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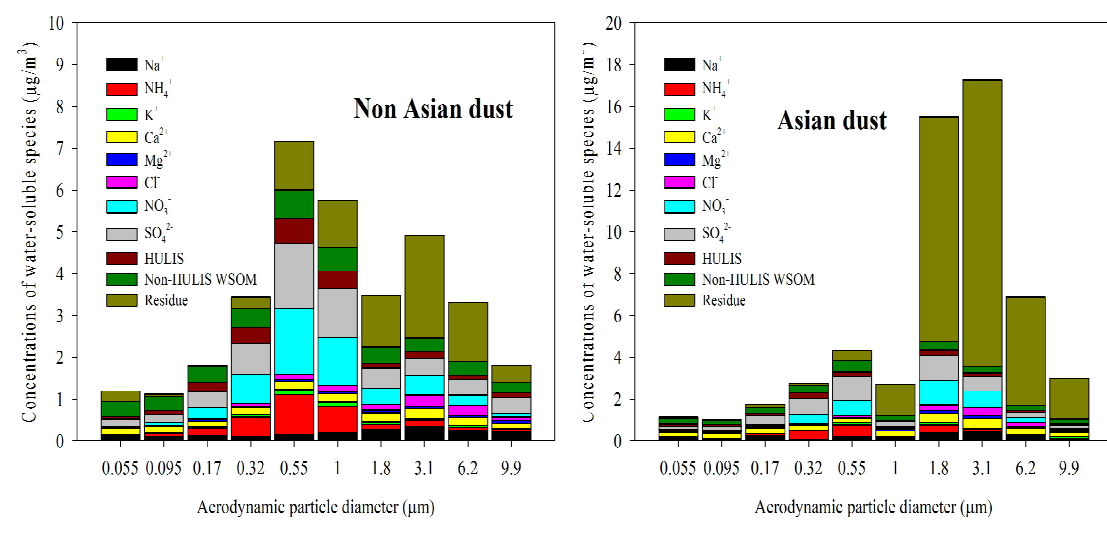
Size distribution and sources of humic-like substances in particulate matter at an urban site during winter

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Highest contribution of HULIS-C to WSOC was observed to be in the particle size bins of 0.55-1.0 μm and 1.8-3.1 μm during non-Asian dust (NAD, 45 \pm 6%) and Asian dust (AD, 44 \pm 7%) periods, respectively. HULIS exhibited uni-modal (@0.55 μm) distribution during the NAD and bimodal (@0.32 and 1.8 μm) during AD, respectively,



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60**Abstract**

This study investigates the size distribution and possible sources of humic-like substances (HULIS) in ambient aerosol particles collected at an urban site in Gwangju, Korea during the winter of 2015. A total of 10 sets of size-segregated aerosol samples were collected using a 10-stage Micro-Orifice Uniform Deposit Impactor (MOUDI), and the samples were analyzed to determine the mass as well as the presence of ionic species (Na^+ , NH_4^+ , K^+ , Ca^{2+} , Mg^{2+} , Cl^- , NO_3^- , and SO_4^{2-}), water-soluble organic carbon (WSOC) and HULIS. The separation and quantification of the size-resolved HULIS component from the MOUDI samples was accomplished using a Hydrophilic-Lipophilic Balanced (HLB) solid phase extraction method and a total organic carbon analyzer, respectively. The entire sampling period was divided into two periods: non-Asian dust (NAD) and Asian dust (AD) periods. The contributions of water-soluble organic mass ($\text{WSOM} = 1.9 \times \text{WSOC}$) and HULIS ($=1.9 \times \text{HULIS-C}$) to fine particles ($\text{PM}_{1.8}$) were approximately two times higher in the NAD samples (23.2 and 8.0%) than in the AD samples (12.8 and 4.2%). However, the HULIS-C/WSOC ratio in $\text{PM}_{1.8}$ showed little difference between the NAD (0.35 ± 0.07) and AD (0.35 ± 0.05) samples. The HULIS exhibited a uni-modal size distribution ($@0.55 \mu\text{m}$) during NAD and a bimodal distribution ($@0.32$ and $1.8 \mu\text{m}$) during AD, which was quite similar to the mass size distributions for particulate matter, WSOC, NO_3^- , SO_4^{2-} , and NH_4^+ in both the NAD and AD samples.

The size distribution characteristics and the results of the correlation analyses indicate that the sources of HULIS varied according to the particle size. In the fine mode ($\leq 1.8 \mu\text{m}$), the HULIS composition during the NAD period was strongly associated with secondary organic aerosol (SOA) formation processes similar to those of secondary ionic species (cloud processing and/or heterogeneous reactions) and primary emissions from biomass burning period, and during the AD period, it was only associated with SOA formation. In the coarse mode (3.1-10 μm), it was difficult to identify the HULIS sources during the NAD period, and during the AD period, the HULIS was mostly likely associated with soil-related particles [$\text{Ca}(\text{NO}_3)_2$ and CaSO_4] and/or sea-salt particles (NaNO_3 and Na_2SO_4).

Keywords: Size-resolved ambient aerosols, humic-like substances, solid phase extraction, size distribution, secondary organic aerosol.

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Environmental impact

Results from size-segregated aerosol measurements of HULIS concentrations at an urban site during winter indicate that the HULIS-C (on a carbon basis)/WSOC ratio in fine particles showed little difference between the non-Asian dust (NAD) and AD periods, but the HULIS showed uni-modal during the NAD and bimodal size distributions during AD, respectively. Also the HULIS in the fine mode was strongly associated with both secondary organic aerosol (SOA) formation processes and primary biomass burning emissions during the NAD period, while it was associated with the SOA formation during the AD period.

92 Introduction

93 The organic compounds that account for a significant amount of ambient particulate
94 matter¹ can be categorized into water-insoluble and water-soluble fractions and may affect the
95 radiative balance of the Earth.^{2,3} Moreover, water-soluble organic compounds can act as
96 cloud condensation nuclei to facilitate cloud formation. The water-insoluble organic carbon
97 fraction is primarily generated as a result of the incomplete combustion of fossil fuels and
98 biomass material^{1,4} and is obtained by extraction with n-hexane, dichloromethane, acetone,⁵
99 or benzene.⁶ In contrast, the water-soluble organic carbon (WSOC) fraction is emitted
100 directly from primary combustion sources, or formed in the atmosphere through
101 transformation processes.⁷ The primary sources of WSOC include biomass burning (BB)
102 emission,^{1,8-12} traffic emissions (with small contributions),¹²⁻¹⁵ and residual oil combustion.¹⁶

103 Solid phase extraction techniques can be used to further separate WSOC into more
104 hydrophilic and more hydrophobic fractions.¹⁷⁻¹⁹ Humic-like substances (HULIS) are
105 important class of substances in the hydrophobic WSOC fraction,^{20,21} and these consist of a
106 mixture of high molecular weight compounds²² with the ability to suppress the surface
107 tension of aerosol and fog water.^{23,24} The physical and chemical properties of HULIS are
108 similar to those of humic and fulvic acids that naturally occur in terrestrial and aquatic
109 environments.^{20,25,26} They are thought to consist of polycyclic aromatic and aliphatic
110 structures with oxygenated functional groups, such as hydroxyl, carboxyl, carbonyl, nitrate,
111 and nitroxy organosulfate groups.^{18,27-29} HULIS play an important role in the hygroscopic
112 growth and cloud condensation nuclei formation of aerosols;^{24,30,31} significantly contribute to
113 light absorption;^{32,33} affect the aqueous-phase oxidation of organic pollutants;³⁴ and promote
114 the production of reactive oxygen species under simulated physiological conditions.³⁵

