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Journal:	<i>Environmental Science: Processes & Impacts</i>
Manuscript ID	EM-ART-06-2018-000286.R1
Article Type:	Paper
Date Submitted by the Author:	27-Aug-2018
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Reduction of PCE and TCE by Magnetite Revisited

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Abstract

Here we revisit whether the common mixed-valent Fe mineral, magnetite, is a viable reductant for the abiotic natural attenuation of perchloroethylene (PCE) and trichloroethylene (TCE) in anoxic groundwater plumes. We measured PCE and TCE reduction by stoichiometric magnetite as a function of pH and Fe(II) concentration. In the absence of added Fe(II), stoichiometric magnetite does not reduce PCE and TCE over a three month period. When Fe(II) is added to magnetite suspensions, PCE and TCE are reduced under Fe(II) and pH conditions that appear to be controlled by the solubility of ferrous hydroxide, Fe(OH)₂(s). Reduction rates are slow with only 1 to 30% carbon products (primarily acetylene) accumulating over several months. We conducted a similar set of experiments with Fe(OH)₂(s) alone and found that, compared to in the presence of magnetite, Fe(OH)₂(s) reduces PCE and TCE only at Fe(II) concentrations that are too high (≥ 13 mM, 726 mg/L) to be representative of natural aquifer conditions. Our results suggest that magnetite present in aquifer sediments alone is unlikely to reduce PCE and TCE sufficiently fast to contribute to natural attenuation of PCE and TCE. The lack of compelling evidence for PCE and TCE reduction by magnetite raises important questions regarding the current application of using magnetic susceptibility as a potential indicator for abiotic natural attenuation. Dynamic conditions and high Fe(II) concentrations that favor active precipitation of minerals, such as Fe(OH)₂(s) in the presence of magnetite (or other Fe minerals), however, may lead to PCE and TCE reduction that could help attenuate PCE and TCE plumes.

Environmental significance

We present evidence that suggests magnetite alone is unlikely to reduce PCE and TCE fast enough to significantly contribute to the natural attenuation of PCE and TCE in contaminated aquifer plumes. Under reducing conditions where high concentrations of ferrous iron may be present, however, active precipitation of metastable phases in the presence of magnetite (or other Fe minerals) may contribute to abiotic natural attenuation of PCE and TCE.

Submitted to *Environmental Science: Processes and Impacts*

Revised Manuscript

24 August 2018

40 Introduction

41 Perchloroethene (PCE) and trichloroethene (TCE) are chlorinated ethenes that were used at
42 thousands of industrial and commercial facilities as cleaning and metal degreasing solvents.^{1,2} For
43 decades, PCE and TCE have been, and continue to be, the most prevalent priority pollutants in
44 groundwater at hundreds of sites and a major concern for the environment and human health.³⁻⁷
45 Despite extensive cleanup efforts costing billions of dollars, PCE and TCE are still detected at many
46 groundwater sites at concentrations above regulatory limits.^{4,8,9} Recent concerns regarding energy
47 consumption and carbon emissions have made it even more critical to assess whether it is feasible
48 for site managers to rely on natural biological, chemical, and physical processes (i.e., natural
49 attenuation) to remediate chlorinated ethene plumes.¹⁰

50 Biological natural attenuation of PCE and TCE via reductive dechlorination has been extensively
51 studied. These investigations have resulted in several tools that can be used to provide lines of
52 evidence for biological natural attenuation including methods for determining the numbers of
53 *Dehalococcoides sp. (Dhc)* bacteria, vinyl chloride reductase gene copy numbers (*vcrA*),¹¹⁻¹⁴ and
54 compound-specific isotope analysis of PCE and TCE.^{15,16} In addition, significant evidence has
55 accumulated for biotic oxidation of chlorinated ethenes in aerobic plumes,^{17,18} and there are hints
56 that anaerobic oxidation may also occur coupled to metal reduction.¹⁹ There still remains significant
57 uncertainty, however, about how much abiotic reduction reactions, such as reduction by ferrous iron
58 (Fe(II)) bearing minerals, natural organic matter, and reduced sulfur species contribute to natural
59 attenuation.

60 Of the abiotic natural attenuation processes, reduction by Fe(II)-containing minerals has been
61 discussed as a potentially promising degradation pathway for chlorinated solvents for over two
62 decades.²⁰⁻²³ Abiotic PCE and TCE degradation remains of significant interest largely because abiotic
63 reduction most often occurs by reductive elimination with acetylene observed as the primary end-
64 product. Acetylene is a preferred end-product because it is benign and avoids the dichloroethene
65 and vinyl chloride stall that commonly occurs with biotic reductive dechlorination pathways.²⁰⁻²⁴

66 Of the Fe(II)-containing minerals, magnetite (Fe₃O₄), a common mixed-valent Fe mineral, has
67 been suggested to be responsible for chlorinated ethene attenuation at some field sites²⁴⁻²⁶ despite
68 slow rates of reduction by magnetite observed in laboratory experiments.^{24,27-29} Indeed, the promise
69 of magnetite as a reductant for chlorinated ethenes has recently led some to suggest that high
70 magnetic susceptibility of aquifer sediments could be used as a potential indicator for abiotic natural
71 attenuation at a site.²⁵ Such proxy methods are desirable due to the difficulty in measuring
72 biologically labile products such as acetylene. We also note that more recent evidence alternatively
73 implicates an oxidative abiotic pathway^{30,31} based on a Fenton-like process involving OH radical in
74 chlorinated ethene degradation by pyrite (FeS₂) in the presence of oxygen.

