

Environmental Science Processes & Impacts

Accepted Manuscript



This is an *Accepted Manuscript*, which has been through the Royal Society of Chemistry peer review process and has been accepted for publication.

Accepted Manuscripts are published online shortly after acceptance, before technical editing, formatting and proof reading. Using this free service, authors can make their results available to the community, in citable form, before we publish the edited article. We will replace this *Accepted Manuscript* with the edited and formatted *Advance Article* as soon as it is available.

You can find more information about *Accepted Manuscripts* in the [Information for Authors](#).

Please note that technical editing may introduce minor changes to the text and/or graphics, which may alter content. The journal's standard [Terms & Conditions](#) and the [Ethical guidelines](#) still apply. In no event shall the Royal Society of Chemistry be held responsible for any errors or omissions in this *Accepted Manuscript* or any consequences arising from the use of any information it contains.

Environmental Impacts Statement

Environmental Impacts Statement

Uranium exerts a threat to surface and groundwater across the globe owing to both anthropogenic activities, such as mining and nuclear fuel production, as well as natural sources. Iron oxides serve as principal hosts of uranium through adsorption and co-precipitation pathways, with the ubiquitous iron phase ferrihydrite serve a particularly prominent, and unique, role in uranium retention. Here we examine the unresolved influence of chemical heterogeneities in ferrihydrite composition on uranium retention upon reaction with Fe(II) (common under anaerobic conditions). We observed that uranium incorporation into Al-ferrihydrite transformation product, goethite, declined from ~70% of solid-phase U without Al to ~30% with 20% Al content. Thus, Al within ferrihydrite diminishes a reaction pathway serving as a long-term host of uranium.

1 *Uranium incorporation into aluminum-substituted ferrihydrite during iron(II)-induced*
2 *transformation*

3 MS# EM-ART-03-2014-00014

4

5

6 Michael S. Massey^{a,d,*}

7

8 Juan S. Lezama-Pacheco^{a,c}

9

10 F. Marc Michel^b

11

12 Scott Fendorf^{a,*}

13

14

15 ^a Department of Environmental & Earth System Science, Stanford University, Stanford,
16 California 94305, United States

17

18 ^b Virginia Polytechnic Institute and State University, Blacksburg, Virginia 24061, United States

19

20 ^c SLAC National Accelerator Laboratory, Menlo Park, California 94025, United States

21

22 ^d Present address: Department of Earth and Environmental Sciences, California State University
23 East Bay, 25800 Carlos Bee Boulevard, Hayward, California 94542, United States; e-mail:
24 mike.massey@csueastbay.edu; phone: (510)885-3486.

25

26 * Corresponding authors

27

28

29 Abstract

30 Uranium retention processes (adsorption, precipitation, and incorporation into host
31 minerals) exert strong controls on U mobility in the environment, and understanding U retention
32 is therefore crucial for predicting the migration of U within surface and groundwater. Uranium
33 can be incorporated into Fe (hydr)oxides during Fe(II)-induced transformation of ferrihydrite to
34 goethite. However, ferrihydrite seldom exists as a pure phase within soils or sediments, and
35 structural impurities such as Al alter its reactivity. The presence of Al in ferrihydrite, for
36 example, decreases the rate of transformation to goethite, and thus may impact the retention
37 pathway, or extent of retention, of U. Here, we investigate the extent and pathways of U(VI)
38 retention on Al-ferrihydrite during Fe(II)-induced transformation. Ferrihydrite containing 0%,
39 1%, 5%, 10%, and 20% Al was reacted with 10 μM U and 300 μM Fe(II) in the presence of 0
40 mM and 4 mM Ca^{2+} and 3.8 mM carbonate at pH 7.0. Solid reaction products were characterized
41 using U L_3 -edge EXAFS spectroscopy to differentiate between adsorbed U and U incorporated
42 into the goethite lattice. Uranium incorporation into Al-ferrihydrite declined from ~70% of solid-
43 phase U at 0% and 1% Al to ~30% of solid phase U at 20% Al content. The decrease in U
44 incorporation with increasing Al concentration was due to two main factors: 1) decreased
45 transformation of ferrihydrite to goethite; and, 2) a decrease of the goethite lattice with
46 increasing Al, making the lattice less compatible with large U atoms. Uranium incorporation can
47 occur even with an Al-substituted ferrihydrite precursor in the presence or absence of Ca^{2+} . The
48 process of U incorporation into Al-goethite may therefore be a potential long-term sink of U in
49 subsurface environments where Al-substituted iron oxides are common, albeit at lower levels of
50 incorporation with increasing Al content.

51 1. Introduction

52 Uranium mining, milling, and refining, as well as nuclear weapon and fuel production,
53 has left a substantial legacy of soil and groundwater contamination. Uranium contamination from
54 legacy U production exists on every continent except Antarctica, with an estimated global
55 volume of over 900 million m³ of U mine/mill tailings. These occupy a land area of nearly 6,000
56 ha, and occur with associated contamination of soils, sediments, and groundwater¹. As one
57 example, the United States Department of Energy manages an inventory of 1.5 billion m³ of
58 contaminated groundwater, and 75 million m³ of contaminated soil/sediment², and U is among
59 the most common radionuclide contaminants at United States DOE sites³. Managing U
60 contamination requires an accurate understanding of U biogeochemical processes and retention
61 mechanisms. Thus, establishing a clearer understanding of U biogeochemistry is crucial for
62 mitigating the impact of legacy contamination, as well as understanding current and future
63 environmental impacts of U.

64 Three U retention processes have been the subject of intensive study over the past several
65 decades: 1) U(VI) adsorption to soil and sediment solids, 2) reduction of U(VI) in groundwater
66 to sparingly-soluble U(IV) solids, and, most recently, 3) incorporation of U(VI/V) into iron
67 oxides such as hematite^{4,5}, goethite⁶⁻⁸, and possibly magnetite⁶. The uranyl cation (UO₂²⁺) can
68 adsorb on iron oxide, hydroxide, and oxyhydroxide minerals, hereafter referred to collectively as
69 “iron oxides”⁹⁻¹². Uranyl also adsorbs on clays¹³ and quartz¹⁴. As a long-term retention
70 mechanism, however, adsorption may be limited due to the reversible nature of adsorption
71 reactions with shifts in aqueous chemistry. An increase in Ca²⁺ and carbonate concentrations, for
72 example, can lead to the formation of uranyl-calcium-carbonate ternary complexes which
73 decrease the extent of adsorption and promote desorption^{15,16}.

74 Uranium(VI) reduction to sparingly-soluble U(IV) solids such as uraninite (UO_2) and
75 “monomeric” U(IV) species is another potential U retention pathway. There are a variety of
76 abiotic and biotic pathways of U(VI) reduction, including reaction with Fe(II), hydrogen sulfide,
77 or Fe(II)-bearing solids such as magnetite¹⁷⁻²⁰. Ferrous iron has been shown to reduce U(VI)
78 either through a heterogeneous reaction mediated by, for example, iron oxide surfaces¹⁷ or
79 homogeneously in aqueous solution²¹. In addition to abiotic reductants of aqueous U(VI), biotic
80 pathways have been examined extensively. They have been used in U contamination remediation
81 via *in situ* biostimulation of metal reducing microorganisms²²⁻²⁷. Independent of its formation
82 pathway, U(IV) can persist for long periods of time, such as in U-bearing roll-front deposits.
83 However, re-oxidation and subsequent re-mobilization of U by dissolved oxygen²⁸, dissolved
84 nitrate²⁹⁻³¹, or Fe(III) and Mn(IV/III) minerals^{32,33} is a concern. The dependence of U reduction
85 on aqueous U speciation, and the potential for re-oxidation/remobilization, make the
86 effectiveness of reductive retention dependent on stable long-term geochemical conditions.

