

Structure and Dynamics of Water on Forsterite Surface

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23 Abstract: The behavior of water on mineral surfaces is the key to understanding interfacial and chemical reaction processes. Olivine is one of the major rock-forming minerals and its 24 interaction with water is a ubiquitous phenomenon both on Earth's surface and in the subsurface. 25 26 This work presents a combined study using molecular dynamics (MD) simulations and quasielastic neutron scattering (QENS) experiments conducted at three different instruments to study 27 the structure and dynamics of water on forsterite (Mg- end member of olivine) surface at 270 K. 28 29 The combination of three different QENS instruments probes dynamical processes occurring across broad time scales (~ 1 ps to ~ 1 ns in this study). The water structure on the hydroxylated 30 surface is composed of three distinct water layers, transitioning from well-ordered and nearly 31 immobile closest to the surface to a less structured layer. The energies of three motions 32 (including translation and rotation) derived from simulations agree well with the experiments, 33 covering the energy range from a few to hundreds of micro electron volts. 34

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37 Introduction

The interaction and reaction of water with minerals control most of the chemical and 38 transport processes encountered at the Earth's surface and subsurface ranging deep into the 39 interior. The presence of water in other planetary bodies, comets and interstellar dust amplify the 40 broad importance of water-mineral systems. Despite the extensive length and time scales over 41 which these interactions can occur, there is the inescapable fact that interfacial phenomena - i.e., 42 mineral surfaces as well as buried interfaces such as grain boundaries at the atomic scale -43 control the exchange of matter and energy that dictate the evolution of these systems.^{1–3} One of 44 45 the most important of these is the olivine-water system. Olivine-series minerals compose a major portion of Earth's upper mantle and oceanic crust, and are observed in lunar and Martian rocks. 46 They are very susceptible to near-surface reactions involving water that are generally complex 47 and lead to hydration, silicification, oxidation or carbonation.^{4,5} Furthermore, olivine-water 48 interactions are thought to be a major source of water during planetary accretion.^{6,7} 49

The dominant olivine-series phase is forsterite which is the Mg- end member of a 50 complete solid solution where magnesium and iron as cations substitute for each other 51 ([Mg,Fe]₂SiO₄). Calorimetric measurements of the adsorption energy have been made on 52 forsterite powder interacting with water.⁸ A more recent experimental study observed the effects 53 of water on the carbonation of forsterite with application to CO₂ sequestration.⁹ To date, probing 54 the structural and dynamical behavior of interfacial water on forsterite has been relegated to 55 studies employing different molecular simulation methods. For example, Kerisit et al.¹⁰ used 56 57 classical molecular dynamics (MD) simulations to investigate water on forsterite surfaces, 58 together with CO₂ under the condition of 50°C and 180 bar to simulate subsurface CO₂ sequestration. Their focus was on the form and transport of the supercritical CO₂ on the interface, 59

which depends on the water content. In a related study, Kerisit et al. used density functional theory (DFT) with dispersion corrections to examine adsorption of water on the (010) surfaces of five end members of olivine minerals ($[X]_2SiO_4$ where X = Mg, Ca, Mn, Fe, and Co).¹¹ Collectively, these and related molecular simulation studies^{12–14} suggest that water on such surfaces can exhibit several dynamical processes occurring over a broad range of characteristic times.

In order to quantify the various types of molecular water motion on the surface of 66 forsterite, we have conducted a series of quasi-elastic neutron scattering (OENS) experiments 67 68 complemented by classical MD simulations. The exceptionally large incoherent neutron scattering cross-section of hydrogen compared to that of other elements makes it an ideal 69 technique to probe the dynamics of water, in particular, the self-diffusive motions.¹⁵ Recent 70 studies using MD simulations in concert with OENS data have demonstrated the power of 71 72 integrating outcomes from the two approaches in quantifying the interfacial behavior of water on select metal oxide and mineral surfaces.^{12–14,16} 73