115 HULIS accounts for a significant fraction of WSOC in ambient aerosols, and numerous
116 studies have shown that BB emissions^{16,21,33,36-39} and secondary formation processes^{16,21,39} are
117 important sources of HULIS in ambient aerosol particles. In addition, marine aerosols⁴⁰ and
118 residual oil combustion¹⁶ have also been identified as possible sources of HULIS. Typically,
119 the fraction of HULIS-C (on a carbon basis) in WSOC tends to increase in ambient samples
120 due to the contribution of BB emissions,^{21,38,41} and this fraction has also been found to be
121 higher in the winter than in the summer, possibly due to the increase in BB activities in the
122 winter.^{16,39} Previous studies have shown that the fraction of HULIS-C in ambient PM_{2.5} is

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4 123 highly variable, accounting for 9-72% of the total WSOC.^{16,20,21,36,39,42,43} Also, BB emissions
5 124 have a lower fraction (~30%) of HULIS-C in WSOC in PM_{2.5}.^{21,41}
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7 125 The size distribution of the chemical constituents in aerosol particles can provide
8 126 important information on their sources, formation, and evolution in air. Size-resolved WSOC
9 127 have been extensively measured in ambient environments⁴⁴⁻⁴⁹ and in biomass burning
10 128 emissions in controlled chambers in laboratories.⁵⁰⁻⁵² The size distributions of the WSOC in
11 129 ambient aerosols and fresh/aged BB emissions mostly exhibit a dominant droplet mode
12 130 peaking at 0.55-1.0 μm .^{45,46,52,53} Until now, HULIS measurements have been mainly made for
13 131 bulk aerosol particles, such as TSP, PM₁₀, and PM_{2.5} at various sampling environments.
14 132 However, a limited amount of information is available on the size distribution of HULIS in
15 133 ambient aerosols. In one example from a rural site in Southern China, the size distribution of
16 134 ambient HULIS showed a dominant droplet mode peaking at 0.63-0.87 μm .²¹ Droplet-mode
17 135 HULIS was attributed to secondary formation through cloud-processing and heterogeneous
18 136 reactions or aerosol-phase reactions, and to the growth of fresh BB particles.

19 137 There is a lack of data on the size distribution of HULIS in ambient air, especially for
20 138 East Asia (including Korea), and the chemical characteristics and sources/processes of size-
21 139 resolved HULIS from these regions are poorly understood. Thus, an investigation of the size
22 140 distribution of ambient HULIS particles can improve our understanding of the sources and
23 141 formation processes of HULIS as a hydrophobic part of WSOC. To this end, this study
24 142 collected 48-hr size-segregated aerosol samples from an urban site in Gwangju, Korea during
25 143 February 2015, and the samples were analyzed in terms of mass, water-soluble inorganic ions,
26 144 total WSOC, and HULIS. This study aims to determine the mass size distribution of HULIS
27 145 as well as to explore the likely sources and formation processes for size-resolved HULIS. In
28 146 addition, differences in the size distribution and sources of HULIS for non-Asian dust and
29 147 Asian dust samples are also discussed.

148 149 150 **Experimental**

151 **Measurements of size-segregated ambient aerosols samples**

152 A 10-stage Micro-Orifice Uniform Deposit Impactor (MOUDI, MSP 110; MSP Corp., MN)
153 was used to collect size-segregated ambient aerosol samples between February 3 and
154 February 28, 2015, at an urban site (35°11'N, 126°54'E) in Gwangju, Korea. The sampling
155 site was located on rooftop of a three-story building (54.3 m above sea level) in a university,

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4 156 about 80 m from a traffic road. Each stage of MOUDI has 50% cut sizes ranging from 0.055
5 157 to 18.0 μm in an aerodynamic diameter at a sampling flow rate of 30 l/min. The available cut-
6 158 off diameters of the MOUDI are 0.055, 0.095, 0.17, 0.32, 0.55, 1.00, 1.8, 3.1, 6.2, 9.9, and
7 159 18.0 μm . The MOUDI samples were collected on pre-baked and pre-weighed 47-mm quartz-
8 160 fiber filters. A 47-mm Teflon substrate was used as a back-up filter, but the WSOC and
9 161 HULIS were not analyzed due to very low mass deposited on the back-up filter. Aerosol
10 162 sampling was conducted for about forty-eight hours to provide reliable results for the quantity
11 163 of size-resolved HULIS. According to the time series of hourly PM_{10} mass concentration
12 164 (Figure 1), an Asian dust (AD) storm event occurred between 22 and 24 February at the site
13 165 (<http://web.kma.go.kr/eng/weather/asiandust/timeseries>), so 24-hr aerosol samples were
14 166 collected for AD. A total of 10 sets of MOUDI samples were made, with 2 sets out of 10
15 167 consisting of AD samples.

16 168 Ambient aerosol pollution in the city has been previously found to be a result of local
17 169 emissions, regional pollution, and AD storms.^{15,54-58} Since the Korean peninsula is located
18 170 downwind from China, anthropogenic and natural aerosols from China contribute to
19 171 particulate matter pollution over the Korean peninsula. The details of the sampling site were
20 172 described in our earlier work,^{12,59} and the collected MOUDI samples were thus used to
21 173 determine the mass, water-soluble ionic species, total water-soluble organic carbon (WSOC),
22 174 and HULIS-C ($\mu\text{g C}/\text{m}^3$) concentrations. Field blanks were also analyzed, and their
23 175 background values were corrected to obtain the actual values of ambient aerosol particles at
24 176 each impactor stage.

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26 178 **Chemical analysis of water-soluble components**

27 179 Quartz-fiber and Teflon filters were weighed before and after collecting the samples by using
28 180 a microbalance with a 1- μg sensitivity (Sartorius CP2P-F). The filters were conditioned for
29 181 24 h in a clean chamber that was maintained at a relative humidity of 40% with a temperature
30 182 of 20 °C. The samples were weighed, and then size-segregated samples were extracted with
31 183 40 mL of ultrapure water (18.2 M Ωcm , Barnstead Nanopure ultrapure water system, Thermo
32 184 Scientific, USA) in an ultrasonic bath at room temperature for 60 min. The extracts were
33 185 filtered with a 0.45 μm membrane filter (Millipore) to remove insoluble materials before the
34 186 analysis. The presence of eight ionic species (Na^+ , NH_4^+ , K^+ , Ca^{2+} , Mg^{2+} , Cl^- , NO_3^- , and SO_4^{2-})
35 187 was determined using an ion chromatography (IC) system (Metrohm 861), and the amount of
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4 188 WSOC was also analyzed using a total organic carbon (TOC) analyzer (Sievers 5310C, USA).
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6 189 Detailed descriptions of the IC and TOC measurements were given in our previous
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8 190 studies.^{15,60,61}