87 Uranium(V/VI) incorporated into iron oxides, on the other hand, is resistant to release
88 upon pH change⁷, oxidizing conditions, and short-period redox cycling³⁴. Uranium(VI)
89 incorporated in hematite^{4,5} and U(V) in goethite^{6-8,35} is present in uranate (octahedral)
90 coordination, rather than uranyl coordination (UO_2^{2+}) of the dominant aqueous and adsorbed U
91 species. Gomez et al.³⁶ found that U incorporation or co-precipitation in iron minerals can exert
92 a substantial control over aqueous U concentrations near mine tailings. Additionally, uranium
93 associated with iron oxides may be stable over geologic timescales³⁷, so U incorporation into
94 iron oxide minerals could serve as a stable, long-term sink for contaminant U.

95 Uranium retention processes occur concurrently with iron biogeochemical
96 transformations in subsurface environments. For example, the Fe hydroxide ferrihydrite can be

97 transformed into the more thermodynamically-favorable products lepidocrocite, goethite, or
98 magnetite by reaction with aqueous Fe(II)³⁸⁻⁴⁰. Many different parameters affect ferrihydrite
99 transformation, however. For example, Fe(II) concentration, pH changes, and the presence of
100 chloride, sulfate, or carbonate in solution can alter the transformation products³⁹, while
101 adsorbates such as phosphate and silicate can inhibit transformation^{41,42}.

102 Structural impurities in ferrihydrite, such as Al or Si, can also dramatically alter the
103 reactivity of ferrihydrite and its transformation products. Structural Al can be found at
104 concentrations as high as 30% in synthetic or natural samples of iron oxides⁴³⁻⁴⁵. Aluminum
105 inhibits ferrihydrite transformation to more-crystalline iron oxides at high Al concentrations^{46,47}.
106 Hansel et al.⁴⁷ also found that structural Al influences the reaction products (e.g., lepidocrocite
107 vs. goethite), though the specific mechanisms behind the influence of structural Al on reaction
108 products remain unclear. Although structural Si in ferrihydrite was noted to preserve reactivity
109 upon desiccation⁴², it also inhibited U incorporation during Fe(II)-induced transformation⁷,
110 highlighting the potential importance of co-precipitated ions such as Si and Al in coupled U and
111 Fe biogeochemical processes. The impact of structural Al on U incorporation has not yet been
112 examined, despite the ubiquitous nature of Al-substituted Fe oxide minerals in the environment.
113 The relevance of the U incorporation pathway in natural systems is predicated upon U
114 incorporation into Fe oxides that only rarely exist in a pure state. Accordingly, the objective of
115 this study was to elucidate the impact of co-precipitated Al-ferrihydrite on U incorporation
116 during Fe(II)-induced ferrihydrite transformation.

117 **2. Methods**

118 In order to test the influence of Al on U incorporation into Al-containing ferrihydrite,
119 ferrihydrite slurries with 0%, 1%, 5%, 10%, and 20% (hereafter, “Al-ferrihydrite”) were
120 synthesized. The Al-ferrihydrite slurry was then reacted with U and Fe(II) at pH 7.0 to induce
121 Al-ferrihydrite transformation and U incorporation.

122 **2.1. Synthesis of Al-ferrihydrite slurry**

123 Stock solutions of 150 μM FeCl_3 and AlCl_3 were combined proportionally for Al mole
124 percentages of 0%, 1%, 5%, 10%, and 20%. The combined solutions were stirred continuously
125 and vigorously while undergoing rapid (< 10 minutes) hydrolysis using 1 M NaOH to bring the
126 final pH to 7.2-7.3. The supernatant was decanted and the slurry was centrifuged at 6,000 RPM
127 and washed with de-ionized water (18 M Ω) five times to remove excess salt. The method was
128 similar to that in Masue et al. ⁴⁸. The final slurries were sampled under vigorous stirring,
129 dissolved with 6 M trace metal grade HCl, and analyzed with inductively-coupled plasma optical
130 emission spectrometry (ICP-OES) to obtain Fe and Al concentrations. X-ray powder diffraction
131 illustrated an Al-ferrihydrite diffraction pattern consistent with 2-line ferrihydrite.

132 **2.2. Batch incubation experiments**

133 Batch incubations were performed in 125 mL glass serum bottles with thick rubber
134 stoppers (Bellco Glass, Inc., New Jersey, USA). A total of 100 mL of solution and slurry was
135 added to serum bottles in a 95% N_2 /5% H_2 atmosphere (Coy Laboratory Products, Michigan,
136 USA). The solution consisted of 10 mM PIPES buffer and 3.8 mM KHCO_3 , adjusted to pH 7.0
137 with trace metal grade HCl. Either 0 mM or 4 mM Ca^{2+} (added as $\text{CaCl}_2 \cdot 2\text{H}_2\text{O}$) was used to
138 examine the effect of uranyl-calcium-carbonato complexation, and 0 or 10 μM U (as uranyl

139 acetate) provided both U-containing samples and no-U controls. The 10 μM U concentration was
140 chosen as a reasonable model for U contamination in groundwater, which often ranges from 10^{-7}
141 M to 10^{-4} M. Solutions were made using de-ionized water (18 M Ω) that had been de-oxygenated
142 by boiling and bubbling with a stream of N₂ gas for several hours.

143 (Al-)ferrihydrite was added to each bottle in an amount based on the ferrihydrite slurry
144 density, in order to achieve a final Fe+Al concentration of ~ 1.7 mM (equivalent to 15-20 mg of
145 Al-ferrihydrite per bottle, or a solid concentration of ~ 150 -200 mg L⁻¹). Then, uranyl acetate was
146 added for a concentration of either 0 μM or 10 μM . This mixture was capped and sealed using
147 thick rubber stoppers, and allowed to equilibrate for ~ 1 h. Finally, 0.3 mL of 100 mM FeSO₄ was
148 added using a needle and syringe for an initial Fe(II) concentration of 300 μM . The addition of
149 Fe(II) initiated ferrihydrite and U transformation. Uranium-containing incubations were
150 performed in triplicate.

151 Capped, anoxic samples were incubated at 25°C for 7-8 days on a rotary shaker at 120
152 RPM. After incubation, duplicate 10 mL aliquots of solution were withdrawn using a needle and
153 syringe, and filtered through 0.22 μm nitrocellulose membranes into 15 mL serum vials. These
154 vials were capped for storage and chemical analysis. The remainder of each sample (~ 80 mL)
155 was vacuum-filtered through a 0.22 μm nitrocellulose membrane, scraped from the filter while
156 wet, washed 3-5 times with de-ionized water, air-dried, and ground for analysis. Solution and
157 solid sampling were performed under a 95% N₂/5% H₂ atmosphere.