75 Here we revisit whether the common mixed-valent Fe mineral, magnetite, is a viable reductant to
76 contribute to abiotic natural attenuation of PCE and TCE in anoxic groundwater plumes. Our results
77 suggest that magnetite present in aquifer sediments, alone, is unlikely to reduce PCE and TCE
78 sufficiently fast to contribute to natural attenuation of PCE and TCE in anoxic plumes. Under
79 dynamic conditions that result in active precipitation of minerals such as Fe(OH)₂(s), however, PCE
80 and TCE may be reduced sufficiently fast to help attenuate PCE and TCE plumes.

81 **Materials and Methods**

82 **Chemicals**

83 Tetrachloroethylene (PCE, $\geq 99\%$) and trichloroethylene (TCE, $\geq 99\%$) were purchased from Sigma-
84 Aldrich. Non-chlorinated C_2 gases were used from a certified mixture containing 2.0% ethane, 1.97%
85 ethylene and 1.9% acetylene mixture in N_2 (Praxair). Hexanes and methanol (Fisher Scientific) were
86 pesticide residue grade and ACS reagent grade $\geq 99.8\%$, respectively. PCE and TCE (24, 250 & 500
87 mM stock solutions were gravimetrically prepared in N_2 -sparged methanol, sealed with viton septa
88 and stored in a glovebox.

89 All deionized water was deoxygenated by purging with N_2 and stored for 24 hours in an anoxic
90 glovebox (93% N_2 /7% H_2) before being used in any experiments. Buffer solutions of 10 mM 3-(N-
91 Morpholino)propanesulfonic acid (MOPS, RPI Corp.) buffer solution, or 10 mM Piperazine-N,N'-
92 bis(3-propanesulfonic acid) (PIPPS, GFS Chemicals) with 10 mM sodium chloride (NaCl) background
93 electrolyte were prepared. All pH adjustments were done with deoxygenated hydrochloric acid (HCl)
94 or sodium hydroxide (NaOH).

95 Ferrous chloride stock solutions (~ 1.4 M $FeCl_2 \cdot 4H_2O$) were further purified by adjusting the pH to
96 ~ 4.5 and filtering to remove any Fe(III) precipitates.

98 **Magnetite synthesis**

99 Magnetite was synthesized using iron chloride salts following the method used as previously
100 described.^{32,33} Briefly, 0.1 M ferrous chloride ($FeCl_2 \cdot 4H_2O$) and 0.2 M ferric chloride ($FeCl_3 \cdot 6H_2O$)
101 solutions were prepared in deoxygenated deionized water within the glovebox. Both solutions were
102 combined. The mixture was vigorously stirred and titrated using 10 M NaOH to set the pH between
103 10.0 and 11.5. The magnetite suspension was sealed and left overnight before filtering. The minerals
104 were removed from the glovebox in a sealed vessel and freeze-dried. Freeze dried minerals were
105 ground with a mortar and pestle and sieved through a 150-micron sieve and stored in the glovebox.
106 With this approach, the (~ 20 nm) magnetite particles have surface area values close to the
107 previously reported 63 ± 7 m² g⁻¹ using N_2 -BET analysis.³³

109 **Mineral characterization**

110 Magnetite stoichiometries ($x = Fe^{2+}/Fe^{3+}$) were determined using previously established methods.³²
111 The first approach was by acid dissolution (x_d) of the magnetite in 5 M HCL under glovebox
112 atmosphere. Using the 1,10-phenanthroline method,^{32,34} we evaluated the Fe^{2+} and total Fe
113 concentrations to determine the stoichiometric ratio. Powder X-ray diffraction (x_{xrd}) was the second
114 approach using a Rigaku MiniFlex II system equipped with a Co source. Magnetite powders were
115 mixed into two drops of glycerol to form a well-mixed paste in the glovebox to avoid oxidation of the
116 mineral during analysis. The powder X-ray diffraction stoichiometries (x_{xrd}) were then derived from
117 unit-cell dimension.³² Transmission Mössbauer spectroscopy was performed with a variable
118 temperature He-cooled system with a ⁵⁷Co source. Unless otherwise noted, Mössbauer spectra were
119 collected at 140 K. Spectra were fit using the Recoil software³⁵ and procedures outlined by Gorski
120 and Scherer.³² To prepare samples and avoid oxidation, we sealed samples with Kapton tape in the
121 glovebox. To characterize minerals after reaction, we shook the reactor and filtered a 5 mL aliquot.

1
2
3 122 These post-reaction samples were sealed for XRD using a layer of Kapton tape over the sample to
4 123 prevent rapid oxidation of $\text{Fe}(\text{OH})_2$.

5 124

6 125 **Reactor design**

7 126 Reactors were 160 mL glass serum bottles sealed with Viton fluoroelastomer septas (20 mm x 8 mm
8 127 depth, Wheaton) and contained 150 mL of liquid and 10 mL of headspace. All batch reactors were
9 128 prepared in a N_2/H_2 filled glovebox with an oxygen content below 1 ppm. Each system was covered
10 129 with foil, stored and mixed (~100 rpm) upside down to have the headspace in contact with glass
11 130 rather than the septum. Mixing reactors upside down was important to minimize headspace gas loss
12 131 through the septum. The desired mass of iron oxides was added to the buffer solution, then the
13 132 reactor pH was adjusted (when needed) and the chlorinated solvent added via a spike from the stock
14 133 solution. A PCE and TCE concentration of 50 μM (8,291 and 6,570 $\mu\text{g}/\text{L}$ of PCE and TCE, respectively)
15 134 was used for most experiments. Reactors contained either 10 mM MOPS (pH range of 7.5 to 8.0) and
16 135 10 mM NaCl or 10 mM PIPPS (pH > 8.0) and 10 mM NaCl as buffer and background electrolyte,
17 136 respectively.

