Accepted Manuscript



This is an *Accepted Manuscript*, which has been through the Royal Society of Chemistry peer review process and has been accepted for publication.

Accepted Manuscripts are published online shortly after acceptance, before technical editing, formatting and proof reading. Using this free service, authors can make their results available to the community, in citable form, before we publish the edited article. We will replace this Accepted Manuscript with the edited and formatted Advance Article as soon as it is available.

You can find more information about *Accepted Manuscripts* in the **Information for Authors**.

Please note that technical editing may introduce minor changes to the text and/or graphics, which may alter content. The journal's standard <u>Terms & Conditions</u> and the <u>Ethical guidelines</u> still apply. In no event shall the Royal Society of Chemistry be held responsible for any errors or omissions in this *Accepted Manuscript* or any consequences arising from the use of any information it contains.



www.rsc.org/chemicalscience

1	Precise, Contactless Measurements of
2	the Surface Tension of Picolitre Aerosol Droplets
3	
4	Bryan R. Bzdek, ^{1, a} Rory M. Power, ^{1, 2, a} Stephen H. Simpson, ^{1, 3}
5	Jonathan P. Reid, ^{1, *} and C. Patrick Royall ^{1, 4, 5}
6	
7	¹ School of Chemistry, University of Bristol, Bristol, BS8 1TS, UK
8	² Max Planck Institute of Molecular Cell Biology and Genetics, Dresden, 01307, Germany
9	³ Institute of Scientific Instruments of the ASCR. v.v.i., Krávolopolská 147, 612 64, Brno, Czech Republic
10	⁴ H. H. Wills Physics Laboratory, University of Bristol, Bristol, BS8 1TL, UK
11	⁵ Centre for Nanoscience and Quantum Information, University of Bristol, BS8 1FD, UK
12	
13	
14	
15	
16	
17	
18	
19	
20	
21	
22	
23	
24	Submitted to:
25	Chemical Science
26	
27	^a These authors contributed equally to this work
28	*To whom correspondence should be addressed. Email: j.p.reid@bristol.ac.uk

29 ABSTRACT

30 The surface composition and surface tension of aqueous droplets can influence key aerosol 31 characteristics and processes including the critical supersaturation required for activation to form cloud 32 droplets in the atmosphere. Despite its fundamental importance, surface tension measurements on droplets 33 represent a considerable challenge owing to their small volumes. In this work, we utilize holographic optical 34 tweezers to study the damped surface oscillations of a suspended droplet (<10 µm radius) following the 35 controlled coalescence of a pair of droplets and report the first contactless measurements of the surface 36 tension and viscosity of droplets containing only 1 - 4 pL of material. An advantage of performing the 37 measurement in aerosol is that supersaturated solute states (common in atmospheric aerosol) may be 38 accessed. For pairs of droplets starting at their equilibrium surface composition, surface tensions and 39 viscosities are consistent with bulk equilibrium values, indicating that droplet surfaces respond to changes in 40 surface area on microsecond timescales and suggesting that equilibrium values can be assumed for growing 41 atmospheric droplets. Furthermore, droplet surfaces are shown to be rapidly modified by trace species 42 thereby altering their surface tension. This equilibration of droplet surface tension to the local environmental 43 conditions is illustrated for unknown contaminants in laboratory air and also for droplets exposed to gas 44 passing through a water-ethanol solution. This approach enables precise measurements of surface tension 45 and viscosity over long time periods, properties that currently are poorly constrained.

47 The surface tension and viscosity of droplets are crucial properties that must be understood and quantified in a range of disciplines, ranging from the evaporation and drying of sprays,¹ the deposition of 48 droplets on surfaces in inkjet printing,² digital microfluidics/micromachinery,³ and cloud droplet activation 49 in the atmosphere.⁴ Although readily probed in bulk phases, direct measurements of these properties for 50 51 droplets are challenging. Reconciling bulk phase data with the high surface-to-volume ratios characteristic of 52 dispersed droplet systems is difficult, particularly in highly dynamic systems where the transfer of volatile 53 components across gas-liquid surfaces must be considered. Additionally, states exhibited by droplets may differ markedly from the (equilibrium) phase behaviour of bulk solutions, accessing metastable 54 supersaturated states with solute concentrations above the solubility limit,⁵ amorphous or glassy states,^{6,7} and 55 liquid-liquid phase separated structures.⁸ Measurements on timescales relevant to dispersed droplet systems 56 may be similarly inaccessible through bulk measurements, as values of the surface tension and viscosity are 57 58 often required for surface ages (i.e. the time since the surface was first generated) as short as a millisecond. 59 Finally, conventional techniques require substantial amounts of sample, often precluding measurements on 60 systems where the sample is limited (for example, in the collection of atmospheric aerosol samples).

In particular, the surface tension of aerosols and growing cloud droplets represents a critical 61 62 uncertainty in determining the effect of aerosols on climate.⁹ Atmospheric aerosols are chemically diverse containing a myriad of organic species, and surface tension depression has been observed in aqueous aerosol 63 extracts.^{10, 11} In the absence of direct measurements, however, there has been considerable debate for many 64 65 years as to the composition of particle surfaces and the magnitude of surface tensions for atmospheric aerosols and cloud droplets.¹² The equilibrium size of an aerosol particle at a particular relative humidity 66 67 (RH) is governed by an interplay between a bulk solute effect and surface curvature and is treated by Köhler theory.¹³ The former effect allows for stable particle sizes under sub-saturated RH conditions while the latter 68 69 governs the magnitude of the characteristic thermodynamic barrier that must be overcome for an aerosol 70 particle to "activate", spontaneously growing by water condensation to form a cloud droplet. This barrier can only be overcome under supersaturated conditions with respect to gaseous water (RH > 100 %), and 71 72 estimating this critical supersaturation is central to predicting the fraction of aerosol particles that are 73 activated to form cloud droplets, impacting on cloud droplet number and size, cloud albedo and persistence, 74 and therefore radiative forcing. The fraction of the aerosol population that can act as cloud condensation

75 nuclei (CCN) is also important for "cloud invigoration" effects, where a larger number of CCN prolong cloud development and lifetime.¹⁴ Some studies suggest that a surface tension equivalent to that of pure 76 water may be appropriate for representing an activating CCN and for determining the critical super-77 saturation from Köhler theory.¹⁵ However, this assumption directly contravenes expectations for droplets 78 79 containing organic solutes, the interplay between bulk and surface partitioning of organic components, and the dependence on droplet size.^{16, 17} More recently, studies have suggested that organics species, which can 80 contribute to more than half of the mass growth of nanoparticles,¹⁸ may reduce nanoparticle surface tension, 81 82 thereby reducing the energy barrier to further growth by vapour condensation.¹⁹ Additionally, surface tension may play a crucial role in determining the thermodynamic state of nucleation mode aerosol, highlighting the 83 need for reliable surface tension data for solute solutions at supersaturated concentrations.²⁰ 84

85 Organic surface coatings on aerosols have been observed by a number of indirect methods. Surface 86 coatings on aerosol have been shown to play a significant role in controlling the mass transport kinetics of volatile components, such as in the transport of water across the droplet surface^{20, 21} or in heterogeneous 87 88 reaction rates with gas-phase species (e.g. N_2O_5 and O_3) that must similarly pass into the bulk.²² Indeed, even 89 substantial discrepancies between measurements of water condensation kinetics on droplets with fresh or 90 aged surfaces have been attributed to contamination by organic species, which may form highly ordered 91 surface films.²¹ Hygroscopicity studies have inferred surface tension through measurements of critical particle radii for activation and have observed surface tension depression on the order of 10% relative to 92 water due to uptake of the small surface-active molecules methylglyoxal and acetaldehyde.²³ Indirect 93 94 inference of surface tension in this manner, however, only provides averaged values over large particle 95 ensembles and requires solute and curvature effects to be separable. Direct and unambiguous measurements 96 of droplet surface tensions are crucial to refine our understanding of CCN activation, the growth of cloud 97 droplets, and the interactions that occur across gas-liquid surfaces in the atmosphere.