158 **2.3. Solution analysis**

159 Aliquots of solutions were diluted using 3% trace metal grade HNO₃ for chemical
160 analysis. Uranium was measured using inductively-coupled plasma mass spectrometry (ICP-MS,

161 Thermo Scientific XSERIES 2, Thermo Fisher Scientific, Waltham, MA), and Fe, Ca, K, Na,
162 and Al were measured using inductively-coupled plasma optical emission spectrometry (ICP-
163 OES, Thermo Scientific ICAP 6300 Dual View, Thermo Fisher Scientific, Waltham, MA).

164 **2.4. Solid analysis**

165 **2.4.1. X-ray absorption spectroscopy**

166 Uranium L₃-edge and Fe K-edge X-ray absorption spectra were collected at beamlines
167 11-2 and 4-1 at the Stanford Synchrotron Radiation Lightsource (SSRL). Dried solid powder
168 (~15 mg) was diluted with deoxygenated BN powder (~70 mg) and homogenized by grinding
169 with an agate mortar and pestle. Samples were sealed with three layers of Kapton film, and
170 placed under vacuum to isolate the samples from oxygen during analysis at room-temperature.
171 The X-ray beam incident energy was controlled using a Si(220) double crystal monochromator
172 in the $\Phi=0^\circ$ orientation (for U) and $\Phi=90^\circ$ orientation (for Fe). An in-line Y or Fe foil was used
173 for energy calibration; the Y K-edge was calibrated to 17038.4 eV, and the Fe K-edge was
174 calibrated to 7111.0 eV. Transmission spectra were collected using an in-line ion chamber, and
175 fluorescence spectra were collected simultaneously with either a 13- or 30-element Ge solid-state
176 detector (Canberra, Connecticut, USA) for U, or a Lytle detector for Fe.

177 Data calibration and averaging were performed using SixPack⁴⁹. Background
178 subtraction, normalization, and linear combination fitting analyses were performed with Athena
179⁵⁰. Standard spectra for linear combination fits were collected and processed under identical
180 conditions and fit using Artemis and FEFF 6.0 or FEFF 8.4^{51,52}. Detailed normalization and
181 fitting parameters, as well as EXAFS spectra, are given in the supporting information. The

182 uncertainty associated with EXAFS linear combination fitting was approximately 5-15%, as
183 detailed in the Supplemental Information.

184 **2.4.2. High-resolution synchrotron X-ray powder diffraction**

185 Ground Al/Fe oxide/U powder was placed in 0.3 mm diameter borosilicate glass
186 capillaries (Hampton Research, Aliso Viejo, CA). Capillaries were sealed using five-minute
187 epoxy (ITW Devcon, Danvers, MA, USA) in a 95% N₂/5% H₂ atmosphere and analyzed at
188 beamline 7-2 at SSRL. Capillaries were contained in sealed plastic containers with Kapton
189 windows for radionuclide containment; the container was purged with He gas to minimize
190 oxygen exposure and decrease scattering background. The incident beam energy was maintained
191 at 16.5 keV, and precise calibration was achieved using a powdered LaB₆ calibration standard in
192 a borosilicate glass capillary. High-resolution synchrotron X-ray powder diffraction data were
193 collected over a Q-space range of ~0.8-12 by scanning a Vortex Si solid state detector (SII
194 Nanotechnology USA Inc., Northridge, CA, USA) in steps of Q=0.05.

195 Diffraction patterns were analyzed using the General Structure Analysis System (GSAS)
196 software package⁵³ with the EXPGUI interface⁵⁴. Lattice parameters and coherent scattering
197 domain size (nominal crystallite size) were determined using Rietveld refinement. Capillary
198 background and a residual ferrihydrite phase were fit separately and included in the goethite
199 Rietveld refinement in order to accurately refine peak shapes and peak intensities of the goethite
200 diffraction pattern. Instrument-specific and experiment-specific parameters were determined
201 using the LaB₆ diffraction pattern, as in Campbell et al.⁵⁵. Lattice parameters and Lorentzian
202 broadening parameters (related to coherent scattering domain size) were then refined. Full details
203 of the refinement are given in Supplemental Information.

204 **3. Results**

205 **3.1. Uranium solid/solution partitioning**

206 Using batch reactors, we investigated the retention of U(VI) upon Fe(II) reaction with
207 ferrihydrite having varying amounts of structural Al. Increasing levels of Al within ferrihydrite,
208 from 0 to 20 mol%, decreased the extent of U retention (Fig. 1). Uranium retention was further
209 diminished by the presence of Ca^{2+} . With Ca^{2+} in solution, U retention decreased by an order of
210 magnitude compared to the no-Ca system. More specifically, in the system with Ca, U in
211 solution increased from 4% (0% Al-ferrihydrite) to 16% (1% Al-ferrihydrite) and up to > 30%
212 (20% Al-ferrihydrite). In absence of Ca, the trend of increasing solution-phase U with increasing
213 Al in the solid generally held, but the impact was small and in all cases greater than 98% of U
214 was associated with the solid phase (Fig. 1).

215 **3.2. Uranium solid phase speciation**

216 Uranium L_3 -edge EXAFS linear combination fitting was performed to determine the
217 proportions of U incorporated in versus adsorbed on the Al-ferrihydrite transformation products
218 (Fig. 2). Greater amounts of Al in ferrihydrite decreased U incorporation into transformation
219 products (Fig. 3). With 4 mM Ca^{2+} , the 0% and 1% Al-ferrihydrite transformation products were
220 very similar, with 72% and 63% incorporated U, respectively. The 5% and 10% Al-ferrihydrite
221 transformation products contained 52% and 54% incorporated U, respectively, while the 20%
222 Al-ferrihydrite transformation product contained 25% incorporated U and 75% adsorbed U (Fig.
223 3a). Calcium decreased the proportion of incorporated U and decreased the proportion of
224 adsorbed U at all Al loadings. No U(IV) solids, such as UO_2 , were observed via EXAFS
225 spectroscopy.

226 Taken together, the U retention (Fig. 1) and EXAFS linear combination fitting results
227 (Figs. 2 and 3) indicate that most of the U in the system was incorporated into the goethite solid
228 (see below for iron oxide transformation products) at 0% and 1% Al contents, regardless of Ca^{2+}
229 presence or absence. However, with increasing Al content, adsorbed and aqueous U(VI) became
230 more prevalent. This trend was particularly evident in the presence of 4 mM Ca^{2+} , with less than
231 20% of the U in the system incorporated into the goethite solid at 20% Al content within
232 ferrihydrite. Even with increased U retention in the absence of Ca^{2+} , only about 32% of the initial
233 U was incorporated into the 20% Al-goethite.

234 3.3. Initial Al-ferrihydrite and Coherent Scattering Domain Size

235 The initial Al-ferrihydrite solids were examined using X-ray diffraction and high-energy
236 total scattering and pair distribution function (PDF) analysis. Initial materials were all two-line
237 ferrihydrite (Fig. S1). The PDF of the initial 1-20% Al-ferrihydrite solids were very similar (Fig.
238 S2), with a coherent scattering domain (CSD) size of ~ 2 nm. There was a slight decrease in CSD
239 with increasing Al content; a similar decrease in CSD (a proxy for particle size) was also
240 observed by Cismasu et al.⁴⁵. No separate Al oxide domains or phases were detected using X-
241 ray scattering, but they may have been present in the 20% Al-ferrihydrite^{45,56}.