74 Experiments and Simulations

The QENS experiments were carried out at the high flux backscattering spectrometer (HFBS)¹⁷ and the disc chopper spectrometer (DCS)¹⁸ at the NIST Center for Neutron Research (NCNR), and at the backscattering silicon spectrometer (BASIS)¹⁹, Spallation Neutron Source (SNS) at the Oak Ridge National Laboratory (ORNL). The dynamic energy ranges are $\pm 11 \mu eV$, $-1 \sim \pm 0.2 meV$, and $\pm 100 \mu eV$ for HFBS, DCS, and BASIS, respectively, and their corresponding energy resolutions are 0.775 μeV , 10.9 μeV , and 3.3 μeV at their full width at half maximum (FWHM). Data affording the best opportunity for comparison was obtained at 270 K.

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82	The sample loading and experimental setup are the same as described in Mamontov et
83	al. ¹⁶ Synthetic nano-powdered forsterite ²⁰ with a high surface area (66.0 m^2g^{-1} , more
84	characterization images can be found in Fig. S1 in the Supporting Information) is placed in a
85	vacuum oven and then equilibrated with water vapor at 150°C for hydration for 24h and at 50°C
86	for outgassing the physisorbed water molecules afterwards. Fully hydrated forsterite sample
87	(water content of ~ 10% wt. determined by thermogravimetric analysis, see Fig. S2 in the
88	Supporting Information) was sealed in a flat aluminum cell and kept in the path of the neutron
89	beam. To avoid multiple scattering, the thickness of the sample cell was chosen such that it
90	allowed a neutron transmission of more than 90%. The three instruments have overlapping
91	energy windows and resolution allowing for an investigation of relaxation times from ~ 1 ps to \sim
92	1 ns (Fig. S3 in the Supporting Information).

93 To complement the QENS measurements, we applied classical MD simulations using empirical force fields to investigate the structural and dynamical properties of water on a 94 hydroxylated forsterite surface, which represents the real mineral surface at near-ambient 95 conditions. MD simulations were performed using DL POLY 4.06 package.²¹ Among the 96 common crystal planes of olivine mineral²² in concert with material characteriztion for the 97 synthetic nano-powdered forsterite (Fig. S1 in the Supporting Information), we selected the (010) 98 surface because it is relatively stable with a lower surface energy, as evidenced by previous 99 simulation studies.^{7,23} There are two possible terminations of the forsterite(010) surface referred 100 to as dipolar and non-dipolar²³, as shown in Fig. S4 of the Supporting Information. We selected 101 102 the non-dipolar termination for our study, which is a reasonable choice as it has a lower surface energy than the dipolar termination. Furthermore, de Leeuw et al. have pointed out the non-103 104 dipolar hydroxylated (010) surface adsorbs molecular water more favorably than the dipolar

hydroxylated surface.²³ Thus this study employed the non-dipolar hydroxylated surface. The 105 simulation box is composed of a forsterite slab (*Pbnm* space group, a = 4.753 Å, b = 10.199 Å, c 106 = 5.981 Å, experimental unit cell determined by Lager et al.²⁴) with a unit cell replicated 6×1.5 107 \times 6 times (1512 atoms) and an additional free space (~ 85 Å) along the *b*-axis to allow 392 water 108 molecules to move. Sec. SI in the Supporting Information indicates how this number of 109 molecules is equivalent to the hydration layers achieved by experiments. This large spacing 110 ensures that the surface effect dominates over the slit-pore confinement. Two sides of the surface 111 are cleaved to be non-dipolar and Si-O tetrahedron is kept intact. More simulation details are 112 presented in the Sec. SI of the Supporting Information. 113

