



Cite this: *Chem. Sci.*, 2024, **15**, 736

All publication charges for this article have been paid for by the Royal Society of Chemistry

Iodide oxidation by ozone at the surface of aqueous microdroplets†

Alexander M. Prophet,^{ab} Kritanjan Polley,^{ab} Gary J. Van Berkel,^b David T. Limmer^{abde} and Kevin R. Wilson^{b*}

The oxidation of iodide by ozone occurs at the sea-surface and within sea spray aerosol, influencing the overall ozone budget in the marine boundary layer and leading to the emission of reactive halogen gases. A detailed account of the surface mechanism has proven elusive, however, due to the difficulty in quantifying multiphase kinetics. To obtain a clearer understanding of this reaction mechanism at the air-water interface, we report pH-dependent oxidation kinetics of I^- in single levitated microdroplets as a function of $[O_3]$ using a quadrupole electrodynamic trap and an open port sampling interface for mass spectrometry. A kinetic model, constrained by molecular simulations of O_3 dynamics at the air-water interface, is used to understand the coupled diffusive, reactive, and evaporative pathways at the microdroplet surface, which exhibit a strong dependence on bulk solution pH. Under acidic conditions, the surface reaction is limited by O_3 diffusion in the gas phase, whereas under basic conditions the reaction becomes rate limited on the surface. The pH dependence also suggests the existence of a reactive intermediate $IOOO^-$ as has previously been observed in the $Br^- + O_3$ reaction. Expressions for steady-state surface concentrations of reactants are derived and utilized to directly compute uptake coefficients for this system, allowing for an exploration of uptake dependence on reactant concentration. In the present experiments, reactive uptake coefficients of O_3 scale weakly with bulk solution pH, increasing from 4×10^{-4} to 2×10^{-3} with decreasing solution pH from pH 13 to pH 3.

Received 14th August 2023
Accepted 3rd December 2023

DOI: 10.1039/d3sc04254e
rsc.li/chemical-science

1. Introduction

Oxidation reactions of iodine in the environment constitute a set of important pathways for mediating global atmospheric oxidant concentrations and reactive emissions.^{1,2} Iodine oxides, the primary products from such reactions, play a particularly significant role in catalytic destruction cycles of ozone in the troposphere, affect the overall atmospheric budget of HO_x and NO_x species, and serve as higher molecular-weight precursors for new particle formation.^{3–7} The iodide anion, I^- , resides in seawater where oxidation at the sea-surface initiates the release of reactive iodine into the atmosphere. Furthermore, the $I^- + O_3$ reaction on the sea-surface makes up a major fraction of overall ozone deposition from the marine boundary layer.⁸ Since discovery of the environmental impact of iodine chemistry through field-based measurements, a host of laboratory-based

experiments have aimed to understand the fundamentals underlying this multiphase reaction. Here we employ laboratory measurements on individual microdroplets together with molecular and continuum modeling to clarify the mechanism of iodide oxidation by ozone.

The earliest laboratory measurements of the heterogenous $I^- + O_3$ reaction found evidence for a bimolecular rate coefficient that approaches the diffusion limit with $k_{rxn} \sim 1 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$.^{9–11} This rate coefficient is similar in magnitude to what has been observed in both gas phase experiments^{12,13} and in the bulk aqueous phase using fast-mixing techniques.¹⁴ This chemistry has been investigated by a number of researchers studying the oxidation of bulk aqueous solutions,^{15–18} micron-scale aerosol,^{19–21} and single iodide-water clusters.^{22,23} Recent theoretical work has aimed to understand the kinetics and energetics of this reaction using a variety of modeling approaches ranging from numerical simulations to *ab initio* calculations of aqueous iodide oxidation by ozone.^{24–27} Although this system has been under investigation for decades, questions remain regarding the detailed reaction mechanism at the air-water interface, under what conditions surface reactions dominate over bulk, and how such processes are coupled to mass transport of reactants, emissions of volatile products, and the overall uptake of ozone from the gas phase. As such, novel experimental and modeling approaches are warranted. Willis &

^aChemical Sciences Division, Lawrence Berkeley National Laboratory, Berkeley, CA 94720, USA. E-mail: krwilson@lbl.gov

^bDepartment of Chemistry, University of California, Berkeley, CA 94720, USA

^cVan Berkel Ventures, LLC, Oak Ridge, TN 37830, USA

^dMaterials Science Division, Lawrence Berkeley National Laboratory, Berkeley, California 94720, USA

^eKavli Energy NanoScience Institute, Berkeley, California 94720, USA

† Electronic supplementary information (ESI) available. See DOI: <https://doi.org/10.1039/d3sc04254e>



Wilson²⁸ and Wilson *et al.*²⁹ introduced a model framework for analyzing multiphase ozone oxidation reactions in aqueous microdroplets by describing ozone partitioning at the air–water interface and subsequent solvation of O₃ into the droplet interior. This approach, implemented through stochastic reaction–diffusion simulations, allows for dynamic concentrations of reactants to be computed at the surface and bulk regions of the droplet using a set of detailed elementary steps describing adsorption, desorption, and solvation of O₃. As we aim to show in this work by combining experiments, kinetic models, and molecular simulations—the dynamics of ozone at the air–water interface are critical for understanding the multiphase chemical mechanism driving surface reactions of I[−] and O₃.