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10 191 In this study, a Hydrophilic-Lipophilic Balanced (HLB) solid phase extraction (SPE)
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12 192 method based on the one used by Lin *et al.*²¹ was applied to isolate the HULIS from the water
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14 193 extracts. HULIS was isolated in the water extracts by using a HLB SPE cartridge (Oasis HLB,
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16 194 30 μm , 60 mg/cartridge, Waters, USA) by following three procedures: pretreatment of the
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18 195 SPE cartridge, loading the extracts on the cartridge, and concentrating the eluate. First, the
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20 196 cartridge was pre-treated with 1 mL of methanol (Honeywell Burdick & Jackson, HPLC
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22 197 grade >99.9%) and 1 mL of ultrapure water, and this pretreatment cycle was repeated two
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24 198 times. It was then acidified with HCl (Daejung, Korea, purity 35.0%) to a pH \approx 2. Second,
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26 199 the water extracts were acidified to a pH \approx 2 by using HCl, and these were loaded onto the
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28 200 SPE cartridge. The cartridge sorbent was rinsed with 2 mL ultrapure water, and then the
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30 201 HULIS fraction was eluted from the SPE cartridge with 1.5 mL methanol containing 2%
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32 202 ammonia (w/w) (Daejung, Korea, purity 25-28%). Finally, the HULIS eluate was evaporated
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34 203 to dryness under a gentle stream of nitrogen and was re-dissolved in ultrapure water. The
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36 204 isolated HULIS fraction was then quantified using a TOC analyzer.

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205 206 **Recovery test of HULIS using the SPE method**

207 Water-soluble inorganic ions, low molecular weight organic acids (e.g., oxalic acid), and
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209 208 sugars (e.g., levoglucosan) are hydrophilic substances that pass through the SPE resin
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211 209 cartridge with little retention, and those that are retained on the cartridge are strongly related
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213 210 to HULIS, including the hydrophobic fraction of WSOC.^{16,19,21,59,60} The performance of the
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215 211 HLB SPE isolation method used in this study was evaluated by testing the water-soluble
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217 212 inorganic ions in ambient aerosol particles taken from a PM_{2.5} low-volume sampler at our
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219 213 sampling site. Aqueous solutions were analyzed three times before and after passing through
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221 214 the SPE cartridge for eight inorganic ions by using an IC to check that they had been retained
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223 215 by the SPE cartridge. The concentrations of the inorganic ionic species (Cl⁻, NO₃⁻, SO₄²⁻, Na⁺,
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225 216 K⁺, Ca²⁺, and Mg²⁺) were compared before and after the cartridge separation of water extracts
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227 217 for ambient aerosol samples, as shown in Figure 2. More than 99.0% of these ionic species
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229 218 (NO₃⁻, SO₄²⁻, Na⁺, K⁺, Ca²⁺, and Mg²⁺) had passed through the SPE cartridge, but low
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231 219 penetration efficiency (\sim 85%) for Cl⁻ was obtained. The low penetration efficiency of Cl⁻

220 compared to those of the other ions is probably due to the acidification of both the SPE
221 cartridge and the aqueous sample solution due to HCl.

222 In this study, the Suwannee River Fulvic Acid (SRFA) and Suwannee River Humic Acid
223 (SRHA) HULIS standards were used to evaluate the recovery on the HLB SPE resin cartridge.
224 The recovery experiments were carried out three times for each of the HULIS standard
225 solution concentrations that were tested. The results of the recovery tests for the two HULIS
226 standards are summarized in Table 1. The recovery efficiencies of the SRFA concentrations
227 that were tested (620-5070 $\mu\text{g C/L}$) from the HLB SPE cartridge ranged from 90.0 to 106.1%,
228 depending on the initial concentrations. For SRHA, the recovery efficiency also varied with
229 the initial concentration that was tested (530-3600 $\mu\text{g C/L}$), ranging from 90.0 to 100.9%.
230 Based on triplicate analyses of the same HULIS standard solution, the relative standard
231 deviation for SRFA and SRHA was estimated at <5.5% and <5.7%, respectively.

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234 Results and discussion

235 General characteristics of water-soluble components in fine and coarse particles

236 As previously mentioned, Asian dust storms from China covered the entire Korean peninsula
237 from February 22 to 24. Thus, the entire sampling period was divided into non-Asian dust
238 (NAD) and Asian dust (AD) periods. Since MOUDI does not have a cut-off diameter of 2.5
239 μm , a diameter of 1.8 μm was defined as the cut-off point to distinguish fine from coarse
240 particles. Therefore, $\text{PM}_{1.8}$ and $\text{PM}_{3.1-10}$ respectively indicate fine and coarse particles. The
241 mass fraction of organic aerosol (OA) is important when assessing the chemical mass balance
242 closure of aerosol particles and the contribution of OA to the total PM mass. In this study, the
243 size-resolved reconstructed PM mass from the MOUDI measurements was calculated as
244 follows: Reconstructed PM mass ($\mu\text{g/m}^3$) = $\text{WSOM} + \text{Na}^+ + \text{NH}_4^+ + \text{K}^+ + \text{Ca}^{2+} + \text{Mg}^{2+} + \text{Cl}^-$
245 $+ \text{SO}_4^{2-} + \text{NO}_3^-$. Since composition data for elemental carbon (EC), organic carbon (OC), and
246 elemental constituents is not available, the size-resolved PM mass was determined by using
247 the concentration of water-soluble organic and inorganic species only. Based on the results
248 reported in the literature,^{20,21,42} a conversion factor of ~ 1.9 was applied to estimate the
249 concentrations ($\mu\text{g/m}^3$) of WSOM (water-soluble organic mass) and HULIS from the WSOC
250 and HULIS-C concentrations in $\mu\text{g C/m}^3$. With the exception of the MOUDI data collected
251 during the AD period, a comparison of the measured and reconstructed size-resolved PM
252 mass concentrations showed an excellent correlation with a slope of 0.74 and an R^2 of 0.85

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4 253 (not shown here). The low regression slope (0.74) was attributed to the absence of EC, OC,
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6 254 and crustal elements.