114 **Results and Discussion**

Figs. 1a and 1b show the surface water structure and its corresponding density profile at 115 270 K obtained by MD simulations. Three layers of water molecules on the surface are visible in 116 the density profile. The first layer (closest to the surface designated as L1) is formed by the 117 interaction between water molecules and the surface OH and H groups. The higher density seen 118 in Fig. 1b for L1 is due to the ordered structure, which is mainly controlled by the morphology of 119 the hydroxylated surface. All the three layers render higher density than bulk water as shown in 120 Fig. 1b. The second layer of water molecules (L2) interacts with L1 and the surface OH groups. 121 The molecules in L2 are less ordered than L1, thus the density is lower. L1 and L2 span a total 122 distance of ~ 2.5 Å. Beyond L2, the third layer of molecules (denoted as L3) interacts with L2 123 and each other in a disordered network. The maximum duration for molecules residing in L2 or 124 L3 is described in Sec. SII of the Supporting Information. Fig. 1c shows trajectories of selected 125 water molecules only residing in a particular layer for their maximum duration. Fig. 1d plots the 126 trajectories perpendicular to the surface of two representative molecules that exhibit their 127

maximum duration in each layer over the entire production time period (1.5 ns). Molecules
residing in L1 show very limited mobility based on their trajectories, while molecules in L2 and
L3 are free to move and communicate between the two layers. In L2, molecules are less mobile
than in L3. It is important to reiterate that L1, L2, and L3 represent separate interfacial layers
occupied by a collection of molecules exhibiting similar structural and dynamical properties.







highlighted areas with boxes indicate L1 (solid red line) and L2 (dashed red line). (b)

136 Corresponding density profile for adsorbed water molecules. (c) Trajectories of four

representative molecules residing in L1, L2, L3 along *a-b* plane. (d) Trajectories perpendicular to

the surface over time using two molecules from L1, L2 and L3. The forsterite slab is shown as a

139 grey rectangle.

Mean squared displacement (MSD) is used to describe the translational diffusive motion
of molecules. The diffusion coefficient (*D*) determined from MSD is based on the Einstein
relation:

$$D = \frac{MSD}{6t} \tag{1}$$

The *D* value determined for water on the forsterite surface is $0.42 \times 10^{-9} \text{ m}^2\text{s}^{-1}$. Using MD 144 simulations, Lee and Rossky placed water with a density of 0.9963 g·cm⁻³ at 25°C on three type 145 of surfaces, i.e., a flat hydrophobic surface, a rough hydrophobic surface, and a hydrophilic silica 146 surface, and obtained the water diffusion coefficients of ~ $3.5 \times 10^{-9} \text{ m}^2 \text{s}^{-1}$, ~ $2.7 \times 10^{-9} \text{ m}^2 \text{s}^{-1}$, ~ 147 $0.96 \times 10^{-9} \text{ m}^2 \text{s}^{-1}$, respectively.²⁵ The value of D in our study is comparable to the hydrophilic 148 149 surface mentioned above, as a hydroxylated surface is expected to strongly interact with water molecules. Details about MSD of water molecules residing in different layers can be found in Fig. 150 S5 in the Supporting Information. Using the trajectory of molecules residing only in L3, the 151 diffusion coefficient is calculated to be $0.58 \times 10^{-9} \text{ m}^2 \text{s}^{-1}$, which is smaller than that of bulk water 152 calculated in this study ($2.39 \times 10^{-9} \text{ m}^2\text{s}^{-1}$ at 298 K). Related to the higher density than bulk 153 water shown in Fig. 1b, a lower D value can be expected. 154

The QENS data collected from the three instruments were reduced using the data analysis package DAVE.²⁶ Representative reduced spectra converted to the dynamical structure factor $S(Q,\omega)$ are shown in Fig. 2. $S(Q,\omega)$ is a property of the sample and is a function of momentum ($\hbar Q$) and energy transferred ($\hbar \omega$) between the neutron and the sample in a scattering event and thus records information on the structure and dynamics of the sample. It is noted that the $S(Q,\omega)$ in these experiments represents the incoherent dynamical structure factor of adsorbed water because the incoherent neutron scattering cross-section of hydrogen dominates the scattering.