242 3.4. Iron oxide transformation products

243 Crystalline solid phases were identified using high-resolution synchrotron X-ray powder
244 diffraction, and Fe oxides were further examined using Fe K-edge EXAFS. Goethite was the
245 only crystalline Fe oxide transformation product observed in either the 4 mM Ca or 0 mM Ca
246 systems (Fig. S3). Rietveld refinement of fits of goethite diffraction patterns indicated that the
247 lattice parameters changed with increasing Al loading (Fig. 4). Typical crystallographic values

248 for pure goethite are $a=4.61 \text{ \AA}$, $b=9.96 \text{ \AA}$, and $c=3.02 \text{ \AA}$ ⁴³. However, the a lattice parameter
249 decreased from $\sim 4.625 \text{ \AA}$ to $\sim 4.615 \text{ \AA}$ as Al substitution increased from 1% Al to 20% Al (Fig.
250 4). Similarly, the b lattice parameter decreased from $\sim 9.98 \text{ \AA}$ to $\sim 9.92 \text{ \AA}$, and the c lattice
251 parameter decreased from $\sim 3.035 \text{ \AA}$ to $\sim 3.010 \text{ \AA}$ as Al content increased from 1% to 20% (Fig.
252 4). Iron K-edge EXAFS linear combination fitting indicated that unreacted ferrihydrite accounted
253 for 28% to 48% of the Fe oxide in the Al-ferrihydrite transformation products of the 4 mM Ca^{2+}
254 system; 23-41% remained as ferrihydrite in the no-Ca system (Figs. S4 and S5).

255 4. Discussion

256 Across all Al loadings, with either 4 mM or 0 mM Ca, ~ 70 - 99% of the U in the system
257 was retained on the solid phase (Fig. 1), consistent with previous findings^{15,16}. The dominance of
258 the uranyl-calcium-carbonato ternary complex in solution decreased U retention at 5-20% Al
259 content. The presence of 4 mM Ca, and the corresponding shift of U aqueous speciation to a
260 regime dominated by the uranyl-calcium-carbonato complex, resulted in an order of magnitude
261 more U in the aqueous phase in comparison to the 0 mM Ca system (Fig. 1). The decrease of U
262 retention was accompanied by a shift in the dominant solid-phase U retention pathway from U
263 incorporation into goethite to U adsorption on the Al-containing oxides (Figs. 2 and 3).
264 However, an increase in both structural Al in ferrihydrite (and Al in the subsequent goethite
265 transformation product), coupled with a shift in uranyl speciation to the ternary uranyl-calcium-
266 carbonato complex, diminishes U incorporation; with 20% Al content and 4 mM Ca, only $\sim 17\%$
267 of total U is incorporated into goethite in the system with Ca.

268 The dominant effect on the U retention mechanisms under the conditions of this study
269 resulted from the structural Al content of the ferrihydrite precursors and goethite transformation

270 products. An increase in Al content had the overriding effect of decreasing U incorporation and
271 shifting the retention mechanism toward uranyl adsorption. Uranium incorporation into Al-
272 ferrihydrite transformation products decreased from 63-88% in 0% and 1% Al-ferrihydrite to
273 25-34% with 20% Al-ferrihydrite (Fig. 3). Uranium incorporation into goethite during Fe(II)-
274 induced ferrihydrite transformation proceeds by the reduction of adsorbed U(VI) to U(V) by
275 Fe(II)⁸. There are two possible reasons for the non-linear decrease of U incorporation into
276 goethite and increased dominance of U adsorption in systems with increasing Al content: 1)
277 inhibition of ferrihydrite transformation to goethite (which may be coupled to a limitation in
278 reduction of adsorbed U(VI)); or, 2) structural incompatibility with incorporated U(V) resulting
279 from a decreased crystal lattice in Al-bearing ferrihydrite and goethite. It is also possible that Al
280 creates a greater availability of adsorption sites by limiting ferrihydrite transformation to goethite
281 but we do not believe this would limit Fe(II)-induced reductive incorporation of U into goethite.

282 The first explanation for the decrease in U incorporation is the decreased transformation
283 of ferrihydrite to goethite (Fig. S4). Formation of goethite via Fe(II)-induced transformation of
284 ferrihydrite is necessary for U(V) incorporation⁸. Masue-Slowey et al.⁴⁶ and Hansel et al.⁴⁷
285 both demonstrated that structural Al decreases the extent of Al-ferrihydrite transformation;
286 Masue-Slowey et al.⁴⁶ attributed this to inhibition by Al of electron hopping and bulk
287 conduction⁵⁷⁻⁶⁰ in the Fe (hydr)oxide structure. Bazilevskaya et al.⁵⁶ noted that at Al content
288 greater than 8%, structural Al³⁺ is clustered rather than independently distributed throughout the
289 Fe/Al (hydr)oxide structure, which further suggests inhibition of electron hopping in the bulk
290 solid, and partially explains the decreased ferrihydrite transformation and U incorporation
291 observed at Al loadings of 10% and 20%. Cismasu et al.⁴⁵ also observed Al-rich clusters at Al

292 contents as low as 15%. Masue-Slowey et al. ⁴⁶ reported only goethite and ferrihydrite products
293 from an Al-ferrihydrite slurry precursor, a finding consistent with the results of the present study.

294 A second possibility to explain the observed decrease in U incorporation is that of
295 structural limitations caused by Al (and U) in goethite, the dominant transformation product (Fig.
296 S3). Several investigators have found that U incorporates into octahedral Fe³⁺ sites in goethite ^{6,7}.
297 Atomistic modeling supports the possibility of U(V) incorporation in these cation sites, with
298 local charge balance achieved through protonation/de-protonation of nearby hydroxyls, or the
299 introduction of structural vacancies at cation sites ³⁵. However, in addition to local charge
300 balance, lattice parameters and relief of lattice strain are also of concern. Rietveld refinement of
301 high-resolution synchrotron X-ray powder diffraction patterns indicated a decrease in lattice
302 parameters with increasing Al content in goethite (Fig. 4). The presence of Al results in
303 decreased unit cell size in the goethite, due to the smaller size of the Al³⁺ cation with respect to
304 Fe³⁺. ^{VI}U⁵⁺ is a large cation (~0.75-0.80 Å when incorporated in goethite as U⁵⁺) compared to
305 ^{VI}Fe³⁺ (0.65 Å), and certainly compared to ^{VI}Al³⁺ (0.54 Å). Uranium substitution for Fe³⁺/Al³⁺ is
306 likely increasingly less favorable with greater Al in the goethite lattice.

