


 Cite this: *Chem. Commun.*, 2025, 61, 17926

 Received 8th August 2025,  
Accepted 13th October 2025

DOI: 10.1039/d5cc04540a

rsc.li/chemcomm

## High stabilization of pentavalent uranium on magnetite nanoparticles evidenced by high-energy-resolution X-ray absorption spectroscopy

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**We observed that U(v) is more stable on magnetite nanoparticles than expected, across a wide range of pH and redox conditions and even after 10 days of reaction, as revealed by high-energy-resolution fluorescence-detection X-ray absorption spectroscopy combined with multivariate curve resolution.**

Uranium(U), widely used as a fuel for nuclear power generation, is a key element in actinide chemistry, nuclear technology, and environmental science due to its usefulness as an energy resource and its chemical toxicity.<sup>1–3</sup> In the environment, it has traditionally been considered to exist as U(IV) and U(VI). Although long overlooked,<sup>4,5</sup> recent studies have highlighted the significance of U(V), notably when U(VI) interacts with Fe(II)-bearing minerals, such as magnetite.<sup>6–11</sup> Magnetite is a ubiquitous mineral in natural settings and may also form in artificial barriers for nuclear waste disposal as a result of steel corrosion.<sup>12</sup> However, the role of U(V) remains unresolved as it is either considered as a metastable reaction intermediate during U(VI) reduction to U(IV) at magnetite surfaces<sup>9</sup> or as a stable species when incorporated *via* coprecipitation with magnetite.<sup>11</sup> As a result, the retention of U by magnetite has been intensively studied in recent years.<sup>9,11,13–15</sup> Nonetheless, the mechanisms controlling U interaction with magnetite in aqueous solution relevant to natural systems remain poorly

understood. In particular, many studies have employed high U concentrations, in which the precipitation of UO<sub>2(s)</sub> nanoparticles is often observed. In contrast, single-atom surface species formation may be a more relevant mechanism to investigate in many environments where U occurs at trace levels,<sup>5</sup> as was also recently shown for rhenium.<sup>16</sup>

Recently, high-energy-resolution fluorescence detection (HERFD)-XANES has been applied to a variety of U compounds as a more sensitive technique for analysing electronic structures.<sup>17–27</sup> The spectral features of HERFD-XANES are more distinct than those of conventional XANES spectra. For example, our group observed a specific peak splitting of U(V) in FeUO<sub>4</sub> with U L<sub>3</sub>-edge HERFD-XANES, due to the strong crystal field splitting of the 6d orbital, which could not be detected in normal XANES spectra.<sup>28</sup> This specific peak splitting is not observed in HERFD-XANES spectra of U(IV) in UO<sub>2</sub> because of its weaker crystal-field splitting than that of U(V) in FeUO<sub>4</sub>. Therefore, these promising results suggest that HERFD-XANES spectra could be a powerful technique to unravel the role played by U(V) when interacting with Fe-bearing minerals such as magnetite nanoparticles.

This paper presents insights into the influence of magnetite stoichiometry ( $0 \leq R = \text{Fe(II)}/\text{Fe(III)} \leq 0.5$ ) on the surface reduction of U(VI) to U(V) and U(IV), as a key parameter controlling U redox speciation in natural settings. Although R can readily change due to the oxidation of structural Fe(II) or proton/ligand-promoted dissolution,<sup>29–31</sup> prior studies have not quantified U(V) when assessing these effects.<sup>4,5</sup> To address this gap, we employed U L<sub>3</sub>-edge HERFD-XANES spectroscopy to investigate the electronic structure of U on magnetite with varying stoichiometries and observed a peak splitting in the HERFD-XANES spectrum of U on magnetite, an effect not previously detected using conventional XANES. In this study, we explore whether U(V) stabilization correlates with structural Fe(II) in magnetite by comparing U reference materials of U(IV)O<sub>2</sub>, FeU(V)O<sub>4</sub>, and U(VI)O<sub>2</sub>(NO<sub>3</sub>)<sub>2</sub>(H<sub>2</sub>O)<sub>6</sub>, for which HERFD-XANES analyses have already been reported.<sup>28</sup>