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An attempt to fit the quasi-elastic broadening (qe) with a single Lorentzian function did not result in satisfactory fits. Therefore, a combination of two Lorentzians was used to model the ge broadening. The whole spectra were thus fit using the equation shown below:

165
$$S(Q,\omega) = \{A(Q)\delta(\omega) + [1 - A(Q)][pL_1(\Gamma_1, \omega) + (1 - p)L_2(\Gamma_2, \omega)] + [B_1 + B_2(Q, \omega)]\} \otimes R(Q, \omega)$$
(2)

166 The elastic signal $A(\mathbf{0})$, two Lorentzian components $L_1(\Gamma_1, \omega)$ and $L_2(\Gamma_2, \omega)$ (p as the prefactor), and a linear background $B_1 + B_2(\mathbf{Q}, \omega)$ are convoluted by the resolution function $R(\mathbf{Q}, \omega)$. The 167 elastic signal derived from motions slower than the resolution of the instrument and weak 168 scattering from the forsterite are modeled using a convoluted delta function ($\delta(\omega)$) centered on 169 zero energy transfer. The quasi-elastic components, modeled by the Lorentzian functions, are 170 separated into a fast and slow component on the time scale of the instrument (HFBS only allows 171 one Lorentzian fitting due to its narrow energy range, see Figs. 2a, 2b). By fitting the spectra 172 with eqn (2), we determined the parameters A, p, and the half width at half maximum (HWHM) 173 of the two Lorentzians, Γ_i^e (e denotes experiment, i = 1, 2). Good quality fits were obtained in 174 this manner as seen in Fig. 2 where the individual components along with the overall fit function 175 176 described in eqn (2) are shown together with the experimental spectra. The two Lorentzians representing the qe broadening have different HWHM that correspond to different time scales of 177 motion, denoted as "fast" and "slow" hereafter. Here the designation of "fast" and "slow" is 178 limited to the energies accessible at one instrument. Both BASIS and DCS have their own "fast" 179 and "slow" components and due to an overlap of the energies accessible at these instruments, the 180 "fast" component of BASIS might be comparable to the "slow" component of DCS. Fitted 181 parameters A and p among the instruments are described in Fig. S11 in the Supporting 182 183 Information. In Fig. S11, the variation of A parameter with Q suggests the presence of a localized

- 184 motion; and *p* values indicate that in the BASIS spectra the fast component dominates, whereas
- in the DCS the slower one dominates, both of which fall in the intermediate dynamic range.
- 186 Detailed analysis can be found in the Supporting Information.



Fig. 2. Representative QENS spectra obtained at 270 K at HFBS (a, b), BASIS (c, d) and DCS (e,
f). Experimental data are represented by hollow circles. Total fit function is shown in red (-);

elastic component in blue (-); two Lorentzian components in magenta (-) and green (-), and
background in light grey (-). Spectra are shown in limited quasi-elastic energy range for clarity.

Normally, a comparison is made between the diffusion coefficients obtained from QENS 193 experiments and those obtained using MSD from MD simulations. However, MSD is a function 194 of time alone and does not discriminate between motions at different length scales, thus it might 195 incorporate motions occurring at length and time scales that are beyond the energy window of 196 the instruments.²⁷ A more direct comparison between the simulations and QENS experiments 197 can be made by calculating the intermediate scattering function (ISF, $I(\mathbf{0},t)$) that is the inverse 198 Fourier transform of the dynamic structure factor $S(\mathbf{Q}, \omega)$ measured in the QENS experiment. 199 Several such direct comparisons have been attempted.^{28–30} In an MD simulation, the ISF can be 200 201 calculated as

