



Cite this: *Dalton Trans.*, 2024, **53**, 2252

Incorporation of U(IV) in monazite–cheralite ceramics under oxidizing and inert atmospheres[†]

Alison El Monjid, ^a Stéphanie Szenknect, ^{*a} Adel Mesbah, ^b Myrtille O. J. Y. Hunault, ^c Denis Menut, ^c Nicolas Clavier ^a and Nicolas Dacheux ^a

This work is the first attempt to prepare $\text{Nd}_{1-x}\text{Ca}_x\text{U}_x\text{PO}_4$ monazite–cheralite with $0 < x \leq 0.1$ by a wet chemistry method. This method relies on the precipitation under hydrothermal conditions ($T = 110^\circ\text{C}$ for four days) of the $\text{Nd}_{1-x}\text{Ca}_x\text{U}_x\text{PO}_4 \cdot n\text{H}_2\text{O}$ rhabdophane precursor, followed by its thermal conversion for 6 h at 1100°C in air or Ar atmosphere. The optimized synthesis protocol led to the incorporation of U and Ca in the rhabdophane structure. After heating at 1100°C for 6 h in air, single-phase monazite–cheralite samples were obtained. However, $\alpha\text{-UP}_2\text{O}_7$ was identified as a secondary minor phase in the samples heated under Ar atmosphere. The U speciation in the samples converted in an oxidising atmosphere was carefully characterized using synchrotron radiation by combining HERFD-XANES and XRD. These results showed the presence of a minor secondary phase containing hexavalent uranium and phosphate with a stoichiometry of $\text{U}:\text{P} = 0.78$. This highly labile uranyl phosphate phase incorporated 21 mol% of the uranium initially precipitated with the rhabdophane precursor. This phase was completely removed by a washing protocol. Thus, single-phase monazite–cheralite was obtained through the wet chemistry route described in this work with a maximum U loading of $x = 0.08$.

Received 4th October 2023,
Accepted 21st November 2023

DOI: 10.1039/d3dt03251e
rsc.li/dalton

Introduction

Monazite is a natural phosphate-based mineral that contains light rare earth elements uranium and thorium with amounts up to 16 wt% UO_2 and 29 wt% ThO_2 , respectively.^{1–3} In spite of the presence of mobile elements incorporated in the monazite structure and high doses of self-irradiation, the chemical durability of natural monazite under weathering conditions is remarkable.^{4–8} These properties meet the requirements of ceramic matrices for the long-term storage of high level nuclear waste (HLW) including actinides or long life fission products. Therefore, LnPO_4 monazite and its derivative have been studied for several decades for the specific conditioning of minor actinides (Am, Cm, and Np) and Pu.^{8–11} These studies have been conducted in several countries (*i.e.* France, Netherlands, Germany, Russia, USA, UK, Japan) where the separation of long-life radionuclides, especially minor actinides

and plutonium derived from dismantled nuclear weapons, was being considered. Indeed, the chemical flexibility of the monazite-type structure allows the incorporation of actinide elements in the trivalent and tetravalent oxidation states either by the direct substitution of Ln^{3+} or by coupled substitution of 2 Ln^{3+} by Ca^{2+} and An^{4+} .¹² The coupled substitution leads to the formation of $\text{Ln}_{1-2x}\text{Ca}_x\text{An}_x\text{PO}_4$ monazite–cheralite solid solutions. This mechanism is the most documented in the literature and allows complete Ca substitution resulting in the formation of $\text{Ca}_{0.5}\text{An}_{0.5}\text{PO}_4$ ($\text{An} = \text{Th, U, Np}$).^{13–17}

Thus, the monazite nine-coordinate site is able to incorporate large amounts of tetravalent cations through appropriate charge compensation with a divalent element like Ca. Podor and Cuney¹⁸ gave two steric criteria for the existence of such a $\text{Ln}_{1-2x}\text{Ca}_x\text{An}_x\text{PO}_4$ monazite–cheralite solid solution,

$$1.107 \text{ \AA} \leq r_{\text{average}} \leq 1.216 \text{ \AA} \text{ with} \quad (1)$$
$$r_{\text{average}} = (1 - 2x)^{\text{IX}} r_{\text{Ln}^{3+}} + x^{\text{IX}} r_{\text{Ca}^{2+}} + x^{\text{IX}} r_{\text{An}^{4+}}$$

$$1 \leq r_{\text{ratio}} \leq 1.238 \text{ with}$$

$$r_{\text{ratio}} = [(1 - 2x)^{\text{IX}} r_{\text{Ln}^{3+}} + x^{\text{IX}} r_{\text{Ca}^{2+}}] / [(1 - 2x)^{\text{IX}} r_{\text{Ln}^{3+}} + x^{\text{IX}} r_{\text{An}^{4+}}] \quad (2)$$

where ${}^{\text{IX}}r$ is the ionic radius of the element in the ninefold coordination.¹⁹

^aICSM, Univ. Montpellier, CNRS, CEA, ENSCM, Site de Marcoule, Bagnols/Cèze, France. E-mail: stephanie.szenknect@cea.fr

^bUniv. Lyon, Université Claude Bernard Lyon 1, CNRS, IRCELYON, F-69626 Villeurbanne, France

^cSynchrotron SOLEIL, L'Orme des Merisiers, Départementale 128, 91190 Saint-Aubin, France

[†]Electronic supplementary information (ESI) available. See DOI: <https://doi.org/10.1039/d3dt03251e>



Using these rules with ${}^{IX}r_{Nd^{3+}} = 1.163 \text{ \AA}$, ${}^{IX}r_{Ca^{2+}} = 1.18 \text{ \AA}$, and ${}^{IX}r_{An^{4+}} = 1.09 \text{ \AA}$, 1.05 \AA , 1.03 \AA , and 1.01 \AA ¹⁵ for Th, U, Np and Pu, respectively, allows the calculation of their maximum incorporation in NdPO_4 . This calculation shows that the coupled substitution can be complete for Th and U (*i.e.* $x = 0.50$) whereas it is limited to $x_{\text{max}} = 0.48$ and $x_{\text{max}} = 0.41$ for Np and Pu, respectively.

These steric considerations indicate that high incorporation rates of tetravalent An are achievable in the monazite/cheralite solid solution. Nevertheless, such high doping rates are generally obtained using dry chemistry routes, which could present difficulties for the multivalent elements (U, Np and Pu in particular)^{17,20} due to the relative instability of tetravalent actinide ions. Compared to Np(IV), the stability domains of U(IV) and Pu(IV) are less extended, as U(IV) is oxidised to U(VI) at lower potential and Pu(IV) is reduced to Pu(III) at higher potential.

Terra *et al.*²¹ synthesized $\text{La}_{1-2x}\text{Ca}_x\text{Th}_{x-y}\text{U}_y\text{PO}_4$ solid solutions (with $0.1 \leq x \leq 0.4$ and $y = x/5$) as well as end-members by a dry chemistry route. For that purpose, AnO_2 was mixed with La_2O_3 , $\text{Ca}(\text{HPO}_4)_2\text{H}_2\text{O}$ and $\text{NH}_4(\text{H}_2\text{PO}_4)$ and the mixture underwent several cycles of milling and calcination at $1200 \text{ }^\circ\text{C}$ under Ar atmosphere. At least two cycles were required to obtain homogeneous single-phase samples. Bregiroux *et al.*¹⁵ also investigated the formation of $\text{Ca}_{0.5}\text{U}_{0.5}\text{PO}_4$ prepared by a similar solid-solid reaction. They performed analyses of samples fired at several temperatures from 400 to $1400 \text{ }^\circ\text{C}$ under Ar atmosphere. Below $1100 \text{ }^\circ\text{C}$, $\alpha\text{-UP}_2\text{O}_7$ was found to co-exist with the cheralite phase, whereas the phosphates decomposed at temperatures higher than $1300 \text{ }^\circ\text{C}$, leading to the formation of CaO and UO_2 and the volatilization of $\text{P}_4\text{O}_{10}(\text{g})$. Thus, the range of temperature that allows the preparation of single-phase cheralite incorporating U was found to be rather constrained. It is worth noting that this synthesis route did not allow the preparation of $\text{Ca}_{0.5}\text{Ce}_{0.5}\text{PO}_4$ cheralite. The authors showed that the reduction of tetravalent Ce occurred during the thermal treatment, leading to the formation of a mixture of CePO_4 monazite and $\text{Ca}_2\text{P}_2\text{O}_7$. Similar results were obtained by Bregiroux *et al.*²⁰ using the same protocol to prepare $\text{Ca}_{0.5}\text{Pu}_{0.5}\text{PO}_4$. However, they also observed the partial incorporation of Pu(IV) in the cheralite when they fired the mixtures under air.

On the other hand, dustless wet chemistry methods have also been widely used to easily prepare LnPO_4 ²² and AnPO_4 monazite (with An = Pu, Am-Bk).²³⁻²⁷ Thus, the precipitation of $\text{LnPO}_4 \cdot n\text{H}_2\text{O}$ or $\text{AnPO}_4 \cdot n\text{H}_2\text{O}$ rhabdophane precursor followed by its thermal conversion leads to the preparation of single-phase monazite incorporating trivalent actinides. In contrast, the incorporation of high loadings of tetravalent actinides in monazite-cheralite appeared difficult using wet chemistry routes. Qin *et al.*²⁸ developed a wet chemistry method for the preparation of $\text{Ln}_{1-2x}\text{Ca}_x\text{Th}_x\text{PO}_4 \cdot n\text{H}_2\text{O}$ ($\text{Ln} = \text{Pr}$ and Nd) Th-rhabdophane solid solutions. Using this hydrothermal protocol, under acidic conditions, the limit of incorporation of Th in the solid solution was around 14 mol%. Beyond this limit, thorium phosphate-hydrogen phosphate hydrate (TPPHH),

$\text{Th}_2(\text{HPO}_4)(\text{PO}_4)_2 \cdot \text{H}_2\text{O}$, was formed as a secondary phase. The thermal conversion of $\text{Nd}_{1-2x}\text{Ca}_x\text{Th}_x\text{PO}_4 \cdot n\text{H}_2\text{O}$ (with $0 \leq x \leq 0.15$) to monazite/cheralite occurred between 750 and $800 \text{ }^\circ\text{C}$ under air.²⁹ Single phase monazite-cheralite solid solutions were obtained only for $x \leq 0.1$. For higher incorporation rates, $\alpha\text{-ThP}_2\text{O}_7$ was found as a secondary phase. Additionally, Qin *et al.*³⁰ demonstrated the feasibility of the direct sintering of $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{Th}_{0.1}\text{PO}_4 \cdot n\text{H}_2\text{O}$ rhabdophane and succeeded in obtaining dense monazite pellets at a temperature (*i.e.* $1400 \text{ }^\circ\text{C}$) significantly lower than those reported in the literature. Finally, the chemical durability of the $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{Th}_{0.1}\text{PO}_4$ monazite was investigated from a thermodynamic point of view.³¹ Solubility experiments demonstrated the exceptional chemical stability of this phase in environmental conditions. The results obtained guarantee the long-term efficient immobilization of Th in the monazite-cheralite matrix.