In the current study, we examine oxidation kinetics of iodide by ozone in individual microdroplets by trapping and reacting charged microdroplets in a quadrupole electrodynamic trap (QET). Reaction kinetics are obtained by measuring droplet composition using an open-port sampling interface for mass spectrometry (OPSI-MS) as recently demonstrated by Kaur Kohli *et al.*³⁰ Iodide decay kinetics are monitored as a function of [O₃] and pH. Using the framework of Willis & Wilson,²⁸ a kinetic model is constructed to explain the observed kinetics. Key model constraints that include the desorption and solvation rates of O₃ at the air–water interface are obtained directly from analysis of molecular simulations. Kinetic modeling results suggest the I[−] + O₃ reaction occurs almost exclusively at the droplet surface under our experimental conditions where surficial O₃ may become substantially depleted by the reaction due to diffusion limitations in the gas phase. Experimental kinetics also show a strong pH dependence, related in part to the different chemical mechanisms relevant for destruction of I[−] as a function of acidity. However, to fully explain the observed pH dependence, we postulate the existence of a relatively short-lived reaction intermediate IOOO[−]. This ozonide intermediate has been previously proposed to exist through both experiment and theory, and is analogous to the intermediate observed in Br[−] oxidation at the air–water interface.³¹ A steady-state analysis of the surface concentrations is also provided which, in conjunction with recently developed expressions for uptake, are used to compute uptake coefficients of O₃ over a range of reactant concentrations.

2. Experimental

To investigate droplet reaction kinetics, a quadrupole electrodynamic trap (QET) is used to react arrays of individual microdroplets, which are then analyzed with mass spectrometry to monitor droplet composition as recently described by Willis & Wilson²⁸ and elsewhere.^{32,33} As this technique has been previously described, we provide an overview with a particular focus on the features relevant to the current experiments, including the recently implemented open-port sampling interface (OPSI) for single droplet mass spectrometry.

2.1. Quadrupole electrodynamic trap (QET)

The QET is used to charge and trap individual microdroplets under a controlled flow of humidified zero air (600 cm³ min^{−1}

for all experiments) at room temperature (295 K). Microdroplets are generated by a piezo-electric dispenser (Microfab, MJ-ABP-01, 30 μm orifice) oriented co-axially with the QET trapping rods (Fig. 1A). Droplets are charged by applying ±200–500 V DC bias to an induction plate located directly below the dispensing region. An array of 10–100 droplets is trapped in an upper balancing region of the trap during a typical experiment. Single droplets can be individually transferred from the upper region to a lower trapping region. Microdroplets in the lower trapping region are sized by illuminating the droplet with a 532 nm laser focused axially across the QET interior. Mie scattered light from the microdroplet is then collected and analyzed as previously described by Davies³⁴ to obtain a droplet radius. Once sized, single microdroplets are ejected from the QET and into the analysis region. Ozone is generated using a corona discharge ozone generator supplied with a low flow of oxygen (20–100 cm³ min^{−1}), which is then diluted into 1.5–3 L min^{−1} of nitrogen. 50 cm³ min^{−1} of the resulting dilution flow is combined with the humidified airflow and directed through the QET. Ozone concentrations in the QET are monitored at the reactor outlet using a UV ozone analyzer (2B Technologies, model 106-M).

2.2. Open-port sampling interface (OPSI) mass spectrometry

Droplets ejected into the analysis region are carried by the gas-flow and collide with the inlet of an open-port sampling interface (OPSI) for analysis by mass spectrometry (MS). A schematic of the OPSI is shown in Fig. 1B, outlining the major components of the assembly. A number of studies utilizing a similar design have been recently published for liquid droplet and particle capture,^{30,35–40} and thus, we briefly review the general construction of the OPSI and provide details on the specific components used. The OPSI design consists of a PEEK 3-way tee connected to a stainless 1/8" tube on the top port. A smaller (360 μm OD, 100 μm ID) PEEK tube runs axially inside the outer tube and tee, connecting the top of the assembly to a commercial heated-electrospray source (Thermo-Fischer HESI-II probe) separated by a length of ~20 cm. An additional length of PEEK tubing delivers solvent (methanol in the current experiments) into the perpendicular tee port, allowing solvent to flow through the outer stainless 1/8" tubing and around the inner PEEK tubing. This solvent then flows to the top of the OPSI where it enters the inner tube given sufficient sheath gas flow is used in the HESI to generate a pressure differential between the top of the inner tube and the electrospray needle.

By balancing the solvent delivery flow and sheath gas pressure, a stable flow of solvent through the inlet tube and into the ionization region is achieved. For the current experiments, the OPSI is operated in a slightly overfilled mode wherein the solvent delivery rate is marginally larger than the flowrate through the inner tube. A typical flow rate is ~2.5 mL h^{−1}. Contents of the microdroplets landing in the solvent pool on top of the OPSI are first diluted before traveling through the inner tube and analyzed using electrospray ionization mass spectrometry (MS, Q Exactive Orbitrap, Thermo Fischer Scientific, Inc.) Measuring the iodide signal as a function of analysis