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8 255 The average PM₁₀ concentrations for the NAD and AD periods were 34.2±13.6 and
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10 256 57.7±24.4 µg/m³, respectively (Table 2). The average fine particle (PM_{1.8}) concentration
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12 257 accounted for 69.2% (65.6-77.2%) of the PM₁₀ for NAD and 53.9% (50.7-57.2%) for AD,
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14 258 respectively. The low fraction of the fine mode particle concentration during the AD period
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16 259 was a result of the higher concentration of crustal material. The mass fractions of WSOC (as
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18 260 WSOM), HULIS-C (as HULIS), and secondary inorganic species (SIS=NO₃⁻+SO₄²⁻+NH₄⁺)
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20 261 in the PM_{1.8} are summarized in Table 2. The contributions of WSOM and HULIS to PM_{1.8}
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22 262 were approximately two times higher in the aerosol samples collected during NAD than those
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24 263 collected during AD, and their contributions were 23.2 and 8.0% for NAD and 12.8 and 4.2%
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26 264 for AD, respectively. The HULIS fraction during NAD was comparable to the results
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28 265 reported in other studies. The HULIS in PM_{2.5} from the Pearl River Delta (PRD) region in
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30 266 Southern China had been found to constitute 8.5-11.7% of the samples.^{16,21} Little difference
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32 267 was observed in the fraction of HULIS-C to WSOC in PM_{2.5} between the NAD (35.0±6.9%)
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34 268 and AD (34.7±5.2%) periods. Results from previous studies indicate that the high HULIS
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36 269 levels and high fraction of HULIS-C in WSOC could be attributed to significant influences of
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38 270 the BB aerosol emissions.^{21,38,41} The HULIS-C fraction was observed to have a lower value
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40 271 of ~30% in fresh BB aerosols,^{21,41} and our results have values that are quite lower than those
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42 272 for the urban and suburban sites (48-60%) in the PRD region,^{16,21} where the influence of BB
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44 273 emissions is significant during the sampling period, and those in urban and rural sites (54-
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46 274 68%) in Europe.^{20,38,62} The low HULIS fractions during the NAD and AD periods in this
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48 275 study suggest a small contribution from BB emissions, and the concentration of SIS in fine
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50 276 particles was higher in the NAD (51.2±3.7%) than in AD (34.4±14.8%), suggesting that
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52 277 aerosol particles were more atmospherically processed during NAD than during AD.

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278 279 **Size-resolved chemical composition of water-soluble components**

280 Figure 3 shows the average size-resolved chemical composition of water-soluble components
281 from the NAD and AD samples. The “residue” in the Figure 3 indicates the sum of EC,
282 water-insoluble OC, and mineral dust mass concentrations, which is the size-resolved total
283 water-soluble matter concentration subtracted from the measured size-resolved PM
284 concentration. The non-HULIS WSOM is calculated as the difference between WSOM and

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4 285 HULIS. In the case of the NAD period, the contribution of non-HULIS WSOM in the
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6 286 ultrafine mode ($<0.1 \mu\text{m}$) was dominant. The secondary NO_3^- , SO_4^{2-} , and NH_4^+ components
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8 287 were dominant contributors in the fine mode ($0.17\text{-}1.8 \mu\text{m}$), followed by non-HULIS WSOM
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10 288 and HULIS. In the coarse mode ($>3.1 \mu\text{m}$), NO_3^- , SO_4^{2-} , and non-HULIS WSOM were
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12 289 significant water-soluble components, with the exception of the unresolved chemical
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14 290 components. However, for the AD period, non-HULIS WSOM and SO_4^{2-} were the main
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16 291 contributors to the ultrafine mode PM while SO_4^{2-} , NO_3^- , NH_4^+ , non-HULIS WSOM, and
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18 292 HULIS contributions were dominant in the fine mode ($0.17\text{-}1.8 \mu\text{m}$). In the coarse mode,
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20 293 NO_3^- , SO_4^{2-} , non-HULIS WSOM, and Ca^{2+} were significant contributors. Figure 4 shows the
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22 294 size-resolved relative contributions of HULIS-C to WSOC for the NAD and AD periods. The
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24 295 average contribution of HULIS-C to WSOC in the ambient particles was of $33.5\pm 6.1\%$ for
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26 296 NAD and $33.6\pm 7.3\%$ for AD. However the highest contribution of HULIS-C to WSOC for
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28 297 the NAD and the AD periods was observed for a particle size cut of $0.55\text{-}1.0 \mu\text{m}$ ($44.9\pm 6.4\%$)
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30 298 and for a size cut of $1.8\text{-}3.1 \mu\text{m}$ ($43.7\pm 7.4\%$), respectively. In the case of the NAD period, a
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32 299 high HULIS-C/WSOC existed mostly for fine particles, but the high contribution of HULIS-
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34 300 C for the AD period occurred for coarse mode particles as well as for fine particles due to the
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36 301 increased concentration of crustal material. Further details of this issue are provided below.

302 303 **Size distribution of HULIS**

304 Figure 5 shows the size distributions of the PM, WSOC, HULIS-C, and major water-soluble
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306 inorganic species (NO_3^- , SO_4^{2-} , NH_4^+ , and K^+) for both the NAD and AD periods. Similar size
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308 distributions in the PM and in other water-soluble chemical species were found for the NAD
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310 and AD samples in the fine particle mode, but in the coarse mode, different shapes of the size
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312 distribution were found. As shown in Figure 5, the mass size distributions of the WSOC and
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314 HULIS-C were similar to those of NO_3^- , SO_4^{2-} , NH_4^+ , and K^+ , suggesting that they were
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316 likely produced through similar atmospheric transformation processes and/or primary
emissions. For the NAD and AD periods, the WSOC and HULIS-C exhibited size
distributions with a dominant droplet mode peaking at $0.55 \mu\text{m}$ or $0.31 \mu\text{m}$. Also the NO_3^- ,
 SO_4^{2-} , NH_4^+ , and K^+ concentrations exhibited dominant droplet modes peaking with particle
sizes of $0.55 \mu\text{m}$. During the AD period, the PM and all chemical species peaked in the size
ranging from 1.8 to $3.1 \mu\text{m}$, as well as in the droplet mode, probably as a result of the long-
range transport of anthropogenic aerosols in polluted and/or industrialized regions of China

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4 317 and crustal material from the source regions of dust storms. Previous studies indicated that
5 318 Asian dust storms carry significant quantities of anthropogenic aerosols during transport in
6 319 addition to large amounts of crustal elements, resulting in the enrichment of anthropogenic
7 320 pollution in downwind locations.^{55,58,63}

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10 321 For NAD, the mass fractions of WSOM and HULIS in particle cut sizes of 0.55 and 0.32-
11 322 1.0 μm accounted for 19.0 and 45.7% of the total WSOM, respectively, and 25.1 and 58.4%
12 323 of the total HULIS, respectively. NO_3^- , SO_4^{2-} , and NH_4^+ contributions were higher than
13 324 WSOC and HULIS, accounting for 30.7, 25.1, and 34.0% of their total concentration at 0.55
14 325 μm , respectively, and for 66.6, 56.2, and 71.7% at 0.32-1.0 μm , respectively. When compared
15 326 to the contribution of water-soluble species concentrations collected from NAD samples, the
16 327 contributions to AD samples were observed to be rather low for fine mode particles. For the
17 328 AD, the size resolved mass fraction of WSOM to the total WSOM contributed 15.6 and 35.4%
18 329 at 0.55 and 0.32-1.0 μm , respectively. The HULIS at 0.55 and 0.32-1.0 μm contributed 14.2
19 330 and 35.8% of the total HULIS, respectively. NO_3^- , SO_4^{2-} , and NH_4^+ contributions to their total
20 331 concentrations were 19.3, 21.1, and 28.1% at 0.55 μm and 32.9, 39.9, and 51.3% at 0.32-1.0
21 332 μm , respectively.