Thus, this type of wet process makes it possible to limit the risks of contamination of the facilities by the dissemination of radioactive dust associated with dry route processes. It also guarantees the preparation of dense and homogeneous pellets of monazite-cheralite incorporating up to 10 mol% of thorium. However, a similar protocol has not been tested for actinide elements prone to oxidation. Therefore, this work is a first attempt to prepare monazite-cheralite incorporating U(IV) by adapting the protocol developed by Qin *et al.*²⁹ Using uranium allows anticipating potential difficulties in the preparation of Np-doped monazite-cheralite by a wet chemistry route, without the difficulties linked to the handling of these highly radioactive radioelements. As the solubility of U(VI) is much higher than that of U(IV) in solution, particular attention was paid to characterizing the oxidation state of uranium in order to guarantee the long-term stability of uranium incorporated in the monazite-cheralite structure.

Results and discussion

Synthesis of U-rhabdophane

The chemical composition of the synthesized precursors was determined by elemental analyses of the solution after complete dissolution of an aliquot of each sample (Table 1). Regardless of the uranium content, the analyses indicated a slight deviation from the target composition. The $\text{P}/(\text{Nd} + \text{Ca} + \text{U})$ ratio was constant regardless of the value of x ($0.025 \leq x \leq 0.1$)

Table 1 Chemical composition of rhabdophane-type samples of formula $\text{Nd}_{1-2x}\text{Ca}_x\text{U}_x\text{PO}_4 \cdot n\text{H}_2\text{O}$, synthesized by hydrothermal route at $110 \text{ }^\circ\text{C}$ for 4 days

Target value	Nd		Ca		U		P	
	Ca + Nd + U	Ca + Nd + U	Ca + Nd + U	Ca + Nd + U	Ca + Nd + U	Ca + Nd + U	Ca + Nd + U	Ca + Nd + U
x	$1 - 2x$	x	x	x	x	x	x	1.00
0.025	0.96 ± 0.05	0.017 ± 0.001	0.026 ± 0.001	0.11 ± 0.06				
0.05	0.90 ± 0.05	0.042 ± 0.002	0.056 ± 0.003	1.12 ± 0.06				
0.075	0.85 ± 0.04	0.064 ± 0.003	0.082 ± 0.004	1.13 ± 0.06				
0.1	0.82 ± 0.04	0.085 ± 0.004	0.099 ± 0.005	1.12 ± 0.06				



and was systematically slightly higher than 1. The incorporation of uranium in the rhabdophane structure implies the coupled substitution of two Nd^{3+} by one Ca^{2+} and one U^{4+} . The results presented in Table 1 show a slight deviation from this mechanism. Whereas the neodymium and uranium mole fractions were close to the expected values, the Ca content was found to be lower than the target value despite the large excess of Ca introduced in the initial cation mixture. One assumption to explain these results would be the partial oxidation of uranium to the hexavalent state during the hydrothermal treatment. In that case, uranium introduced in the system that oxidised could also have partly precipitated with phosphate ions as a secondary phase. Another possibility might be the direct substitution of four Nd^{3+} by three U^{4+} and a vacancy. This type of direct substitution has been already reported in natural samples³² but never evidenced in synthetic ones.

The PXRD patterns of the prepared samples are plotted in Fig. 1 and compared to the rhabdophane reference.³³ They confirmed that the prepared samples were single-phase and only composed of the rhabdophane structure under its monoclinic form or that the amount of secondary phase is below the detection limit of this technique.

Rietveld refinements of the PXRD diagrams were carried out in order to show the impact of uranium incorporation on the unit cell parameters. The results displayed in Table 2 show a contraction of the unit cell volume with the increase of the uranium mole content. By considering the effective ionic radii of Nd^{3+} , Ca^{2+} and U^{4+} in the 9-fold coordination (1.163, 1.18 and 1.05 Å, respectively),¹⁹ this observation indicated that uranium was incorporated in the rhabdophane structure.

Finally, the prepared rhabdophanes were all characterized by SEM. A micrograph of the rhabdophane sample with $x = 0.1$ is presented in Fig. 2a. The needle-like morphology of the grains was similar to the one observed by Qin *et al.*²⁸ for Th-rhabdophane. Grains exhibiting a different morphology were not evidenced. These observations show the absence of any significant amount of secondary phases. Similar results were obtained for the other samples, regardless of the uranium content.

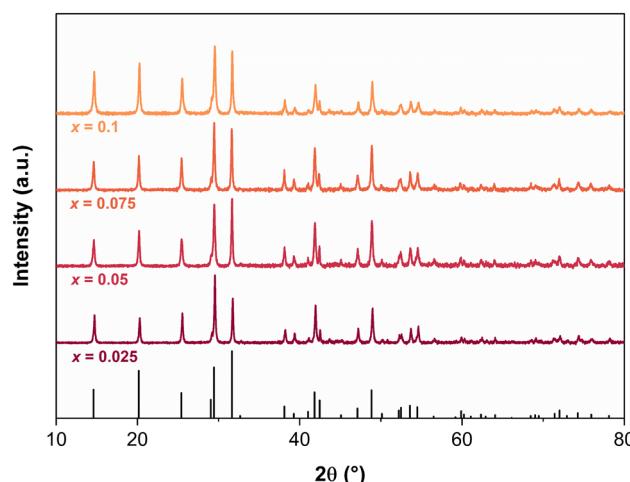


Fig. 1 PXRD patterns of the prepared $\text{Nd}_{1-x}\text{Ca}_x\text{U}_x\text{PO}_4\cdot n\text{H}_2\text{O}$ samples. XRD lines associated with $\text{SmPO}_4\cdot 0.667\text{H}_2\text{O}$ rhabdophane are included as reference³³ (black vertical bars).

Table 2 Unit cell parameters extracted from .cif file after Rietveld refinement of the structure of rhabdophane samples ($\text{Nd}_{1-x}\text{Ca}_x\text{U}_x\text{PO}_4\cdot n\text{H}_2\text{O}$) and monazite–cheralite samples ($\text{Nd}_{1-x}\text{Ca}_x\text{U}_x\text{PO}_4$) prepared in air ① or under Ar atmosphere ② after heating for 6 h at 1100 °C. Data reported in italics and bold were obtained from the SPXRD pattern

	x	a (Å)	b (Å)	c (Å)	β (°)	V (Å ³)
Rhabdophanes	0.025	28.3216(1)	7.0062(1)	12.1320(1)	115.47(1)	2173.36(1)
	0.05	28.3231(1)	7.0039(1)	12.1251(1)	115.46(1)	2171.62(1)
	0.075	28.3174(1)	7.0012(1)	12.1222(1)	115.40(1)	2170.94(1)
	0.1	28.3192(1)	6.9979(1)	12.1210(1)	115.40(1)	2169.87(1)
①	0.025	6.7383(1)	6.9523(1)	6.4083(1)	103.69(1)	291.68(1)
	0.05	6.7359(1)	6.9490(1)	6.4087(1)	103.69(1)	291.45(1)
	0.075	6.7319(1)	6.9437(1)	6.4048(1)	103.71(1)	291.01(1)
	0.1	6.7293(1)	6.9409(1)	6.4064(1)	103.70(1)	290.71(1)
	0.1	6.7241(1)	6.9365(1)	6.3994(1)	103.70(1)	289.97(1)
②	0.025	6.7380(1)	6.9520(1)	6.4078(1)	103.69(1)	291.63(1)
	0.05	6.7356(1)	6.9489(1)	6.4084(1)	103.69(1)	291.42(1)
	0.075	6.7327(1)	6.9453(1)	6.4079(1)	103.71(1)	291.10(1)
	0.1	6.7287(1)	6.9410(1)	6.4053(1)	103.72(1)	290.61(1)