Two approaches have been taken to determine droplet surface tensions. One approach is the oscillating droplet method, which has been used to infer surface tension and viscosity for droplets consisting of Newtonian and non-Newtonian fluids primarily in falling droplet chains.^{2, 24, 25} The damped oscillatory motion of a droplet following generation can be compared with the expected natural oscillation modes of a droplet initially described by Rayleigh²⁶ and extended to include viscous damping by Lamb.²⁷ Using this approach, Yang et al. have recently shown that surface tensions inferred from measurements of the

104 oscillation frequencies of droplets several μ L down to ~100 pL in volume over 10s of ms can be reconciled with dynamic surface tension measurements made on bulk samples.² Over the short timescale of these 105 106 measurements, droplet composition is necessarily limited to sub-saturated solutions and the interaction with 107 the surrounding gas-phase (crucial to determine aerosol surface properties) is minimal. Another approach 108 adopted by Morris et al. uses atomic force microscopy (AFM) to probe the surfaces of aerosol particles sampled and deposited on a substrate by relating the probe tip retention force to the surface tension.²⁸ This 109 110 approach allows equilibrium surface tension measurements on the same droplet with varying RH (solute 111 concentration). The benefit of this approach is that measurements are performed on individual particles 112 deposited onto a substrate. However, this approach is necessarily more invasive as it requires contact of the 113 AFM probe tip to the droplet, which can result in crystallization on the tip and impact the quality of the 114 measurement.

115 In this work, we present contactless, non-invasive measurements of the surface tensions and 116 viscosities of aqueous solution droplets with volumes typically 1 - 4 picolitres. We show that the surface tension and viscosity of a droplet can be determined with accuracies better than $\pm 1 \text{ mN} \cdot \text{m}^{-1}$ and $\pm 1 \times 10^{-3} \text{ Pa} \cdot \text{s}$. 117 118 respectively, a consequence of the extremely high accuracy with which droplet size $(\pm 2 \text{ nm})$ and refractive index (RI, ± 0.0005) are determined.²⁹ The excitation and damping of oscillatory modes are observed for the 119 120 smallest droplets to date (radii down to 6 μ m) and approaching the fundamental hydrodynamic limit as 121 determined by Chandrasekhar.³⁰ This approach has the additional advantage that the composition is known at 122 the moment of measurement (by virtue of the precise RI determination). Further, measurements can be made 123 over a wide range in surface age extending from seconds to hours with the exposure of the droplet to varying 124 gas phase conditions (for example, RH) permitting direct investigation of the dependence of surface tension 125 and viscosity on particle composition. Performing such measurements on trapped, airborne droplets allows 126 access to supersaturated solute states that are characteristic of aerosol. We show that surface tensions 127 measured on picolitre droplets agree well with bulk measurements of sub-saturated solutions. Additionally, 128 we find that that droplet surfaces become contaminated on timescales of minutes unless the most stringent 129 precautions are taken to control the purity of the local gas phase environment of the droplet. Finally, we 130 show that the response of droplet surface tension to changes in its local gas phase environment (resulting in 131 small changes to droplet chemical composition) can be monitored by this approach. Combined, these results 132 describe a versatile new technique to study a crucial aerosol property that currently is poorly constrained.

133

Chemical Science

Chemical Science Accepted Manuscript

134 EXPERIMENTAL SECTION

135 Holographic Optical Trapping and Controllable Coalescence of Aerosol Droplets

The basic optical tweezers approach used here has been described previously^{5, 31} and a schematic of 136 137 the apparatus is shown in Fig. S1 of the Electronic Supplementary Information (ESI). Briefly, the optical 138 tweezers are configured in the standard inverted microscope geometry. To form multiple steerable optical 139 traps, the phase front of a continuous wave 532 nm laser (Laser Quantum, Opus 3W) is dynamically shaped 140 using a liquid crystal on silicon spatial light modulator (LCOS-SLM, Hamamatsu X10468). The beam is 141 expanded to fill the SLM display, which is conjugated to the back focal plane of a high numerical aperture 142 microscope objective (Olympus ACH, $100 \times / 1.25$, oil) by a pair of condensing 4f telescopes. A pre-calculated 143 sequence of kinoforms provides a simple and reproducible method by which the trapping positions can be 144 reconfigured. Upon initiation of the user, the trap separation is varied leading to eventual droplet 145 coalescence. The capture and relative position of trapped droplets is monitored by a camera (Dalsa Genie 146 HM 640, CMOS). Brightfield contrast is provided by widefield illumination with a high power LED (Thorlabs, 470 nm). Similar to previous studies,^{5, 31} backscattered Raman light is imaged onto the entrance 147 148 slit of a 0.5 m focal length spectrograph (Princeton Instruments, Action Spectra Prop SP-2500), dispersed by 149 a 1200 line pairs/mm grating onto a cooled CCD camera. The resulting Raman spectrum from a spherical 150 droplet consists of a broad underlying Stokes band with superimposed resonant structure at wavelengths 151 commensurate with whispering gallery modes (WGMs), from which the radius, RI, and dispersion can be determined with accuracies better than ± 2 nm, ± 0.0005 and $\pm 3 \times 10^{-8}$ cm respectively.²⁹ The intensity of the 152 153 elastically scattered component is measured using a silicon photodetector (Thorlabs, DET 110) and recorded 154 using a low-load, high bit-rate oscilloscope (LeCroy, HDO 6034-MS). The signal is sampled with a 155 bandwidth >100 MHz, typically > 3 orders of magnitude higher than the oscillation frequency. A high frame 156 rate camera (Vision Research, Phantom v. 7.3) is used to obtain the images shown in Fig. 1a, which have a 157 time resolution around 8 µs. The camera acquisition is synchronized to the trigger of the oscilloscope to 158 capture the coalescence on the camera buffer.

To isolate trapped droplets and allow environmental control, the objective focuses the beams through a coverslip (Chance Glass, #0 thickness) into a custom built trapping chamber. To populate the optical traps a fine mist of aerosol is produced using an ultrasonic nebulizer (Omron NE U22) from an aqueous solution

162 containing involatile (or very low volatility) solutes. All chemicals are used without further purification 163 (sodium chloride, Sigma Aldrich, 99.9999%; glutaric acid, Acros Chemicals, 99%) and dissolved in 164 ultrapure water. The RH in the trapping cell can be controlled by varying the relative flow rates of dry and 165 humidified nitrogen (BOC, 99.998%, 2 ppm THC max), using paired mass flow controllers (Bronkhorst), 166 allowing fine control over the water activity of the solution droplet and control over the solute 167 concentrations. The initial concentration of a trapped droplet has been shown to be comparable to that of the 168 nebulized solution,³² which provides a coarser method for tuning the solute concentration.

169

170 Determination of Droplet Composition and Density

171 Fitting of the wavelengths of the WGMs in the droplet Raman spectra to Mie theory provides a 172 precise determination of droplet radius and RI. RI is then related to solute concentration, droplet density, 173 expected surface tension, and expected viscosity through different parametrizations. For NaCl, these 174 parametrizations were obtained through polynomial fits of RI to concentration and then from concentration 175 to density, surface tension, and viscosity across the entire range of water activity using the Aerosol Diameter Dependent Equilibrium Model (ADDEM)^{33, 34} and the Aerosol Inorganics Model (E-AIM).³⁵ For glutaric 176 acid, thermodynamic data are not available, so these parametrizations were based on the relationship 177 178 between measured RI (at 589 nm, Misco Digital Refractometer PA203) and solution concentration and 179 measured RI, dispersion, and solution density (Mettler Toledo, Densito 30PX) for a range of subsaturated 180 solutions. Parametrizations used are available in ESI.