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23 334 **Resolving the sources of fine and coarse mode HULIS**

24 335 Figure 6 shows the correlation of WSOC and HULIS-C with water-soluble inorganic species
25 336 for all size-resolved measurements during the NAD and AD periods. HULIS-C exhibited a
26 337 very high correlation ($R^2=0.73-0.94$) with WSOC for two periods, indicating their similar
27 338 chemical characteristics. However, a higher regression slope was obtained for HULIS-
28 339 C/WSOC from the NAD samples (0.59) than from the AD samples (0.39), suggesting
29 340 additional sources and/or further atmospheric processing of HULIS in size-segregated aerosol
30 341 samples measured during NAD. The measurement data indicates that the HULIS for the
31 342 NAD period was strongly associated with the biomass burning (BB) emission marker (K^+)
32 343 and the secondary inorganic species (NO_3^- , SO_4^{2-} , and NH_4^+), suggesting BB emissions and
33 344 secondary formation processes could be important sources of HULIS during NAD. However
34 345 the HULIS for the AD period was associated with secondary ionic species (NO_3^- , SO_4^{2-} , and
35 346 NH_4^+), suggesting its association with secondary formation processes.

36 347 To further identify HULIS sources in fine and coarse modes, the relationships between
37 348 HULIS-C and water-soluble inorganic ions (Na^+ , Cl^- , K^+ , Ca^{2+} , SO_4^{2-} , NO_3^- , and NH_4^+) were

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4 349 examined, and the results are summarized in Table 3. For the NAD period, HULIS-C was
5 350 strongly correlated with NH_4^+ , NO_3^- , SO_4^{2-} , and K^+ in the fine particle mode at an R^2 of 0.92,
6 351 0.88, 0.75 and 0.78, but was weakly correlated with water-soluble ions in the coarse mode.
7
8 352 The droplet modes of SO_4^{2-} and NO_3^- are known to be predominantly produced through
9 353 aqueous-phase reactions of SO_2 and NO_2 in clouds or in aerosol surfaces.⁶⁴ The strong
10 354 correlations of HULIS-C with NO_3^- , SO_4^{2-} , and K^+ in the fine mode for the NAD indicates
11 355 that the fine mode HULIS was likely produced from secondary transformation processes
12 356 similar to those of NO_3^- and SO_4^{2-} (i.e., aqueous-phase oxidation and/or cloud processing), as
13 357 well as from biomass burning emissions.^{21,41,45,53} However, it is hard to identify the sources of
14 358 low HULIS in the coarse mode during NAD due to poor correlations ($R^2=0.00-0.34$) of
15 359 HULIS-C with water-soluble ionic species in the coarse mode.

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22 360 For the AD period, fine mode HULIS-C was also correlated with secondary NH_4^+ , NO_3^- ,
23 361 and SO_4^{2-} at an R^2 of 0.56, 0.66, and 0.48, respectively, while the correlation between
24 362 HULIS-C and K^+ was not significant, suggesting that fine mode HULIS is mainly associated
25 363 with secondary formation processes similar to those of NO_3^- and SO_4^{2-} . However, unlike the
26 364 HULIS-C observed during the NAD, coarse mode HULIS-C observed during AD was
27 365 strongly correlated with Na^+ , Cl^- , Ca^{2+} , NO_3^- , and SO_4^{2-} at an R^2 of 0.84, 0.83, 0.67, 0.79, and
28 366 0.63, respectively. The coarse mode Na^+ and Cl^- are mostly from marine aerosols, and the
29 367 coarse mode Ca^{2+} is derived from soil CaCO_3 particles. Once they are emitted in the
30 368 atmosphere, the reaction of NaCl and CaCO_3 with HNO_3 and H_2SO_4 may occur during
31 369 transport.^{45,65} Previous studies indicated that the coarse mode WSOC is associated with either
32 370 sea salt particles (Na^+) or soil-related particles (Ca^{2+}).^{46,59} In the coarse mode, strong
33 371 correlations of Na^+ ($R^2=0.88$, $p<0.01$) and Ca^{2+} ($R^2=0.76$, $p<0.01$) with NO_3^- and good
34 372 correlations of Na^+ ($R^2=0.73$, $p<0.01$) and Ca^{2+} ($R^2=0.60$, $p<0.05$) with SO_4^{2-} suggest that the
35 373 coarse mode NO_3^- and SO_4^{2-} were likely associated with reactions from NaCl(s) and $\text{CaCO}_3(s)$
36 374 with $\text{HNO}_3(g)$ and $\text{H}_2\text{SO}_4(g)$. As a consequence, the results from these correlations and from
37 375 previous studies suggest that the coarse mode HULIS during AD was possibly condensed on
38 376 the surface of the NaNO_3 and Na_2SO_4 and/or the $\text{Ca(NO}_3)_2$ and CaSO_4 particles.

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52 378 **Summary and conclusion**

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54 379 Forty-eight hours size-resolved measurements of ambient aerosol particles were made at an
55 380 urban site during the period from February 3 through 28, 2015 in order to investigate the size
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4 381 distribution and sources of HULIS. The size-resolved HULIS from water extracts was
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6 382 isolated by using an HLB solid phase extraction and was quantified with a TOC analyzer. The
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8 383 HULIS-C concentration in fine particles (PM_{1.8}) accounted for 35±7% and 35±5% of the
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10 384 WSOC concentration during the NAD and AD periods, respectively, with the highest HULIS-
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12 385 C/WSOC ratio in the particle cut sizes from 0.55 to 1.0 µm for the NAD (45±6%) and 1.8 to
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14 386 3.1 µm for AD (44±7%). The size distribution for HULIS were similar to those in PM,
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16 387 WSOC, NO₃⁻, SO₄²⁻, NH₄⁺, and K⁺. HULIS exhibited a dominant droplet mode peaking at
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18 388 0.55 µm during the NAD period and showed a bimodal size distribution peaking at both 0.32
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20 389 and 1.8 µm during the AD period.

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22 390 The correlation of the size-resolved HULIS with other water-soluble ionic species
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24 391 strongly indicates that fine mode HULIS is produced by formation processes of SOA similar
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26 392 to those of secondary ionic species and primary BB emissions during NAD but is directly
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28 393 associated with SOA formation processes during the AD period. In the coarse mode, strong
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30 394 correlations of HULIS with Na⁺, Cl⁻, Ca²⁺, NO₃⁻, and SO₄²⁻ during the AD period suggest that
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32 395 the coarse mode HULIS during the AD period was likely to condense on the surface of sea-
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34 396 salt particles (NaNO₃ and Na₂SO₄) and/or soil-related particles [Ca(NO₃)₂ and CaSO₄].
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36 397 However, the processes that could lead to the small presence of HULIS in the coarse mode
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38 398 during the NAD period are not known.

