[117] as determined by SEC with DOC concentration detection. A multilinear regression was presented to predict the concentration of low molecular weight neutral compounds using $HIX_{em,2002}$ and β/α as predictors.[303]

4.4.2.2 Aromaticity. Like HIX, data from Kellerman *et al.* (2018)[35] presents the strongest evidence across a large gradient of environmental systems of an inverse relationship between BIX and relative abundance of condensed aromatic and polyphenolic formulae (Fig. 10b). A similar inverse correlation between BIX and ^{13}C aromaticity was observed for aquatic isolates (Fig. 10c).[39] In contrast to Fig. 7F for AQY, neither YHPOA nor PPFA were outliers compared to the other isolates (Fig. 10c). YHPOA followed the same inverse correlation as other aquatic isolates. Unlike $HIX_{em,1999}$, the soil-derived fulvic acid (PPFA) followed a similar trend as the humic acids. Notably, the slope was shallower for humic acids compared to aquatic, fulvic acid-dominated isolates. The disconnect between sample groups in Fig. 10c urges caution between generalizing interpretations across contexts.

Although these analyses provide strong evidence for BIX as a surrogate for aromatic carbon, the context is dominated by isolates (Fig. 10b and c), with NOM isolates obtained using reverse osmosis underrepresented compared to XAD isolates.[35,39] Since BIX was originally defined as a surrogate for microbial exudates, differences in DOM recovery by XAD resins may hinder generalizations to whole-waters due to material not recovered during XAD isolation. For example, fulvic acids have been attributed to <20% of whole-water DOM (measured as DOC) in autochthonous Antarctic lakes compared to 30–60% of DOM more broadly across freshwaters.[1,2,306] As a result, aromaticity generalizations to whole-waters may not be well-supported unless strong agreement is observed between paired BIX measurements of whole-waters and isolates.

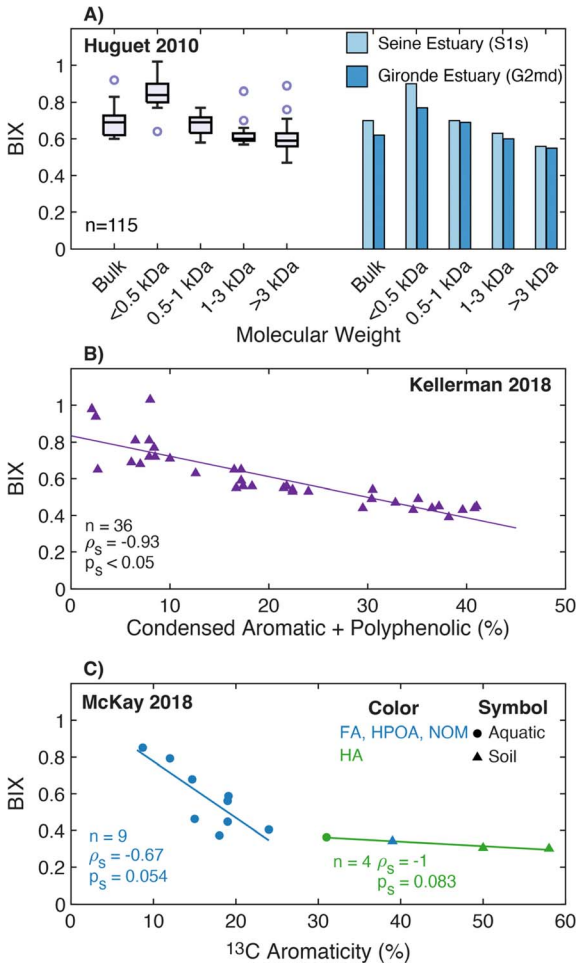

Fig. 10 Relationship between biological index (BIX) and either molecular weight or different measures of aromatic carbon. (A) BIX *versus* UF fraction including a boxplot with all data and bar chart with two samples.[297] (B) BIX *versus* the relative abundance of condensed aromatic and polyphenolic formulae (%) per sample calculated by AI_{mod}[132] from FT-ICR MS data.[35] (C) BIX *versus* % aromaticity per sample calculated from ^{13}C NMR.[39–43]

4.5 Fluorescence index (FI)

4.5.1 Definition and genesis. McKnight *et al.* (2001) originally proposed fluorescence index (FI) as a metric to assess the origin and aromaticity of aquatic fulvic acids using an emission scan at excitation 370 nm.[232] There are two calculation variations (eqn (12) and (13)). The 2001 paper[232] followed eqn (12) calculating the ratio of emission intensities at 450 to 500 nm. Since fulvic acids from Antarctica were blue-shifted relative to SRFA, 450 nm was proposed as a peak emission wavelength between the two endmembers. A 2010 follow-up study[84] revised the emission wavelengths to 470 and 520 nm (eqn (13)) with the incorporation of instrument-specific correction factors. Since spectral corrections are standard practice,[79] the 2010 calculation method is recommended (Table 2).