181

182 Bulk Tensiometry and Viscometry

183 The surface tension of bulk aqueous solutions of sodium chloride and glutaric acid are also reported 184 for the same solutions studied by optical tweezers. For sodium chloride, bulk measurements of surface 185 tension allow comparison to both the optical tweezers method and calculated E-AIM results. For glutaric 186 acid, no thermodynamic model results are available and (as discussed later) existing parametrizations of 187 concentration to surface tension span a relatively wide range for a given solution concentration. Therefore, 188 the surface tensions of subsaturated glutaric acid solutions were experimentally determined to more directly 189 compare to the results of the optical tweezers approach. The surface tensions of all solutions were measured 190 using a Wilhelmy plate tensiometer (Krüss K100).

191 The concentration dependence of the viscosity of glutaric acid solutions has not been reported in 192 detail, and we report here measurements for subsaturated glutaric acid solutions measured by capillary 193 viscometry using a Cannon-Fenske routine 200 viscometer (psl-Rheotek).

194

195 RESULTS AND DISCUSSION

196 Direct Observation of Oscillatory Modes from Coalescing Droplets

197 The coalescence of two liquid droplets, \sim 5-10 µm in radius, yields an initially dumbbell-shaped 198 particle, which undergoes relaxation in shape via damped surface oscillations to a single sphere. This shape 199 relaxation is driven by capillary forces and the minimization of surface free energy. Above some critical 200 value of viscosity, n_{crit} (typically 10-20 mPa s for the droplet sizes considered here), the surface relaxation is 201 so efficiently damped that purely aperiodic behaviour is observed. Previously we have shown that 202 measurement of characteristic relaxation times spanning from us to days provides a probe of the bulk properties of metastable solution droplets with viscosities spanning from 0.01 to 10⁹ Pa·s, a range of 11 203 orders of magnitude.^{5, 31} Here we consider in detail the relaxation dynamics of droplets with viscosities 204 below η_{crit} within the underdamped regime, demonstrating that accurate measurements of surface tension and 205 206 viscosity can be made.

207 Following coalescence the large initial distortion leads to the excitation of many oscillatory surface 208 modes, with those of highest order damped over <10 us and the longest-lived over ~100 us. A sequence of 209 high frame rate images of the oscillating particle recorded with a time-resolution of 8 µs is shown in Fig. 1a. 210 These images show a dilute sodium chloride droplet doped with the surfactant sodium dodecyl sulphate 211 following coalescence. The composite droplet is axially situated just below the objective focal plane due to reduced buoyancy when trapping in air³⁶ and laterally at the centre-of-mass of the coalescing pair. Due to the 212 comparably weak optical forces confining the particle when compared with capillary forces,³⁷ the optical 213 forces have no influence on the relaxation dynamics and the composite particle is only drawn back into an 214 optical trap over much longer times (>1 ms).^{38, 39} Even after only a few µs, the symmetry of the oscillation 215 216 appears to be consistent with only the lowest order (l = 2) mode for an incompressible droplet. After about 217 50 μ s, the amplitude of the oscillation has diminished significantly. The images capture the approximate 218 turning points in the oscillations clearly, as these are the geometries at which the sinusoidally oscillating 219 droplet spends most time. The damping time can be readily estimated from the maxima in the droplet aspect

Page 9 of 31

Chemical Science

ratio. The main uncertainty in determining the damping time from the droplet image aspect ratios arises from the camera frame rate being of similar order to the droplet oscillation frequency, so the moment of maximum shape distortion is not known with high precision relative to the oscillation period. The time-dependence of the ratio of the two droplet axes (the aspect ratio, a_y/a_x) is shown in Fig. 1b. The decay in the oscillation amplitude with time is evident. From the treatment of Lamb, the characteristic damping time for the *l*-th order mode (τ_1) is related to the dynamic viscosity (η), radius (*a*), and density (ρ) of the droplet by the expression:²⁷

227
$$\tau_{l} = \frac{a^{2}\rho}{(l-1)(2l+1)\eta}$$
(1)

For a droplet where $a = 10 \,\mu\text{m}$, $\rho = 1000 \,\text{kg} \cdot \text{m}^{-3}$, and $\eta = 1 \,\text{mPa} \cdot \text{s}$, the second order mode has a characteristic damping time of 20 μ s, broadly consistent with the timescale of damping shown in Fig. 1b.

230 The natural angular oscillation frequencies (ω_l) of the mode of order *l* can be expressed as:²⁶

$$\omega_l^2 = \frac{l(l-1)(l+2)\sigma}{a^3\rho}$$
(2)

where σ is the surface tension. For a droplet of the size and density considered above and with $\sigma = 75 \text{ mN} \cdot \text{m}^{-1}$ ¹ the expected linear oscillation frequency ($\omega/2\pi$) for the l = 2 mode is ~100 kHz, equivalent to a period of ~10 µs. Even with the highest frame rate accessible in the experiment, the imaging bandwidth imposes a Nyquist-limited resolution of ~50 kHz. Therefore, a more accurate method to measure oscillation frequency and infer surface tension utilising elastic backscattered light is described below.

237 Figure 1b also shows the time dependence of the light intensity backscattered from the coalescing 238 droplet shown in the images. A correspondence is immediately apparent between the distortions in shape 239 apparent in the images (quantified by the aspect ratio) and the oscillations in backscattered intensity. A 240 damping time can be estimated from the backscattered light by fitting an exponential to the peaks in the 241 signal. The backscattered light measurement gives a damping time of 31 ± 3 µs. The origin of the 242 backscattered light signal can be considered analogous to the variation in reflected intensity from a dielectric 243 slab of variable thickness. As the path length of the light through the droplet changes with the distortion in 244 shape, the interference of light reflected from the front and back face of the droplet leads to a modulation of 245 the backscattered light amplitude at the modal oscillation frequencies. A Fast Fourier Transform (FFT) of the 246 backscattered light signal allows the frequency components of the oscillation to be determined. The FFT of

Chemical Science Accepted Manuscript

the backscattered light signal in Fig. 1b is shown in Fig 1c. A single (linear) oscillation frequency corresponding to the $l = 2 \mod (\omega_{l=2}/2\pi)$ is evident at 59.6 ± 0.3 kHz. The linewidth of the resonance peak is a consequence of the damping rate of the oscillator. From this frequency it is possible to determine experimentally the droplet surface tension using Eq. (2). However, first we validate our interpretation of the backscattered light signal using a heuristic model to demonstrate that the form of the light scattering signal and its corresponding frequency spectrum result from the geometries accessed by the oscillating droplet and the optical fields that confine it.