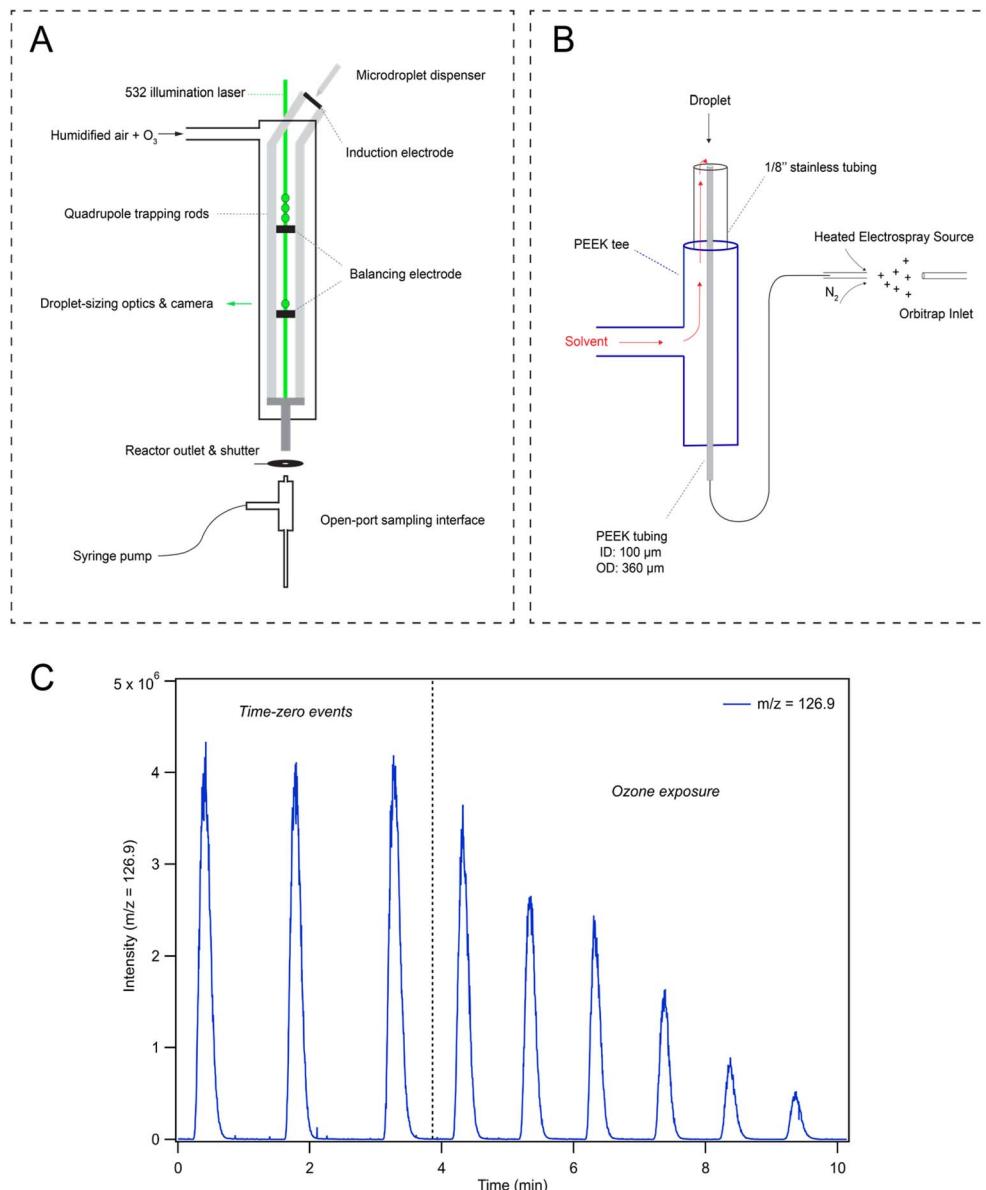


Fig. 1 Experimental schematic of the QET is shown in (A) and an overview of the OPSI-MS assembly is shown in (B). Panel (C) provides example iodide decay kinetics as observed in the MS using an initial 250 mM solution of NaI at pH 3 with droplet radius $r = 24 \mu\text{m}$. Time-zero events demonstrate the stability between successive droplet events before exposure to ozone, while the dashed vertical line denotes the time when the droplets are exposed to $\sim 1 \text{ ppm} [\text{O}_3]$ flow.

time provides a time-series showing single droplet detection events (Fig. 1C), with typical peak widths of $\sim 30 \text{ s}$. Peak areas from these single droplet events are then used to quantitatively determine concentrations of analytes in microdroplets. A calibration curve for droplet concentration response is provided in ESI Section SI-2,[†] demonstrating the linear response of the signal with droplet [NaI]. Fig. 1C shows that the observed $[\text{I}^-]$ decays when droplets are exposed to a gas phase flow of 1 ppm $[\text{O}_3]$ through the QET. After time-zero (marked with a dashed vertical line in Fig. 1C), the peak area of subsequent droplet events is observed to decay as I^- is consumed by O_3 . Typical mass spectra observed in the experiment are provided in Fig. S2 in SI-2,[†] where an initial pre-reaction spectrum is provided

along with a spectrum obtained from a droplet after undergoing ozone exposure.

During reaction, a shutter is placed between the QET outlet and the OPSI inlet to avoid the solvation and buildup of O_3 in the ionization region of the mass spectrometer. During a typical measurement the shutter is opened for $\sim 0.5 \text{ s}$ for each droplet detection event. Without this shutter, gas-phase chemistry (likely ion–molecule reactions) in the ESI source occurs even in the presence of trace O_3 . Using the shutter configuration outlined in Fig. 1A, gas-phase interference can be ruled out by shutting off the ozone flow midway through the experiment and analyzing droplet composition to ensure the iodide ion signal does not recover. No noticeable recovery of iodide signal is

observed with the current approach, indicating all the chemistry observed originates from reactions occurring in the microdroplets and not gas-phase chemistry in the ionization region of the mass spectrometer.