$$FI_{450/500} = \frac{I(370\ \text{nm}, 450\ \text{nm})}{I(370\ \text{nm}, 500\ \text{nm})} \quad (12)$$

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$$FI_{470/520} = \frac{I(370\ nm, 470\ nm)}{I(370\ nm, 520\ nm)} \quad (13)$$

The first objective in McKnight *et al.* (2001) was to leverage differences in the chemical composition of fulvic acids (*e.g.*, aromaticity) to differentiate sources using a fluorescence surrogate. Previous work showed that aromaticity was higher for terrestrial (25–30%) compared to microbial (12–17%) fulvic acids,[306,307] which would change fluorescence spectra.[308] The second objective assessed if FI of the whole-water was a good surrogate for the isolated fulvic acid. Spanning samples collected from the from the contiguous U.S. to the dry valleys of Antarctica, isolates from microbial endmembers (without higher plant inputs) had higher FI values (1.7–2.0) compared to terrestrial endmembers (1.3–1.4) with higher plant inputs.[232]

McKnight *et al.* (2001) presented direct evidence of an inverse relationship between FI and ^{13}C aromaticity for aquatic fulvic acids (Fig. 11a). Although there is a strong relationship, the study urged caution using FI as a generalized, quantitative surrogate for two reasons. First, extrapolating the regression models would yield an unrealistic, negative aromaticity at FI values >2.1, which were also observed in whole-water samples from Antarctic lakes[232] (Fig. 11b) and more recently in wastewater effluents.[38,295] Second, although the correlation was strong for fulvic acids, generalizing to whole-waters would require verifying that the regression holds with the non-humic fraction included.

Both studies[84,232] advocated for reporting FI with the peak emission wavelength to confirm that FI increased as peak emission wavelength decreased. Although not originally plotted, Fig. 11b shows a strong inverse correlation between FI and peak emission wavelength with continuity across the fulvic acids and whole-waters in the 2001 paper.[232] Lastly, a post-hoc analysis of the 2001 data supports a positive correlation between peak emission wavelength and ^{13}C aromaticity for aquatic fulvic acids ($r = 0.86$, $n = 18$). Croue *et al.* (2000)[125] concentrated (RO and nanofiltration) and fractionated (XAD) four surface waters, also reporting a positive correlation between ^{13}C aromaticity and peak emission wavelength at excitation 320 nm ($R^2 = 0.61$, $n = 25$).

4.5.2 Continued inquiry into aromaticity and molecular weight

4.5.2.1 Aromaticity. Using the relative abundance of condensed aromatic and polyphenolic formulae measured by FT-ICR MS, Kellerman *et al.* (2018)[35] also shows that relative abundance decreased as FI increased (Fig. 11a). However, scatter in the data increased for relative abundances <20%. Optical data from McKay *et al.* (2018) shows a similar trend (Fig. 11a). Unlike AQY, HIX, or BIX, Fig. 11a and b for McKay *et al.* (2018) show an internally consistent correlation spanning the entire dataset from aquatic fulvic acids through soil humic acids. Although each of these studies show internally consistent correlations, generalizing to whole-waters requires caution due the limited context of XAD isolation, which does not recover the non-humic fraction known to contribute to whole-water fluorescence.[238]

4.5.2.2 Molecular weight. In the genesis studies,[84,232] FI was not proposed as a surrogate for molecular weight. However, studies have recognized that DOM aromaticity and molecular weight often correlate.[253,256,291] For example, Fig. 11d shows