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Stock suspensions of 10 nm-sized stoichiometric ( $R = 0.5$ ) and non-stoichiometric ( $R = 0.1, 0.2, 0.3$  and  $0.4$ ) magnetite were synthesized and characterized as described in previous studies.<sup>31,32</sup> The U stock solution (0.5 mM) was prepared from uranyl acetate and dissolved in 0.01 M HCl solution. The uranium concentration (10  $\mu\text{M}$ , 2.3 ppm), background electrolyte concentration (10 mM NaCl), and total Fe concentration in the suspension (20 mM; equivalent to 1.1  $\text{g L}^{-1}$  for stoichiometric magnetite) were set equal for all experiments. As the surface area of 10 nm-sized magnetite nanoparticles is approximately  $100 \text{ m}^2 \text{ g}^{-1}$ , the maximum surface loading corresponds to  $0.1 \mu\text{mol m}^{-2}$  ( $0.06 \text{ atom nm}^{-2}$ ), which ensures that surface complexation was the predominant mechanism over U precipitation.<sup>5</sup> Batch experiments were conducted in 50 mL polypropylene tubes. Dilute HCl or NaOH solutions were used to adjust the solution pH. Adsorption and reduction experiments of U(VI) were performed (i) for 6 distinct pH values (4, 5, 6, 7, 8, and 9) onto stoichiometric magnetite (R0.5) and (ii) at pH 6 and 8 onto non-stoichiometric magnetite (R0.3 and R0.2) at room temperature under anaerobic condition ( $\text{N}_2$ -glovebox, JACOMEX,  $\text{O}_2 < 10 \text{ ppm}$ ). In addition, interaction of U(VI) with R0.1 was studied under ambient atmospheric conditions (*i.e.* in presence of  $\text{O}_{2(\text{g})}$ ); a reference system in which U(VI) is expected stable. After 10 days of reaction time, the solid phase was separated from the liquid phase by magnetic separation and dried in the glovebox. The dried samples were triple sealed using 100  $\mu\text{m}$  high-barrier film (AP-1522, ISO). The HERFD-XANES experiments were performed at beamline BL39XU, SPring-8, as previously reported.<sup>28</sup> A combined energy resolution of  $\sim 1.5 \text{ eV}$  was achieved in the experiments, as determined by measuring the full width at half-maximum (FWHM). Data processing was carried out using the Fastosh software<sup>33</sup> and Athena software.<sup>34</sup> Detailed experimental conditions are described in the SI.

The HERFD-XANES spectra at the U  $L_3$  edge of U interacting with R0.5 at pH 6 (denoted R0.5 pH6), collected by monitoring the  $L\alpha_1$  emission line, and the corresponding XANES spectrum collected in the conventional fluorescence mode are compared in Fig. 1a. The white line at approximately 17176 eV in the conventional XANES spectrum of R0.5pH6 mainly originates from the electron transition from U  $2p_{3/2}$  to unoccupied 6d states. The HERFD-XANES spectrum of the same sample, the



**Fig. 1** (a) U  $L_3$  edge HERFD-XANES spectrum (red line) and XANES spectrum (black line) of U-magnetite samples prepared at pH 6. (b) Solid line U  $L_3$  edge HERFD-XANES spectrum of U-magnetite samples prepared at pH 6. (Dotted line) HERFD-XANES spectra of  $\text{UO}_2$  (black),  $\text{FeUO}_4$  (red), and  $\text{UO}_2(\text{NO}_3)_2 \cdot 6\text{H}_2\text{O}$  (blue), previously published by Yomogida *et al.*<sup>28</sup>

red line in Fig. 1a, revealed more distinct features: the white line was split into two peaks in the HERFD-XANES spectra. The HERFD-XANES spectra at the U  $L_3$  edge of R0.5 pH6,  $\text{UO}_2$ ,  $\text{FeUO}_4$ , and  $\text{UO}_2(\text{NO}_3)_2 \cdot (\text{H}_2\text{O})_6$  are compared in Fig. 1b. Peak splitting was observed in the HERFD-XANES spectrum of  $\text{FeUO}_4$ , whereas no peak splitting was observed in the HERFD-XANES spectra of  $\text{UO}_2$  and  $\text{UO}_2(\text{NO}_3)_2 \cdot (\text{H}_2\text{O})_6$ , indicating that a significant fraction of U is stabilized as U(V) species on magnetite.