307 A further contribution to the shift from incorporation to adsorption at higher Al content is
308 the availability of surface sites for uranyl adsorption. At 10-20% Al content, decreased U
309 incorporation into goethite and increased residual ferrihydrite resulted in U retention primarily as
310 an adsorbed species. This may be partly attributed to differences in residual ferrihydrite, in
311 concert with decreased removal of U from solid-solution exchange by U incorporation. Typical
312 N₂-BET surface area of ferrihydrite slurry ranges from 159-234 m² g⁻¹ ^{61,62}; the available surface
313 area of undried slurry is likely even higher due to particle aggregation from drying prior to
314 measurement. Roden and Zachara ⁶³ reported goethite surface area ranging from 31 to 153 m² g⁻¹,

315 depending on goethite particle size (100-200 nm to 15-30 nm, respectively). Higher surface area,
316 especially from residual Al-ferrihydrite, will favor U adsorption over U incorporation due to
317 greater availability of adsorption sites. Potentially more important, however, is that U retained
318 on ferrihydrite is not available for reduction to U(V) and incorporation into the goethite
319 transformation product, since incorporation of U(V) relies on the mineral transformation.

320 However, factors such as aggregation and competing solutes also impact adsorption. For
321 example, with increasing Al content, Cismasu et al.⁴⁴ found that natural ferrihydrite samples
322 with high Al and Si content tended to form aggregates and have lower surface area than pure
323 ferrihydrite (as low as 65 m² g⁻¹ in ferrihydrite with many impurities, down from 312 m² g⁻¹ in
324 ferrihydrite with fewer impurities). In the present study, higher amounts of Al only substantially
325 decreased total U retention in the presence of Ca, suggesting that aggregation resulting from Al
326 substitution did not appreciably decrease U adsorption site availability. There were also no
327 competing solutes such as phosphate or silicate in this study, which have been shown to limit
328 mineral transformation^{41,42} and U incorporation⁷ by blocking reactive surface sites or
329 preventing recrystallization. Uranium adsorption site availability was therefore not a limiting
330 factor for any of the Al loadings, lending support to the effect of Al on mineral transformation
331 (via lattice changes, stress/strain, etc.) as the primary control on the U retention mechanism.

332 A secondary result of the increased prevalence of U adsorption over incorporation at high
333 (10-20%) Al content was a decrease in overall U retention in the presence of 4 mM Ca²⁺. Our
334 results are consistent with the observation of Stewart et al.¹⁶ that conditions in which the uranyl-
335 calcium-carbonate ternary complex is dominant in solution result in an order of magnitude
336 increase of solution-phase U. Ergo, decreasing U incorporation with increasing Al content is of

337 particular concern in groundwater rich in Ca, due to lower U retention (via adsorption or
338 incorporation).

339 In summary, the presence of Al³⁺ in ferrihydrite did not completely inhibit U
340 incorporation into goethite resulting from Fe(II)-induced ferrihydrite transformation, even at Al
341 contents as high as 20%. However, the extent of U incorporation decreased substantially with
342 increasing Al content. Decreased U incorporation was due to the decrease of Al-ferrihydrite
343 transformation to goethite and incompatibilities between the Al-goethite lattice and incorporated
344 U. The shift toward U adsorption and a corresponding decrease of U incorporation due to
345 structural Al also resulted in increased U in solution in the presence of Ca²⁺ (Fig. 5).
346 Nevertheless, despite the decrease of U incorporation with increased Al, incorporation into
347 goethite was still a significant U retention pathway and have the potential to serve as a long-
348 term, stable sink of U contamination in natural environments.

349

350

351 **Acknowledgements**

352 Support for M.M. was provided partially by the Robert and Marvel Kirby Stanford Graduate
353 Fellowship. Additionally, this research was supported by the U.S. Department of Energy Office
354 of Biological and Environmental Research, through the Subsurface Biogeochemical Research
355 program (grant number DE-SC0006772) and the SLAC Science Focus Area Research Program
356 (FWP #10094). Use of the Stanford Synchrotron Radiation Lightsource, SLAC National
357 Accelerator Laboratory, is supported by the U.S. Department of Energy, Office of Science,
358 Office of Basic Energy Sciences under Contract No. DE-AC02-76SF00515. The contents of this
359 publication are solely the responsibility of the authors and do not necessarily represent the

360 official views of NIGMS, NCCR or NIH. Use of the Advanced Photon Source, an Office of
361 Science User Facility operated for the U.S. Department of Energy (DOE) Office of Science by
362 Argonne National Laboratory, was supported by the U.S. DOE under Contract No. DE-AC02-
363 06CH11357. We wish to thank the technical staff on APS BL ID-11-B for their support. We
364 thank Moses Gonzalez, Guangchao Li, and John Bargar for their assistance with this research,
365 and two anonymous reviewers for their contributions that substantially improved the manuscript.
366 We also appreciate the technical and safety support provided by L. Amoroso, D. Day, A. Gooch,
367 D. Menke, C. Morris, D. Murray, C. Patty, and R. Russ.

368
369
370
371

372 **References**

373
374
375
376
377
378
379
380
381
382
383
384
385
386
387
388
389
390
391
392
393

1. IAEA, *The long term stabilization of uranium mill tailings*, International Atomic Energy Agency, 2004.
2. DOE, *Linking Legacies: Connecting the Cold War Nuclear Weapons Production Processes To Their Environmental Consequences*, US Department of Energy, 1997.
3. R. Riley and J. Zachara, *Chemical contaminants on DOE lands and selection of contaminant mixtures for subsurface science research*, 1992.
4. M. Duff, J. Coughlin, and D. Hunter, *Geochimica et Cosmochimica Acta*, 2002, **66**, 3533–3547.
5. E. S. Ilton, J. S. L. Pacheco, J. R. Bargar, Z. Shi, J. Liu, L. Kovarik, M. H. Engelhard, and A. R. Felmy, *Environ. Sci. Technol.*, 2012, **46**, 9428–9436.
6. P. S. Nico, B. D. Stewart, and S. Fendorf, *Environ. Sci. Technol.*, 2009, **43**, 7391–7396.
7. D. D. Boland, R. N. Collins, T. E. Payne, and T. D. Waite, *Environ. Sci. Technol.*, 2011, **45**, 1327–1333.
8. M. S. Massey, J. S. Lezama-Pacheco, M. E. Jones, E. S. Ilton, J. M. Cerrato, J. R. Bargar, and S. Fendorf, *Geochimica et Cosmochimica Acta*, in press.
9. T. D. Waite, J. A. Davis, T. E. Payne, G. A. Waychunas, and N. Xu, *Geochimica et Cosmochimica Acta*, 1994, **58**, 5465–5478.
10. J. Bargar, R. Reitmeyer, and J. Davis, *Environ. Sci. Technol.*, 1999, **33**, 2481–2484.
11. D. Giammar and J. Hering, *Environ. Sci. Technol.*, 2001, **35**, 3332–3337.