18 137 In magnetite reactors containing $\text{Fe}(\text{II})$, bottles containing buffer and background electrolyte
19 138 were first spiked with $\text{Fe}(\text{II})$ from an 1.4 M FeCl_2 stock, and then the initial $\text{Fe}(\text{II})$ concentration was
20 139 measured. In reactors containing magnetite, magnetite was added, and then the systems were
21 140 titrated to the desired pH using 2.5 or 10 M NaOH. Because titration of $\text{Fe}(\text{II})$ solutions results in a pH
22 141 plateau near pH 8.0, we took care to add the same volume of NaOH to reactors that were prepared
23 142 as replicates. Once the pH was adjusted, a 500 μL sample was collected after filtration and the final
24 143 $\text{Fe}(\text{II})$ concentration ($[\text{Fe}(\text{II})]_i$) of the filtrate was determined. Reactors containing only $\text{Fe}(\text{II})$ were
25 144 prepared as described above, but without the addition of magnetite.

26 145

27 146 **Analytical Procedures**

28 147 Analyses of parent and product analytes were performed using an Agilent 6890 gas chromatograph
29 148 equipped with electron capture (ECD) and flame ionization (FID) detectors. PCE and TCE were
30 149 quantified with GC-ECD after a liquid-liquid extraction of 0.25 to 1 mL of sample containing both the
31 150 aqueous and solid phases added to 2 mL of hexanes. The daughter products ethane, ethylene,
32 151 acetylene, the dichloroethenes, and vinyl chloride were detected using GC-FID. Further details on
33 152 analytical methods are provided in the Supporting Information.

34 153

35 154 **Results and Discussion**

36 155 **Reduction of PCE and TCE by Magnetite**

37 156 To evaluate whether magnetite reduces PCE and TCE, we measured PCE and TCE reduction by
38 157 stoichiometric magnetite ($x = \text{Fe}^{2+}/\text{Fe}^{3+} \approx 0.5$) over a range of pH values (7.0 – 8.0) and solids loading
39 158 (5 – 20 g/L). In all experiments, we observed negligible loss of both PCE and TCE over 140 days
40 159 (**Figure 1 and Table S1 in the Supporting Information**). Further, no carbon products were observed,
41 160 and we were able to recover nearly all of the carbon initially present in the system (recoveries were
42 161 TCE: $105 \pm 8\%$ for $n = 7$; PCE: $98 \pm 6\%$ for $n = 5$, details in **Table S1**). We originally anticipated that
43 162 magnetite stoichiometry would affect the rate of PCE and TCE reduction by magnetite as we have

163 previously observed that magnetite stoichiometry strongly influenced the rates and extent of
 164 uranium, mercury, and nitroaromatic compound reduction.³⁶⁻³⁸ Here, however, we observed no
 165 measurable reduction of PCE and TCE by magnetite even with stoichiometric magnetite ($x \approx 0.5$) and
 166 therefore did not further explore reduction of PCE and TCE by non-stoichiometric magnetite.

167 The complete lack of PCE and TCE
 168 reduction by stoichiometric magnetite was
 169 unexpected as PCE and TCE reduction by
 170 magnetite has been previously reported.²⁶⁻
 171 ²⁹ A close look at the data, however, reveals
 172 that prior evidence for reduction of PCE and
 173 TCE by magnetite is somewhat limited.
 174 While an early study by Sivavec and Horney
 175 reported fast rates of TCE reduction by
 176 magnetite with a surface-area normalized
 177 rate coefficient (k_{SA}) of $4.5 \times 10^{-4} \text{ L m}^{-2} \text{ d}^{-1}$
 178 (half-life of 19 days),²⁸ more recent
 179 studies²⁶⁻²⁹ reported about 100-fold slower
 180 rates ($k_{SA} \approx 10^{-6} \text{ L m}^{-2} \text{ d}^{-1}$) than those in
 181 Sivavec and Horney.²⁸ In addition, carbon
 182 reduction products were only reported in
 183 one study²⁹, and in that study the
 184 uncertainties on product measurements
 185 were quite large (30 – 200%). Based on the
 186 high uncertainties, the authors
 187 appropriately chose not to quantify rates of
 188 reduction.²⁹ Of the other two studies, the anomalously high rate of reduction was based on TCE loss
 189 alone with no report of products²⁸ and the other study relied on chloride accumulation to indicate
 190 that reduction had occurred.²⁷ Given the rather limited observations of reduced carbon products
 191 and high uncertainties in the previously reported data, we suggest that our observation of negligible
 192 reduction of PCE and TCE by magnetite is not so surprising.