254

255 Modelling the Intensity of Scattered Light from Oscillating Droplets

256 Semi-analytical T-matrix calculations have been performed to give insight into the origin of the 257 detected signal. Full details of the model are provided in ESI. It is assumed that the signal derives primarily 258 from the l = 2 mode and that the centre-of-mass of the composite droplet remains stationary over the course of the measurement,³⁸ centred between the propagation axes of the two beams, which are separated by a 259 260 distance equivalent to the radii of the initial droplets. This understanding of the position of the composite 261 droplet relative to the trapping beams, made possible by the high frame rate images, is the key advance that enables more accurate light scatter simulations relative to our previous work.⁵ Since the optical and inertial 262 263 forces are negligible, the droplet has a symmetry axis that is directed between the centres of the droplets 264 prior to coalescence. In the model, the axis of rotational symmetry coincides with the x-axis, and the beams 265 propagate along the z axis (out of the page). The separation between the beam axes is equivalent to the diameter of the initial droplets, i.e. the beams pass through the points on the x-axis at $y = +(a/2)^{1/3}$ and y = -266 $(a/2)^{1/3}$, where a is the radius of the combined droplet. 267

268 Figure 2a shows the droplet shape for a series of relative amplitudes, A/a, of the l = 2 mode. For $A_2/a = \pm 0.5$, the droplet shape resembles a dumbbell. For $A_2/a = 0$, the droplet shape is spherical (i.e. the 269 270 shape oscillation has no amplitude so no shape deformation occurs). For $A_2/a = \pm 0.25$, the droplet has an 271 elliptical shape, corresponding to an intermediate deformation between the two extremes. Figure 2b shows 272 the change in the integrated backscatter with variation in A_2/a for an assumed droplet radius of 5.621 µm with $\sigma = 74.5 \text{ mN} \cdot \text{m}^{-1}$, $\rho = 1059 \text{ kg} \cdot \text{m}^{-3}$, and $\eta = 1.28 \text{ mPa} \cdot \text{s}$. For comparison, integrated backscatter for two 273 274 additional droplet radii are included: one where the radius is larger by $\lambda/4$, where λ is the wavelength of the 275 trapping beam (532 nm), and another where the radius is smaller by $\lambda/4$. The inclusion of these two

10

Page 11 of 31

Chemical Science

additional droplet sizes illustrates the sensitivity in the light scatter signal to relatively small (~2%) changes in droplet radius. From this figure it is clear that at positive relative amplitudes ($A_2/a > 0$) the intensity of the backscatter is much larger than the backscatter at negative relative amplitudes ($A_2/a < 0$). The difference results directly from the position of the trapping beams relative to the composite droplet: when $A_2/a > 0$ the composite droplet intercepts the trapping beams while it does not when $A_2/a < 0$.

281 The time resolved signal (shown in Fig. 2c) is determined by considering the evolution of $A_{2/a}$ governed by the damped oscillation of the surface.^{5, 31} Again, the three radii were considered. The result is a 282 283 backscattered light signal with additional features superimposed on an underlying oscillating signal. These 284 features are broadly similar to those observed in the experiment (compare to Fig. 1b). The dominant 285 influence on the signal amplitude arises from the extent to which the beams intercept the oscillating droplet 286 surface. Higher frequency features superimposed on the underlying structure arise from the Fabry-Perot type 287 resonance caused by optical path length changes through the droplet and the concomitant interference fringes 288 expected in reflected light. Moreover, for the three droplet radii shown, although there are differences in the 289 intensities of various higher frequency features, the general form of the signal is consistent, indicating that 290 these higher frequency features should not complicate the interpretation of the experimental results.

Figure 2d shows the corresponding FFTs of the simulated signals. Again, the dominant feature is a single resonance peak at the frequency corresponding to the l = 2 mode (calculated from the input parameters). Clearly, small changes in the droplet radius lead to resolvable shifts in the resonance peaks, highlighting the need for highly accurate particle size measurements. Such differences are readily resolvable in our measurements.

296

297 Measurements of Droplet Viscosity and Surface Tension

Inferring the droplet surface tension from the modal frequencies estimated from the FFT of the elastic light scattering signal requires an accurate determination of the radius of the droplet and the density, apparent from Eq. (2). We have shown previously that the radius and RI can be determined with high accuracy from the fingerprint of WGMs apparent in the Raman spectrum.²⁹ Once the RI is known, the droplet composition can be inferred and the density estimated. The relationship between density, RI, and mass fraction of solute can be treated by the molar refraction mixing rule.^{40, 41} We now return to consider the oscillation frequency measured for the FFT shown in Fig. 1d.

Chemical Science Accepted Manuscript

When viscous damping is significant ($\eta > \sim 1 \times 10^{-3}$ Pa·s for the droplet sizes considered here), the influence of this dissipation on the natural oscillation frequencies is non-negligible and the measured oscillation frequency (ω_l^*) is lower than that predicted using Eq. (2) and is more accurately given by:²⁷

$$\omega_l^* = \sqrt{\omega_l^2 - \tau_l^{-2}} \tag{3}$$

308

309 With the signal dominated by the l = 2 mode of oscillation, the characteristic damping time $\tau_{l=2}$ (and 310 consequently viscosity) can determined from the exponentially decaying light scatter signal. Once $\tau_{l=2}$ is 311 determined, the frequency and surface tension can be determined in the non-dissipative limit. While 312 neglecting the influence of damping may be increasingly valid as the particle size increases, for the particles 313 sizes investigated here this can lead to systematically low values for the inferred surface tension, which may 314 be as large as 4 mN \cdot m⁻¹ for the smallest, most viscous droplets studied. Note that if the viscosity is larger than the critical viscosity ($\eta_{crit} \ge 0.76 \times (a\sigma p)^{1/2}$), then surface oscillations are so efficiently damped that only 315 aperiodic relaxation is observed, i.e. a slow merging of two droplets.³⁰ In this limit, the surface tension 316 317 cannot be obtained by this method.

318 The viscosity of aqueous organic aerosols is less well understood than for inorganic components. Numerous studies suggest that organic aerosols may exist in a highly viscous state.^{6, 7, 42, 43} The viscosities for 319 320 glutaric acid droplets inferred from the damping times are reported in Fig. 3 and compared with bulk phase 321 measurements. To determine the viscosity, the maxima in the backscattered light signal were fit to an 322 exponential decay to determine the damping time. Note that whereas initially the backscattered light does not 323 follow an exponential decay, at later times it does (see Fig. 1b). Therefore, the number of maxima included 324 in the fit was systematically varied to produce the fit most consistent with the later portion of the backscatter 325 signal. The damping times and additional droplet parameters (a, ρ) were then used to determine η from Eq. 326 (1). As shown in Fig. 3, the agreement between the optical tweezers approach and bulk measurements of 327 solution viscosity is very good.

328 Strictly, the linear theory of Rayleigh is valid only for small amplitude oscillations. Non-linear 329 effects including mode coupling and frequency modulation may dominate when the distortion of the droplet 330 exceeds 10% of the unperturbed droplet radius.^{44, 45} The shape distortion is increasingly dominated by the l =331 2 mode as the amplitude of surface oscillations decreases; the oscillation frequency then asymptotically

approaches the values predicted by the linear theory. Thus, the oscillation frequency is determined in thisasymptotic limit to exclude possible hydrodynamic non-linearities (see Fig. S2 in ESI).

334 Droplet surface tensions inferred from the coalescence measurements are shown in Fig. 4 for 335 droplets containing sodium chloride (Fig. 4a) or glutaric acid (Fig. 4b). Both sodium chloride and glutaric 336 acid are atmospherically relevant compounds. Sodium chloride makes up a significant fraction of 337 atmospheric aerosol, generated by mechanical processes at the marine surface.⁴⁶ Similarly, the dicarboxylic 338 acids are atmospherically prevalent and glutaric acid is often chosen as a surrogate for water soluble, 339 surface-active organic aerosol.

340 In Fig. 4, the lower x-axis shows solute concentration, which is then used to infer water activity, a_w 341 (shown on the upper x-axis). The relationship between concentration and a_w can be determined directly from 342 E-AIM calculations for sodium chloride. For glutaric acid, a_w was determined through parametrizations of 343 bulk concentration and density coupled with E-AIM calculations to determine the solute mass fraction as a 344 function of RH. For the optical tweezers results, individual measurements are aggregated into 0.2 M 345 concentration bins. Uncertainties reported on the x- and y-axes are the standard deviation of the 346 measurements in any given bin. If no uncertainty is provided, then only one data point fell into that specific 347 concentration bin.