- 394 12. T. Hiemstra, W. H. Van Riemsdijk, A. Rossberg, and K.-U. Ulrich, *Geochimica et*
395 *Cosmochimica Acta*, 2009, **73**, 4437–4451.
- 396 13. D. M. Singer, K. Maher, and G. E. Brown Jr., *Geochimica et Cosmochimica Acta*, 2009,
397 **73**, 5989–6007.
- 398 14. E. S. Ilton, Z. Wang, J.-F. Boily, O. Qafoku, K. M. Rosso, and S. C. Smith, *Environ. Sci.*
399 *Technol.*, 2012, **46**, 6604–6611.
- 400 15. P. M. Fox, J. A. Davis, and J. M. Zachara, *Geochimica et Cosmochimica Acta*, 2006, **70**,
401 1379–1387.
- 402 16. B. D. Stewart, M. A. Mayes, and S. Fendorf, *Environ. Sci. Technol.*, 2010, **44**, 928–934.
- 403 17. E. Liger, L. Charlet, and P. Van Cappellen, *Geochimica et Cosmochimica Acta*, 1999, **63**,
404 2939–2955.
- 405 18. B. Hua, H. Xu, J. Terry, and B. Deng, *Environ. Sci. Technol.*, 2006, **40**, 4666–4671.
- 406 19. B. Hua and B. Deng, *Environ. Sci. Technol.*, 2008, **42**, 8703–8708.
- 407 20. D. E. Latta, C. A. Gorski, M. I. Boyanov, E. J. O'Loughlin, K. M. Kemner, and M. M.
408 Scherer, *Environ. Sci. Technol.*, 2012, **46**, 778–786.
- 409 21. X. Du, B. Boonchayaanant, W.-M. Wu, S. Fendorf, J. Bargar, and C. S. Criddle, *Environ.*
410 *Sci. Technol.*, 2011, **45**, 4718–4725.
- 411 22. D. R. Lovley and E. J. P. Phillips, *Environ. Sci. Technol.*, 1992, **26**, 2228–2234.
- 412 23. R. Anderson, H. Vrionis, I. Ortiz-Bernad, C. Resch, P. Long, R. Dayvault, K. Karp, S.
413 Marutzky, D. Metzler, A. Peacock, D. White, M. Lowe, and D. Lovley, *Appl. Environ.*
414 *Microbiol.*, 2003, **69**, 5884–5891.
- 415 24. J. R. Lloyd, *Fems Microbiol. Rev.*, 2003, **27**, 411–425.
- 416 25. W.-M. Wu, J. Carley, T. Gentry, M. Ginder-Vogel, M. Fienen, T. Mehlhorn, H. Yan, S.
417 Caroll, M. Pace, J. Nyman, J. Luo, M. Gentile, M. Fields, R. Hickey, B. Gu, D. Watson,
418 O. Cirpka, J. Zhou, S. Fendorf, P. Kitanidis, P. Jardine, and C. Criddle, *Environ. Sci.*
419 *Technol.*, 2006, **40**, 3986–3995.
- 420 26. J. D. Wall and L. R. Krumholz, *Annu. Rev. Microbiol.*, 2006, **60**, 149–166.
- 421 27. S. B. Yabusaki, Y. Fang, P. E. Long, C. T. Resch, A. D. Peacock, J. Komlos, P. R. Jaffe,
422 S. J. Morrison, R. D. Dayvault, D. C. White, and R. T. Anderson, *J. Contam. Hydrol.*,
423 2007, **93**, 216–235.
- 424 28. W.-M. Wu, J. Carley, J. Luo, M. Ginder-Vogel, E. Cardenas, M. Leigh, C. Hwang, S.
425 Kelly, C. Ruan, L. Wu, J. Van Nostrand, T. Gentry, K. Lowe, S. Carroll, W. Luo, M.
426 Fields, B. Gu, D. Watson, K. Kemner, T. Marsh, J. Tiedje, J. Zhou, S. Fendorf, P.
427 Kitanidis, P. Jardine, and C. Criddle, *Environ. Sci. Technol.*, 2007, **41**, 5716–5723.
- 428 29. H. R. Beller, *Appl. Environ. Microb.*, 2005, **71**, 2170–2174.
- 429 30. H. S. Moon, J. Komlos, and P. R. Jaffe, *Environ. Sci. Technol.*, 2007, **41**, 4587–4592.
- 430 31. W.-M. Wu, J. Carley, S. J. Green, J. Luo, S. D. Kelly, J. V. Nostrand, K. Lowe, T.
431 Mehlhorn, S. Carroll, B. Boonchayanant, F. E. Löffler, D. Watson, K. M. Kemner, J.
432 Zhou, P. K. Kitanidis, J. E. Kostka, P. M. Jardine, and C. S. Criddle, *Environ. Sci.*
433 *Technol.*, 2010, **44**, 5104–5111.
- 434 32. M. Ginder-Vogel, B. Stewart, and S. Fendorf, *Environ. Sci. Technol.*, 2010, **44**, 163–169.
- 435 33. Z. Wang, S.-W. Lee, P. Kapoor, B. M. Tebo, and D. E. Giammar, *Geochimica et*
436 *Cosmochimica Acta*, 2013, **100**, 24–40.
- 437 34. B. D. Stewart, P. S. Nico, and S. Fendorf, *Environ. Sci. Technol.*, 2009, **43**, 4922–4927.
- 438 35. S. Kerisit, A. R. Felmy, and E. S. Ilton, *Environ. Sci. Technol.*, 2011, **45**, 2770–2776.
- 439 36. P. Gómez, A. Garralón, B. Buil, M. J. Turrero, L. Sánchez, and B. de la Cruz, *Science of*