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Table 1. Recovery results of two HULIS standards using HLB SPE cartridge method

HULIS type	Initial conc. ($\mu\text{g C/L}$)	HULIS-C conc. ($\mu\text{g C/L}$)	% Recovery
Suwannee River Fulvic Acid (SRFA)	618 \pm 8	587 \pm 32	95.0 \pm 4.2
	1053 \pm 40	1054 \pm 17	100.1 \pm 2.4
	2417 \pm 38	2564 \pm 79	106.1 \pm 1.6
	2563 \pm 6	2701 \pm 235	101.3 \pm 8.3
	5073 \pm 131	4564 \pm 89	90.0 \pm 1.4
Suwannee River Humic Acid (SRHA)	529 \pm 10	507 \pm 25	95.9 \pm 3.1
	1750 \pm 20	1604 \pm 92	95.1 \pm 1.7
	1843 \pm 6	1861 \pm 55	100.9 \pm 2.7
	3617 \pm 50	3551 \pm 60	90.0 \pm 7.1

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Table 2. Comparison of PM₁₀, PM_{1.8}, PM_{1.8}/PM₁₀, and contributions of WSOM, HULIS and SIC to PM_{1.8} between NAD and AD periods.

Conditions	PM ₁₀ ($\mu\text{g}/\text{m}^3$)	PM _{1.8} ($\mu\text{g}/\text{m}^3$)	PM _{1.8} /PM ₁₀ (-)	WSOM/PM _{1.8} (-)	HULIS/PM _{1.8} (-)	SIC/ PM _{1.8} (-)
Non-Asian dust	34.2±13.6	24.0±11.2	0.69±0.04	0.23±0.04	0.08±0.02	0.51±0.04
Asian dust	57.7±24.4	30.5±10.5	0.54±0.05	0.13±0.08	0.04±0.02	0.34±0.15

Notes) WSOM (=1.9×WSOC): water-soluble organic mass, HULIS (=1.9×HULIS-C), and SIC: sum of secondary inorganic components (=NO₃⁻+SO₄²⁻+NH₄⁺)

Table 3. Coefficient of determination (R^2) between HULIS-C and chemical markers in fine mode ($PM_{1.8}$) and coarse mode ($PM_{3.1-10}$) particles during NAD and AD periods

Dependent Variable	NAD period		AD period		
	$PM_{1.8}$	$PM_{3.1-10}$	$PM_{1.8}$	$PM_{3.1-10}$	
HULIS-C	WSOC	0.98^{**}	0.57	0.72^{**}	0.72^{**}
	Na ⁺	0.01	0.00	0.14	0.84^{**}
	Cl ⁻	0.28	0.05	0.24	0.83^{**}
	K ⁺	0.78^{**}	0.04	0.25	0.04
	Ca ²⁺	0.08	0.34	0.02	0.67[*]
	NH ₄ ⁺	0.92^{**}	0.08	0.56^{**}	0.36
	NO ₃ ⁻	0.88^{**}	0.19	0.66^{**}	0.79^{**}
	SO ₄ ²⁻	0.75^{**}	0.06	0.48^{**}	0.63[*]

Note) Superscripts * and ** are significant at p values of < 0.05 and < 0.01 , respectively.

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List of Figure Caption

Figure 1. Time-series of hourly PM₁₀ mass concentration over the study period

Figure 2. Interference of water-soluble ionic species to measured HULIS concentrations in ambient aerosol samples

Figure 3. Average chemical composition of size-resolved PM for (a) NAD and (b) AD periods

Figure 4. Size-resolved HULIS-C/WSOC ratio for NAD and AD periods

Figure 5. Size distributions of PM, WSOC, HULIS-C, secondary ionic species, and K⁺ concentrations for NAD and AD periods

Figure 6. Correlations of HULIS-C and WSOC with K⁺, NO₃⁻, SO₄²⁻, and NH₄⁺ for NAD and AD periods