In Fig. 4a, surface tensions of sodium chloride droplets obtained by the optical tweezers approach are compared with bulk tensiometry measurements and E-AIM model predictions. Additionally, results from a recent study of the surface tension of submicron aerosol by AFM are included.²⁸ Overall, there is very good agreement between the optical tweezers measurements and the other measurements, validating the experimental approach. The solubility limit of NaCl is approximately 6 M, so the optical tweezers results extend into the supersaturated solute regime and appear to agree well with expected values (from E-AIM).

In Fig. 4b, surface tensions of glutaric acid droplets obtained by the optical tweezers approach are compared with bulk tensiometry measurements and a number of different parametrizations of surface tension and concentration. These parametrizations come from several independent measurements⁴⁷⁻⁵¹ but we use the parameters of the Szyszkowski equation for each that are summarized in Lee and Hildemann.⁴⁹ In addition, we show results from the AFM approach only for a_w close to the saturation limit, which is closest to our measured concentrations. Overall, our experimental results agree very well with bulk measurements. Among the various parametrizations, our measurements agree best with the parametrization of Gaman et al.⁴⁸ and 13

appear to be consistent with the AFM results at high a_w . Overall, the precision of the optical tweezers measurements typically is < 1 mN·m⁻¹. Notably, the inorganic salt and organic dicarboxylic acid show the anticipated increase⁵² and decrease in surface tension, respectively, relative to water (72.6 mN·m⁻¹ at 294 K).

A related parameter encountered in surface science is the dynamic surface tension. When forming 364 365 new surface area, diffusion to the interface is non-instantaneous resulting in a surface age dependent surface 366 tension. This phenomenon is readily observed in macroscopic measurements but also in oscillating droplets. 367 Using the falling droplet approach, Yang et al. found that for droplets with volumes of 100s of picolitres and 368 surface ages on the order of $10 - 100 \,\mu$ s, the surface tensions of droplets were considerably higher than expected based on equilibrium values.² This observation is consistent with an expectation that a surfactant 369 370 film (depending on surface excess and concentration) achieves an equilibrium composition on a timescale 371 considerably longer than the surface age in such measurements, typically ~ 20 ms. No such effect is seen here 372 despite the extremely short surface ages during droplet oscillation. However, it should be noted that the 373 relevant timescale governing the surface compositions of the two droplets is the time frame prior to 374 coalescence when they can be assumed to reach their equilibrium values.

375 During the relaxation process, the surface-to-volume ratio must relax by a factor of 1.26 with initial 376 oscillations that approach a change of this order. Over one period of oscillation in surface area ($\sim 10 \mu s$), a 377 typical solute with a diffusion constant of 2×10^{-9} m²·s⁻¹ may diffuse over ~200 nm. Thus, the expansion and 378 contraction of the surface area remains sufficiently slow that the near-surface concentration gradients 379 established by the inorganic salt double layer and the surface excess for the organic component can rapidly 380 compensate, maintaining a near-surface equilibrium distribution of solute molecules throughout the 381 relaxation process. Such an observation may provide insight into the value of the surface tension during water condensation and CCN activation in the atmosphere.^{53, 54} Moreover, turbulent mixing is avoided under 382 the low Reynolds number conditions encountered here, similarly maintaining the radial concentration 383 gradients that exist in the pre-coalescence pair.⁵⁵ In fact, the essentially static values of surface tension 384 385 measured are more comparable with those of Yamada et al. for coalescence under large shear deformations²⁵ and Apfel et al. for super-deformed droplets under microgravity,²⁴ where static values of the surface tension 386 387 are observed for oscillating droplets. Thus, the surface tensions inferred from these measurements match the 388 expected equilibrium values and provide a robust method for determining the surface tension of 389 atmospherically relevant droplet compositions.

390

391 The Contamination of Droplet Surfaces

392 It has long been assumed that liquid surfaces can become contaminated through adsorption of trace 393 impurities in the gas phase and this can have consequences, in particular, for understanding mass transfer rates across the surface of liquid droplets.²¹ Despite this presumption, direct confirmation has not been 394 395 possible. Using the droplet coalescence approach it is possible to directly explore the possible time-396 dependent contamination of liquid surfaces. Up to this point, coalescence has been induced within 10 s of 397 droplet capture. To study surface adsorption from the gas-phase, pairs of sodium chloride droplets are 398 retained in the optical traps for well-defined times ranging from 10 s to > 7000 s and representing more than 399 2 orders of magnitude in surface age.

400 In Fig. 5, we report the surface tension depression recorded for variably surface aged droplet pairs. A clear decrease in surface tension is observed with surface age, converging to $\sim 30 \text{ mN} \cdot \text{m}^{-1}$ when pairs of 401 402 droplets were captured and held in the trapping chamber opened to laboratory air (no gas flow introduced 403 into the cell). The time scale for the decrease in surface tension is characteristic of adsorption from the gas-404 phase rather than the surface partitioning of organic impurities from the droplet bulk, which would be very 405 rapidly salted out under conditions of high ionic strength. In addition, time dependencies of two different 406 starting salt concentrations (0.7 and 5.1 M) were found to be similar, which would not be expected if the 407 bulk solution concentration of impurities were increased by a factor of 7.5. The bulk surface tension data 408 recorded using a tensiometer corroborates this interpretation. It should be stressed that the level of contamination may remain very small: forming a complete monolayer would require $\sim 10^{10}$ molecules or 10^{-10} 409 ¹⁴ moles. Droplet surface tension relaxes to an approximately constant value around 30 mN \cdot m⁻¹ after about 410 411 2000 s in the trap. This timescale for equilibration (τ_{σ}) can be related to a Henry's law constant (H) for the 412 condensing species if its molar mass (M) and accommodation coefficient (α) are known, where a is the 413 droplet radius, *R* is the gas constant, and *T* is temperature:

414
$$\tau_{\sigma} = \frac{r \times H \sqrt{2\pi MRT}}{3\alpha}$$
(4)

415 Accommodation coefficients can vary widely for different compounds but are frequently reported in 416 the range of 0.01-1.0. Assuming a compound with molar mass of 100 g·mol⁻¹ condenses on an 8 μ m radius 417 droplet, the Henry's law constant would be expected to be in the range from 10⁴ to 10⁶ M·atm⁻¹, depending 15

Fage 10 01

on the assumed value of α . Such a range for the Henry's law constant could be consistent with absorption and contamination of the droplet surface by variety of ambient species including glyoxal and oxalic acid. Absorption and dissolution would be clearly evident in the fingerprint region (500 – 3000 cm⁻¹, 550 – 625 nm) of the Raman spectrum if the particle were to attain an appreciable bulk concentration of organics. Such a change is not observed here, indicating a change in composition below the Raman detection limit (typically ~100 mM).

424 Further evidence that the depression in surface tension can be attributed to adsorption from the gas 425 phase is obtained by generating droplets from a bulk solution intentionally doped with an ionic surfactant 426 (sodium dodecyl sulphate). In each case, coalescence was induced quickly following capture (< 10 s). In 427 contrast to the data in Fig. 5, a marked depression of the inferred surface tension is immediately apparent (to 428 60±4 mN·m⁻¹ over 8 individual measurements). (The range of measured surface tensions for the surfactant-429 doped droplets is larger than for the single component droplets due to the variability in the absolute quantity 430 of surfactant in the droplet for each coalescence event.) Sodium chloride remains in a 100:1 excess and the 431 amount of surfactant in the droplet is insufficient to form a full surface monolayer, typically containing 10^7 – 10^8 molecules, assuming the stoichiometry of the solute components in the droplet remains the same as in the 432 433 bulk.

Also shown in Fig. 5 are the results of two other experiments where we examine the response of droplet surface tension to the environmental conditions. Initially we use a cylinder of ultrapure nitrogen, which shows only a minor improvement over laboratory air in terms of the time required for the surface tension depression. Using the boil-off from a liquid nitrogen dewar is the most reliable way of avoiding the presence of trace quantities of contaminants in the gas phase. In this case, suppression of droplet surface tension over many thousands of seconds becomes possible.