- 440 *The Total Environment*, 2006, **366**, 295–309.
- 441 37. J. C. Pett-Ridge, V. M. Monastera, L. A. Derry, and O. A. Chadwick, *Chem. Geol.*, 2007,
442 **244**, 691–707.
- 443 38. C. Hansel, S. Benner, J. Neiss, A. Dohnalkova, R. Kukkadapu, and S. Fendorf,
444 *Geochimica et Cosmochimica Acta*, 2003, **67**, 2977–2992.
- 445 39. C. M. Hansel, S. G. Benner, and S. Fendorf, *Environ. Sci. Technol.*, 2005, **39**, 7147–7153.
- 446 40. H. D. Pedersen, D. Postma, R. Jakobsen, and O. Larsen, *Geochimica et Cosmochimica*
447 *Acta*, 2005, **69**, 3967–3977.
- 448 41. T. Borch, Y. Masue, R. K. Kukkadapu, and S. Fendorf, *Environ. Sci. Technol.*, 2007, **41**,
449 166–172.
- 450 42. A. M. Jones, R. N. Collins, J. Rose, and T. D. Waite, *Geochimica et Cosmochimica Acta*,
451 2009, **73**, 4409–4422.
- 452 43. R. M. Cornell and U. Schwertmann, *The Iron Oxides: Structure, Properties, Reactions,*
453 *Occurrences, and Uses*, Wiley-VCH GmbH & Co. KGaA, 2nd edn. 2003.
- 454 44. A. C. Cismasu, F. M. Michel, A. P. Tcaciuc, T. Tyliszczak, and J. Brown Gordon E,
455 *Comptes Rendus Geoscience*, 2011, **343**, 210–218.
- 456 45. A. C. Cismasu, F. M. Michel, J. F. Stebbins, C. Levard, and G. E. Brown Jr., *Geochimica*
457 *et Cosmochimica Acta*, 2012, **92**, 275–291.
- 458 46. Y. Masue-Slowey, R. H. Loeppert, and S. Fendorf, *Geochimica et Cosmochimica Acta*,
459 2011, **75**, 870–886.
- 460 47. C. M. Hansel, D. R. Learman, C. J. Lentini, and E. B. Ekstrom, *Geochimica et*
461 *Cosmochimica Acta*, 2011, **75**, 4653–4666.
- 462 48. Y. Masue, R. H. Loeppert, and T. A. Kramer, *Environ. Sci. Technol.*, 2007, **41**, 837–842.
- 463 49. S. M. Webb, *Physica Scripta*, 2005, 1011.
- 464 50. B. Ravel and M. Newville, *Journal of Synchrotron Radiation*, 2005, **12**, 537–541.
- 465 51. S. Zabinsky, J. Rehr, A. Ankudinov, R. Albers, and M. Eller, *Phys. Rev. B*, 1995, **52**,
466 2995–3009.
- 467 52. A. L. Ankudinov, J. J. Rehr, and S. D. Conradson, *Phys. Rev. B*, 1998, **58**, 7565–7576.
- 468 53. A. C. Larson and R. B. Von Dreele, *GSAS*, Los Alamos National Laboratory, 2000.
- 469 54. B. Toby, *J. Appl. Crystallogr.*, 2001, **34**, 210–213.
- 470 55. K. M. Campbell, H. Veeramani, K.-U. Ulrich, L. Y. Blue, D. E. Giammar, R. Bernier-
471 Latmani, J. E. Stubbs, E. Suvorova, S. Yabusaki, J. S. Lezama-Pacheco, A. Mehta, P. E.
472 Long, and J. R. Bargar, *Environ. Sci. Technol.*, 2011, **45**, 8748–8754.
- 473 56. E. Bazilevskaya, D. D. Archibald, M. Aryanpour, J. D. Kubicki, and C. E. Martínez,
474 *Geochimica et Cosmochimica Acta*, 2011, **75**, 4667–4683.
- 475 57. S. V. Yanina and K. M. Rosso, *Science*, 2008, **320**, 218–222.
- 476 58. R. M. Handler, B. L. Beard, C. M. Johnson, and M. M. Scherer, *Environ. Sci. Technol.*,
477 2009, **43**, 1102–1107.
- 478 59. K. M. Rosso, S. V. Yanina, C. A. Gorski, P. Larese-Casanova, and M. M. Scherer,
479 *Environ. Sci. Technol.*, 2010, **44**, 61–67.
- 480 60. D. E. Latta, J. E. Bachman, and M. M. Scherer, *Environ. Sci. Technol.*, 2012, **46**, 10614–
481 10623.
- 482 61. S. A. Crosby, D. R. Glasson, A. H. Cuttler, I. Butler, D. R. Turner, M. Whitfield, and G.
483 E. Millward, *Environ. Sci. Technol.*, 1983, **17**, 709–713.
- 484 62. J.-H. Jang, B. A. Dempsey, and W. D. Burgos, *Environ. Sci. Technol.*, 2007, **41**, 4305–
485 4310.

- 486 63. E. E. Roden and J. M. Zachara, *Environ. Sci. Technol.*, 1996, **30**, 1618–1628.
487
488

Figures

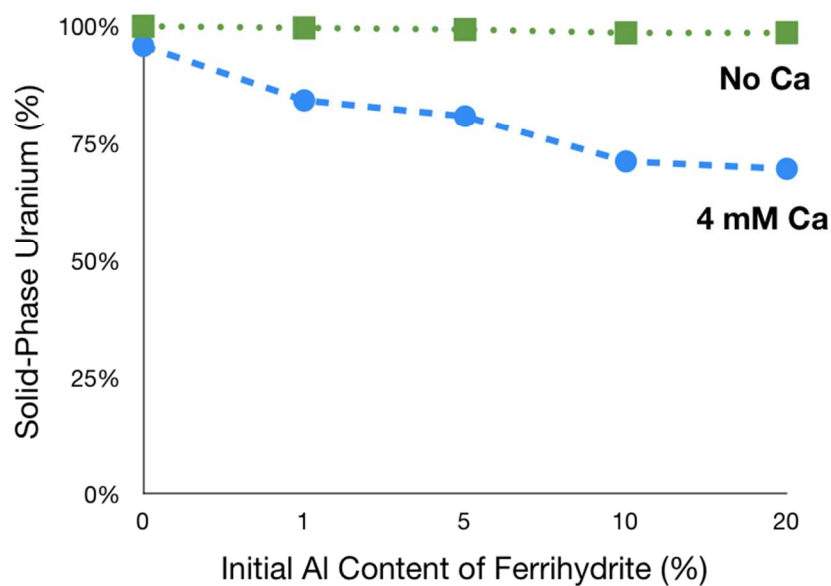


Figure 1. Uranium partitioned into the solid phase as a function of Al content of initial ferrihydrite slurry (0-20 mol% Al substitution for Fe). Incubations were performed with an initial $U(VI)_{(aq)}$ concentration of $10 \mu M$, in the presence or absence of $Ca^{2+}_{(aq)}$ (0 mM or 4 mM Ca), upon reaction with (Al-)ferrihydrite slurry, 0.3 mM Fe(II), and 3.8 mM carbonate at a pH of 7.0. Error bars are smaller than the data symbols.

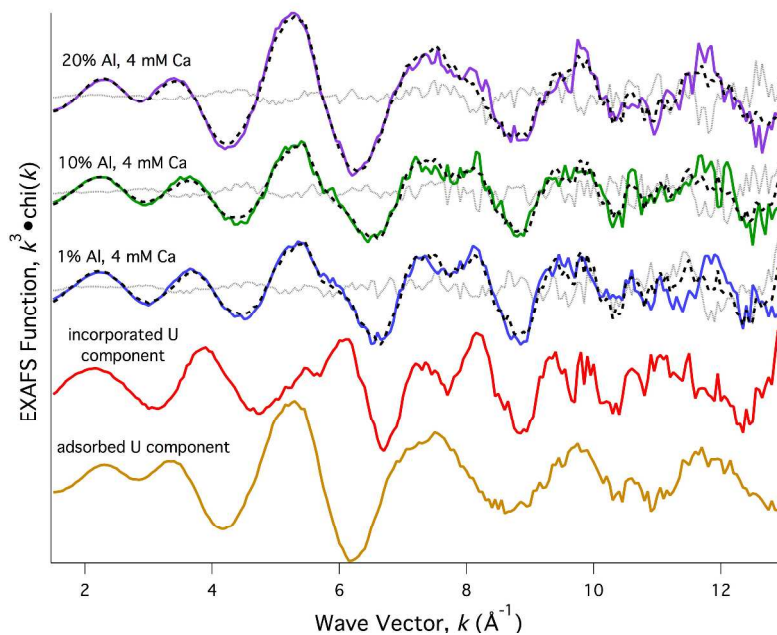


Figure 2. EXAFS linear combination fitting results for k^3 -weighted U L_3 -edge EXAFS spectra for Al-ferrihydrite (0-20 mol% Al substituted for Fe) reacted with $10 \mu\text{M}$ U(VI), 0.3 mM Fe(II), 3.8 mM carbonate, and 0 mM Ca, at pH 7.0. Data (colored lines), fitting components (incorporated U, adsorbed U), linear combination fits (black dotted lines), and residuals (light grey dotted lines) are shown.