Fig. 2a shows the HERFD-XANES spectra of U-R0.5 samples prepared at different pH values. Only one peak was observed at 17175 eV at pH 4. As the pH increased, the intensity of the first peak decreased, while the second peak around 17180 eV increased. Fig. 2b and c show the HERFD-XANES spectra of U-magnetite samples with different stoichiometries at pH 6 and 8, respectively. HERFD-XANES analysis revealed a gradual decrease in the second peak intensity (17180 eV) with decreasing stoichiometry. This trend was observed at both pH 6 (Fig. 2b) and pH 8 (Fig. 2c). These observations confirm previous observations that the U redox speciation is influenced by  $R$ , with higher extent of U(VI) reduction in Fe(II)-rich system.<sup>4</sup> The effect of pH, which jointly affects U redox speciation in solution and  $R$  with the redox potential ( $E_h$ ),<sup>30</sup> will be discussed in more detail later in this manuscript.

To quantify the number of U species on stoichiometric and non-stoichiometric magnetite, singular value decomposition (SVD) analysis was applied to HERFD-XANES spectra of U on magnetites using FASTOSH software.<sup>33</sup> Fig. S1 shows the dependence of the singular values on the number of components ( $n$ ), obtained by the SVD of the data matrices. Because the singular values were nearly zero for  $n > 4$ , we considered that 3 components were sufficient to analyze the experimental data. Fig. 3a–c show the pure XANES spectra of components 1–3, obtained by multivariate curve resolution with alternating least squares (MCR-ALS) analysis of the data.<sup>35</sup> Fig. 3a shows the XANES of component 1, which displays a broad white line peak.



**Fig. 2** U  $L_3$ -edge HERFD-XANES spectra of U reacted with (a) stoichiometric magnetite ( $R = 0.5$ ) nanoparticles at different pH, (b) magnetites with different stoichiometries at pH 6 and (c) pH 8.



**Fig. 3** Pure XANES spectra of (a) component 1, (b) component 2, and (c) component 3 in MCR-ALS analysis of the data matrix using the PureS method. (Dotted line) HERFD-XANES spectra of  $\text{UO}_2$  (black),  $\text{FeUO}_4$  (red), and  $\text{UO}_2(\text{NO}_3)_2 \cdot 6\text{H}_2\text{O}$  (blue), as previously published in Yomogida *et al.* Percentage of fractions of  $\text{U}(\text{iv})$ ,  $\text{U}(\text{v})$ , and  $\text{U}(\text{vi})$  versus (d) different pH, (e) different magnetite stoichiometry at pH 6, and (f) different magnetite stoichiometry at pH 8.

The features of the XANES of the component 1 are consistent with those of the  $\text{UO}_2$ . Fig. 3b shows the XANES the component 2, whose main peak is split and agrees with that of  $\text{FeUO}_4$ . Fig. 3c shows the XANES spectrum of component 3, whose main peak is the narrowest but agrees relatively well with that of  $\text{UO}_2(\text{NO}_3)_2(\text{H}_2\text{O})_6$ . Thus, we assign the U surface species on magnetite to the 3 relevant oxidation states of U.