193 Despite the limited laboratory evidence for PCE and TCE reduction by magnetite, degradation
 194 rates extrapolated from fate and transport modelling of monitoring well concentrations in the field
 195 have sometimes been attributed to abiotic attenuation of chlorinated ethenes by magnetite.^{24-26,39}
 196 For example, aquifer degradation rate coefficients ($k_{aquifer}$) on the order of $\sim 1 \text{ yr}^{-1}$ were extrapolated
 197 from contaminated groundwater at Twin Cities Army Ammunition Plant (TCAAP) to describe
 198 chlorinated ethene plumes that were smaller than expected based on dilution.^{26,44,45} More extensive
 199 data was collected for cis and 1,1-DCE and the combination of plume attenuation beyond dilution,
 200 microcosm data showing similar degradation rates with sterile and autoclaved sediments, and the
 201 presence of magnetite in the sediments led the authors to conclude that the loss of DCE may be due
 202 to reaction with magnetite.²⁶ While this a reasonable conclusion to draw, it is not, as was noted by
 203 Ferrey et. al., the only potential explanation for their observations.²⁶

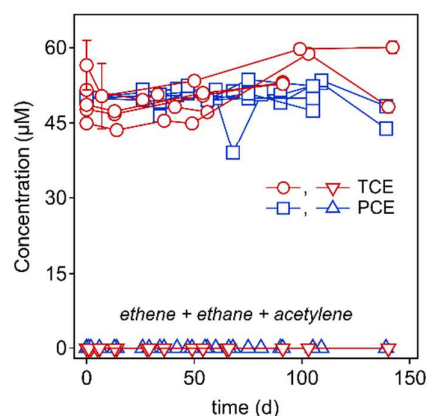


Fig. 1. PCE and TCE concentration versus time in the presence of stoichiometric magnetite ($x = \text{Fe}^{2+}/\text{Fe}^{3+} = 0.46 - 0.50$). Experimental conditions: 50 μM PCE/TCE, 10 mM MOPs/NaCl at pH 7.0, 7.5, and 8.0 for TCE reactors and pH 7.5 for PCE reactors, mass loading 5 - 20 g/L. Average carbon recoveries, TCE ($105 \pm 8\%$) for ($n = 7$), and PCE ($98 \pm 6\%$), for ($n = 5$). Error bars represent 1σ for a set of triplicate reactors (pH 8.0, 5 g/L, TCE)

Nevertheless, this work²⁶ is often cited as evidence for chlorinated ethene (PCE, TCE, and DCE) reduction by magnetite^{24-26,39} despite the data only being for DCEs, a need to use rate constants from batch reactors containing high Fe(II) concentrations, and the authors carefully stating that the loss of DCE *may be* due to reaction with magnetite and further research into the mechanisms of loss is needed. Collectively, the limited field and laboratory data for reduction of PCE and TCE by magnetite, and our observation of negligible reduction of PCE and TCE by magnetite, suggests that reduction by magnetite alone is unlikely responsible for field extrapolated degradation rates of PCE and TCE that have been previously attributed to magnetite. We note that there are some important differences between our experiments and aquifer plume conditions (e.g., buffers, magnetite crystallinity, flow environment, etc.). Taken together however, the existing literature and our findings provide no rigorous evidence that magnetite reduces PCE and TCE under anoxic conditions, and in fact, suggest that magnetite does not reduce PCE and TCE under a variety of conditions.

Reduction of PCE and TCE by Magnetite and Fe(II)

To investigate whether magnetite plus aqueous Fe(II) can abiotically degrade PCE and TCE, we measured the reduction of PCE and TCE by magnetite in the presence of Fe(II) over a range of Fe(II) concentrations and pH values (Table S2 and S3).

Consistent with previous chlorinated ethene work,^{27,40} we found that

adding Fe(II) to magnetite suspensions did, in some cases, result in PCE and TCE reduction. For example, we observed measurable loss of PCE and TCE and accumulation of 25% carbon products

(primarily as acetylene) with 5 g/L magnetite and 9.3 mM Fe(II) or 33 mM Fe(II), for PCE and TCE, respectively (Figure 2). Our observation of acetylene as the primary product in all reactors suggests that reductive β -elimination of PCE and TCE was likely the primary mechanism for reduction.⁴¹ Carbon recovery was higher with TCE (99 \pm 4.3%) than PCE (78%). In an attempt to better close the PCE mass balance, we measured for reductive dechlorination products, including dichloroethenes and vinyl chloride, but did not detect any. It is unclear if the mass balance loss of PCE is due to reduction to an unknown product or loss due to sorption or volatilization, however, we did observe up to 30% PCE and TCE loss in controls sampled more frequently than these reactors (Table S1), suggesting volatilization was likely the main contributor to loss.

To avoid including potential sorption and/or volatilization losses in estimated rates of PCE and TCE reduction, we quantified reduction rates based solely on accumulation of carbon reduction products (primarily acetylene). We modeled product accumulation over time using an exponential product in-growth equation (Equation 1), where $P(t)$ is the mass of products (in μ moles C) at time t ,

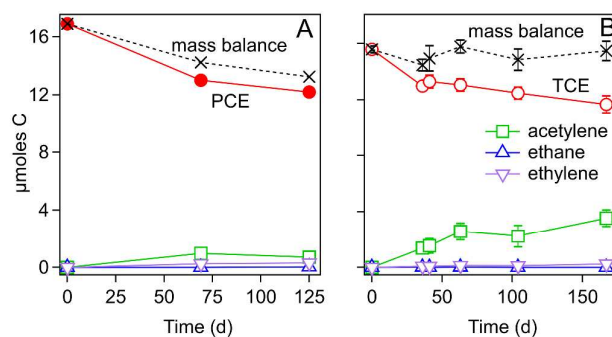


Fig. 2. Reduction of (A) PCE and (B) TCE over time with magnetite and aqueous Fe(II). Experimental conditions: (A) 54 μ M PCE, 5 g/L Fe_3O_4 (s), 33 mM Fe(II), 10 mM MOPs/NaCl, pH 7.9, single reactor. (B) 50 μ M TCE, 5 g/L Fe_3O_4 (s), 9.3 \pm 0.6 mM Fe(II), 10 mM MOPs/NaCl, pH 8, for n from 3 to 12.