440

441 Time-Dependent Observations of the Adsorption of a Water Soluble Organic and the Change in
442 Surface Tension

We have demonstrated that the experimental approach is capable of performing precise measurements of droplet surface tension and that it is possible to control the local environment surrounding the trapped droplets such that undesired contamination of the droplet surface is minimized over long time periods. As a next step a gas phase "contaminant" is intentionally introduced and the response of droplet 16

Page 17 of 31

Chemical Science

Chemical Science Accepted Manuscript

surface tension explored; that is, the response of droplet surface tension to changes in the gas phase is monitored. This process has relevance to atmospheric aerosol microphysics where aerosol chemical composition can change substantially, likely impacting surface tension. A proof-of-concept example of such an experiment is shown in Fig. 6. In this experiment, aqueous sodium chloride droplets were trapped in the cell and subjected to a vapour flow from air bubbled through a solution of water and ethanol (97.5%/2.5% volume/volume). The salt concentration in the droplets was approximately 1 M.

453 In the upper panel of Fig. 6, we report the change in droplet surface tension as a function of time 454 exposed to the vapour flow. Droplet surface tension at time t = 0 s (i.e. a purely aqueous NaCl droplet) is about 73 mN·m⁻¹. By comparison, a surface tension of 74 mN·m⁻¹ was measured for a bulk 1 M aqueous 455 456 NaCl solution (dotted line in the figure), in close agreement with the droplet measurements (where salt 457 concentration is not precisely controlled but can be inferred from measurements of RI). As the droplet 458 equilibrates to the vapour flow, the surface tension decreases and after ~ 200 s stabilizes around a value of ~63 mN·m⁻¹. If the liquid reservoir from which the vapour flow arises is converted into a 1 M NaCl solution 459 460 (i.e. 1 M NaCl in 97.5% water, 2.5% ethanol by volume), bulk measurements of surface tension yield a value of 64 mN \cdot m⁻¹ (dotted line in the figure), corresponding well with the droplet measurements and indicating 461 462 that the droplet has equilibrated with the vapour flow.

463 Although the change in surface tension is substantial ($\sim 15\%$), such a change only corresponds to a 464 very minor change in droplet bulk composition. The lower panel of Fig. 6 shows the accompanying change 465 in Raman CH/OH band ratio. The CH/OH ratio can provide a quantitative measure of ethanol uptake. The 466 CH/OH band ratio changes on the same timescale as droplet surface tension, confirming that the change in 467 surface tension is indeed coming from a change in the droplet composition due to absorption of ethanol. 468 Further, the change in droplet composition is very small and barely resolvable from the change in Raman 469 intensity. In short, this proof-of-concept experiment demonstrates that our experimental approach is capable 470 of resolving changes in droplet surface tension for very small changes in droplet chemical composition and 471 promises to allow us to quantitatively resolve the impact of atmospherically relevant chemistry on droplet 472 surface tension.

473

474 CONCLUSIONS

475 We have demonstrated for the first time that contactless measurements of surface tension can be 476 made on micrometre-sized droplets containing only 1 - 4 pL of material. Holographic optical tweezers 477 provide a versatile platform for conditioning aerosol, controllably inducing coalescence and probing the 478 dynamics of the subsequent relaxation in shape via damped oscillations. By freeing the droplet from contact 479 with a substrate or external probe and in concert with spectroscopic measurements of particle size and RI, 480 this platform can be utilised as an all optical tensiometer. We demonstrated that measurements of surface 481 tension can be made on simple organic/inorganic solute droplets with a comparable level of accuracy (± 1) 482 mN·m⁻¹) to bulk phase measurements. Additionally, measurements of the droplet viscosity are also possible 483 and agree well with bulk values, essentially enabling a simultaneous determination of both surface tension 484 and viscosity in one measurement. Therefore, this approach allows for a full exploration of the relaxation 485 dynamics following droplet coalescence and provides complementary data to previous studies of aerosol 486 viscosity.

487 Measurement of the droplet surface tension is limited to droplets with viscosities below the critical 488 value, above which surface oscillations become overdamped (generally 10-20 mPa·s for droplet radii studied 489 here). However, droplet surface tension is key to the activation of aerosol to cloud droplets, which occurs at 490 high water activities. As a result, activating droplets would be expected to have substantial water content and 491 viscosities below this critical value. Droplets analyzed by this approach are typically between 5-10 µm 492 radius, larger than the typical diameter of a CCN at activation ($\sim 1 \mu m$), but allowing direct measurements of 493 comparably sized droplets for the first time. In principle, surface tension measurements over the typical 494 range accessible to aqueous droplets are measurable by this approach.

495 Light scattering simulations reproduce the form of the oscillatory light scattering signal, 496 simultaneously confirming our interpretation of the data and providing an experimental validation of models 497 used to describe light scattering from non-spherical dielectric objects. Surface tension measurements are 498 performed on microsecond timescales (the period during which the droplet shape is oscillating). Since these 499 measurements agree well with bulk values, the agreement suggests that equilibrium is retained in the surface 500 composition during shape relaxation when starting from droplets with pre-equilibrated composition. Such a 501 result has implications for activating aerosol growing into cloud droplets from an equilibrium state, as the 502 surface tension at any given point along the growth can be taken as its equilibrium value.

503 Building on the ability to measure precisely both surface tension and viscosity, we have explored the 504 contamination of droplets due to adsorption of organics from the gas-phase, which represents a key 505 mechanism by which droplets may obtain a surfactant coating. Recent studies have suggested the possibility 506 of co-condensation of organic vapours during hygroscopic growth as a route to enhanced cloud droplet 507 number concentrations.⁵⁶ The surface tension of aqueous droplets was observed to decrease, tending towards 508 a value for a saturated surfactant coated surface over a timescale of minutes to hours. These data have a 509 significant implication for the assumed surface tension values of aqueous droplets in any environment and, 510 more specifically, on the assumed value in atmospheric models of cloud droplet number. Although further 511 studies are necessary to explore the contamination of droplet surfaces under tropospherically relevant 512 conditions, this study suggests that assuming a surface tension of a surfactant coated surface ($\sim 30 \text{ mN} \cdot \text{m}^{-1}$) 513 may be appropriate. Moreover, we show that we are able to control the local environment and monitor the 514 response of droplet surface tension. Previous studies on aerosol extracts have shown that the organic fraction 515 of aerosol certainly contains surfactants and a surface tension depression on the order of $\sim 60\%$ could correspond to an increase in cloud droplet number concentration as large as ~40%.¹⁰ Regardless, we have 516 517 shown that surface partitioning and adsorption from the gas-phase can be probed in micrometre sized aerosol 518 droplets which should provide fertile ground for further studies aiming to disentangle the various 519 contributors to aerosol activation and the effect of aerosol on climate.

520

521 ELECTRONIC SUPPLEMENTARY INFORMATION

Electronic supplementary information (ESI) available: parametrizations used to infer concentration, density,
viscosity, and surface tension from refractive index for sodium chloride and glutaric acid; description of the
semi-analytical T-matrix calculations; Figs. S1-S2. See DOI: 10.1039/XXXXXX.

525

526 ACKNOWLEDGEMENTS

527 The authors acknowledge Jean-Pierre Wolf and Jérôme Kasparian for loan of the high frame rate camera.
528 J.P.R., B.R.B., R.M.P., and C.P.R. acknowledge support from the Engineering and Physical Sciences
529 Research Council (EPSRC) through grant EP/L010569/1. S.H.S. was supported by the Ministry of
530 Education, Youth and Sports of the Czech Republic (project L01212). C.P.R. acknowledges the Royal
531 Society and the European Research Council (ERC Consolidator Grant NANOPRS, project number 617266).