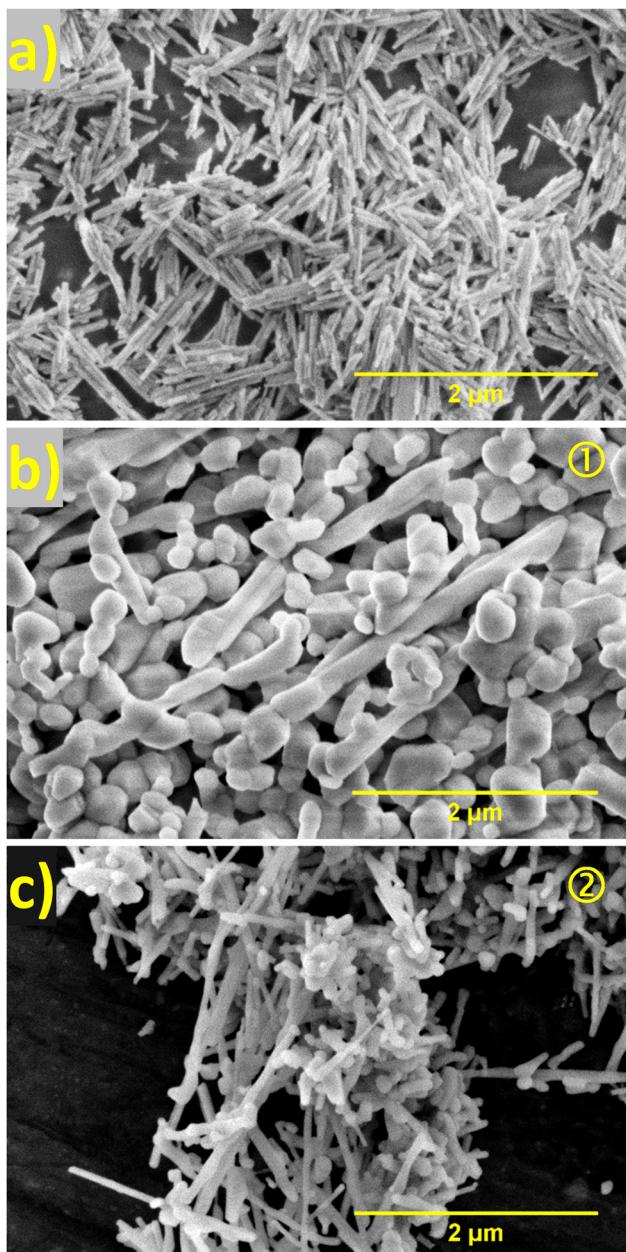


Fig. 2 SEM images recorded in SE mode for $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{U}_{0.1}\text{PO}_4\cdot n\text{H}_2\text{O}$ sample (a) and the corresponding monazite obtained after conversion in air (b) or in Ar atmosphere (c).

endothermic peak occurred at a very low temperature (*i.e.* below 100 °C) and was attributed to the formation of the hemihydrate form $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{U}_{0.1}\text{PO}_4\cdot 0.5\text{H}_2\text{O}$,³⁴ while the second deep endothermic peak (around 200 °C) corresponds to the complete dehydration of the rhabdophane structure leading to the hexagonal form.³⁵ The measured water content was found in the range of 0.64 to 0.79 H₂O molecule per formula unit in the samples prepared (Table 3). It was found to be close to the amount of structural water proposed by Mesbah *et al.*³³ (*i.e.* 0.667 H₂O per formula unit) and in the range usually observed for Th-rhabdophane or LnPO₄·nH₂O rhabdophanes due to the

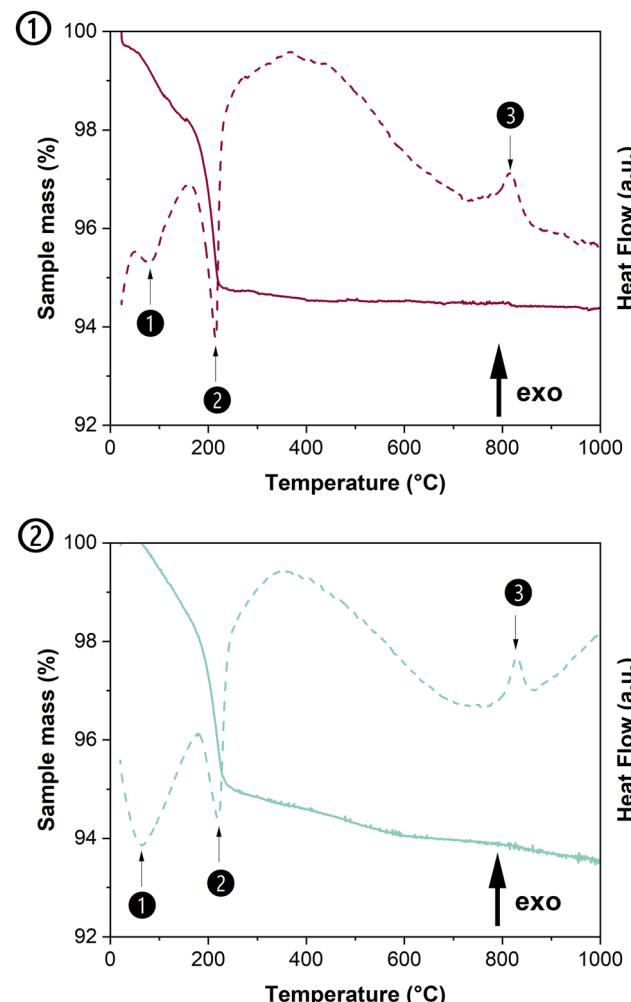


Fig. 3 TGA and DTA scans obtained when heating $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{U}_{0.1}\text{PO}_4\cdot n\text{H}_2\text{O}$ sample in air (1) or inert conditions (Ar, 2).

adsorption of variable amounts of water molecules onto the surface of the samples or to their incorporation in the zeolitic channels.^{29,31,36} The exothermic peak observed between 690 and 780 °C corresponded to the irreversible phase transition from U-rhabdophane to U-monazite–cheralite. The temperature of full dehydration of rhabdophane was determined from the derivative of the TGA curve, whereas the temperature of phase transition was that of the maximum of the DTA exothermic peak. The temperatures of dehydration and phase transition obtained for all the prepared samples are listed in Table 3 and compared to the data reported in the literature. From these results, it appeared that the incorporation of uranium and calcium into the structure had low impact on the dehydration temperature of rhabdophanes.

To investigate the influence of uranium incorporation on the conversion temperature of the phase transition from U-rhabdophane to U-monazite–cheralite, the data obtained by DTA were plotted as a function of the average ionic radius in Fig. S2 of the ESI.† The effective ionic radii used to calculate the average ionic radius are reported in Table S1 of the ESI.†

Table 3 Dehydration and transition temperatures obtained from TGA/DTA experiments for $\text{Nd}_{1-2x}\text{Ca}_x\text{U}_x\text{PO}_4\cdot n\text{H}_2\text{O}$ solid solutions heated under air ① or Ar ②

<i>X</i>	Dehydration temperature (°C)	Conversion temperature (°C)	H_2O per formula unit (total amount)	Ref.	
① 0	244	784	0.8–1.1	35	
	230 ± 5	700 ± 50		37	
	210 ± 3	693 ± 1	0.68 ± 0.05	29	
	0.025	223 ± 5	0.79 ± 0.05	This work	
	0.050	221 ± 5	0.64 ± 0.05	This work	
	0.075	217 ± 5	0.67 ± 0.05	This work	
	0.100	214 ± 3	0.79 ± 0.05	29	
②	213 ± 5	817 ± 5	0.72 ± 0.05	This work	
	0	230 ± 5	700 ± 50	37	
	0.025	214 ± 5	805 ± 5	0.77 ± 0.05	This work
	0.050	217 ± 5	837 ± 5	0.75 ± 0.05	This work
	0.075	216 ± 5	832 ± 5	0.65 ± 0.05	This work
	0.100	221 ± 5	828 ± 5	0.65 ± 0.05	This work

The temperature of phase transition from U-rhabdophane to U-monazite–cheralite solid solution was found to be higher than the temperature of conversion of Th-rhabdophane of equivalent average ionic radius obtained by Qin *et al.*²⁹ and closer to the data obtained by Kijkowska³⁵ for $\text{LnPO}_4\cdot n\text{H}_2\text{O}$ (with $\text{Ln} = \text{La}$ to Gd). This observation led to the conclusion that the thermal field of existence of U-rhabdophanes was larger than that obtained for Th-rhabdophanes. This result suggested that the respective thermodynamic properties of the U- and Th-rhabdophanes slightly differ. Kijkowska³⁵ prepared rhabdophane samples by crystallisation from boiling 2 mol L⁻¹ H_3PO_4 solution containing 0.02 mol L⁻¹ of lanthanide elements. The data obtained by Kijkowska³⁵ were also systematically higher than those measured by Jonasson and Vance³⁷ for the conversion of rhabdophane to monazite. Not only was the protocol of synthesis of the hydrated rhabdophanes different (refluxing $\text{Ln}_4(\text{P}_2\text{O}_7)_3$ in 0.1 mol L⁻¹ HNO_3 solution), but also Jonasson and Vance³⁷ did not specify the nature of the atmosphere used in their DTA experiments. This comparison showed that several factors could affect the temperature of transition, although there is a general decrease of the conversion temperature with the increase of the average ionic radius of the cations.

In order to validate the final objective of this study, $\text{Nd}_{1-2x}\text{Ca}_x\text{U}_x\text{PO}_4\cdot n\text{H}_2\text{O}$ rhabdophanes were heated for 6 hours under oxidizing (air) or inert (Ar) atmosphere at 1100 °C. This temperature was chosen to ensure the complete conversion of U-rhabdophane into U-monazite–cheralite (based on the results of TG analyses), allow the complete decomposition of the $\alpha\text{-UP}_2\text{O}_7$ reaction intermediate¹⁵ and avoid the decomposition of phosphates and the formation of CaO , UO_2 and $\text{P}_4\text{O}_{10}(\text{g})$ that occurred at 1300 °C.¹⁵

The thermally fired samples were first observed by SEM (Fig. 2b and c). After thermal conversion in air, the morphology of the crystals evolved markedly, whereas the changes were less pronounced under Ar atmosphere. The width of the needle-like crystals increased (about 6 times wider) and their

progressive rounding was observed. The observation of both samples did not reveal the presence of secondary phases exhibiting a different morphology.