Fig. 3d–f show the distribution (%) of the 3 components obtained by MCR-ALS. Linear combination fitting (LCF) of XANES spectra of U on magnetite, using reference spectra of  $\text{UO}_2$ ,  $\text{FeUO}_4$  and  $\text{UO}_2(\text{NO}_3)_2(\text{H}_2\text{O})_6$ , also provided results similar to MCR-ALS (Fig. S2). On stoichiometric magnetite, approximately half of the U bound to magnetite at  $\text{pH} > 6$  was  $\text{U}(\text{v})$  (Fig. 3d). Thermodynamic speciation calculations in solution may support this observation,  $E_{\text{H}}$  values determined in R0.5 suspensions<sup>30</sup> approach the predominant field of  $\text{U}(\text{v})$  for  $\text{pH} 5.5\text{--}7.5$  (Fig. S3). However, such calculations do not account for the formation of surface species, which generally favour tetra- and hexavalent oxidation states of actinides ( $\text{An} = \text{U}, \text{Np}, \text{Pu}$ ) over the pentavalent ones. Indeed, actinides(V) weakly sorb to iron oxide surfaces because of the smaller effective charge of cation ( $\text{U}(\text{v})\text{O}_2^+ < \text{U}(\text{vi})\text{O}_2^{2+} < \text{U}^{4+}$ ).<sup>36,37</sup> In addition, for  $\text{pH} > 7.5$ , the theoretical prevalence of aqueous  $\text{U}(\text{iv})$  contrasts with the observed prevalence of  $\text{U}(\text{v})$  at the magnetite surface. Therefore, this unexpectedly high stability of  $\text{U}(\text{v})$  is difficult to explain solely by surface complexation processes. This may

instead result from its favorable incorporation into the magnetite structure. Indeed, spectroscopic studies evidenced that  $\text{U}(\text{v})$  was highly stable when coprecipitated with Fe during magnetite synthesis,<sup>6</sup> or that other ions such as  $\text{Ni}(\text{ii})$ ,<sup>38</sup>  $\text{Co}(\text{ii})$ ,<sup>39</sup> or  $\text{Cr}(\text{iii})$ <sup>32</sup> incorporated into magnetite during sorption experiments. The XANES spectra of  $\text{U}(\text{v})$  (Fig. 3b) is almost consistent with  $\text{U}(\text{v})$  in  $\text{FeUO}_4$ ,  $\text{U}(\text{v})$  may partially substitute octahedral site in magnetite.<sup>6</sup> At  $\text{pH} < 6$ , magnetite can partially reduce  $\text{U}(\text{vi})$  to  $\text{U}(\text{iv})$ , but  $\text{U}(\text{v})$  is absent. In fact, significant amounts of structural  $\text{Fe}(\text{ii})$  are released from magnetite as  $\text{Fe}^{2+}$  ion to solution below  $\text{pH} 6$ ,<sup>30</sup> leading to a decrease in the magnetite stoichiometry. Since non-stoichiometric magnetites are less efficient at incorporating metal ions (e.g.  $\text{Ni}(\text{ii})$ ,<sup>38</sup>  $\text{Co}(\text{ii})$ <sup>39</sup> and  $\text{Cr}(\text{iii})$ <sup>32</sup>), this process may explain the absence of  $\text{U}(\text{v})$  stabilisation. At  $\text{pH} 6$  (Fig. 3e), the presence of  $\text{U}(\text{iv})$  at intermediate stoichiometries, but its absence replaced by  $\text{U}(\text{v})$  under the most reducing conditions ( $R = 0.5$ ), supports this interpretation. At  $\text{pH} 8$  (Fig. 3f), strong interaction of  $\text{U}(\text{iv})$  with magnetite, combined with unfavourable redox conditions for  $\text{U}(\text{v})$  aqueous species, allows its occurrence. However, these mechanisms and conditions are not sufficient to overcome  $\text{U}(\text{v})$  stabilisation, which occurs in similar proportions to  $\text{U}(\text{iv})$ . These results indicate that the stabilization of  $\text{U}(\text{v})$  species on magnetite is strongly related to  $\text{Fe}(\text{ii})$  in magnetite structures, not only for the reduction process, but also for the  $\text{U}(\text{v})$ -magnetite binding mechanism.