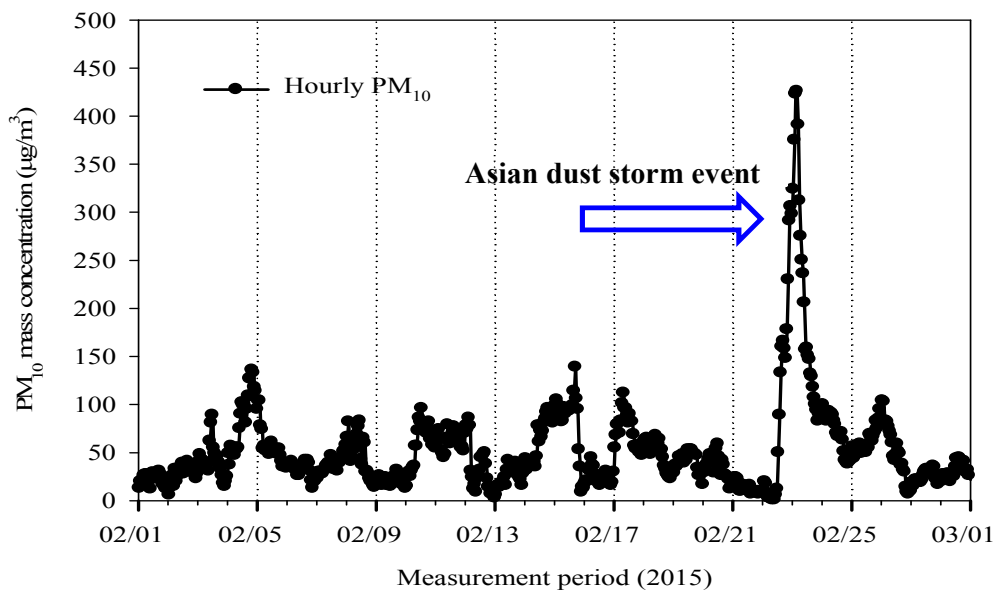


Figure 1. Time series of hourly PM₁₀ mass concentration over the study period

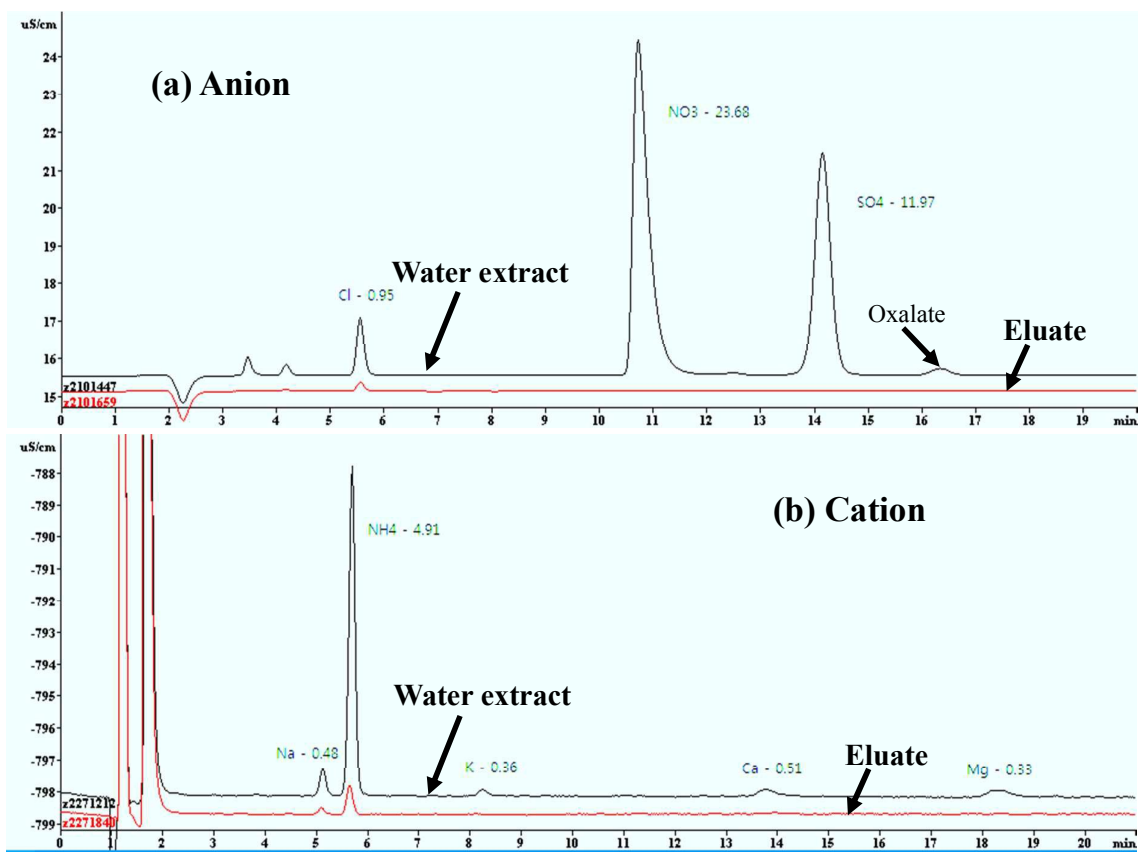


Figure 2. Interference of water-soluble ionic species to measured HULIS concentrations in ambient aerosol samples.

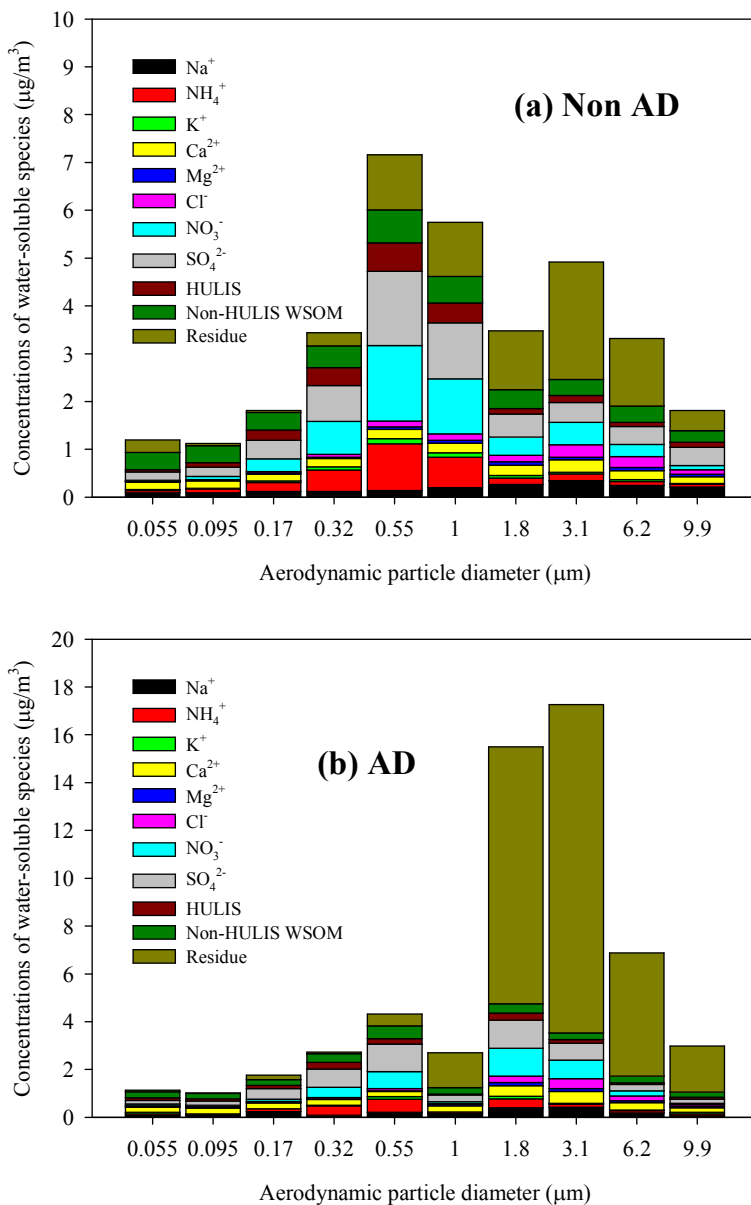


Figure 3. Average chemical composition of size-resolved PM for (a) NAD (b) AD periods

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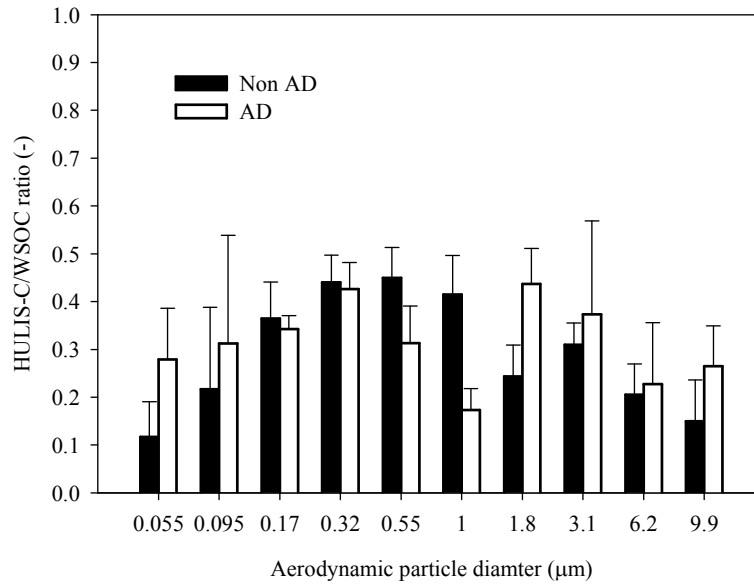


Figure 4. Size-resolved HULIS-C/WSOC ratio for NAD and AD periods.