246 C_0 is the initial PCE or TCE amount (in $\mu\text{moles C}$), and k_{obs} is the first-order rate coefficient. Data for
 247 total products was linearized (Equation 2), and the slope of the linearized product accumulation
 248 versus time equation was used to determine k_{obs} .

249

$$P(t) = C_0(1 - e^{-k_{\text{obs}}t}) \quad (1)$$

251

$$\frac{\ln\left(1 - \frac{P(t)}{C_0}\right)}{t} = -k_{\text{obs}} \quad (2)$$

253

254 First-order rate coefficients for PCE and TCE reduction determined from product accumulation
 255 varied considerably depending on the Fe(II) concentration and pH and ranged from zero to 5.2×10^{-8}
 256 s^{-1} (**Table S2** and **S3**).

257 We plotted k_{obs} values versus initial
 258 Fe(II) concentration and pH to explore
 259 trends in reduction rate, as we anticipated
 260 reduction rates might increase with both
 261 pH and Fe(II) concentration (**Figure 3**).
 262 Although our data are concentrated
 263 around our most common condition (\sim pH
 264 8 and \sim 10 mM Fe(II)), there are still
 265 sufficient data to conclude that there is no
 266 clear trend with Fe(II) concentration. With
 267 pH, there is a pattern of higher pH
 268 reactors having measurable reduction
 269 ($k_{\text{obs}} > 0$), but no obvious trend of
 270 increasing rates at higher pH values. In
 271 addition, there was visual evidence from
 272 our reactors that experimental conditions
 273 had some influence on whether products
 274 were observed. More specifically, we
 275 observed a white precipitate forming in
 276 reactors in which products were observed
 277 (**Figure S1**). Given the pH and Fe(II)
 278 concentrations we used, we suspected
 279 that the white precipitate was ferrous
 280 hydroxide, $\text{Fe}(\text{OH})_2(\text{s})$, which was
 281 confirmed by X-ray diffraction of the
 282 filtered reactor solids (**Figure S2**).

283 To quantitatively evaluate if $\text{Fe}(\text{OH})_2(\text{s})$ precipitation was necessary for PCE and TCE reduction, we
 284 plotted k_{obs} values versus the saturation index (SI) for $\text{Fe}(\text{OH})_2(\text{s})$ (**Figure 4**). We calculated the SI as
 285 $\log(\text{IAP}/K_{\text{sp}})$ using the initial Fe(II) concentration and an $\text{Fe}(\text{OH})_2(\text{s})$ K_{sp} of 5×10^{-15} .⁴²⁻⁴⁴ A noticeable
 286 pattern emerges with PCE and TCE reduction corresponding to conditions where the initial SI > 0 .

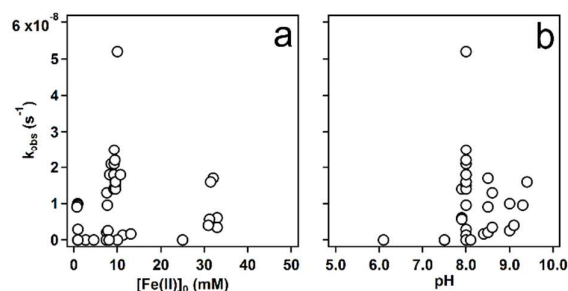


Fig. 3. Plots of k_{obs} for PCE and TCE reduction as a function of $[\text{Fe}(\text{II})]_0$ (initial Fe(II) concentration) (a) and pH (b)

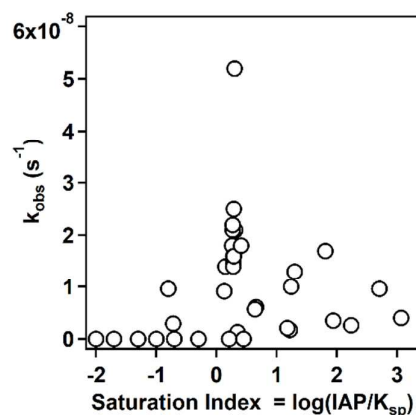


Fig. 4. Plot of k_{obs} for PCE and TCE reduction vs. the $\text{Fe}(\text{OH})_2(\text{s})$ initial saturation index (< 0 undersaturated, 0 at saturation, > 0 oversaturated).

287 This pattern implies that PCE and TCE reduction only occurs under experimental conditions that are
 288 saturated (or super-saturated) with respect to $\text{Fe}(\text{OH})_2(\text{s})$ based on the amount of $\text{Fe}(\text{II})$ added and
 289 the pH value the reactor was set at. We note that after precipitation of $\text{Fe}(\text{OH})_2$, the SI is likely closer
 290 to 0. For experiments that were saturated
 291 or super-saturated, 93% (27 of 29) had
 292 measurable reduction products form.
 293 Conversely, for experiments that were
 294 undersaturated only 18% (2 of 11) had
 295 measurable reduction products form.