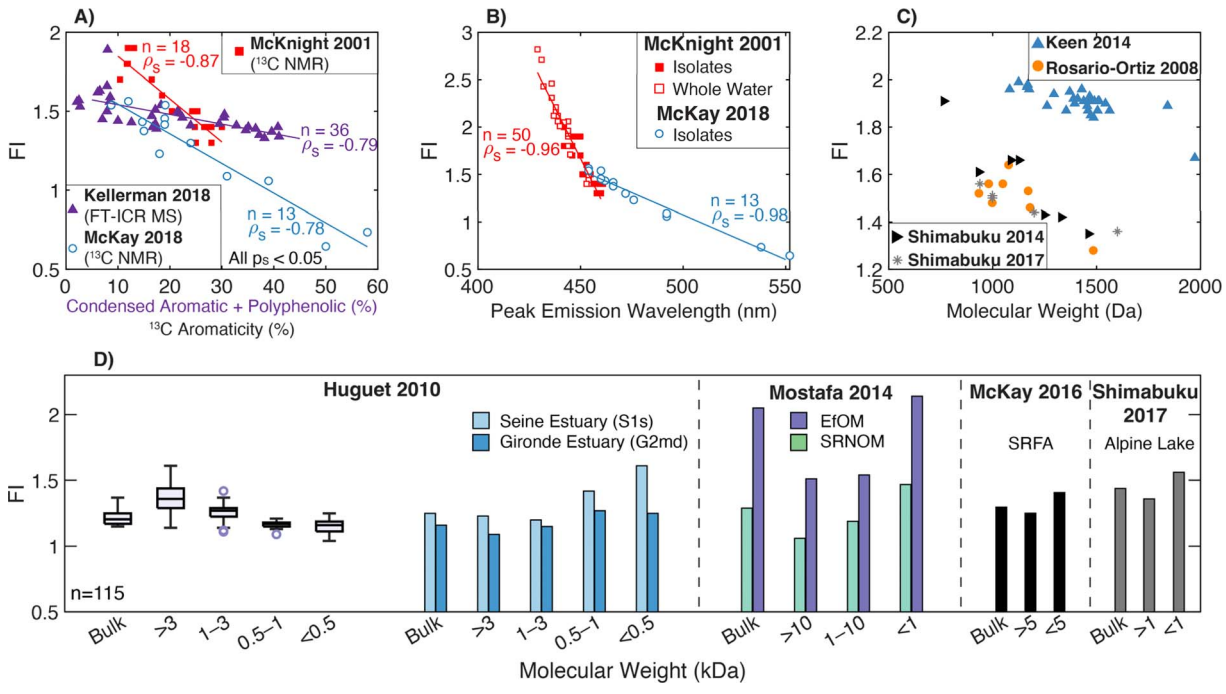

Fig. 11 Relationship between fluorescence index (FI) and either molecular weight, peak emission wavelength, or different measures of aromatic carbon. (A) FI *versus* % aromaticity per sample calculated from ^{13}C NMR[39–43,232] or the relative abundance of condensed aromatic and polyphenolic formulae (%) per sample calculated by AI_{mod}[132] from FT-ICR MS data.[35] (B) FI *versus* the peak emission wavelength at excitation wavelength of 370 nm.[39,232] (C) FI *versus* weight-averaged molecular weight (M_w) by SEC with a UV_{254} detector.[293,295,309,323] (D) FI *versus* nominal molecular weight fraction following UF fractionation.[38,44,293,297]

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internal consistency within single samples fractionated by UF, where FI increased as molecular weight decreased for riverine and estuarine samples,[297] an alpine lake,[293,309] SRNOM,[38] and wastewater effluent.[38] Other papers showed similar consistency.[253,310–312] Despite internal consistency for single samples, Fig. 11d illustrates that within a nominal size fraction, a wide range of FI values should be expected between source materials, and inverse interpretations of molecular weight between waters should not be made using FI. For example, the 1–10 kDa fraction of wastewater effluent had the same FI as the <1 kDa fraction from SRNOM. Fig. 11c contrasts two studies that used the same SEC instrument (UV detection) but different calibration standards.[295,309] In Shimabuku *et al.* (2014), there was an internally consistent gradient across a fractionated sample that also aligned with other surface waters.[309] However, wastewater samples from 8 facilities in Keen *et al.* (2014) had a comparatively smaller change in FI relative to the range of M_w values.[295] Romera-Castillo *et al.* (2014) used preparative SEC to fractionate 4 samples from the Florida Everglades into 7 fractions. FI increased with decreasing size except for the smallest fraction where signals were near baseline.[313]

Interestingly, two studies by Shimabuku and colleagues[293,309] showed that FI may contradict other lines of evidence regarding aromaticity and molecular weight. Using granular activated carbon (GAC) column tests, effluent at the beginning of the run (low bed volumes) had both lower $SUVA_{254}$ and FI, which contradicts expectations of an inverse correlation between surrogates.[293] In addition, SEC characterization of influent and effluent samples found that M_w was inversely related to FI ($R^2 = 0.91$) but not $SUVA_{254}$. Connecting optical surrogates to the role of molecular sieving in GAC separation, the study postulated that FI may be more dependent on molecular weight than aromaticity. Another study[309] drew similar conclusions using coagulation and powdered activated carbon after observing greater competition between DOM and micropollutants in source waters with high FI. These studies corroborated results from Schreiber *et al.* (2005) that also defied conventional interpretations, where GAC preferentially removed small molecular weight chromophores but spectral slope increased.[314] As a whole, the interplay of chemical sorption *versus* physical sieving by GAC may offer a unique opportunity to question causation *versus* correlation for optical surrogates and their dependence on aromaticity and molecular weight.