19

532 Author contributions: B.R.B. undertook final measurements presented in the paper and the lead in writing the 533 manuscript. R.M.P. undertook preliminary measurements and analysis. S.H.S. provided all modeling and 534 computational support. The experimental data are provided through the University of Bristol data repository 535 Reid, J.P. (2015): Hygroscopicity, University of DOI: at Aerosol Bristol, 536 10.5523/bris.162fotifg0d961a0gb3nogipwe.

537	REFERENCES	
538 539	1.	J. Eggers and E. Villermaux, <i>Rep. Prog. Phys.</i> , 2008, 71 , 036601, doi: 10.1088/0034-4885/71/3/036601.
540	2.	L. Yang, B. K. Kazmierski, S. D. Hoath, S. Jung, WK. Hsiao, Y. Wang, A. Berson, O. Harlen, N.
541	2	Napul and C. D. Dalli, <i>Phys. Fluids</i> , 2014, 20 , 115105, doi: 10.1005/1.4901625.
542	5. 1	N. J. Cha, A. Dehusigho and M. Plakash, Nature, 2015, 519 , 440-450.
545	4.	C. K. Kueni and K. K. Wilson, J. Phys. Chem. A, 2014, 118, 3952-3966.
544	5.	R. M. Power, S. H. Simpson, J. P. Reid and A. J. Hudson, <i>Chem. Sci.</i> , 2013 , 4, $259/-2004$.
545	0.	D. L. Bones, J. P. Keld, D. M. Llennard and U. K. Krieger, <i>Proc. Natl. Acaa. Sci. U. S. A.</i> , 2012, 100, 11(12, 11(12))
546	7	
547	1.	A. Virtanen, J. Joutsensaari, I. Koop, J. Kannosto, P. Yil-Pirila, J. Leskinen, J. M. Makela, J. K.
548 549		827.
550	8.	Y. You, L. Renbaum-Wolff, M. Carreras-Sospedra, S. J. Hanna, N. Hiranuma, S. Kamal, M. L.
551 552		Smith, X. Zhang, R. J. Weber, J. E. Shilling, D. Dabdub, S. T. Martin and A. K. Bertram, <i>Proc. Natl. Acad. Sci. U. S. A.</i> , 2012, 109 , 13188-13193.
553	9.	A. Nenes, R. J. Charlson, M. C. Facchini, M. Kulmala, A. Laaksonen and J. H. Seinfeld, <i>Geophys.</i>
554		Res. Lett. 2002. 29. 1848. doi: 10.1029/2002g1015295.
555	10.	M. C. Facchini, M. Mircea, S. Fuzzi and R. J. Charlson, <i>Nature</i> , 1999, 401 , 257-259.
556	11.	B. Noziere, C. Baduel and JL. Jaffrezo, Nature Communications, 2014, 5, 3335, doi:
557		10.1038/ncomms4335.
558	12.	V. F. McNeill, N. Sareen and A. N. Schwier, in Atmospheric and Aerosol Chemistry, eds. V. F.
559		McNeill and P. A. Ariya, 2014, vol. 339, pp. 201-259.
560	13.	H. Kohler, Trans. Faraday Soc., 1936, 32 , 1152-1161.
561	14.	D. Rosenfeld, U. Lohmann, G. B. Raga, C. D. O'Dowd, M. Kulmala, S. Fuzzi, A. Reissell and M. O.
562		Andreae, <i>Science</i> , 2008, 321 , 1309-1313.
563	15.	H. Wex, M. D. Petters, C. M. Carrico, E. Hallbauer, A. Massling, G. R. McMeeking, L. Poulain, Z.
564		Wu, S. M. Kreidenweis and F. Stratmann, Atmos. Chem. Phys., 2009, 9, 3987-3997.
565 566	16.	C. R. Ruehl, P. Y. Chuang, A. Nenes, C. D. Cappa, K. R. Kolesar and A. H. Goldstein, <i>Geophys. Res. Lett.</i> , 2012, 39 , L23801, doi: 10.1029/2012gl053706.
567	17.	D. O. Topping, G. B. McFiggans, G. Kiss, Z. Varga, M. C. Facchini, S. Decesari and M. Mircea,
568		Atmos. Chem. Phys., 2007, 7, 2371-2398.
569	18.	B. R. Bzdek, M. J. Lawler, A. J. Horan, M. R. Pennington, J. W. DePalma, J. Zhao, J. N. Smith and
570		M. V. Johnston, Geophys. Res. Lett., 2014, 41, 6045-6054, doi: 10.1002/2014gl060160.
571	19.	J. Wang and A. S. Wexler, <i>Geophys. Res. Lett.</i> , 2013, 40 , 2834-2838, doi: 10.1002/grl.50455.
572	20.	Y. Cheng, H. Su, T. Koop, E. Mikhailov and U. Poeschl, Nat. Comm., 2015, 6, 5923, doi:
573		10.1038/ncomms6923.
574	21.	J. F. Davies, R. E. H. Miles, A. E. Haddrell and J. P. Reid, Proc. Natl. Acad. Sci. U. S. A., 2013, 110,
575		8807-8812.
576	22.	V. F. McNeill, J. Patterson, G. M. Wolfe and J. A. Thornton, Atmos. Chem. Phys., 2006, 6, 1635-
577		1644.
578	23.	N. Sareen, A. N. Schwier, T. L. Lathem, A. Nenes and V. F. McNeill, <i>Proc. Natl. Acad. Sci. U. S. A.</i> ,
579	~ /	2013, 110 , 2723-2728.
580	24.	R. E. Aptel, Y. Tian, J. Jankovsky, T. Shi, X. Chen, R. G. Holt, E. Trinh, A. Croonquist, K. C.
581		Thornton, A. Sacco, C. Coleman, F. W. Leslie and D. H. Matthiesen, <i>Phys. Rev. Lett.</i> , 1997, 78, 1012–1015
582 582	25	1912-1913. T. Varrada and K. Sakai, Dhun, Ehvida 2012 24, 022102 dai: 10.10(2/1.2(01010)
583	25. 26	1. Yamada and K. Sakai, <i>Phys. Fluids</i> , 2012, 24 , 022103, doi: 10.1063/1.3681810.
504 505	∠0. 27	L. Rayleigii, Froc. R. Soc. London, 10/9, 29, /1-9/. U. Lamb. Hydrodynamias 6 th Edition Combridge University Dress, Combridge, 1022
585 586	∠1. 28	H. Lamo, <i>Hydrodynamics</i> , 0 <i>Lamon</i> , Camonage University Press, Camonage, 1952.
587	∠o. 20	11. 5. IVIOLLES, V. 11. OLASSIALI ALL A. V. LIVALISKI, CHEM. SCI., 2013, 0, 5242-5247. T. C. Dreston and I. D. Deid, I. Ont. Soc. Am. P. 2012, 20, 2112, 2122.
588	29. 30	1. C. 1105001 dilu J. 1. Kelu, J. Opt. 50C. All. D. 2015, 30 , 2115-2122. S. Chandrasekhar Proc London Math. Soc. 1050 30 , 141-140
580	30. 31	B. Chandrasekhal, 1700 London Main. 500., 1757, 5-7, 141-147. R. M. Dower and I. D. Reid Rep. Drog. Dhys. 2014. 77, 074601. doi: 10.1000/0024
590	51.	4885/77/7/074601.
591	32.	J. R. Butler, J. B. Wills, L. Mitchem, D. R. Burnham, D. McGloin and J. P. Reid. Lab on a Chin.
592		2009, 9, 521-528.