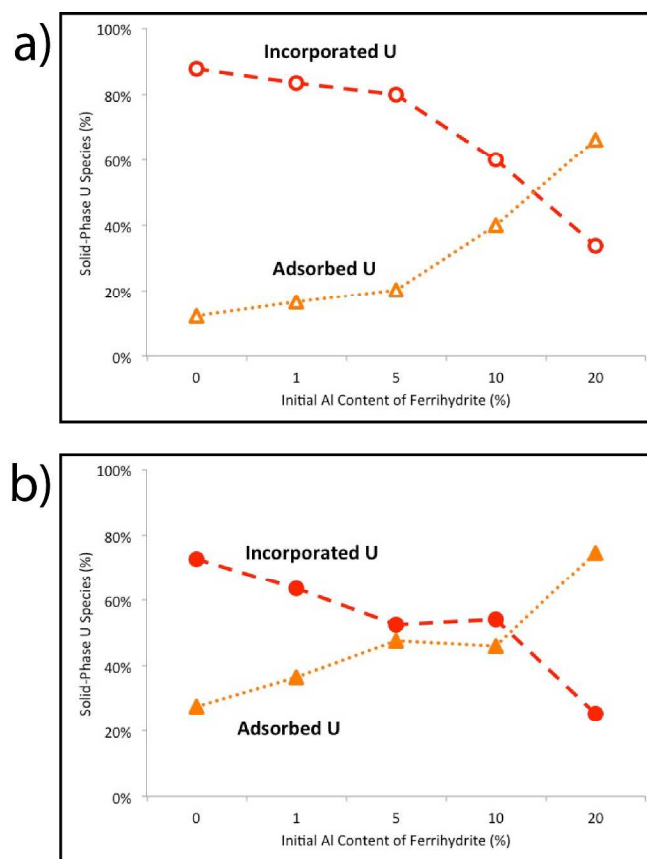


Figure 3. Solid phase U speciation as a function of initial Al content of ferrihydrite slurry (0-20 mol% Al substitution for Fe) after reaction with 10 μM initial $\text{U(VI)}_{(aq)}$, 0.3 mM Fe(II), 3.8 mM carbonate at pH 7.0 and a) 0 mM Ca, or b) 4 mM Ca. Linear combination fitting EXAFS percentages are shown for incorporated U(V) (red circles) and adsorbed U(VI) (yellow triangles).

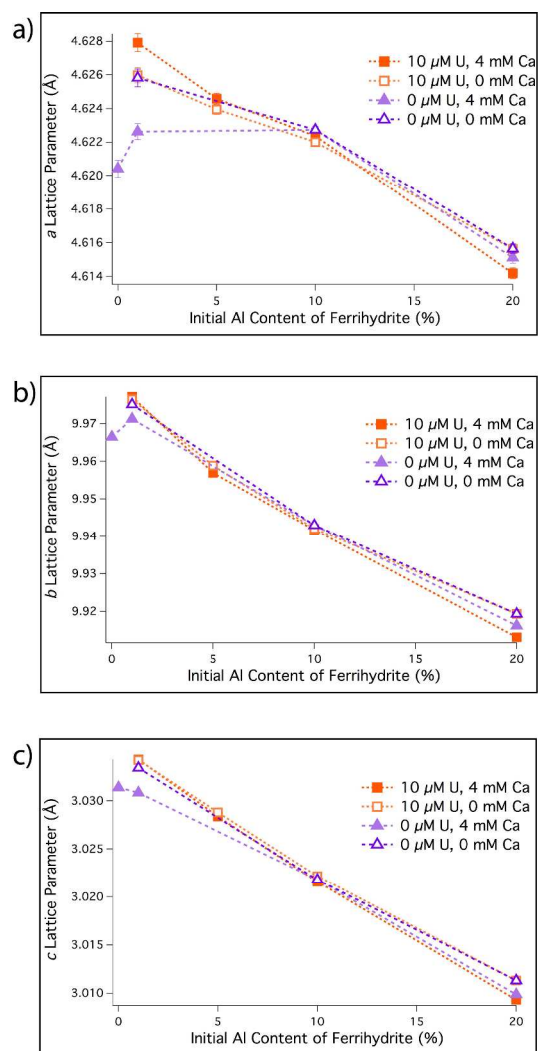


Figure 4. Lattice parameters of goethite as a function of Al content in ferrihydrate slurry (0-20 mol% Al substituted for Fe). Goethite was the result of the reaction of Al-ferrihydrate slurry with 10 μM initial $\text{U(VI)}_{(aq)}$, 0.3 mM Fe(II) , 3.8 mM carbonate, and 0 mM Ca or 4 mM Ca, at pH 7.0. Panels a, b, and c correspond to *a*, *b*, and *c* lattice parameters of goethite, respectively.

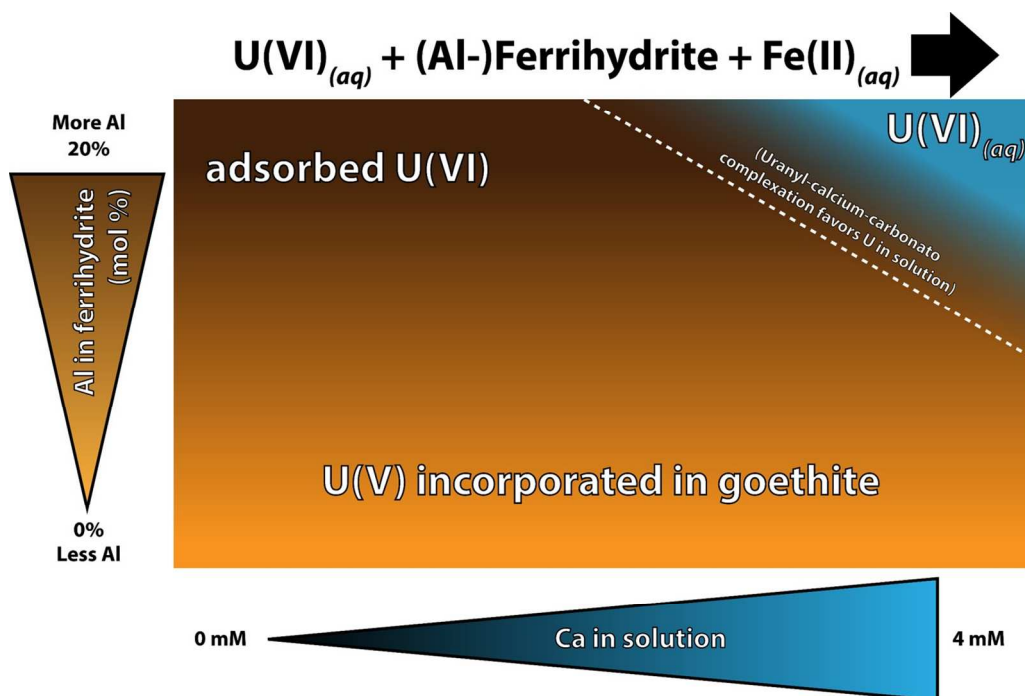


Figure 5. Conceptual illustration of the effects of Al substitution in ferrihydrite/goethite and aqueous Ca concentration on U fate for the reaction of U(VI) with (Al)ferrihydrite and Fe(II). The amount of U (as U(V)) incorporated into Al-goethite decreases with increasing Al content. Increased Ca concentration decreases U(VI) adsorption and when coupled with Al incorporation with ferrihydrite/goethite leads to great aqueous concentration of U (i.e., less retention).