296 To better visualize the influence of
 297 $\text{Fe}(\text{OH})_2(\text{s})$ precipitation on PCE and TCE
 298 reduction, we also plotted the initial pH and
 299 $\text{Fe}(\text{II})$ concentration of these magnetite plus
 300 $\text{Fe}(\text{II})$ experiments on an $\text{Fe}(\text{OH})_2(\text{s})$
 301 solubility diagram (**Figure 5**). In Figure 5,
 302 solid, red markers indicate conditions
 303 where reduction products were observed,
 304 whereas open markers indicate conditions
 305 where no reduction products were
 306 observed. In addition, we scaled the size of
 307 the red markers to the relative amount of
 308 products formed.

309 A clear visual picture emerges
 310 highlighting that products were observed
 311 (i.e., red markers) only when $\text{Fe}(\text{II})$
 312 concentration and pH values were such that
 313 $\text{Fe}(\text{OH})_2(\text{s})$ was expected (and visually
 314 observed) to precipitate. Of the forty
 315 experiments we conducted, thirty-six (90%)
 316 followed the trend of reduction occurring
 317 only when conditions were such that
 318 $\text{Fe}(\text{OH})_2(\text{s})$ was expected to precipitate.
 319 These results provide compelling evidence
 320 that, under our experimental conditions,
 321 precipitation of ferrous hydroxide is
 322 necessary for PCE and TCE reduction to be
 323 observed in the presence of magnetite.
 324 Note that in the presence of different
 325 buffers, such as carbonate, other minerals would likely precipitate (such as siderite or carbonate
 326 green rust).

327 Although adding $\text{Fe}(\text{II})$ to the magnetite suspensions resulted in reduction of PCE and TCE when
 328 $\text{Fe}(\text{OH})_2(\text{s})$ precipitated, the rates of reduction are still quite slow. Only 1 to 30% carbon products

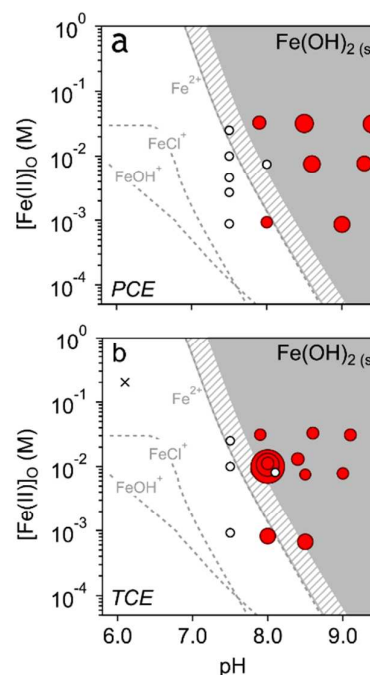


Fig. 5. $\text{Fe}(\text{OH})_2(\text{s})$ solubility diagram with magnetite plus $\text{Fe}(\text{II})$ reactor conditions overlaid for (a) PCE and (b) TCE reactors. $\text{Fe}(\text{II})$ is plotted as the initial $\text{Fe}(\text{II})$ concentration – $[\text{Fe}(\text{II})]_0$. Red markers represent PCE and TCE reactors where carbon products were observed and the markers are scaled relative to the amount of products produced (ranging up to 46%). Black open markers represent reactors where no products were observed. The grey hatched area represents a range of ferrous hydroxide solubility product K_{sp} ($[\text{Fe}(\text{OH})_2(\text{s})] = [\text{Fe}^{2+}] [\text{OH}^-]^2$) of $10^{-14.51}$ to $10^{-15.11}$. The grey hatched and shaded areas represent the region where $\text{Fe}(\text{OH})_2$ is expected to precipitate. Speciation diagrams were calculated with Visual MINTEQ for 10 mM Na^+ and 30 mM Cl^- to represent Cl^- added with a nominal $\text{Fe}(\text{II})$ spike of 10 mM.

329 (primarily acetylene) accumulated over a three to five month time period. Surface-area normalized
 330 first-order rate coefficients (k_{SA}) estimated from carbon product accumulation ranged from 3.7×10^{-7}
 331 to $1.5 \times 10^{-5} \text{ L m}^{-2} \text{ d}^{-1}$, corresponding to half-lives ($t_{1/2}$) ranging from 0.42 to 17 years (**Table S2 and**
 332 **S3**). These rates are much slower than those for TCE reported by Sivavec in experiments conducted
 333 at pH 6.0 and 200 mM Fe(II) with 217 g/L magnetite ($t_{1/2} \sim 3 \text{ d}$).⁴⁰ We attempted to reproduce the
 334 experimental conditions used in the Sivavec
 335 patent but observed no measurable TCE
 336 reduction (**Figure S3**). One possible
 337 explanation for the rapid TCE reduction
 338 observed by Sivavec is that the conditions
 339 may have led to formation of green rust,
 340 which has been shown to reduce PCE and
 341 TCE.⁴⁵ Note that our replication of Sivavec's
 342 conditions (marked by \times in **Figure 5**) are
 343 well below ferrous hydroxide saturation
 344 and therefore consistent with our finding
 345 that no reduction by magnetite occurs
 346 under these conditions.
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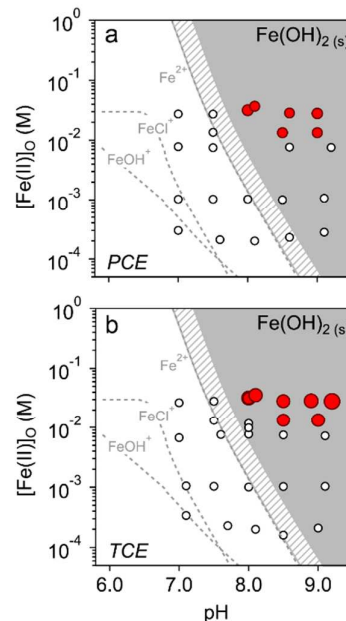