202
$$I(Q,t) = \frac{1}{N} \sum_{i} \langle e^{iQ \cdot (r_i(t) - r_i(0))} \rangle$$
(3)

where Q is the momentum transfer, and r_i is the position vector of the *i*th hydrogen atom in the 203 simulation. The summation is over hydrogen atoms of all water molecules. The angular brackets 204 denote an ensemble average and integrate the contribution from *Q* in all directions with the same 205 magnitude. Fig. 3 shows the calculated ISF function from one example Q value (= 0.525 Å⁻¹) at 206 207 270 K using trajectories of 1.5 ns. Another simulation with 10 ns production run after 20 ns of equilibration was carried out to confirm if 1.5 ns production run is sufficient to produce 208 converged quantities. ISF and MSD calculated from this longer simulation of 10 ns yielded very 209 close results to those calculated from 1.5 ns simulation. (see in Fig. S6 in the Supporting 210 Information). Because experimental $S(Q, \omega)$ profiles were modeled with combinations of 211 Lorentzian functions, the corresponding ISF calculated from the simulation can be modeled by a 212

combination of exponential functions.²⁷ In order to model the total ISF over the entire time range,
a function composed of three exponential functions and a constant is employed:

215
$$I(Q,t) = A_1 e^{-t/\tau_1} + A_2 e^{-t/\tau_2} + A_3 e^{-t/\tau_3} + c$$
(4)

where τ_1 , τ_2 , and τ_3 are the decay times that define three distinct decay stages. The fitting curve in Fig. 3 is the result of fitting calculated ISF data using eqn (4). The three exponential decay terms in eqn (4), when inverse Fourier transformed, would result in three Lorentzian functions $L_i^s(\Gamma_i^s,\omega)$ (*i*=1, 2, 3) with HWHM (Γ_i^s) related to the corresponding decay times according to the relation

220
$$\Gamma_i^s = \frac{\hbar}{\tau_i}$$
(5)

where s denotes simulation. The Lorentzians $L_i^s(\Gamma_i^s,\omega)$ represent the same quantities as the

experiment-derived Lorentzians $L_i^e(\Gamma_i, \omega)$ (*i*=1, 2) in eqn (2) and are amenable to direct

comparison. The four components of the fit (three Lorentzians and one constant) are also plotted

in Fig. 3. In addition, the trajectories of ten representative molecules from L1, L2 and L3 with a

certain time period facilitates the respective layer-specific ISF which are also shown in Fig. 3.



Fig. 3. Calculated total ISF using one *Q* value at 270 K and fitting curve obtained using eqn (4).
Four components in eqn (4) are plotted in dashed lines. The ISF curves representing L1, L2 and L3 are represented by open circles.

In Fig. 3, the three exponential decay components of the fit (dashed lines) show different 230 decaying rates; however, the constant component represents extremely slow motions that may 231 remain unresolved in an experiment due to the finite resolution of the instrument used. This slow 232 motion matches the characteristics of the molecules in L1 and L2 (black and red circles), which 233 show restricted mobility. Only molecules in L3 exhibit motions that can be described on the time 234 scales of the three exponential components used to fit the overall ISF. The ISF calculated for L3 235 236 (green circles) can be fitted by a linear combination of these three exponentials (Fig. S7 in the Supporting Information). From the weight percentage of the three exponentials in the fits to the 237 238 L3 ISF, we conclude that the intermediate time scale with a τ_2 value of 73 ps dominates the dynamics in L3. Mamontov et al. examined the dynamics of molecular water on the $TiO_2(110)$ 239

surface using QENS and MD, and found that while the chemisorbed water in L1 is too slow to be
observed in their experiment, water molecules in both the (L2+L3) above this L1 contributed to
the experimental spectra.¹²

In addition, average residence time (τ_R), a measure of how long a typical molecule remains in a given region of space,^{31,32} was also estimated for all the three layers (Fig. S10 in the Supporting Information). Estimates for the average residence time are 956 ps, 49.5 ps, and 335 ps for L1, L2, and L3, respectively. The motions of molecules in L1 are severely restricted, thus the τ_R is longer for L1 than that for the other two layers. Due to the thicknesses of L2 and L3 shown in Fig. 1b, we can expect the molecules in L2 have shorter τ_R and easily move beyond the narrow region and molecules in L3 experience a longer time to escape the thicker layer.