The powders obtained after conversion in air or under Ar atmosphere were then analyzed by PXRD (Fig. 4). According to the PXRD diagrams of the samples converted in air, the obtained powders were all single-phase monazite–cheralite, whereas the presence of two additional peaks (at 17.9° and 20.7°) was detected for $x \geq 0.05$ in the samples converted under Ar. These two peaks could be attributed to small amounts of $\alpha\text{-UP}_2\text{O}_7$ phase, which was observed as a reaction intermediate during the irreversible phase transition from rhabdophane to monazite–cheralite by Bregiroux *et al.*¹⁵ This phase was also frequently formed as an intermediate in the synthesis of uranium phosphates such as $\text{U}(\text{UO}_2)(\text{PO}_4)_2$ or $\text{U}_2\text{O}(\text{PO}_4)_2$.^{38,39} However, under Ar atmosphere, Bregiroux *et al.*¹⁵ observed the complete disappearance of $\alpha\text{-UP}_2\text{O}_7$ between 900 and 1000 °C and prepared monophasic $\text{Ca}_{0.5}\text{U}_{0.5}\text{PO}_4$ cheralite above 1100 °C. Several explanations can be suggested to explain the presence of $\alpha\text{-UP}_2\text{O}_7$ in our prepared samples. First, part of the uranium could segregate into $\alpha\text{-UP}_2\text{O}_7$ phase during the thermal treatment of the rhabdophane precursor due to the lack of Ca to fully compensate the charge in the monazite–cheralite structure. Second, part of the uranium might have precipitated in a phase other than rhabdophane during the hydrothermal synthesis. This secondary phase would be converted into $\alpha\text{-UP}_2\text{O}_7$ during the thermal treatment under Ar and not completely decompose after the heat treatment at 1100 °C. However, the presence of $\alpha\text{-UP}_2\text{O}_7$ phase was not evidenced after conversion in air, which indicates different mechanisms of phase transition, probably involving different reaction intermediates or different behaviour of $\alpha\text{-UP}_2\text{O}_7$ in air due to U(IV) oxidation.³⁹

The PXRD diagrams of the obtained monazite–cheralite were refined using the Rietveld method (Fig. S3 of the ESI†). The unit cell parameters are listed in Table 2. Fig. 5 shows the variation of the unit cell volume of the U-monazite–cheralite solid solutions *versus* the target uranium mole fraction x . (Fig. S4 of the ESI† shows the difference between the experimental cell volume obtained by Rietveld refinement of the diffractograms and that calculated from the proportional combination of the end-members.) The unit cell volumes of the U-monazite–cheralite samples converted either under air or Ar atmosphere decreased linearly with the incorporation rate. When comparing the unit cell volumes of the prepared monazite–cheralite with $x \leq 0.1$ with the data reported in the literature for the end-members, a slight deviation from Retger's law was evidenced. Nevertheless, considering the low incorporation rates of uranium in the prepared samples, it is difficult to draw conclusions on the ideality of the solid solution. Moreover, considering the presence of the additional $\alpha\text{-UP}_2\text{O}_7$ phase, the real uranium mole fraction in the monazite–cheralite phase converted under Ar is expected to be lower than the target x value and so does the deviation from Retger's law.

The cell volumes obtained by refinement of the PXRD diagrams obtained using a laboratory X-ray source revealed that uranium was well incorporated into the structure, since the



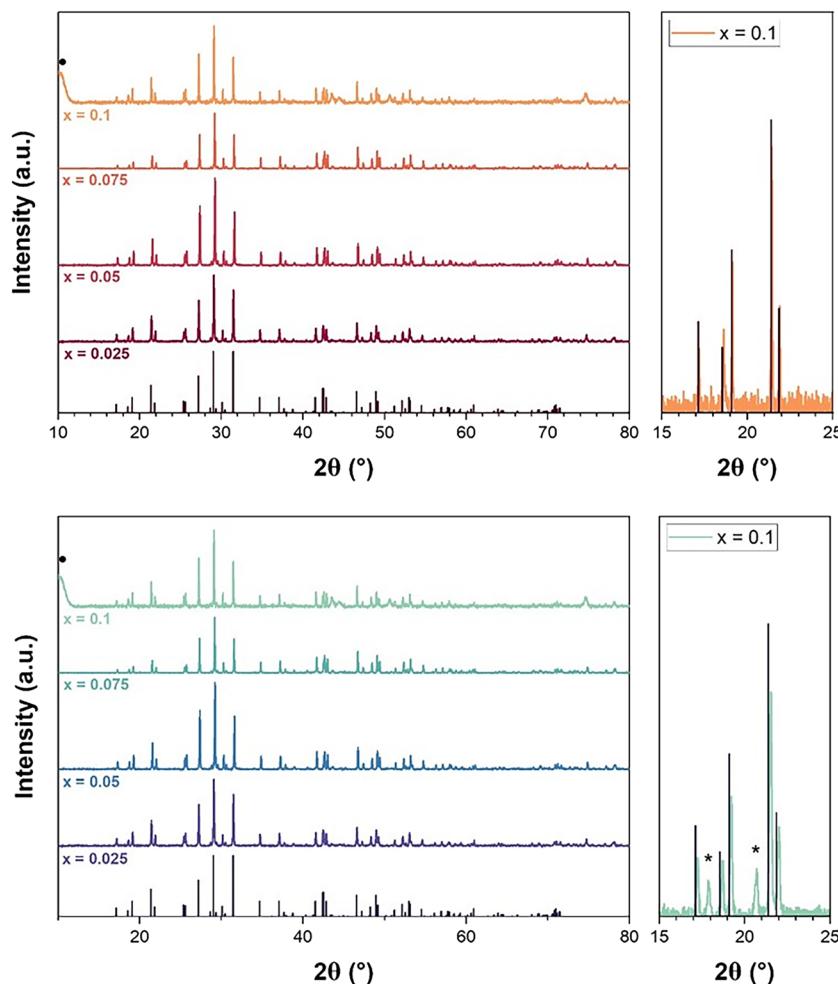


Fig. 4 PXRD patterns obtained after thermal conversion of U-rhabdophanes to U-monazite–cheralites at 1100 °C for 6 h in air (a) or under argon atmosphere (b). XRD lines associated with monazite NdPO₄ are indicated by black vertical bars,⁴⁰ while black dot and stars indicate sample holder and additional peaks attributed to α-UP₂O₇ secondary phase, respectively.

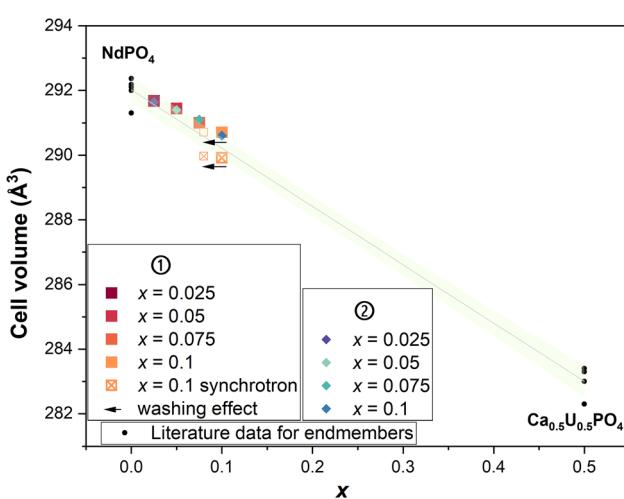


Fig. 5 Refined unit cell volumes of U-monazite–cheralite solid solutions $\text{Nd}_{1-2x}\text{Ca}_x\text{U}_x\text{PO}_4$ prepared in air ① or under argon atmosphere ②. Black dots correspond to the unit cell volumes of the endmembers reported in ref. 15, 21, 36, 41–46. Continuous line corresponds to the fitted Retger's law, whereas green area corresponds to the 95% confidence level band.

unit cell volume decreased when the incorporation rate increased. Nevertheless, a difference can be observed between the diffractogram obtained in the laboratory and that recorded on the MARS beamline of the SOLEIL synchrotron for the $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{U}_{0.1}\text{PO}_4$ sample after heat treatment in air. This difference could be due to sample preparation (plate *vs.* capillary), the two Kapton confinements or the possible evolution of the sample under the beam.

Raman and FTIR spectroscopies

The monazite–cheralite samples obtained by thermal conversion in air and Ar atmosphere were also characterized by Raman and FTIR spectroscopies. The main objective was to check that U(vi) (in the uranyl form) was not present in the prepared samples.⁴⁷ The Raman and FTIR spectra of monazite–cheralites converted in air or under Ar-atmosphere are presented in Fig. 6.

The Raman spectra (Fig. 6a) of the $\text{Nd}_{1-2x}\text{Ca}_x\text{U}_x\text{PO}_4$ samples (with $x \leq 0.1$) presented the deformation modes of

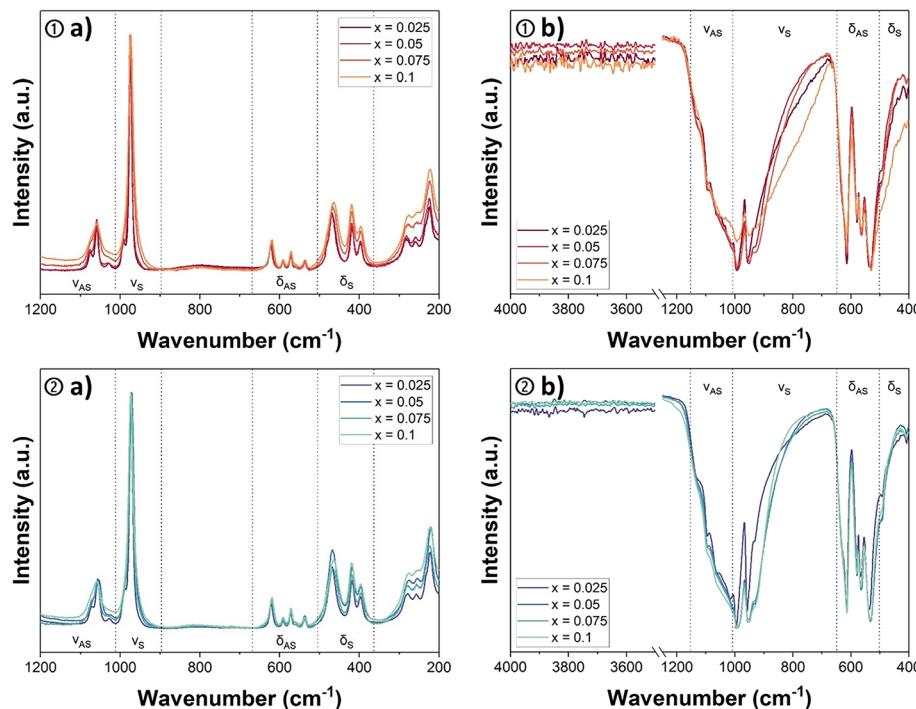


Fig. 6 Raman spectra recorded of U-monazite–cheralite solid solutions $\text{Nd}_{1-2x}\text{Ca}_x\text{U}_x\text{PO}_4$ prepared by conversion in air ① or under argon atmosphere ② in the $200\text{--}1200\text{ cm}^{-1}$ range (a) and associated FTIR spectra in the $400\text{--}4000\text{ cm}^{-1}$ range (b).