2.3. Droplet composition

All reactions are performed using an initial $[NaI]$ of 250 mM and $[NaCl]$ between 500 mM and 700 mM. NaCl is used to decrease the water activity of solution to $a_w = 0.95$, matching the gas phase relative humidity of $95 \pm 1\%$ in the QET. Maintaining the same water activity in both the dispenser solution and the gas phase prevents droplets from significantly changing size when equilibrating inside the QET. Using this approach, the 30 μ m orifice dispenser reliably provides droplets of radius $24 \pm 1 \mu\text{m}$. The use of NaCl also ensures that the droplet size remains relatively stable during the reaction, with droplet radius changing only $\sim 1 \mu\text{m}$ during reaction as I^- is consumed. pH 3 droplets are dispensed from a bulk aqueous solution containing 300 mM citric acid/sodium citrate, buffered at pH 3, and $[NaCl] = 500$ mM. Likewise, pH 8 droplets are dispensed from a bulk solution containing 300 mM Tris base (tris(hydroxymethyl) aminomethane), buffered at pH 8, and $[NaCl] = 700$ mM. The particular organic buffers used were chosen due to the presence of multiple hydroxyl groups in both buffer systems, which decreases the interfacial preference of the organics and therefore minimizes any interference with iodide surface chemistry. However, we anticipate that most buffers in this pH range would not have a strong surface propensity and therefore, our results would not be sensitive to the particular buffer used.

Slightly different NaCl concentrations are used to compensate for the differences in the hygroscopicity of the citrate/citric acid and Tris base salt buffers. The $[NaCl]$ is selected for each solution to give the same final droplet size of 24 μm , yielding a $[NaI]$ concentration that is constant across pH given the same QET relative humidity. Unlike the pH 8 and pH 3 droplets, the pH 13 droplets are unbuffered due to the lack of viable buffers for this region and the constraint on the overall salt concentration to maintain stable droplet sizes. However, the consequence of an unbuffered solution in this case is minor as the reaction mechanism under basic conditions does not incorporate production or consumption of H^+ or OH^- . Fresh solutions were prepared daily using HPLC-grade water (Sigma-Aldrich), NaI (Sigma-Aldrich, $\geq 99.5\%$), NaCl (Sigma-Aldrich, $\geq 99.5\%$), citric acid (Sigma-Aldrich, $\geq 99.5\%$), sodium citrate monobasic ($\geq 99.5\%$), tris(hydroxymethyl)aminomethane (Sigma-Aldrich, $\geq 99.8\%$) and NaOH (Sigma-Aldrich, $\geq 98\%$).

3. Experimental results & discussion

3.1. Iodide decay kinetics

Droplet reaction kinetics for a series of pH and $[O_3]$ are shown in Fig. 2. Experimental results (shown as points) are compared with kinetic simulations (lines), described below in Sections 4 and 5. Error bars for individual data points are estimated using the standard deviation of droplet peak areas from a set of unreacted droplet events prior to each experiment (e.g., droplet

events in the “time-zero” section in Fig. 1C). Fig. 2A and B show iodide decay in microdroplets dispensed from solutions buffered at pH 3 and pH 8, respectively. A number of qualitative features are observed as a function of pH and $[O_3]$. For each droplet pH, the iodide consumption rate is first order in $[O_3]$. This is summarized in Fig. 3, where initial rates of decay are shown to increase linearly with increasing ozone concentration. A strong dependence on dispensed-solution pH is also observed. The overall rate of the reaction increases significantly with decreasing pH. From the most basic to acidic conditions, the observed initial decay rate increases by almost an order of magnitude. As considered further in the model discussion Section 6.2 and as noted previously from product emission studies,^{15–17} a major contribution to the apparent reaction acceleration with decreasing pH originates from the reaction of I^- with its primary oxidation product HOI to produce I_2 . However, an additional pH dependence of the reaction is proposed below to account for the complete set of pH dependent observations shown in Fig. 2 and 3.

Beyond a difference in the overall rate, changes in the shape of the decay kinetics in Fig. 2 are also observed as a function of pH and even $[O_3]$ under strongly basic conditions. The decay of iodide in acidic solution appears mostly linear in time, with a small but significant tail as the iodide concentration approaches zero. Conversely, the decay of iodide under strongly basic conditions appears exponential for $[O_3] > 1$ ppm but becomes increasingly bi-exponential or linear in time as $[O_3]$ decreases below 1 ppm. The origin of these functional forms for the decay kinetics appears to be driven by the complex interplay of reaction and surface adsorption of iodide, reaction intermediates, and products and will be examined in more detail below in Section 6.

3.2. Reaction product yields and kinetics

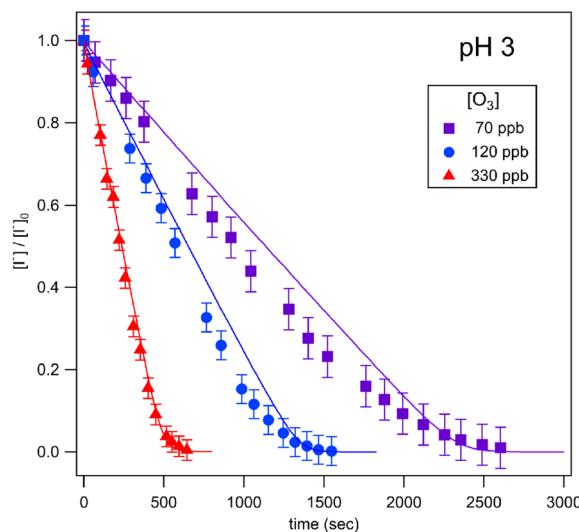
Utilizing droplet composition analysis with the OPSI-MS, products formed and remaining in the condensed phase of the droplet are detected in positive- or negative-mode electrospray ionization. For experiments done using pH 3 and 8, however, no condensed-phase products are observed as all I^- in solution is quickly oxidized to I_2 , which evaporates into the continuous gas flow through the QET. However, under strongly basic conditions (notably, above the pKa of HOI of 10.8) IO_3^- is observed in the mass spectrometer at $m/z = 174$ as shown in SI-2 and Fig. S2.† Peak areas from time-traces of the ion signal at m/z 174 are monitored throughout the reaction under basic conditions and provided in Fig. S3† for three example ozone concentrations. As discussed further in SI-2,† the iodate yield appears to decrease with decreasing $[O_3]$. While the origin of this shift in product yield is unknown, this may suggest a change in surface mechanism with decreasing availability of O_3 .