Fig. 4 compares the Γ_i values obtained by fitting the experimental QENS spectra using 250 eqn (2) and those obtained from eqn (4) by converting the decay constants of exponentials 251 derived from ISF fitting from the simulations. In Fig. 4, we observe the three components (Γ_1^s , 252 Γ_2^s, Γ_3^s) determined by simulation span a wide energy range, whose HWHM ranges from a few 253 to hundreds of µeV. Each component matches the combination of motion from different 254 instruments, which is due to the specific energy range accessible at the instrument. In the three 255 energy domains in Fig. 4, the fastest component (Γ_1^s) from simulation matches the fast 256 component observed on DCS; the intermediate component (Γ_2^s) matches the fast component 257 observed on BASIS and the slow component from DCS; and the slowest component (Γ_3^s) 258 259 matches the slow component from BASIS and the single component from HFBS. We observe that Γ_1^s becomes much faster in higher **Q** region. However, in the **Q** range that is consistent with 260





Fig. 4. Comparison of the QENS data with MD simulation data. Three representative energy domains are outlined in grey areas. All the error bars are shown in orange.

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277	Rotational behavior observed from the second and third stages of decay of the rotational
278	ISF (see Fig. S9 in the Supporting Information) suggest the corresponding time scales are within
279	the ranges accessible at DCS and BASIS. Therefore, the quasi-elastic broadening observed in the
280	DCS and BASIS spectra can be expected to have a contribution from the rotational motion of
281	water. This is consistent with A values whose trends with Q indicates localized motion in the
282	experimental spectra (see Fig. S11a in the Supporting Information).

283 **Conclusions**

In conclusion, this study examines structure and dynamics of water on hydroxylated 284 forsterite(010) surface at 270 K using MD simulation and QENS experiments. The density 285 profile indicates water molecules on the surface compose three distinct layers: L1+L2 include 286 nearly immobile molecules with relatively less structural disorder in L2; and L3 is composed of 287 disordered water molecules. Motion of water on the forsterite surface consists of several 288 dynamical processes that occur at different time scales. The ISF calculated from the simulations 289 can be described by four components that represent motions at different time scales. Out of these 290 four, the energies associated with three components match the energies of three components of 291 motions observed in the QENS experiments. Simulation results indicate water molecules in L3 292 293 are primarily responsible for the observable dynamics, while the restricted molecules in L1+L2 cannot be detected by the experiments. We have provided an excellent example of how MD 294 simulation in concert with neutron scattering experimental data provides precise and 295 complementary information for understanding the structure and dynamics of water on a mineral 296 surface. 297

298 **Conflicts of interest**

299 Authors do not declare any conflict of interest.

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315 **References**

316 317	1	A. Lüttge and R. S. Arvidson, in <i>Kinetics of Water-Rock Interaction</i>, eds. S. L. Brantley, J. D. Kubicki and A. F. White, Springer New York, New York, NY, 2008, pp. 73–107.
318 319	2	A. Striolo, From Interfacial Water to Macroscopic Observables: A Review, <i>Adsorpt. Sci. Technol.</i> , 2011, 29 , 211–258.
320 321	3	C. V. Putnis and E. Ruiz-Agudo, The Mineral-Water Interface: Where Minerals React with the Environment, <i>Elements</i> , 2013, 9 , 177–182.
322 323	4	M. A. Velbel, Dissolution of olivine during natural weathering, <i>Geochim. Cosmochim. Acta</i> , 2009, 73 , 6098–6113.
324 325	5	M. J. Wilson, Weathering of the primary rock-forming minerals: processes, products and rates, <i>Clay Miner.</i> , 2004, 39 , 233–266.
326 327 328 329	6	H. E. King, M. Stimpfl, P. Deymier, M. J. Drake, C. R. A. Catlow, A. Putnis and N. H. de Leeuw, Computer simulations of water interactions with low-coordinated forsterite surface sites: Implications for the origin of water in the inner solar system, <i>Earth Planet. Sci. Lett.</i> , 2010, 300 , 11–18.
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330 7 M. Stimpfl, A. M. Walker, M. J. Drake, N. H. de Leeuw and P. Deymier, An ångström-