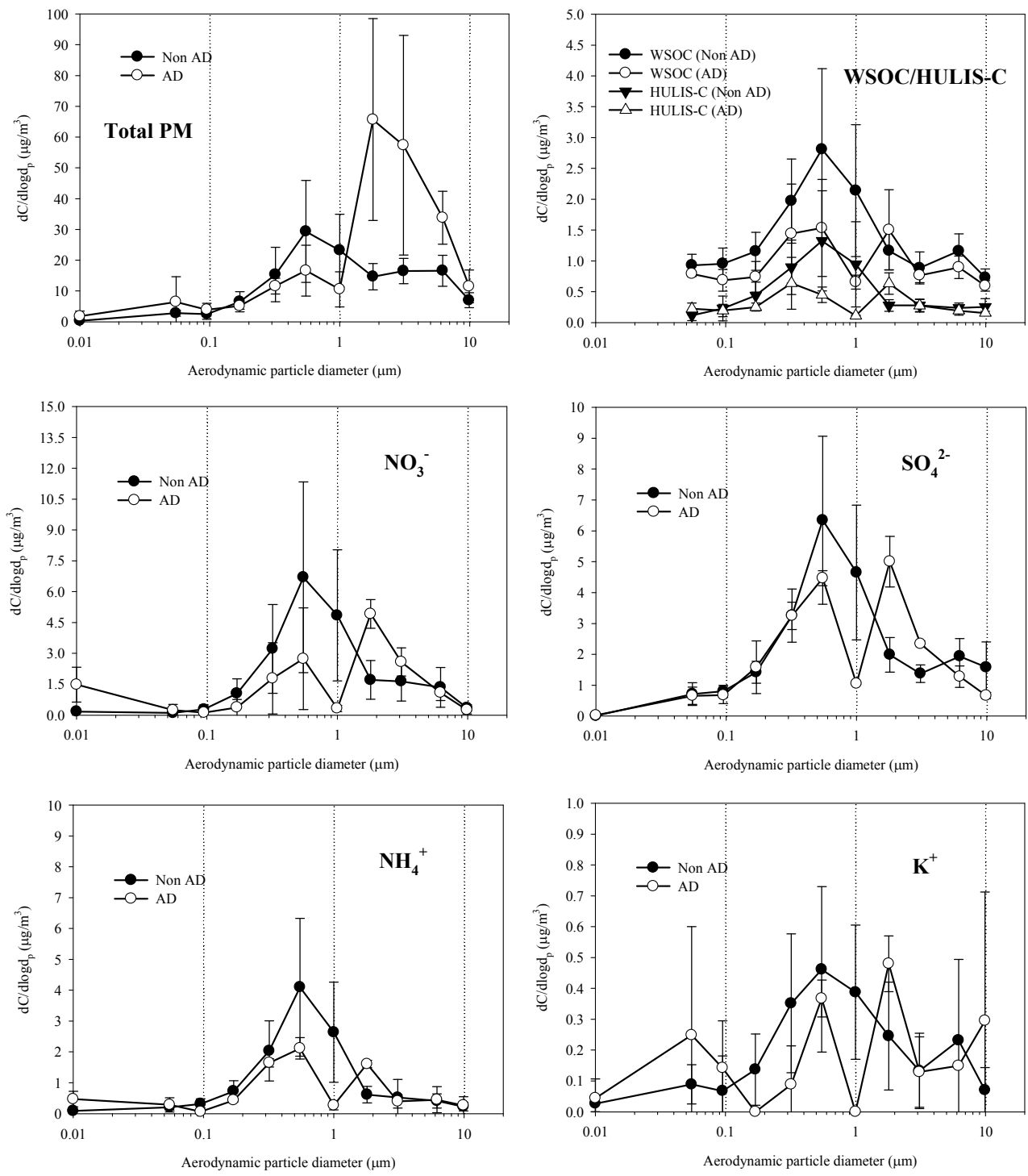


Figure 5. Size distributions of PM, WSOC, HULIS-C, secondary ionic species, and K⁺ concentrations for NAD and AD periods

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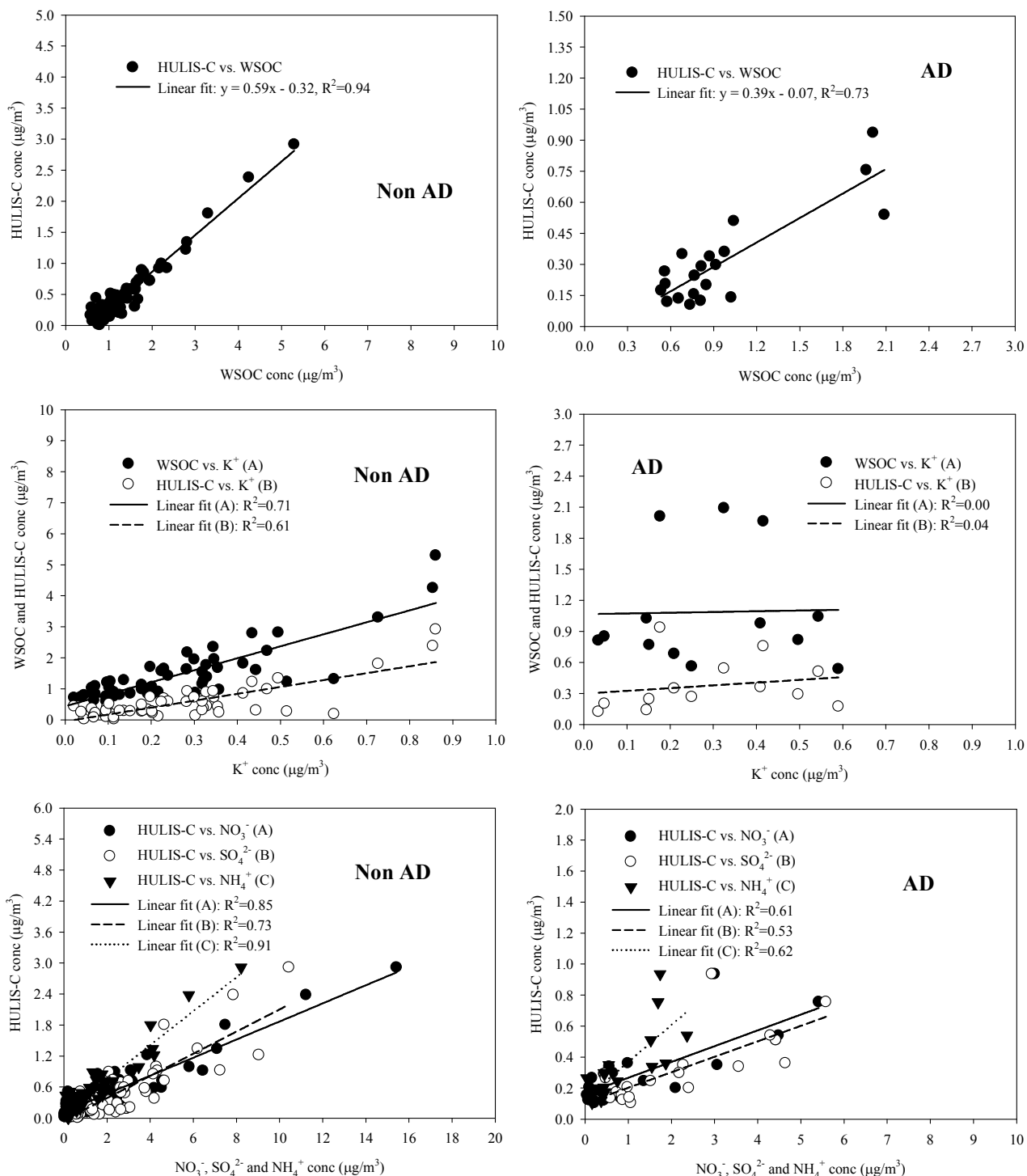


Figure 6. Correlations of HULIS-C and WSOC with K^+ , NO_3^- , SO_4^{2-} , and NH_4^+ for NAD and AD periods