Fig. 6. $\text{Fe}(\text{OH})_2(\text{s})$ solubility diagram with aqueous Fe(II) reactors without magnetite conditions overlaid for PCE and TCE reactors. Fe(II) is plotted as the initial Fe(II) concentration – $[\text{Fe}(\text{II})]_0$. Red markers represent PCE and TCE reactors where products were observed and the markers are scaled relative to the amount of products produced (ranging up to 46%). Black open markers represent reactors where

348 **Reduction of PCE and TCE by Fe(II) and** 349 **ferrous hydroxide**

350 Our observation that $\text{Fe}(\text{OH})_2(\text{s})$
351 precipitation is necessary for PCE and TCE
352 reduction to occur in the presence of
353 magnetite led us to wonder whether
354 $\text{Fe}(\text{OH})_2(\text{s})$ or even aqueous Fe(II) alone can
355 reduce TCE. To evaluate whether $\text{Fe}(\text{OH})_2(\text{s})$
356 alone or aqueous Fe(II) alone could reduce
357 TCE, we measured TCE reduction over a wide range of Fe(II) concentrations and pH values (in the
358 absence of magnetite) and plotted the results on an $\text{Fe}(\text{OH})_2(\text{s})$ solubility diagram (**Figure 6**). As
359 expected, no reduction of PCE and TCE was observed by aqueous Fe(II) (denoted by open markers).
360 For most conditions, no PCE and TCE reduction were observed even when $\text{Fe}(\text{OH})_2(\text{s})$ had
361 precipitated. However, at *very* high concentrations of initial Fe(II) (> 13 mM, 726 mg/L) some slow
362 PCE and TCE reduction was observed with 0.3 to 13% products accumulating over a five to six month
363 time period. First-order rate coefficients for PCE and TCE reduction determined from product
364 accumulation for these high Fe(II) experiments (in absence of magnetite) ranged from $1.8 \times 10^{-10} \text{ s}^{-1}$
365 to $1.8 \times 10^{-8} \text{ s}^{-1}$ (**Table S4** and **S5**). We thought the high Fe(II) concentrations might have resulted in
366 precipitation of an additional phase possibly via secondary mineral transformation of the $\text{Fe}(\text{OH})_2(\text{s})$.
367 However, XRD and Mössbauer spectroscopy of the solids after 150 days indicated no additional
368 phases present other than $\text{Fe}(\text{OH})_2(\text{s})$ (**Figures S4** and **S5**). While these results are interesting, we
369 would like to emphasize that the high concentrations of Fe(II) (> 13 mM, 726 mg/L) make these
370 conditions unlikely to be relevant in groundwater aquifers where Fe(II) concentrations rarely exceed
371 50 mg/L.⁴⁶⁻⁴⁹ Finally, we caution that we are using $\text{Fe}(\text{OH})_2$ saturation index and the plot in Figures 5
372 and 6 as a graphical aid to explain our data. The figures should not be used as a predictive tool
373 because if $\text{Fe}(\text{OH})_2$ precipitates the saturation index will likely be near 0.

374 **Conclusions and Environmental Implications**

375 **Environmental Implications**

376 Our results, combined with previous laboratory studies,^{27,29} suggest that magnetite in aquifer
377 sediments is, on its own, unlikely to reduce PCE and TCE and contribute significantly to natural
378 attenuation of PCE and TCE in anoxic plumes. We did, however, find that precipitation of ferrous
379 hydroxide in the presence of aquifer minerals such as magnetite might provide a mechanism for
380 abiotic attenuation of chlorinated ethenes. To evaluate whether PCE and TCE reduction by
381 magnetite in the presence of ferrous hydroxide could be important in natural (or engineered)
382 attenuation strategies at contaminated sites, we scaled our laboratory rate coefficients (k_{SA}) to
383 reflect aquifer conditions (k_{field}) (example calculations in **Supporting Information**). Using a field
384 magnetite content of 1 g kg^{-1} , we estimated field-scaled first-order rate coefficients (k_{field}) ranging
385 from 0.070 to 2.8 yr^{-1} ($t_{1/2}$ from 0.25 to 9.9 years) (**Table S2**). These rates are comparable to both
386 sediment microcosm rates and field rates that have been attributed to abiotic degradation.^{24,26}
387 However, in nearly all of our experiments the amount of Fe(II) added exceeds typical concentrations

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3 388 of Fe(II) in groundwater (typically < 50 mg/L or 1 mM),⁴⁶⁻⁴⁹ making the precipitation of ferrous
4 389 hydroxide unlikely to be relevant to field conditions. Furthermore, precipitation of ferrous hydroxide
5 390 is unlikely in natural aquifers due the ubiquitous presence of carbonate, making siderite and ferrous
6 391 hydroxy-carbonate species more likely candidates for Fe(II) precipitates. Our work does, however,
7 392 highlight that active precipitation of reactive Fe(II) phases may be important in assessing abiotic
8 393 natural attenuation.