331 332		sized window on the origin of water in the inner solar system: Atomistic simulation of adsorption of water on olivine, <i>J. Cryst. Growth</i> , 2006, 294 , 83–95.
333 334	8	S. Chen and A. Navrotsky, Calorimetric study of the surface energy of forsterite, <i>Am. Mineral.</i> , 2009, 95 , 112–117.
335 336 337	9	C. J. Thompson, P. F. Martin, J. Chen, P. Benezeth, H. T. Schaef, K. M. Rosso, A. R. Felmy and J. S. Loring, Automated high-pressure titration system with in situ infrared spectroscopic detection, <i>Rev. Sci. Instrum.</i> , 2014, 85 , 044102.
338 339	10	S. Kerisit, J. H. Weare and A. R. Felmy, Structure and dynamics of forsterite–scCO2/H2O interfaces as a function of water content, <i>Geochim. Cosmochim. Acta</i> , 2012, 84 , 137–151.
340 341	11	S. Kerisit, E. J. Bylaska and A. R. Felmy, Water and carbon dioxide adsorption at olivine surfaces, <i>Chem. Geol.</i> , 2013, 359 , 81–89.
342 343 344 345	12	E. Mamontov, D. J. Wesolowski, L. Vlcek, P. T. Cummings, J. Rosenqvist, W. Wang and D. R. Cole, Dynamics of Hydration Water on Rutile Studied by Backscattering Neutron Spectroscopy and Molecular Dynamics Simulation, <i>J. Phys. Chem. C</i> , 2008, 112 , 12334–12341.
346 347 348	13	A. G. Stack, J. M. Borreguero, T. R. Prisk, E. Mamontov, HW. Wang, L. Vlcek and D. J. Wesolowski, Precise determination of water exchange on a mineral surface, <i>Phys. Chem. Chem. Phys.</i> , 2016, 18 , 28819–28828.
349 350 351 352	14	L. J. Michot, A. Delville, B. Humbert, M. Plazanet and P. Levitz, Diffusion of Water in a Synthetic Clay with Tetrahedral Charges by Combined Neutron Time-of-Flight Measurements and Molecular Dynamics Simulations, <i>J. Phys. Chem. C</i> , 2007, 111 , 9818–9831.
353 354 355	15	M. Bee, M. Foulon, J. P. Amoureux, C. Caucheteus and C. Poinsignon, The glassy phase of 1-cyanoadamantane; a study of the molecular reorientations by high-resolution quasi- elastic neutron scattering, <i>J. Phys. C Solid State Phys.</i> , 1987, 20 , 337–349.
356 357 358 359	16	E. Mamontov, L. Vlcek, D. J. Wesolowski, P. T. Cummings, W. Wang, L. M. Anovitz, J. Rosenqvist, C. M. Brown and V. G. Sakai, Dynamics and Structure of Hydration Water on Rutile and Cassiterite Nanopowders Studied by Quasielastic Neutron Scattering and Molecular Dynamics Simulations, <i>J. Phys. Chem. C</i> , 2007, 111 , 4328–4341.
360 361 362	17	A. Meyer, R. M. Dimeo, P. M. Gehring and D. A. Neumann, The high-flux backscattering spectrometer at the NIST Center for Neutron Research, <i>Rev. Sci. Instrum.</i> , 2003, 74 , 2759–2777.
363 364	18	J. R. D. Copley and J. C. Cook, The Disk Chopper Spectrometer at NIST: a new instrument for quasielastic neutron scattering studies, <i>Chem. Phys.</i> , 2003, 292 , 477–485.
365 366	19	E. Mamontov and K. W. Herwig, A time-of-flight backscattering spectrometer at the Spallation Neutron Source , BASIS, <i>Rev. Sci. Instrum.</i> , 2011, 82 , 085109.
367	20	L. M. Anovitz, A. J. Rondinone, L. Sochalski-kolbus, J. Rosenqvist and M. C. Cheshire,