In a previous study involving conventional XANES, Latta *et al.*<sup>4</sup> reported that reduction of  $\text{U}(\text{vi})$  to  $\text{UO}_2$  occurred only for  $R > 0.38$  at  $\text{pH} 7.2$ . Our results show the presence of  $\text{U}(\text{v})$  species on magnetite and that the  $\text{U}(\text{vi})$  to  $\text{U}(\text{v})$  reduction occurred at lower magnetite stoichiometry ( $R > 0.2$ ). Besides the more difficult detection of  $\text{U}(\text{v})$  by using normal XANES spectroscopy when  $\text{U}(\text{vi})$  and  $\text{U}(\text{iv})$  co-occur, this difference might be caused by the experimental conditions, such as the use of pH-buffer (50 mM MOPS or  $\text{NaHCO}_3$ ) that might limit  $\text{U}(\text{v})$  adsorption, and the higher U/Fe ratio (5 g  $\text{L}^{-1}$  magnetite and 500  $\mu\text{M}$  U) that may favour the prevalence of  $\text{U}(\text{iv})$  in the solid phase due to the enhanced precipitation of  $\text{UO}_2$ . Our results suggest that the chemical species of U and its concentration in solution may also contribute to differences in the U adsorption mechanism on magnetite.

In conclusion, the present study provides new insights into interactions between U and magnetite nanoparticles regarding the influence of magnetite stoichiometry and pH on the U speciation. HERFD-XANES spectra of  $\text{U}(\text{v})$  on magnetites show a peak splitting that has not been observed in conventional XANES, which enable us to demonstrate the high stability of  $\text{U}(\text{v})$  species under a wide range of conditions, after 10 days on magnetite by the combination of MCR-ALS analysis. A key finding is that structural  $\text{Fe}(\text{ii})$ , whose abundance depends on pH and redox conditions, plays a critical role in the stabilization of  $\text{U}(\text{v})$  on magnetite. Compared with previous studies using conventional XANES, our results with the novel application of HERFD-XANES reveal that the reported  $\text{U}(\text{vi})$  reduction of  $\text{UO}_2$  only occurs for  $R > 0.38$  at  $\text{pH} 7.2$ ,<sup>4</sup> highlighting the importance of high-resolution spectroscopy in accurately

determining complex redox interactions at mineral surfaces. Revealing U(v) species, potentially acting a *meta*-stable species such as temporary sink or a reservoir for slow re-oxidation, on a widespread mineral like magnetite may be important to better predict the fate of U in reducing environments.

This work was also supported by the Grants-in-Aid for Scientific Research (KAKENHI) from JSPS (Grant No. 25K08534). The study was conducted with the approval of JASRI/SPRING-8 (Proposal No. 2021B1821, 2022A1733, 2024B1496, and 2024B1905). Additional support was provided by the COLOSSAL project funded by ANR (project number ANR-23-CE01-0001; PI: R.M.), as well as the French Brittany Region SAD program COLORED.

## Conflicts of interest

There are no conflicts to declare.

## Data availability

The data supporting this article have been included in the IPGP Research Collection See DOI: <https://doi.org/10.18715/IPGP.2025.mgb3600v>.

Supplementary information (SI): details on sample preparation, XANES data acquisition, data analysis, and speciation modeling results. See DOI: <https://doi.org/10.1039/d5cc04540a>.