4. Model description

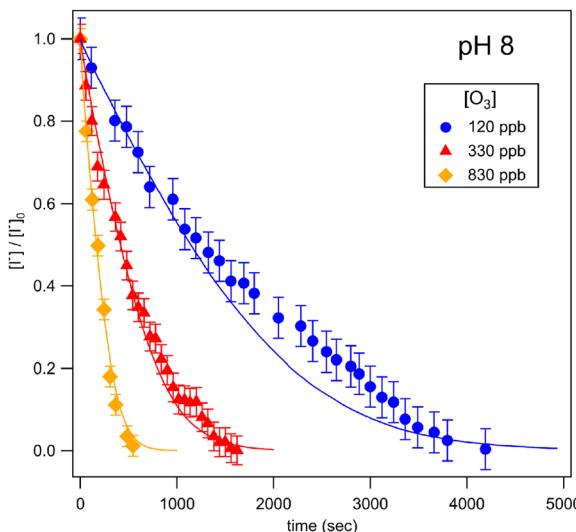
To understand the reactivity observed in the droplet experiments, a kinetic model is constructed. This model describes the



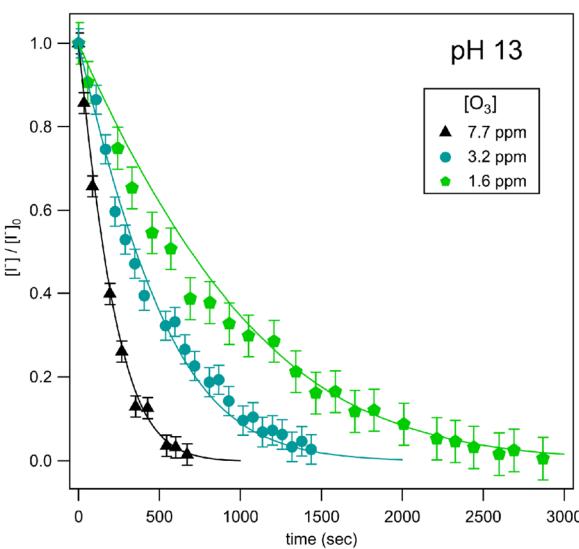
A



B



C



D

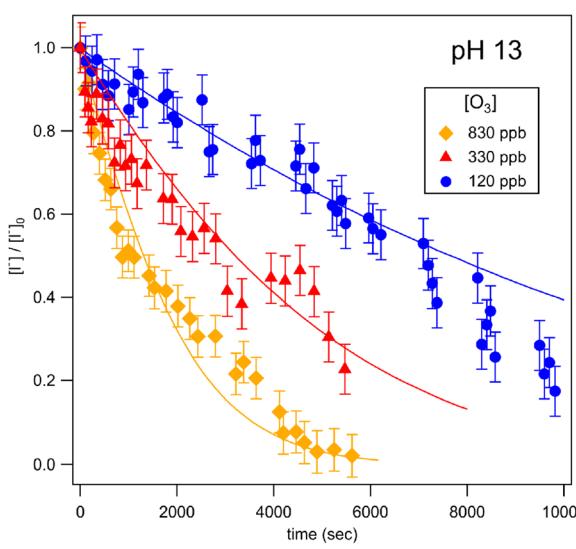


Fig. 2 Iodide decay kinetics showing ozone oxidation of $24\text{ }\mu\text{m}$ droplets containing 250 mM initial $[\text{NaI}]$ for varying $[\text{O}_3]$ and dispensed solution pH. Experimental data is shown as points in (A)–(D) and simulation results as solid lines. Panel (A) shows results using droplets dispersed from pH 3 solution buffered with citrate/citric acid. Panel (B) shows results using pH 8 solution buffered with Tris base. Panels (C) and (D) shows results for unbuffered solution with an initial pH of 13, with $[\text{O}_3]$ above 1 ppm given in (C) and below 1 ppm given in (D).

partitioning and reaction of gas- and liquid-phase species at the droplet surface and within the bulk interior. The model framework is based on work from Willis & Wilson²⁸ where a kinetic description of O_3 adsorption and desorption at the microdroplet surface and subsequent solvation into the bulk liquid is presented and benchmarked using a set of ozone oxidation reactions. Further analysis of this approach in Wilson *et al.*²⁹ provides closed-form expressions for predicting uptake and reaction of trace gases into microdroplets. In the current work, molecular dynamics (MD) simulations are used to study the solvation of ozone at the air–water interface and the kinetics of both the solvation and desorption processes. A water slab with 768 water molecules and sodium halide salts, modeled with a classical polarizable force field,⁴¹ was used to represent

an aqueous droplet at 300 K . A snapshot from the simulation of O_3 adsorption is shown in Fig. 4A. The free energy profile for transferring an ozone molecule through the air–water interface, with sodium halide salts ($0.28\text{ M NaI} + 0.84\text{ M NaCl}$) in the solution, is shown in Fig. 4B where the shaded blue region is the (scaled and shifted) water density profile. Results and details of the MD simulations and the procedures for obtaining the solvation and desorption rates are included in the ESI Section SI-3.† In the following modeling sections, we briefly outline the kinetic model and detail the specific novel components relevant for the application of this framework to the analysis of the $\text{I}^- + \text{O}_3$ reaction.