Chemical Science Accepted Manuscript

Chemical Science

- 593 33. D. O. Topping, G. B. McFiggans and H. Coe, Atmos. Chem. Phys., 2005, 5, 1205-1222.
- 594 34. D. O. Topping, G. B. McFiggans and H. Coe, Atmos. Chem. Phys., 2005, 5, 1223-1242.
- 595 35. S. L. Clegg, P. Brimblecombe and A. S. Wexler, J. Phys. Chem. A, 1998, 102, 2155-2171.
- 596 36. R. M. Power, D. R. Burnham and J. P. Reid, *Appl. Opt.*, 2014, **53**, 8522-8534.
- 597 37. M. Guillon, R. E. H. Miles, J. P. Reid and D. McGloin, New J. Phys., 2009, 11, 103041, doi: 10.1088/1367-2630/11/10/103041.
- 38. R. Power, J. P. Reid, S. Anand, D. McGloin, A. Almohamedi, N. S. Mistry and A. J. Hudson, J. Phys. Chem. A, 2012, 116, 8873-8884.
- 601 39. A. D. Ward, M. G. Berry, C. D. Mellor and C. D. Bain, Chem. Comm., 2006, 4515-4517.
- 602 40. Y. Liu and P. H. Daum, J. Aerosol Sci., 2008, **39**, 974-986.
- 603
 41.
 I. N. Tang and H. R. Munkelwitz, J. Geophys. Res. Atmos., 1994, 99, 18801-18808, doi: 10.1029/94jd01345.
- 42. T. Koop, J. Bookhold, M. Shiraiwa and U. Poeschl, *Phys. Chem. Chem. Phys.*, 2011, 13, 19238-19255.
- 607 43. B. Zobrist, C. Marcolli, D. A. Pedernera and T. Koop, *Atmos. Chem. Phys.*, 2008, **8**, 5221-5244.
- 608 44. E. Becker, W. J. Hiller and T. A. Kowalewski, *J. Fluid Mech.*, 1991, **231**, 189-210.
- 609 45. E. Becker, W. J. Hiller and T. A. Kowalewski, *J. Fluid Mech.*, 1994, **258**, 191-216.
- 46. K. A. Prather, T. H. Bertram, V. H. Grassian, G. B. Deane, M. D. Stokes, P. J. DeMott, L. I.
 Aluwihare, B. P. Palenik, F. Azam, J. H. Seinfeld, R. C. Moffet, M. J. Molina, C. D. Cappa, F. M.
 Geiger, G. C. Roberts, L. M. Russell, A. P. Ault, J. Baltrusaitis, D. B. Collins, C. E. Corrigan, L. A.
 Cuadra-Rodriguez, C. J. Ebben, S. D. Forestieri, T. L. Guasco, S. P. Hersey, M. J. Kim, W. F.
 Lambert, R. L. Modini, W. Mui, B. E. Pedler, M. J. Ruppel, O. S. Ryder, N. G. Schoepp, R. C.
 Sullivan and D. Zhao, *Proc. Natl. Acad. Sci. U. S. A.*, 2013, 110, 7550-7555.
- 616 47. E. Aumann, L. M. Hildemann and A. Tabazadeh, Atmos. Environ., 2010, 44, 329-337.
- 48. A. I. Gaman, M. Kulmala, H. Vehkamaki, I. Napari, M. Mircea, M. C. Facchini and A. Laaksonen, *J. Chem. Phys.*, 2004, **120**, 282-291.
- 619 49. J. Y. Lee and L. M. Hildemann, Atmos. Environ., 2014, 89, 260-267.
- 50. M. L. Shulman, M. C. Jacobson, R. J. Carlson, R. E. Synovec and T. E. Young, *Geophys. Res. Lett.*, 1996, 23, 277-280, doi: 10.1029/95gl03810.
- 622 51. R. Tuckermann, Atmos. Environ., 2007, 41, 6265-6275.
- 623 52. L. M. Pegram and M. T. Record, Jr., Proc. Natl. Acad. Sci. U. S. A., 2006, 103, 14278-14281.
- 624 53. M. D. Petters and S. M. Kreidenweis, *Atmos. Chem. Phys.*, 2013, **13**, 1081-1091.
- 625 54. R. Sorjamaa, B. Svenningsson, T. Raatikainen, S. Henning, M. Bilde and A. Laaksonen, *Atmos. Chem. Phys.*, 2004, 4, 2107-2117.
- 627 55. K. Avila, D. Moxey, A. de Lozar, M. Avila, D. Barkley and B. Hof, *Science*, 2011, **333**, 192-196.
- 628 56. D. Topping, P. Connolly and G. McFiggans, *Nat. Geosci.*, 2013, **6**, 443-446. 629

630 FIGURE CAPTIONS

Figure 1: Coalescence of an aqueous ~1 M NaCl droplet doped with the surfactant sodium dodecyl sulphate. a) High frame rate images of the composite droplet for the first several microseconds after coalescence. Time ranges underneath each image provide the time period during which the image was taken. b) Backscattered light collected after coalescence (left axis, time t = 0 µs corresponds with the moment of coalescence) and droplet aspect ratios (a_y/a_x) determined from the high frame rate imaging (right axis). c) Fast Fourier Transform of the backscattered light gives the frequency of the shape oscillation.

638 Figure 2: Computational modeling of droplet coalescence with the following composite droplet properties: a = 5.621 μ m, σ = 74.5 mN·m⁻¹, ρ = 1059 kg·m⁻³, and η = 1.28 mPa·s. a) droplet shapes at various l = 2639 640 distortion amplitudes (A_2) . The trapping beams are perpendicular to the plane of the page and pass through the points on the x-axis at $y = +(a/2)^{1/3}$ and $y = -(a/2)^{1/3}$. b) intensity of the integrated 641 642 backscatter as a function of relative amplitude (A_2/a) . c) Simulated time-resolved backscattered light 643 for three droplet radii (offset for clarity). d) Fast Fourier Transform of the time-resolved 644 backscattered light in part c) to give the frequency of the shape oscillation. Colors in c) and d) follow 645 the legend in b).

Figure 3: Viscosity of glutaric acid as a function of molar concentration. Symbols give measurements on
 droplets using the optical tweezers approach. The line corresponds to bulk capillary viscometry
 measurements over the same concentration range. The upper axis gives water activity (a_w), which is
 inferred from the solute concentration.

Figure 4: Surface tension measurements on two atmospherically relevant systems: a) sodium chloride and b)
 glutaric acid. Also included are bulk tensiometry measurements, E-AIM modeling (sodium chloride),
 surface tension parametrizations performed by others (glutaric acid),⁴⁷⁻⁵⁰ and measurements on
 submicron aerosol using atomic force microscopy.²⁸ The upper axis gives water activity (a_w), which is
 inferred from the solute concentration.

Figure 5: Surface tension of dilute sodium chloride droplets held in the trapping cell for varying lengths of time. Red circles correspond to measurements where no gas flow was introduced to the cell. Green circles correspond to a gas flow at 80 mL·min⁻¹ using a high purity N_2 cylinder. Blue circles correspond to a gas flow at the same flow rate but using the boiloff flow from a liquid N_2 dewar.

Figure 6: Exposure of aqueous sodium chloride to a gas flow bubbled through a 2.5% ethanol/97.5% water (volume/volume) solution. In the upper panel, symbols give measured surface tension with an assigned $\pm 1 \text{ mN} \cdot \text{m}^{-1}$ uncertainty. Dotted lines correspond to a 1 M aqueous NaCl solution (74 mN·m⁻ 1) and a 1 M 2.5% ethanol solution (64 mN·m⁻¹). The lower panel gives the Raman CH/OH band ratio. Uncertainty bars correspond to the standard deviation of the ratio obtained from the spectra over the 20-50 seconds after coalescence.

665 666





671



672 **Figure 3**

















681

683 Table of Contents Entry



- 685 Descriptive Statement: Precise measurements of the surface tension and viscosity of airborne picolitre
- 686 droplets can be accomplished using holographic optical tweezers.