the PO_4 tetrahedron in the range between 300 and 800 cm^{-1} and the characteristic bands of the stretching modes in the $900\text{--}1200\text{ cm}^{-1}$ range. The positions of the different bands and their assignments, based on the work of Clavier *et al.*,⁴⁸ are listed in Table 4. The comparison of the Raman spectra of the monazite–cheralites obtained by thermal treatment under air or Ar atmosphere did not reveal any particular differences. The positions of the vibration bands were found to be very close. For instance, the strongest band, assigned to the symmetric stretching vibration of the P–O bonds in phosphate groups $\nu_{\text{S}}(\text{P–O})$, was located in the narrow wavenumber range between 972 and 976 cm^{-1} irrespective of the atmosphere of conversion. The absence of an intense and narrow vibration band in the domain of $820\text{--}880\text{ cm}^{-1}$, characteristic of the symmetric stretching vibration ν_{S} (frequently labelled ν_1 in the literature) of the uranyl bond ($\text{U}=\text{O}$),^{49,50} suggested the absence of UO_2^{2+} entities and the stabilization of tetravalent uranium during the thermal conversion, even in air.

The FTIR spectra of the monazite–cheralite samples converted under air or Ar did not differ strongly (Fig. 6b). The sharp band observed between 950 and 957 cm^{-1} was assigned to the $\nu_{\text{S}}(\text{P–O})$ symmetrical stretching mode. The position of this band was not significantly influenced by the uranium content or the atmosphere of conversion. The infrared active antisymmetric stretching vibration of UO_2^{2+} , ν_{AS} (also called ν_3), is usually observed in the range of 910 to 962 cm^{-1} .^{49–51} The potential overlap of the $\nu_{\text{S}}(\text{P–O})$ and $\nu_{\text{AS}}(\text{UO}_2^{2+})$ characteristic bands in infrared spectroscopy could lead to misinterpretation in the case of uranyl phosphates. However, the absence

of a strong active ν_{S} (often referred to as ν_1) band in the Raman spectra and the similarity between the spectra of samples converted under oxidizing or inert atmosphere led to the conclusion that the uranyl moiety was not present in the monazite–cheralite samples prepared or that the quantity was below the detection limit of this technique. Infrared spectroscopy also shows the absence of vibrations associated with symmetric and anti-symmetric elongation of the O–H bond of water molecules (normally located between 3650 and 3750 cm^{-1}). This confirms that the solids characterized are totally dehydrated. Additionally, the symmetric and anti-symmetric stretching vibrations of the (P–O–P) bond observable at 737 and 955 cm^{-1} , respectively, in the FTIR spectra of $\alpha\text{-UP}_2\text{O}_7$ were not detected in the samples calcined under argon,⁵² although this phase was identified by XRD. The absence of these bands in the FTIR spectra presented in Fig. 6b ② shows that the amount of $\alpha\text{-UP}_2\text{O}_7$ was below the detection limit of this technique.

Synchrotron analyses

HERFD-XANES analysis of the $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{U}_{0.1}\text{PO}_4$ sample prepared by thermal conversion of U-rhabdophane at $1100\text{ }^{\circ}\text{C}$ for 6 hours in air was performed at the M_4 edge of uranium. Thanks to its high spectral resolution, this technique allowed the detection of the presence of several oxidation states of uranium in the monazite–cheralite sample. Fig. 7 shows the measured spectra of the two references as well as that of $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{U}_{0.1}\text{PO}_4$. Metatorbernite, $\text{Cu}(\text{UO}_2)_2(\text{PO}_4)_2\cdot n\text{H}_2\text{O}$, was chosen as a pure UO_2^{2+} bearing phosphate reference whose



Table 4 Infrared and Raman spectra recorded for the prepared U-monazite–cheralites $\text{Nd}_{1-2x}\text{Ca}_x\text{U}_x\text{PO}_4$ prepared by thermal conversion of U-rhabdophane in air ^① or under Ar-atmosphere ^② and data reported for $\text{GdPO}_4\text{·}n\text{H}_2\text{O}$ as a reference material⁴⁸

	Raman spectroscopy (cm^{-1})			FTIR spectroscopy (cm^{-1})			
	$\delta_{\text{S}}(\text{P-O})$	$\delta_{\text{AS}}(\text{P-O})$	$\nu_{\text{S}}(\text{P-O})$	$\nu_{\text{AS}}(\text{P-O})$	$\delta_{\text{S}}(\text{P-O})$	$\delta_{\text{AS}}(\text{P-O})$	
Ref. 48	404, 428, 476	539 (567), 577, 598, 632	987	1004, 1042, 1071, 1092	492	538, 570, 585, 623 (642)	
①	$x = 0.025$ $x = 0.05$ $x = 0.075$ $x = 0.100$	398, 419, 469 398, 419, 467 398, 419, 467 396, 419, 464	538, 572, 592, 622 537, 573, 592, 622 536, 572, 591, 621 536, 571, 591, 619	974 975 975 974	992, 1029, 1059, 1075 1030, 1060, 1076 1059 1060	493 493 493 492	933, 960, 995 928, 957, 992 930, 951, 990 928, 950, 991 928, 951, 993
②	$x = 0.025$ $x = 0.050$ $x = 0.075$ $x = 0.100$	397, 419, 469 396, 418, 468 397, 418, 466 394, 419, 467	537, 572, 592, 622 537, 572, 592, 621 537, 573, 590, 621 536, 572, 591, 622	973 972 973 976	990, 1028, 1057, 1072 1055 1058 1064	491 493 493 491	537, 567, 579, 615 532, 565, 579, 615 533, 565, 580, 615 532, 567, 580, 614

associated HERFD-XANES spectrum exhibited three distinct spectral features characteristic of UO_2^{2+} compounds. The three main fluorescence peaks were assigned to the transitions of $3\text{d}_{3/2}$ electrons to the $5\text{f}_{\delta/\phi}$, 5f_{π^*} and 5f_{σ^*} valence orbitals.^{53,54} These peaks appeared at 3728.3, 3730.4, and 3734.3 eV, respectively, in the metatorbernite spectrum. The second reference was a pure U(iv) bearing phosphate phase, diuranium oxophosphate, $(\text{U}_2\text{O})(\text{PO}_4)_2$. This spectrum presented only one large peak at 3726.9 eV assigned to U(iv). Due to different local environments, the position of the U(iv) peak in the $(\text{U}_2\text{O})(\text{PO}_4)_2$ reference was shifted towards higher energy compared to that of UO_2 (reported at 3723 eV by Massonnet *et al.*)⁵⁵. In order to determine the relative proportion of U(iv) and U(vi) oxidation states within the monazite–cheralite converted in air, the HERFD-XANES spectrum of the $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{U}_{0.1}\text{PO}_4$ sample was fitted by a linear combination of the two reference spectra. Fig. 7 shows the calculated spectrum and the difference from the measured spectrum. Even if the quality of the fit was disputable due to variations of the local environment of uranium in the monazite–cheralite compared to the two references, it gave a semi-quantitative indication of the proportion of U(iv) and U(vi) in the monazite–cheralite sample, with obtained contents of $84 \pm 2\%$ and $16 \pm 2\%$, respectively.

Since the uranium content of the prepared sample was equal to 10 wt%, the mass fraction of U(vi) in the monazite–cheralite was less than 2 wt%. This low amount might explain the absence of the bands characteristic of the vibrations of UO_2^{2+} entities in the FTIR and Raman spectra of this sample (*a fortiori*, in the other samples with lower uranium contents). It is noteworthy that the incorporation of uranyl in the monazite–cheralite structure is very unlikely. Thus, the presence of a minor secondary phase was considered. As the PXRD pattern obtained using a laboratory X-ray source did not evidence the presence of a secondary phase, synchrotron X-ray diffraction analysis of the $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{U}_{0.1}\text{PO}_4$ sample was performed after its thermal conversion in air. Fig. 8 presents the SPXRD diagram of this sample obtained with the CX3 station of the MARS beamline at the SOLEIL synchrotron.

The SPXRD pattern recorded at the MARS beamline was also refined by the Rietveld method (Fig. S3†). The obtained unit cell parameters of the monazite–cheralite are reported in Table 2 and Fig. 5. The unit cell volume was found to be lower than the one obtained by refinement of the laboratory PXRD pattern. However, thanks to the high sensitivity of the experimental device, several small additional peaks were detected between 8° and 17° 2θ (zoom in Fig. 8). The nature of the U(vi) secondary phase was not identified, although the HERFD-XANES spectrum confirmed that it likely incorporated the uranyl moiety rather than uranate.⁵⁶ This phase could have been formed by precipitation during the hydrothermal synthesis of the U-rhabdophane precursor or during the thermal treatment of conversion under air.

In order to evaluate the composition of this secondary phase, 100 mg of the $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{U}_{0.1}\text{PO}_4$ sample was washed for 5 days in 100 mL of a 0.1 mol L^{-1} HNO_3 solution at 40°C . The washed sample was also characterized at the MARS beamline.