## Notes and references

- 1 P. L. Smedley and D. G. Kinniburgh, *Appl. Geochem.*, 2023, **148**, 105534.
- 2 A. Y. Romanchuk, I. E. Vlasova and S. N. Kalmykov, *Front. Chem.*, 2020, **8**, 630.
- 3 Y. Takahashi, A. Yamaguchi and T. Yomogida, in *Treatise on Geochemistry*, ed A. Anbar and D. Weis, Elsevier, Oxford, 3rd edn, 2025, pp. 105–150.
- 4 D. E. Latta, C. A. Gorski, M. I. Boyanov, E. J. O'Loughlin, K. M. Kemner and M. M. Scherer, *Environ. Sci. Technol.*, 2012, **46**, 778–786.
- 5 Z. Wang, K.-U. Ulrich, C. Pan and D. E. Giammar, *Environ. Sci. Technol. Lett.*, 2015, **2**, 227–232.
- 6 H. E. Roberts, K. Morris, G. T. W. Law, J. F. W. Mosselmans, P. Bots, K. Kvashnina and S. Shaw, *Environ. Sci. Technol. Lett.*, 2017, **4**, 421–426.
- 7 O. Stagg, K. Morris, A. Lam, A. Navrotsky, J. M. Velázquez, B. Schacherl, T. Vitova, J. Rothe, J. Galanzew, A. Neumann, P. Lythgoe, L. Abrahamsen-Mills and S. Shaw, *Environ. Sci. Technol.*, 2021, **55**, 16445–16454.
- 8 S. Kushwaha, B. Sreedhar and P. Padmaja, *Langmuir*, 2012, **28**, 16038–16048.
- 9 K. Yuan, E. S. Ilton, M. R. Antonio, Z. Li, P. J. Cook and U. Becker, *Environ. Sci. Technol.*, 2015, **49**, 6206–6213.
- 10 C. Dewey, D. Sokaras, T. Kroll, J. R. Bargar and S. Fendorf, *Environ. Sci. Technol.*, 2020, **54**, 6021–6030.
- 11 Z. Pan, B. Bartova, T. LaGrange, S. M. Butorin, N. C. Hyatt, M. C. Stennett, K. O. Kvashnina and R. Bernier-Latmani, *Nat. Commun.*, 2020, **11**, 4001.
- 12 C. J. Dodge, A. J. Francis, J. B. Gillow, G. P. Halada, C. Eng and C. R. Clayton, *Environ. Sci. Technol.*, 2002, **36**, 3504–3511.
- 13 K. Yuan, D. Renock, R. C. Ewing and U. Becker, *Geochim. Cosmochim. Acta*, 2015, **156**, 194–206.
- 14 H. Veeramani, D. S. Alessi, E. I. Suvorova, J. S. Lezama-Pacheco, J. E. Stubbs, J. O. Sharp, U. Dippon, A. Kappler, J. R. Bargar and R. Bernier-Latmani, *Geochim. Cosmochim. Acta*, 2011, **75**, 2512–2528.
- 15 Y. Wang, J. Wang, P. Li, H. Qin, J. Liang and Q. Fan, *Environ. Tech. Innov.*, 2021, **23**, 101615.
- 16 R. Ding, C. Guida, C. I. Pearce, E. Arenholz, J.-M. Grenèche, A. Gloter, A. C. Scheinost, K. O. Kvashnina, K. Wang, A. Fernandez-Martinez, Y. Mu, K. M. Rosso and L. Charlet, *Sci. Adv.*, 2025, **11**, eadq3650.
- 17 T. Vitova, K. O. Kvashnina, G. Nocton, G. Sukharina, M. A. Denecke, S. M. Butorin, M. Mazzanti, R. Caciuffo, A. Soldatov, T. Behrends and H. Geckeis, *Phys. Rev. B: Condens. Matter Mater. Phys.*, 2010, **82**, 235118.
- 18 K. O. Kvashnina, S. M. Butorin, P. Martin and P. Glatzel, *Phys. Rev. Lett.*, 2013, **111**, 253002.
- 19 K. O. Kvashnina and F. M. F. de Groot, *J. Electron Spectrosc. Relat. Phenom.*, 2014, **194**, 88–93.
- 20 K. O. Kvashnina, Y. O. Kvashnin and S. M. Butorin, *J. Electron Spectrosc. Relat. Phenom.*, 2014, **194**, 27–36.