368 369		Nano-scale synthesis of the complex silicate minerals forsterite and enstatite, <i>J. Colloid Interface Sci.</i> , 2017, 495 , 94–101.
370 371 372	21	I. T. Todorov, W. Smith, K. Trachenko and M. T. Dove, DL_POLY_3: new dimensions in molecular dynamics simulations via massive parallelism, <i>J. Mater. Chem.</i> , 2006, 16 , 1911–1918.
373 374	22	W. A. Deer, R. A. Howie and J. Zussman, Eds., <i>Rock-forming minerals: Vol. 1A. Orthosilicates</i> , Geological Society Publishing House, London, 2nd edn., 1997.
375 376 377	23	N. H. de Leeuw, S. C. Parker, C. R. a. Catlow and G. D. Price, Modelling the effect of water on the surface structure and stability of forsterite, <i>Phys. Chem. Miner.</i> , 2000, 27 , 332–341.
378 379 380	24	G. A. Lager, F. K. Ross, F. J. Rotella and J. D. Jorgensen, Neutron powder diffraction of forsterite, Mg2SiO4: a comparison with single-crystal investigation, <i>J. Appl. Cryst.</i> , 1981, 14 , 137–139.
381 382 383	25	S. H. Lee and P. J. Rossky, A comparison of the structure and dynamics of liquid water at hydrophobic and hydrophilic surfaces—a molecular dynamics simulation study, <i>J. Chem. Phys.</i> , 1994, 100 , 3334–3345.
384 385 386 387	26	R. T. Azuah, L. R. Kneller, Y. Qiu, P. L. W. Tregenna-Piggott, C. M. Brown, J. R. D. Copley and R. M. Dimeo, DAVE : A Comprehensive Software Suite for the Rduction, Visualization, and Analysis of Low Energy Neutron Spectroscopic Data, <i>J. Res. Natl. Inst. Stand. Technol.</i> , 2009, 114 , 341–358.
388 389 390	27	S. Gautam, T. Le, A. Striolo and D. Cole, Molecular dynamics simulations of propane in slit shaped silica nano-pores: direct comparison with quasielastic neutron scattering experiments, <i>Phys. Chem. Chem. Phys.</i> , 2017, 19 , 32320–32332.
391 392 393	28	N. Morelon, G. R. Kneller, M. Ferrand, A. Grand, J. C. Smith and M. Bée, Dynamics of alkane chains included in an organic matrix: Molecular dynamics simulation and comparison with neutron scattering experiment, <i>J. Chem. Phys.</i> , 1998, 109 , 2883–2894.
394 395	29	S. Gautam, S. Mitra, R. Mukhopadhyay and S. L. Chaplot, Diffusion of acetylene inside Na-Y zeolite: Molecular dynamics simulation studies, <i>Phys. Rev. E</i> , 2006, 74 , 041202.
396 397 398	30	S. Gautam, S. Mitra, S. L. Chaplot and R. Mukhopadhyay, Dynamics of 1,3-butadiene adsorbed in Na-Y zeolite: A molecular dynamics simulation study, <i>Phys. Rev. E</i> , 2008, 77 , 061201.
399 400	31	S. Kerisit, E. S. Ilton and S. C. Parker, Molecular Dynamics Simulations of Electrolyte Solutions at the (100) Goethite Surface, <i>J. Phys. Chem. B</i> , 2006, 110 , 20491–20501.
401 402	32	U. Terranova and N. H. de Leeuw, Structure and dynamics of water at the mackinawite (001) surface, <i>J. Chem. Phys.</i> , 2016, 144 , 094706.
403		