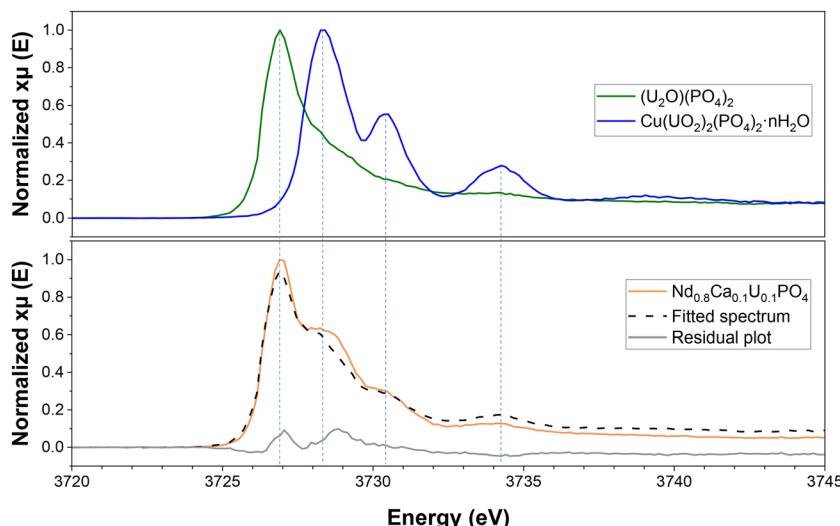


Fig. 7 U M₄-edge HERFD-XANES spectra of Nd_{0.8}Ca_{0.1}U_{0.1}PO₄ converted under air. Reference spectra of U(IV) and U(VI) phosphates were obtained using (U₂O)(PO₄)₂ and Cu(UO₂)₂(PO₄)₂·nH₂O, respectively. Dashed black line corresponds to the spectrum of Nd_{0.8}Ca_{0.1}U_{0.1}PO₄ fitted by linear combination of the two references.

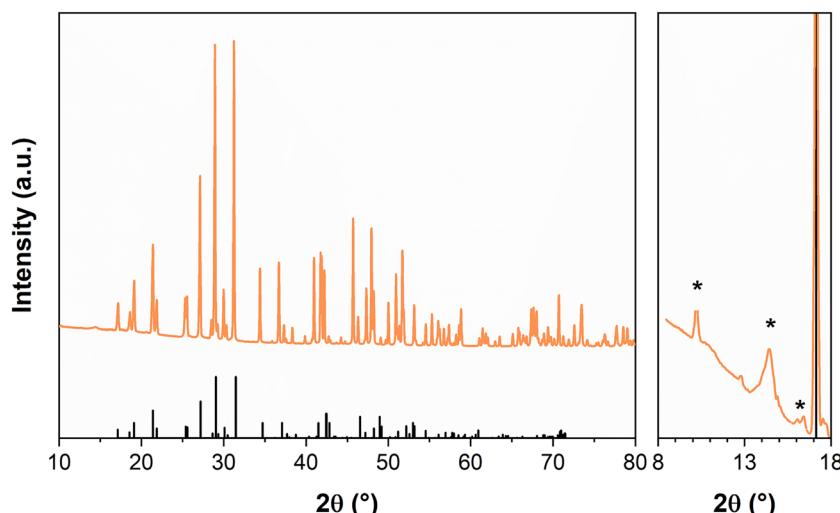


Fig. 8 SPXRD pattern obtained for Nd_{0.8}Ca_{0.1}U_{0.1}PO₄ monazite–cheralite converted in air at 1100 °C for 6 hours. PXRD patterns reported for monazite NdPO₄ (black vertical bars);⁴⁰ * indicates additional peaks observed in the experimental pattern.

The resulting XRD diagram showed the disappearance of the additional peaks attributed to the presence of a secondary phase (Fig. S5 of the ESI†). Additionally, the U M₄-edge HERFD-XANES spectrum of the washed Nd_{0.8}Ca_{0.1}U_{0.1}PO₄ sample was also recorded at the MARS beamline (Fig. S6 of the ESI†). It showed the complete disappearance of the U(VI) component. The Ca, U, Nd and P elemental concentrations in the washing solution were determined by ICP-AES analysis. The results indicated that 21.3 mol% of the uranium initially present in the solid phase was dissolved in 5 days, in addition to 2.4 mol% of phosphorus and 4.8 mol% of Ca. The mole fraction of Nd released in the solution was negligible, indicating that the monazite–cheralite phase was virtually not dis-

solved during the washing step. Assuming that the released elements came only from a single additional phase, its stoichiometry would be approximately Ca:P = 0.15 and U:P = 0.78. Alternatively, Ca and U phosphates could have precipitated in two separate phases of unknown stoichiometry. Nevertheless, these results demonstrated that the U(VI) contained in the sample before washing was completely incorporated as a secondary phase. This more soluble uranyl phosphate phase was removed by the washing protocol. 21 mol% of the uranium precipitated initially with the rhabdophane precursor was incorporated in the secondary phase, which was evidenced by synchrotron measurements only. Thus, the limit of incorporation of tetravalent uranium in the monazite–chera-

lite structure through the wet chemistry route described in this work reached $x = 0.08$ rather than 0.1, as initially stated. Additionally, using the real value of x in the washed monazite–cheralite brings the cell volume calculated by Rietveld refinement closer to the value estimated using Retger's law (Fig. 5). This observation indicates that the main reason for deviation from Retger's law is an overestimation of the uranium mole fraction.

Experimental

Synthesis

To avoid the oxidation of tetravalent U by the introduction of nitrate ions in the system, analytical grade metal chloride salts and H_3PO_4 (ACS Reagent, ≥ 85 wt%, Sigma Aldrich) and HCl (ACS Reagent, 37%, Carlo Erba) solutions were used for the preparation of the monazite–cheralite hydrated precursor of the rhabdophane structure $\text{Nd}_{1-2x}\text{Ca}_x\text{U}_x\text{PO}_4 \cdot n\text{H}_2\text{O}$ (with $x \leq 0.1$). $\text{NdCl}_3 \cdot 6\text{H}_2\text{O}$ and $\text{CaCl}_2 \cdot n\text{H}_2\text{O}$ (Sigma Aldrich) were first dissolved in 0.1 mol L^{-1} hydrochloric acid to obtain a cation concentration ranging from 1 to 1.5 mol L^{-1} . The uranium solution was obtained by dissolving uranium metal chips (kindly provided by CETAMA) in hydrochloric acid.⁵⁷ In this very acidic hydrochloric medium, U(IV) is stabilized against oxidation by the formation of UCl_6^{2-} in the starting mixture and then by complexation with phosphates during the further mixture with phosphoric acid. The metal chips were washed with dichloromethane, acetone and ethanol to remove the traces of oil required for their long-term storage. Then, the dried uranium metal chips were weighed and dissolved in 6 mol L^{-1} HCl solution cooled in an ice bath to obtain a final elemental concentration of 0.6 mol L^{-1} . The accurate elemental concentrations of the solutions containing each cation were then measured using inductively coupled plasma optical emission spectroscopy (ICP-OES, Thermo Fisher Scientific, iCAP 7000 series).

The hydrothermal synthesis protocol of the Th-rhabdophane developed by Qin *et al.*^{28,29} was adapted to allow the incorporation of uranium. Specifically, Qin *et al.*²⁸ showed that a Ca:Th molar ratio of 10 was mandatory to obtain single-phase $\text{Nd}_{1-2x}\text{Ca}_x\text{Th}_x\text{PO}_4 \cdot n\text{H}_2\text{O}$ samples. Thus, the synthesis method consists of mixing adequate volumes of each solution in a beaker under magnetic stirring in order to obtain a mixture with the stoichiometrically expected proportions of Nd and U in the solid phase. The Ca:U stoichiometric ratio varied with the U mole fraction of the sample, x , so that $\text{Ca}/\text{U} = 100x$. Once the solutions were mixed, 5 mol L^{-1} H_3PO_4 solution was added dropwise to the previous batch with a 3 mol% excess with regard to Nd, Ca and U. The mixture was stirred for 15 minutes before being introduced into the autoclave (Parr 4744). In order to guarantee an identical autogenous pressure for all the syntheses, deionized water was added to the mixture to obtain a final volume of 25 mL and a constant dead volume of 20 mL. The autoclave was then sealed and placed in an oven at 110 °C for four days. The precipitate obtained was washed

twice with water and once with ethanol through cycles of centrifugation at 14 500 rpm for 5 minutes. The powder was then dried in an oven at 90 °C overnight.

Finally, the monazite–cheralite hydrated precursor underwent thermal treatment under air or under Ar at 1100 °C for 6 hours in a muffle furnace or a tubular furnace, respectively.

ICP-OES

The chemical composition of the synthesized rhabdophanes was determined by ICP-OES (Thermo Fisher Scientific, iCAP 7000 series) measurements after their complete dissolution. For this purpose, about 20 mg of powder was weighed, then dissolved in fresh aqua regia. After complete dissolution, the solutions were diluted in 0.2 mol L^{-1} HNO_3 solution (ACS Reagent, 70%, Sigma-Aldrich) and analyzed by ICP-OES. The ICP-OES was initially calibrated using several standard solutions prepared by the dilution of Th, U, Ca, Nd and P elemental solutions certified at 1000 ppm (SCP Science) in 0.2 mol L^{-1} HNO_3 solution. 3 to 4 of the recommended emission wavelengths were selected for each element. Three replicates for each measurement were performed. The stoichiometry of each compound was then calculated using the average elemental concentrations.

PXRD/SPXRD

Powder X-ray diffraction patterns (PXRD) of the synthesized compounds were obtained using a Bruker D8 Advance diffractometer equipped with a Cu $\text{K}\alpha_{1,2}$ X-ray source ($\lambda = 1.5418 \text{ \AA}$), a 60 mm sealed Göbel mirror and a LynxEye linear position sensitive detector in θ/θ step scan mode. The patterns were acquired at room temperature in the range of 5 to 100° (2θ) with a step size of 0.02°. The counting time for one sample was about 3 hours.