5 Conclusions and future opportunities

This critical review traced the genesis and assessed the continued inquiry of absorbance and fluorescence spectroscopy as surrogates for DOM aromaticity and molecular weight. Our analysis raises several findings that warrant further study.

First, more caution is needed when interpreting specific optical surrogates as unique indicators of either aromaticity or molecular weight. There are patterns in the literature of specific interpretations gaining popularity, such as $SUVA_{254}$ as a surrogate for aromaticity *versus* E2:E3 as a surrogate for molecular weight. However, many of the genesis papers (*e.g.*, Chin *et al.* (1994)[154] and Peuravuori and Pihlaja (1997)[155]) and other studies present evidence that the optical surrogates correlated with both characteristics. For example, several reports have shown that goethite, a common iron mineral, preferentially sorbed high molecular weight, chromophoric DOM. The non-sorbed fraction had lower molecular weight, lower SUVA, and higher E2:E3 values.[315,316] Sharpless *et al.* (2014) found that photooxidation decreased both $SUVA_{280}$ for aquatic fulvic acid isolates and the amount of spectral tailing (*i.e.*, increased E2:E3).[317] A lack of specificity between surrogates and DOM chemistry suggests that aromaticity and molecular weight are correlated variables in many environmental systems. The expectation that optical surrogates can distinguish aromaticity from molecular weight is unreasonable.

Second, many optical surrogates have been developed using isolates, and generalizing to whole-waters needs further investigation. For ^{13}C NMR and FT-ICR MS analyses of aromatic carbon, isolation processes are necessary to concentrate DOM and remove inorganic constituents. No isolation method recovers 100% of organic carbon, and there are criticisms that exposure to high pH transforms organic matter (*e.g.*, ester hydrolysis).[209] For example, Croue *et al.* (2000) showed that raw waters with low $SUVA_{254}$ had lower DOC recovery using XAD resins.[125] Using both XAD-8 and XAD-4 resins, recovery of whole-water DOC was 60–70% for low $SUVA_{254}$ waters ($\sim$2.5 L mg_C^{-1} m^{-1}), which was lower than the 80–90% recovery for high $SUVA_{254}$ waters (4.5–5.0 L mg_C^{-1} m^{-1}).[125] For FT-ICR MS analysis, solid phase extraction with PPL cartridges is a common isolation method, but DOC recoveries, reported by Dittmar *et al.* (2008)[149] as only $\sim$60%, are not always reported across the research field. There is an opportunity to investigate optical surrogate relationships with both aromatic carbon and molecular weight for PPL extracted DOM. Across solid phase extraction methods (*i.e.*, XAD, PPL), the absence of the unrecovered fraction in advanced chemical characterization (*i.e.*, NMR, FT-ICR MS) questions whether the chemical characteristics are truly explanatory. Surrogate relationships established using isolates may be over generalized if applied to contexts that would have low recovery by similar isolation methods (*e.g.*, wastewater).[318] Future research can focus on isolation methods with higher DOC recoveries (*e.g.*, RO) and evaluate how well surrogate relationships hold across waters from diverse sources and processes, incorporating both endmembers and a gradient in between.

Third, most optical surrogates used to describe molecular weight are derived from fractionating individual samples and not comparing different sources. For example, the study by Peuravuori and Pihlaja (1997)[155] is highly cited, but generalizations are drawn by fractionating only three starting materials. In this review, many of the bar charts depicting UF fractionated samples show systematic trends between optical surrogates and fractions from one starting material. However, these trends do not transcend across different source materials to enable inverse interpretations of chemistry from optical surrogates. For example, the biggest (>10 kDa) fraction of PLFA had a similar spectral slope as the intermediate (5–10 kDa) and

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smallest (<3 kDa) fractions of SRFA (Fig. 6e). There were also fewer studies in the literature that use chromatography-based size fractionation, like SEC or field-flow fractionation. Across the field, there is an opportunity to challenge these generalizations by interrogating if whole-water optical surrogates can be predictive measures of whole-water molecular weight distributions across a gradient of diverse source materials.