The model developed in this section is implemented in Kinetiscope⁴², a software package previously used to simulate



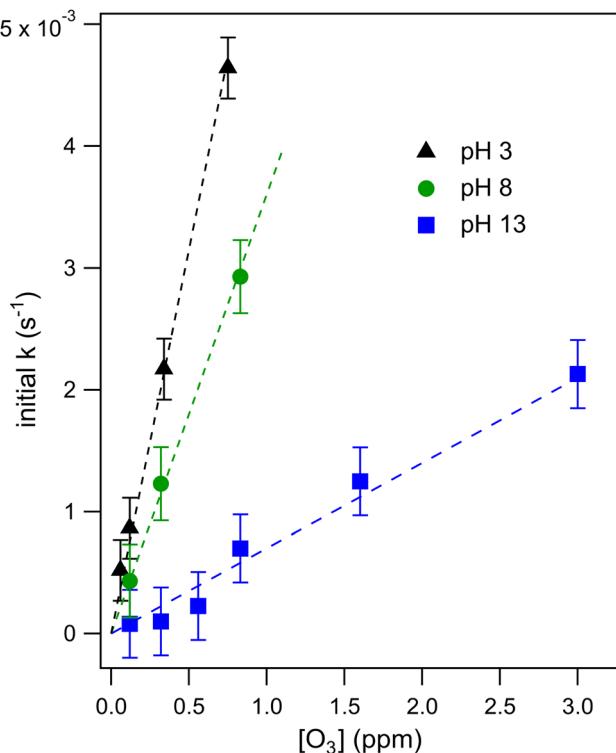


Fig. 3 Kinetic plot showing initial iodide decay rate vs. $[O_3]$. Initial rates from experiments using solutions with pH 3, 8, 13 are shown as points and simulation results are shown as dashed lines. Linear scaling with $[O_3]$ is observed for all cases, with dramatic acceleration between pH 13 and pH 3. Error bars for data points reflect a combined uncertainty in reactor $[O_3]$ concentration and deviation between individual droplet measurements.

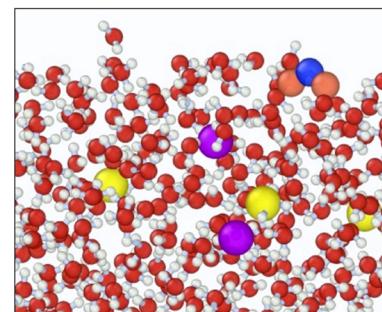
kinetics in a variety of systems such as organic aerosol, aqueous microdroplets, and emulsions.^{43–45} Droplet kinetics are simulated by conceptualizing the droplet bulk interior and surface as two separate compartments that have a rectangular prism geometry. Both compartments have the same $1 \times 1 \text{ nm}^2$ cross-sectional area but have different compartment lengths. The length of the bulk compartment is chosen to be $r/3$, in order to preserve the correct surface-to-volume ratio of a sphere with radius, r . The surface compartment length (δ) is 1 nm and represents the thickness of the air–water interface. This surface depth corresponds to the approximate length scale over which the water density changes across the air–water interface as observed by the MD results in Fig. 4B and previous simulations.⁴⁶ Molecules move between the surface and bulk compartments in Kinetiscope by Fickian diffusion, which is governed by the relative concentrations in each compartment and the compartment lengths. Solutes that diffuse into the surface compartment from the bulk can then adsorb to the air–water interface through a kinetic description of solvation/desolvation to establish surface concentrations. Similarly, gas-phase species partition to the surface through a set of adsorption–desorption steps described in the following sections and summarized in the multiphase framework in Fig. 5A.

4.1. Ozone partitioning scheme

The ozone partitioning scheme is conceptualized as two parallel processes: one describing the kinetic adsorption process at the gas–liquid interface, and one describing diffusion of ozone through the gas phase to the droplet surface. As the kinetic partitioning description draws from the work of Willis & Wilson,²⁸ we revisit the mechanics of this description before introducing the description of ozone diffusion.

4.1.1 Ozone gas–liquid kinetic equilibrium. The kinetic description of O_3 equilibration between phases begins with decomposing the overall dimensionless Henry's law coefficient (H_{cc}) for ozone solvation into gas-to-surface (gs) and surface-to-bulk (sb) components, the product of which preserves the overall gas-to-bulk (gb) coefficient $H_{cc}^{gb} = H_{cc}^{gs} \cdot H_{cc}^{sb}$. The individual components H_{cc}^{gs} and H_{cc}^{sb} , which link the gas, interface and bulk O_3 concentrations, are computed²⁸ from solvation free energies obtained from MD simulation results shown in Fig. 4 and discussed in SI-3,† yielding $H_{cc}^{gs} = 9.3$ and $H_{cc}^{sb} = 0.0156$ with $H_{cc}^{gb} = 0.145$. The value of H_{cc}^{gb} is consistent with literature values for solvation of O_3 in 1 M sodium chloride solutions.^{47,48} Surface components H_{cc}^{gs} and H_{cc}^{sb} can in turn be expressed kinetically (eqn (1) and (2)), relating the ozone partitioning steps shown in Fig. 5A, *i.e.*, adsorption/desorption from the gas-phase and solvation/desolvation from the liquid phase:

A



B

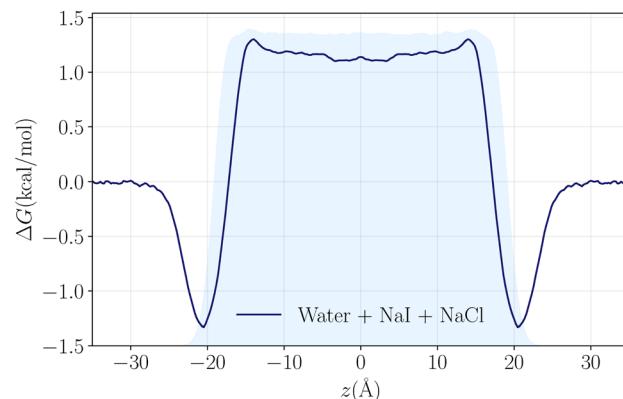


Fig. 4 (A) Snapshot of the MD simulation where an ozone molecule is adsorbed near the air–water interface. (B) The free energy profile for transferring an ozone molecule through a water slab with 0.28 M NaI and 0.84 M NaCl is displayed. The shaded blue region shows the (scaled and shifted) water density profile.