9
10 394 The lack of compelling evidence for PCE and TCE reduction by magnetite raises important
11 395 questions regarding whether magnetic susceptibility of aquifer sediments is a useful indicator for
12 396 abiotic degradation of chlorinated ethenes by magnetite.²⁵ Recent work has suggested that *in situ*
13 397 magnetic susceptibility measurements might be used along with chlorinated ethene concentration
14 398 decreases in monitoring wells as a line of evidence for the occurrence of natural attenuation by
15 399 sediment magnetite in an aquifer.²⁵ Our findings, however, suggest that magnetic susceptibility may
16 400 not be a useful indicator for abiotic natural attenuation of chlorinated ethenes by reductive
17 401 elimination. While magnetite and maghemite have magnetic susceptibilities of 2-3 orders of
18 402 magnitude greater than the other Fe oxides, making magnetic susceptibility a reasonable proxy for
19 403 sediment magnetite and/or maghemite content, the magnetic susceptibility of magnetite and fully-
20 404 oxidized maghemite are within 20% of each other.⁵⁰ The similarity between maghemite and
21 405 magnetite makes magnetic susceptibility measurements at the field level nearly insensitive to Fe
22 406 redox speciation. Furthermore, the correlation shown in Weidemier et al.²⁵ between chlorinated
23 407 ethene degradation and magnetic susceptibility is weak, with $r^2 = 0.18$ and a Spearman's rank-order
24 408 correlation coefficient of 0.41, making the correlation statistically non-significant at even a 90%
25 409 confidence interval (critical $p = 0.49$).

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29 410 Our work, together with the poor mechanistic and statistical correlation between magnetic
30 411 susceptibility and abiotic chlorinated ethene reduction by magnetite suggests that further
31 412 measurements are needed to link iron mineralogy to abiotic natural attenuation. Although our work
32 413 suggests that magnetite is not likely to contribute significantly to abiotic degradation of PCE and TCE,
33 414 there is substantial laboratory evidence that other Fe(II)-containing Fe minerals, such as
34 415 mackinawite (FeS) and green rust, reduce chlorinated ethenes much faster (see reviews by He;
35 416 Fan).^{24,51} The faster PCE and TCE reduction rates for mackinawite and green rust suggest that
36 417 reduction by these minerals may be an important degradation process in contaminated plumes.
37
38 418 Indeed, sulfate amendments to induce biological formation of mackinawite in-situ has been
39 419 demonstrated in the field and in the laboratory.⁵²⁻⁵⁶ Consistent with active precipitation of FeS being
40 420 important for continued degradation of PCE and TCE in sulfate-reducing field sites,⁵⁴ it is possible
41 421 that Fe(II) precipitation as Fe(OH)₂ and green rusts might occur in dynamic environments receiving a
42 422 constant flux of Fe(II) from dissimilatory Fe reduction, or in zones where a change in pH occurs. We
43 423 suggest that field screening methods for acid volatile sulfides targeting FeS⁵⁷ and citrate-bicarbonate
44 424 (CB) extractable Fe targeting green rusts and labile Fe(II) phases^{58,59} might provide a measure of the
45 425 potential for abiotic chlorinated ethene reduction by Fe(II) and sulfide minerals, although further
46 426 study is needed. Zones of active Fe(II) precipitation in anoxic aquifer could result in PCE and TCE
47 427 reduction that is sufficiently fast to help attenuate PCE and TCE plumes.

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52 429 **Conflicts of interest**

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3 430 There are no conflicts to declare.

4 431

5 432 **Acknowledgements**

6 433 This research has been funded entirely with funds from the Strategic Environmental Research and
7 434 Development Program – (SERDP) project ER-2532, and an NSF – GRFP to Johnathan Culpepper. The
8 435 contents do not necessarily reflect the views and policies either of the sponsors nor does mention of
9 436 trade names or commercial products that constitute endorsement or recommendation for use. The
10 437 authors would like to thank Luiza Notini (University of Iowa) for designing the table of contents art.

11 438

12 439 **Electronic Supplementary Information**

13
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15 440 † Electronic Supplementary Information (ESI) available: See DOI: 10.1039/x0xx00000x

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18 443 **References**

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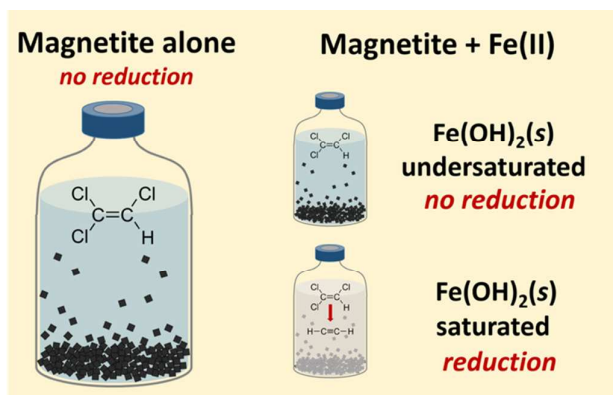
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597 **Table of Contents Artwork**

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600 Magnetite is unlikely to be as important as previously thought for abiotic reduction of PCE and TCE in
601 groundwater plumes.