Fourth, optical surrogates broadly divide into two groups, as illustrated by the McKay *et al.* (2018)[39] data. One group shows a continuous trend from aquatic, fulvic acid-dominated isolates through soil humic acids (*e.g.*, SUVA, E2:E3, $S_{300–600}$, FI) in Fig. 4a, 5a, 6a, and 11a. However, the second group shows discontinuity between fulvic- and humic-acid isolates (*e.g.*, $S_{350–400}$, AQY, SFI for Peak C, $HIX_{em,1999}$, $HIX_{em,2002}$, BIX) in Fig. 6b, 7f, g, 8b, 9d, f, and 10c. This bifurcation has also arisen from experiments evaluating the solvatochromism of DOM fluorescence,[39] absorbance attenuation with sodium borohydride,[319] and fluorescence quenching with cationic nitroxides.[22] These discontinuities highlight the risk of generalizing trends between endmembers. For example, SFI near Peak C is lower for humic acids compared to fulvic acids (Fig. 8b), encouraging an interpretation that SFI decreases as aromaticity increases. However, this approach would miss the trend within the aquatic subset (fulvic acid, HPOA, and NOM) that would support a positive correlation between SFI and aromaticity. These groups identify opportunities for further investigation to understand how differences in underlying chemistry lead to these groupings.

Fifth, continued inquiry is needed to better understand surrogate variability in DOM samples with low aromatic carbon content. In the Kellerman *et al.* (2018)[35] dataset, plots between optical surrogates and formula relative abundance often show tighter relationships for the terrestrial endmembers, as compared to isolates with a low relative abundance of condensed aromatic and polyphenolic formulae (Fig. 5b; 6c; 8c; 10b). For size-fractionated samples, available data are more limited. In several instances, trends that held for terrestrial endmembers did not hold for size-fractionated PLFA or EfOM (Fig. 4c, f; 5c; 6e; 7c; 8f and 11d). The DOM community has recognized the need to invest resources in expanding isolate libraries to include EfOM and other anthropogenically impacted sources.[320] A concerted effort is needed to better understand when interpretations from more conventional water sources can be generalized across contexts, such is the case for the expanding interest in water reuse.

Lastly, data curation for this review identified a need for increased specificity of methods and the value of supporting open science. For several optical surrogates, there are multiple calculation methods that can be sources of ambiguity. There are also cases (*e.g.*, FI, BIX) where the calculation method has evolved since the genesis paper(s). The scope of future critical reviews could also be enhanced with more raw data available in machine-readable formats. These raw data would be most useful when accompanied by appropriate metadata, such as isolation method (and recoveries) if applicable, geographic location, sampling protocol, instruments used, surrogate calculation methods, and QA/QC procedures used during spectral measurements. As a community, we should push for more explicit calculation methods in supporting information and raw data reporting.

Overall, this critical review traced the genesis of two of the most common interpretations (*i.e.*, aromaticity and molecular weight) inferred from optical surrogates. We recognize that it could be quite unsatisfying to end a review without a concrete, convenient set of unambiguous recommendations specifying how to interpret optical surrogates. Aside from not recommending E4:E6, the absence of such broad recommendations reflects on the knowledge gaps that persist in understanding the linkages between DOM chemistry and optical surrogates. More broadly, there is ongoing debate about the underlying biogeochemical processes (*i.e.*, humification) that several optical surrogates were proposed to describe.[209,211] Instead, this review reveals that interpretations need to recognize limitations and be presented with appropriate nuance. One of the most important limitations is the need to validate interpretations before extending optical surrogates to entirely new contexts, like petrochemical contaminations or microplastics. Although FT-ICR MS has gained in popularity in the field, current practice (*i.e.*, negative ESI and PPL extraction) has an analytical window that does not necessarily overlap with chromophoric DOM. Forward progress in FT-ICR MS inquiries could include using multiple ionization modes (positive and negative), ionization techniques (*e.g.*, laser desorption ionization[144]), and solid phase extraction materials. We advocate for continued progress in DOM isolation techniques (existing and emerging) that better represent whole-water DOM in sufficient mass to facilitate solid-state ^{13}C NMR studies. Future research may also benefit from improved liquid-state NMR techniques for structural characterization.[321,322] We hope that this resource is valuable for consolidating past research trajectories and offers insight for future opportunities to strengthen our understanding and application of DOM optical measurements.

Data availability

No primary research results, software or code have been included and no new data were generated or analysed as part of this review. Data from existing publications that support this article have been consolidated and included as part of the ESI.†

Author contributions

JAK: conceptualization (equal), methodology (equal), data curation (equal), formal analysis (equal), visualization (lead), writing-original draft (lead), writing-review & editing (lead). GM: conceptualization (equal), methodology (equal), data curation (equal), formal analysis (equal), visualization (supporting), writing-original draft (supporting), writing-review & editing (supporting).

Conflicts of interest

There are no conflicts to declare.

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