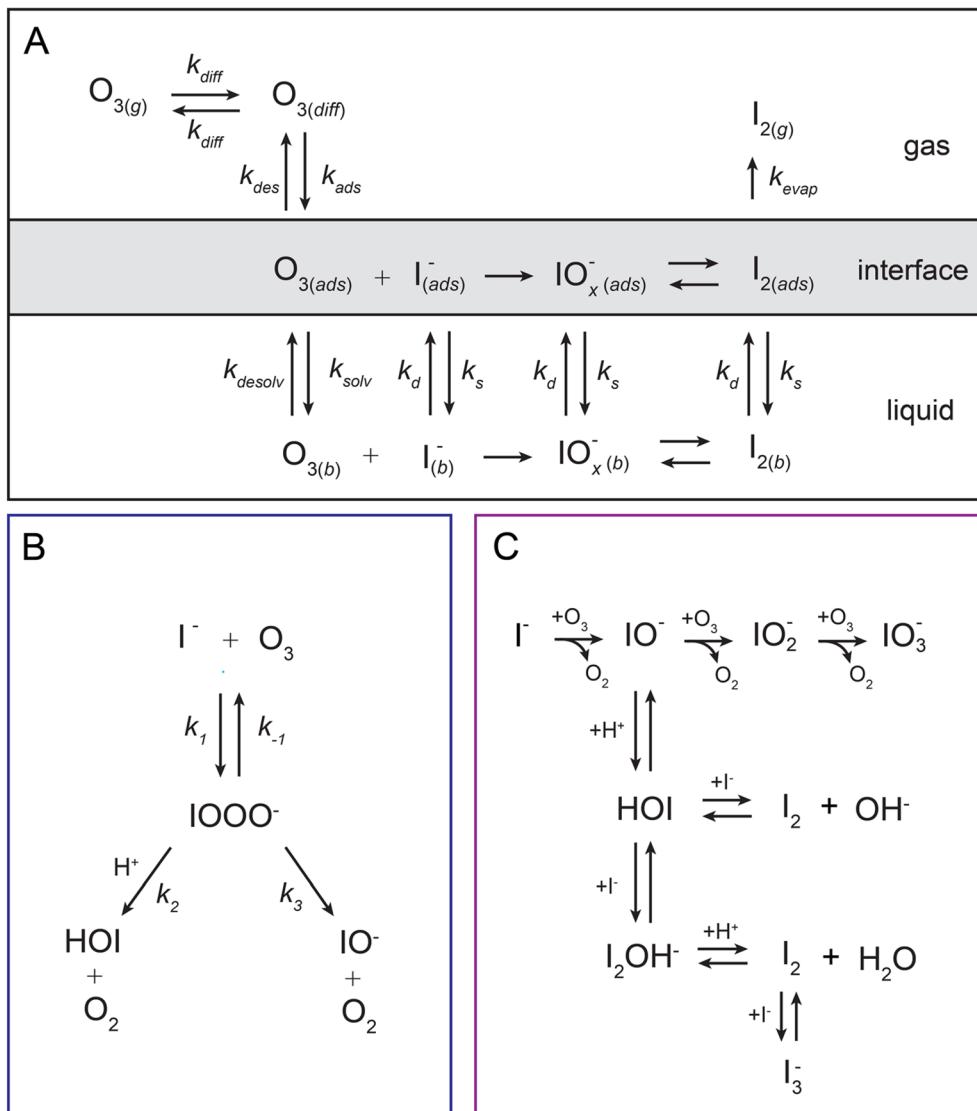


Fig. 5 The multiphase framework and chemical mechanisms implemented in the simulations. Panel (A) outlines the heterogeneous ozone oxidation model framework showing adsorption of liquid- and gas-phase species to the interface. Panel (B) shows the primary $I^- + O_3$ reaction mechanism involving the reactive intermediate $IOOO^-$. Panel (C) shows subsequent chemistry with HOI and IO^- including both further oxidation by O_3 and secondary chemistry with I^- . Chemistry shown in (B) and (C) is included in both the bulk liquid and the droplet interface as denoted in (A) through an abbreviated reaction scheme.

$$H_{cc}^{gs} = \frac{[O_{3(ads)}]}{[O_{3(g)}]} = \frac{k_{ads} \cdot \Gamma_{O_3}^\infty \cdot \sigma}{k_{des} \cdot \delta}, \quad (1)$$

$$H_{cc}^{sb} = \frac{[O_{3(b)}]}{[O_{3(ads)}]} = \frac{k_{solv} \cdot \delta}{k_{desolv} \cdot \Gamma_{O_3}^\infty}. \quad (2)$$