Structural analyses of selected samples were also performed using the SPXRD (synchrotron PXRD) set-up of the CX3 end-station of the MARS beamline at the SOLEIL synchrotron (Saint-Aubin, France).⁵⁸ About 20 mg of each sample was placed in 0.6 mm Kapton capillary tubes (3 cm long). The capillary was inserted in a second Kapton capillary for the confinement of the radioactive powder. Then, the powder was sealed in the tube using epoxy resin. XRD patterns were recorded at 17 keV in the transmission Debye–Scherrer geometry. The monochromatic beam was sized at $300 \times 150 \mu\text{m}^2$ ($H \times V$). The 2D hybrid pixel detector PILATUS3 2M CdTe (Dectris AG, Switzerland) was located at 680 mm from the sample on a 2θ motorized rotation axis. 11 images were recorded using a 5° 2θ step between 0 and 50°. 60 s were needed to record one image and the total signal was thus obtained in about 6 min. The 11 images were merged and integrated using the multi-geometry class of the pyFAI library⁵⁹ to reduce and reconstruct 1D powder diffraction patterns. The setup's geometry correction and instrumental resolution function were done using a NIST SRM LaB₆ 660b powder (lattice parameter: 4.15689 \AA). During the pattern recording, the samples were continuously spun around the capillary axis after mounting and alignment on a goniometer head.



For both the synchrotron and laboratory X-ray sources, PXRD patterns were refined by the Rietveld method using the Fullprof_Suite package.⁶⁰ The rhabdophane precursors as well as the monazite–cheralites were refined in the monoclinic $C2$ ³³ and $P2_1/n$ ⁶¹ space groups, respectively. In both cases, the instrumental function was extracted using Si and LaB₆, respectively. During refinement, the profile parameters, such as the zero-point shift, unit cell parameters, and asymmetry, were refined and a size model was introduced in order to evaluate the microstructural effect. When needed, the atomic positions and thermal displacement parameters were also refined.

SEM

The morphology of the prepared compounds was investigated by scanning electron microscopy using an FEI QUANTA 200 ESEM FEG. For this purpose, a secondary electron detector (SE) and/or a backscattered electron detector (BSE) were used at room temperature with an accelerating voltage of 2 kV in order to obtain high-resolution images. In terms of sample preparation, the powder was deposited on a polished SEM pad, then dispersed with ethanol in order to deagglomerate the powder as much as possible. No further preparation, such as metallization, was carried out before analysis.

TGA/DTA

Thermogravimetric (TGA) and thermodifferential (TDA) analyses were carried out using a SETARAM Setsys Evolution apparatus equipped with a Pt thermocouple. For this purpose, 10 to 15 mg of powder was weighed, then introduced into an alumina crucible. During the analysis, the sample was heated to 1000 °C at a heating rate of 5 °C min⁻¹ and a 20 mL min⁻¹ flow rate of air or Ar. The TGA curves allowed the determination of the water content of the samples, whereas the DTA curves showed the temperature of transition from rhabdophane to monazite–cheralite.

Vibrational spectroscopies (Raman/FTIR)

Fourier transform infrared spectroscopy in attenuated total reflection mode (ATR-FTIR) was performed using a PerkinElmer Spectrum 100 spectrometer operating with a Globar MIR source in combination with a cesium iodide (CsI) beamsplitter and a deuterated triglycine sulphate DTGS detector through CsI transmission windows. To carry out these analyses, the powdered samples were deposited at the surface of a 2 mm diameter top-plate diamond crystal stacked on a KRS-5 (thallium bromide-iodide) bottom-plate microfocus condenser without preparation. High pressure was exerted on the sample and adjusted until stabilisation of the signal. Spectra were recorded from 380 to 4000 cm⁻¹ adding 4 scans and operating with background correction. The nominal resolution of the spectra was 4 cm⁻¹.

The Raman spectra were recorded with a Horiba-Jobin Yvon Labram Aramis equipped with an edge filter and a Nd:YAG laser (532 nm). Before starting the analyses, the instrument was calibrated with silicon, using the first order Si line at 520.7 cm⁻¹. Once the calibration was completed and the

sample was deposited on a glass lamella, the laser was focused on the sample using a confocal microscope with, in optimal conditions, an axial resolution lower than 2 μm and a spot diameter smaller than 1 μm laterally. For the acquisition of spectra, this microscope was equipped with a 532 nm excitation laser (*i.e.* about 60 mW), a $\times 100$ objective and no filter. An accumulation of 5 spectra per sample was carried out for a total analysis time of approximately 5 minutes. Data processing, such as spectral normalization, was carried out using ORIGIN software.

HERFD-XANES

High-energy resolution fluorescence detected XANES (HERFD-XANES) data were measured at the MARS beamline of the SOLEIL synchrotron (Saint-Aubin, France).⁵⁸ The storage ring was operated in the top-up mode at an electron current of 500 mA, 2.5 GeV. The beam size at the sample surface was 300 μm × 150 μm ($H \times V$). Spectra were measured at the U M₄-edge (3728 eV) using the double-crystal monochromator (DCM) equipped with a pair of Si(111) crystals. Higher harmonics rejection and vertical focusing were achieved using the Si strip of each mirror inserted before and after the DCM with a 4 mrad incidence angle. The incident energy was calibrated using the absorption K-edge of potassium from a KBr pellet (3608.4 eV). HERFD-XANES was performed using the crystal-analyzer emission spectrometer in the Rowland geometry and a KETEK single-element silicon solid-state detector. The M_β emission line of U (3339 eV) was analyzed using the 220 reflection of a Si(220) bent, stripped crystal analyzer with a curvature radius of 1 m. The energy resolution of the emission spectrometer was 0.8 eV at 6.676 keV. The samples were oriented at 45° with respect to the incident beam. A He-filled balloon was used to reduce the scattering of the incident and emitted X-rays by the air between the sample, the crystal analyzer and the detector. Samples were prepared as pellets for each compound. All the spectra collected were normalized at the maximum of absorption using Athena software,⁶² then fitted as a linear combination of different references to quantify the contributions of U(iv), U(v), and U(vi). Reference spectra were obtained using (U₂O)(PO₄)₂^{63–66} as the U(iv) standard and Cu(UO₂)₂(PO₄)₂·*n*H₂O⁶⁷ as the U(vi) standard in a phosphate environment.

Conclusion

In this work, Nd_{1–2x}Ca_xU_xPO₄·*n*H₂O (with $x \leq 0.1$) rhabdophane-type precursors were prepared by a wet chemistry route followed by hydrothermal treatment at 110 °C for four days. A large excess of Ca in the initial mixture of cations was necessary to obtain single-phased rhabdophanes using this protocol. Nd_{1–2x}Ca_xU_xPO₄ monazite–cheralite solid solutions were then obtained by thermal conversion of the rhabdophane precursors at 1100 °C for 6 hours either under oxidising (air) or inert (Ar) atmosphere. This temperature was chosen based on TG analyses to guarantee complete phase transition and to avoid



the decomposition of the phosphate group and the formation of CaO , UO_2 and $\text{P}_4\text{O}_{10}(\text{g})$. The obtained monazite–cheralite samples were characterized using a wide set of analytical techniques. When heated in Ar atmosphere, PXRD revealed the presence of an $\alpha\text{-UP}_2\text{O}_7$ secondary phase in the monazite–cheralite samples, whereas the samples converted under air were single-phase. Special attention was paid to the oxidation state of uranium in the monazite–cheralite samples obtained by conversion of the rhabdophane precursors under air. FTIR and Raman spectroscopies did not reveal the presence of uranyl moieties in these samples. However, the U M_4 -edge HERFD-XANES spectrum of the $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{U}_{0.1}\text{PO}_4$ compound showed the coexistence of about 80 mol% U(IV) and 20 mol% U(VI) present as uranyl in the sample. U(VI) was associated with the presence of a minor secondary phase in the $\text{Nd}_{0.8}\text{Ca}_{0.1}\text{U}_{0.1}\text{PO}_4$ sample detected only by SPXRD. This soluble phase was easily removed by washing the sample in a 0.1 mol L^{-1} HNO_3 solution. Finally, the limit of incorporation of uranium in the $\text{Nd}_{1-2x}\text{Ca}_x\text{U}_x\text{PO}_4$ monazite–cheralite prepared by this wet chemistry route reached around $x = 0.08$.

The conversion mechanism of U-rhabdophane to monazite–cheralite under Ar atmosphere was very similar to that studied by Qin *et al.*²⁹ for Th-monazite–cheralites. It is therefore likely that the presence of $\alpha\text{-UP}_2\text{O}_7$ results from the presence of an excess of uranium (relative to Ca) associated with the presence of vacancies in the U-rhabdophane structure. In this case, the excess U segregates upon heat treatment to form $\alpha\text{-UP}_2\text{O}_7$. However, this phenomenon was not observed when heating in air, due to the lower stability of $\alpha\text{-UP}_2\text{O}_7$ under these oxidizing conditions.³⁹ The $\alpha\text{-UP}_2\text{O}_7$ would then be converted into a uranyl phosphate phase with $\text{U:P} = 0.78$ stoichiometry, which coexisted as a minor phase with the monazite–cheralite. This phase, whose presence could affect the long-term behavior of the ceramic in the field of immobilization of actinides, was easily removed by washing in dilute acid to obtain a pure monazite–cheralite phase containing tetravalent uranium only. Using this washing step will allow the study of the chemical durability of $\text{Nd}_{1-2x}\text{Ca}_x\text{U}_x\text{PO}_4$ monazite–cheralite (with $x = 0.08$) from kinetic and thermodynamic points of view in a forthcoming dedicated study.

Conflicts of interest

There are no conflicts to declare.

Acknowledgements

The authors are thankful to ICSM staff for their technical support, especially J. Lautru for his meticulous work with ESEM and EDS analyses, B. Baus-Lagarde for ICP-AES analyses and A. Jonchère for his help in using FTIR and Raman spectroscopies. The authors would also like to thank the SOLEIL synchrotron for providing beamtime, as well as the University of Montpellier for funding Alison El Monjid's PhD work.