- 21 J. G. Tobin, S. W. Yu, C. H. Booth, T. Tylyszczak, D. K. Shuh, G. van der Laan, D. Sokaras, D. Nordlund, T. C. Weng and P. S. Bagus, *Phys. Rev. B: Condens. Matter Mater. Phys.*, 2015, **92**, 035111.
- 22 R. Bes, M. Rivenet, P. L. Solari, K. O. Kvashnina, A. C. Scheinost and P. M. Martin, *Inorg. Chem.*, 2016, **55**, 4260–4270.
- 23 G. Leinders, R. Bes, J. Pakarinen, K. Kvashnina and M. Verwerft, *Inorg. Chem.*, 2017, **56**, 6784–6787.
- 24 I. Pidchenko, K. O. Kvashnina, T. Yokosawa, N. Finck, S. Bahl, D. Schild, R. Polly, E. Bohnert, A. Rossberg, J. Gottlicher, K. Dardenne, J. Rothe, T. Schafer, H. Geckeis and T. Vitova, *Environ. Sci. Technol.*, 2017, **51**, 2217–2225.
- 25 L. Desfougeres, E. Welcomme, M. Ollivier, P. M. Martin, J. Hennuyer, M. Hunault, R. Podor, N. Clavier and L. Favregeon, *Inorg. Chem.*, 2020, **59**, 8589–8602.
- 26 R. Bes, G. Leinders and K. Kvashnina, *J. Synchrotron Radiat.*, 2022, **29**, 21–29.
- 27 K. O. Kvashnina, Y. O. Kvashnin, J. R. Vegelius, A. Bosak, P. M. Martin and S. M. Butorin, *Anal. Chem.*, 2015, **87**, 8772–8780.
- 28 T. Yomogida, D. Akiyama, K. Ouchi, Y. Kumagai, K. Higashi, Y. Kitatsuji, A. Kirishima, N. Kawamura and Y. Takahashi, *Inorg. Chem.*, 2022, **61**, 20206–20210.
- 29 P. Jungcharoen, M. Pédrot, F. Choueikani, M. Pasturel, K. Hanna, F. Heberling, M. Tesfa and R. Marsac, *Environ. Sci.: Nano*, 2021, **8**, 2098–2107.
- 30 P. Jungcharoen, M. Pédrot, F. Heberling, K. Hanna, F. Choueikani, C. Catrouillet, A. Dia and R. Marsac, *Environ. Sci.: Nano*, 2022, **9**, 2363–2371.
- 31 P. Jungcharoen, R. Marsac, F. Choueikani, D. Masson and M. Pédrot, *Nanoscale Adv.*, 2023, **5**, 4213–4223.
- 32 J. Scaria, M. Pédrot, L. Fablet, T. Yomogida, T. T. Nguyen, Y. Sivry, C. Catrouillet, A. E. Pradas Del Real, F. Choueikani, D. Vantelon, A. Dia, A. Groleau and R. Marsac, *Environ. Sci. Technol.*, 2025, **59**, 5747–5755.
- 33 G. Landrot and E. Fonda, *J. Synchrotron Radiat.*, 2025, 1085–1094.
- 34 B. Ravel and M. Newville, *J. Synchrotron Radiat.*, 2005, 537–541.
- 35 W. Windig and D. A. Stephenson, *Anal. Chem.*, 1992, **64**, 2735–2742.
- 36 H. Geckeis, J. Lützenkirchen, R. Polly, T. Rabung and M. Schmidt, *Chem. Rev.*, 2013, **113**, 1016–1062.
- 37 R. Marsac, C. Catrouillet, M. Pédrot, M. F. Benedetti, A. Dia, E. D. van Hullebusch, M. Davranche, Y. Sivry, A.-C. Pierson-Wickmann, M. Tharaud and F. Heberling, *Curr. Opin. Colloid Interface Sci.*, 2024, **72**, 101820.
- 38 L. Fablet, M. Pédrot, F. Choueikani, I. Kieffer, O. Proux, A.-C. Pierson-Wickmann, V. Cagniard, T. Yomogida and R. Marsac, *Environ. Sci.: Nano*, 2025, **12**, 2815–2827.
- 39 L. Fablet, F. Choueikani, M. Pédrot and R. Marsac, *Environ. Sci.: Nano*, 2024, **11**, 2036–2048.