References

- 1 L. A. Boatner, in *Phosphates: Geochemical, Geobiological, and Materials Importance*, ed. M. J. Kohn, J. Rakovan and J. M. Hughes, 2002, vol. 48, pp. 87–121.
- 2 C. M. Gramaccioli and T. V. Segalstad, *Am. Mineral.*, 1978, **63**, 757–761.
- 3 G. R. Lumpkin, *Can. Mineral.*, 1998, **36**, 339–353.
- 4 G. R. Lumpkin, Y. Gao, R. Giere, C. T. Williams, A. N. Mariano and T. Geisler, *Mineral. Mag.*, 2014, **78**, 1071–1095.
- 5 L. A. Boatner and B. C. Sales, *Monazite, Radioactive waste forms for the future*, North Holland, Amsterdam (Netherlands), 1988.
- 6 A. M. Seydoux-Guillaume, X. Deschanels, C. Baumier, S. Neumeier, W. J. Weber and S. Peuget, *Am. Mineral.*, 2018, **103**, 824–827.
- 7 J. M. Montel, Kornprobst and Vielzeuf, *J. Metamorph. Geol.*, 2000, **18**, 335–342.
- 8 N. Dacheux, N. Clavier and R. Podor, *Am. Mineral.*, 2013, **98**, 833–847.
- 9 L. A. Boatner, G. W. Beall, M. M. Abraham, C. B. Finch, P. G. Huray and M. Rappaz, *Sci. Basis Nucl. Waste Manage.*, 1980, **2**, 289–296.
- 10 R. C. Ewing and L. Wang, *Rev. Mineral. Geochem.*, 2002, **48**, 673–699.
- 11 B. Omel'Yanenko, T. Livshits, S. Yudintsev and B. Nikonov, *Geol. Ore Deposits*, 2007, **49**, 173–193.
- 12 N. Clavier, R. Podor and N. Dacheux, *J. Eur. Ceram. Soc.*, 2011, **31**, 941–976.
- 13 N. Dacheux, N. Clavier, A.-C. Robisson, O. Terra, F. Audubert, J.-É. Lartigue and C. Guy, *C. R. Chim.*, 2004, **7**, 1141–1152.
- 14 X. Deschanels, V. Picot, B. Glorieux, F. Jorion, S. Peuget, D. Roudil, C. Jégou, V. Broudic, J. Cachia and T. Advocat, *J. Nucl. Mater.*, 2006, **352**, 233–240.
- 15 D. Bregiroux, O. Terra, F. Audubert, N. Dacheux, V. Serin, R. Podor and D. Bernache-Assollant, *Inorg. Chem.*, 2007, **46**, 10372–10382.
- 16 P. Raison, R. Jardin, D. Bouexiere, R. Konings, T. Geisler, C. Pavel, J. Rebizant and K. Popa, *Phys. Chem. Miner.*, 2008, **35**, 603–609.
- 17 A. Tabuteau, M. Pagès, J. Livet and C. Musikas, *J. Mater. Sci. Lett.*, 1988, **7**, 1315–1317.
- 18 R. Podor and M. Cuney, *Am. Mineral.*, 1997, **82**, 765–771.
- 19 R. D. Shannon, *Acta Crystallogr., Sect. A: Cryst. Phys., Theor. Gen. Crystallogr.*, 1976, **32**, 751–767.
- 20 D. Bregiroux, R. Belin, P. Valenza, F. Audubert and D. Bernache-Assollant, *J. Nucl. Mater.*, 2007, **366**, 52–57.
- 21 O. Terra, N. Dacheux, N. Clavier, R. Podor and F. Audubert, *J. Am. Ceram. Soc.*, 2008, **91**, 3673–3682.
- 22 O. Terra, N. Clavier, N. Dacheux and R. Podor, *New J. Chem.*, 2003, **27**, 957–967.
- 23 J. M. Cleveland, *The Chemistry of Plutonium*, Gordon & Breach Science Publishers, New York, 1970.



24 C. E. Bamberger, R. G. Haire, H. E. Hellwege and G. M. Begun, *J. Less-Common Met.*, 1984, **97**, 349–356.

25 C. W. Bjorklund, *J. Am. Chem. Soc.*, 1957, **79**, 6347–6350.

26 C. Keller and K. H. Walter, *J. Inorg. Nucl. Chem.*, 1965, **27**, 1253–1260.

27 D. E. Hobart, G. M. Begun, R. G. Haire and H. E. Hellwege, *J. Raman Spectrosc.*, 1983, **14**, 59–62.

28 D. Qin, A. Mesbah, C. Gausse, S. Szenknect, N. Dacheux and N. Clavier, *J. Nucl. Mater.*, 2017, **492**, 88–96.

29 D. W. Qin, A. Mesbah, N. Clavier, S. Szenknect and N. Dacheux, *Cryst. Growth Des.*, 2019, **19**, 2794–2801.

30 D. W. Qin, A. Mesbah, J. Lautru, S. Szenknect, N. Dacheux and N. Clavier, *J. Eur. Ceram. Soc.*, 2020, **40**, 911–922.

31 D. Qin, A. Shelyug, S. Szenknect, A. Mesbah, N. Clavier, N. Dacheux and A. Navrotsky, *Appl. Geochem.*, 2023, **148**, 105504.

32 R. Podor, PhD thesis, University of Lorraine, 1994.

33 A. Mesbah, N. Clavier, E. Elkaim, C. Gausse, I. B. Kacem, S. Szenknect and N. Dacheux, *Cryst. Growth Des.*, 2014, **14**, 5090–5098.

34 A. Mesbah, N. Clavier, E. Elkaim, S. Szenknect and N. Dacheux, *J. Solid State Chem.*, 2017, **249**, 221–227.

35 R. Kijkowska, *Thermochim. Acta*, 2003, **404**, 81–88.

36 A. Shelyug, A. Mesbah, S. Szenknect, N. Clavier, N. Dacheux and A. Navrotsky, *Front. Chem.*, 2018, **6**, 604.

37 R. Jonasson and E. Vance, *Thermochim. Acta*, 1986, **108**, 65–72.

38 V. Brandel and N. Dacheux, *J. Solid State Chem.*, 2004, **177**, 4743–4754.

39 V. Brandel, N. Dacheux and M. Genet, *J. Solid State Chem.*, 1996, **121**, 467–472.

40 K. M. Kurbanov, V. A. Efremov and V. P. Orlovsky, *Kristallografiya*, 1986, **31**, 800–802.

41 D. F. Mullica, D. A. Grossie and L. A. Boatner, *J. Solid State Chem.*, 1985, **58**, 71–77.

42 R. Jardin, C. C. Pavel, P. E. Raison, D. Bouëxière, H. Santa-Cruz, R. J. M. Konings and K. Popa, *J. Nucl. Mater.*, 2008, **378**, 167–171.

43 Y. Ni, J. M. Hughes and A. N. Mariano, *Am. Mineral.*, 1995, **80**, 21–26.

44 S. V. Ushakov, K. B. Helean, A. Navrotsky and L. A. Boatner, *J. Mater. Res.*, 2001, **16**, 2623–2633.

45 E. Du Fou de Kerdaniel, PhD thesis, University of Paris Sud-Paris XI, 2007.

46 J. M. Montel, J. L. Devidal and D. Avignant, *Chem. Geol.*, 2002, **191**, 89–104.

47 G. Lu, A. J. Haes and T. Z. Forbes, *Coord. Chem. Rev.*, 2018, **374**, 314–344.

48 N. Clavier, A. Mesbah, S. Szenknect and N. Dacheux, *Spectrochim. Acta, Part A*, 2018, **205**, 85–94.

49 K. W. Bagnall and M. W. Wakerley, *J. Inorg. Nucl. Chem.*, 1975, **37**, 329–330.

50 R. L. Frost, *Spectrochim. Acta, Part A*, 2004, **60**, 1469–1480.

51 J. Čejka, J. Čejka and A. Muck, *Thermochim. Acta*, 1985, **86**, 387–390.

52 N. Dacheux, V. Brandel and M. Genet, *New J. Chem.*, 1995, **19**, 15–25.

53 T. Vitova, I. Pidchenko, S. Biswas, G. Beridze, P. W. Dunne, D. Schild, Z. Wang, P. M. Kowalski and R. J. Baker, *Inorg. Chem.*, 2018, **57**, 1735–1743.

54 Y. Podkoryrina, I. Pidchenko, T. Prüßmann, S. Bahl, J. Göttlicher, A. Soldatov and T. Vitova, *J. Phys.: Conf. Ser.*, 2016, **712**, 012092.

55 M. Massonnet, L. Claparede, J. Martinez, P. M. Martin, M. O. J. Y. Hunault, D. Prieur, A. Mesbah, N. Dacheux and N. Clavier, *Inorg. Chem.*, 2023, **62**, 7173–7185.

56 L. Amidani, M. Retegan, A. Volkova, K. Popa, P. M. Martin and K. O. Kvashnina, *Inorg. Chem.*, 2021, **60**, 16286–16293.

57 N. Dacheux, V. Brandel and M. Genet, *New J. Chem.*, 1995, **19**, 1029–1036.

58 B. Sitaud, P. L. Solari, S. Schlutig, I. Llorens and H. Hermange, *J. Nucl. Mater.*, 2012, **425**, 238–243.

59 G. Ashiotis, A. Deschildre, Z. Nawaz, J. P. Wright, D. Karkoulis, F. E. Picca and J. Kieffer, *J. Appl. Crystallogr.*, 2015, **48**, 510–519.

60 C. Frontera and J. Rodriguez-Carvajal, *Phys. B*, 2003, **335**, 219–222.

61 M. Keskar, G. P. Shelke, M. Shafeeq, R. A. Phatak, S. K. Sali and S. Kannan, *J. Solid State Chem.*, 2019, **278**, 120850.

62 B. Ravel and M. Newville, *J. Synchrotron Radiat.*, 2005, **12**, 537–541.

63 N. Dacheux, PhD thesis, University of Paris Sud-Paris XI, 1995.

64 J. H. Albering and W. Jeitschko, *Z. Kristallogr.*, 1995, **210**, 878.

65 N. Dacheux, N. Clavier, G. Wallez and M. Quarton, *Solid State Sci.*, 2007, **9**, 619–627.

66 P. Bénard, D. Louer, N. Dacheux, V. Brandel and M. Genet, *An. Quim.*, 1996, **92**, 79–87.

67 N. Clavier, F. Cretaz, S. Szenknect, A. Mesbah, C. Poinsot, M. Descostes and N. Dacheux, *Spectrochim. Acta, Part A*, 2016, **156**, 143–150.