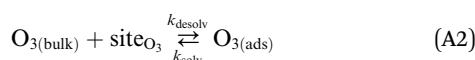
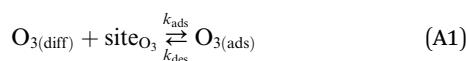
In eqn (1) and (2), the O_3 subscript (ads) denotes surface-adsorbed ozone, (g) denotes gas-phase O_3 and (b) refers to bulk solvated O_3 . In the case of fast surface reactions, the effective $[O_{3(g)}]$ near the interface may become depleted due to gas-phase diffusion limitations. To address this in the model, we introduce the species $[O_{3(diff)}]$ that denotes the gas-phase O_3 concentration that has diffused across a characteristic diffusion length in the gas phase. This characteristic length and the

kinetic steps governing $[O_{3(diff)}]$ are addressed below in Sections 4.1.2 and SI-4.† The diffusional O_3 description is implemented in the model through modification of eqn (1)

$$H_{cc}^{gs} = \frac{[O_{3(ads)}]}{[O_{3(diff)}]} = \frac{k_{ads} \cdot \Gamma_{O_3}^\infty \cdot \sigma}{k_{des} \cdot \delta}, \quad (3)$$

where the diffusion-limited $[O_{3(diff)}]$ description replaces the overall gas phase concentration $[O_{3(g)}]$. Values for the individual coefficients k_{des} and k_{solv} are computed directly from the MD simulations detailed in Section SI-3† with results shown in Fig. S6.† k_{ads} and k_{desolv} can then be calculated using the equilibrium expressions in eqn (1) and (2). The complete set of coefficients is included in Table S1.†

Simulating the adsorption and solvation steps for O_3 within the kinetic model is done using a modified Langmuir adsorption framework⁴⁹ where O_3 adsorbs (from either the gas or liquid phase) to a surface site (steps (A1) and (A2), respectively):



Surface sites for ozone (not pictured in Fig. 5 for brevity), are conceptualized as available regions of the interface where ozone can adsorb. In this case, we assume the area of one surface site is equal to the molecular area of O_3 (18.5 \AA^2) as computed by Vieceli *et al.*⁵⁰ The maximum number of surface sites is $\Gamma_{O_3}^\infty$ and can be expressed in volumetric units using surface thickness $\Gamma_{O_3}^\infty/\delta$. Throughout the simulation, the total number of surface sites is conserved and expressed as,

$$[\text{site}_{O_3}] = \frac{\Gamma_{O_3}^\infty}{\delta} - [O_{3(\text{ads})}]. \quad (4)$$

Eqn (1) includes a dependence on σ , the sticking coefficient (or the thermal accommodation coefficient) for a single O_3 molecule onto a site. The sticking coefficient σ is distinct in the present model from the accommodation coefficient α , defined⁵¹ as the probability of solvation relative to desorption: $\alpha = k_{\text{solv}}/(k_{\text{des}} + k_{\text{solv}})$. From the simulations presented in SI-3,[†] $\alpha = 0.0097$. While Willis & Wilson²⁸ could identify a lower bound for σ as $>10^{-4}$, the conditions they analyzed were insensitive to values of σ above this value and the treatment of O_3 gas-phase diffusion was not addressed. In the current work, a greater sensitivity is observed in the kinetic model, and a value $\sigma = 0.93$ is obtained from molecular simulations, in general agreement with measurements from direct scattering experiments.^{52–54} A value of $\sigma \sim 1$ means that every O_3 molecule that collides with the interface thermalizes before undergoing desorption, solvation, or reaction.

4.1.2 Ozone gas diffusion description. Diffusion limitations for surface reactions become relevant for a select range of droplet sizes and reaction rates, where the reactive loss on the surface exceeds the maximum rate O_3 can diffuse to replenish the droplet interface. In this case, the diffusion limit generates an ozone concentration gradient extending from the droplet surface across some characteristic diffusive length L_{diff} into the gas phase (see SI-4 and Fig. S8[†] for illustration of relevant length scales). Previous work studying gas diffusion limitations to droplet surfaces assumed L_{diff} to be equal to the droplet radius r for droplet sizes large enough that the gas phase can be described by the continuum regime ($\text{Kn} < 0.01$).^{55,56} The diffusion rate for O_3 in air with diffusion coefficient D_g across a length of $L_{\text{diff}} = r$ in one dimension is

$$k_{L_{\text{diff}}} = \frac{2 \cdot D_g}{r^2}. \quad (5)$$

We note that the simulation geometry effectively simulates mass transport using a one-dimensional description, meaning

that higher order geometric diffusional corrections are unnecessary.

As the present model avoids explicitly simulating concentration gradients in preference for discretized spatial compartments, an additional characteristic length (the adsorption length L_{ads}) is necessary for describing the local O_3 concentrations near the interface. Originally introduced in the context of surfactant adsorption,⁵⁷ L_{ads} denotes the length over which O_3 is depleted in the gas-phase directly from the adsorption and desorption kinetics introduced above as step A1. Conceptually, as shown in Fig. S8,[†] a shell of width L_{ads} surrounding the droplet contains the same number of O_3 molecules as the droplet interface when the system comes to equilibrium. When the droplet surface reaches equilibrium, the number of O_3 molecules on the interface can be calculated as $n_{O_3} = [O_{3(\text{g})}] \cdot H_{\text{cc}}^{\text{gs}} \cdot V_{\text{surf}}$ where V_{surf} is the simulated surface volume ($1 \times 1 \times \delta \text{ nm}^3$ in current simulations). The rapid kinetic process at the surface draws n_{O_3} molecules contained in a gas-adsorption volume V_{ads} extending radially from the droplet surface. Maintaining the rectangular prism simulation geometry introduced above then necessitates V_{ads} have length $H_{\text{cc}}^{\text{gs}} \cdot \delta \text{ nm}$. This defines the simulation adsorption length $L_{\text{ads}} = H_{\text{cc}}^{\text{gs}} \cdot \delta \text{ nm}$ (see SI-4 and Fig. S8[†] for summary illustration of characteristic diffusion and adsorption lengths). Since the volume defined by L_{ads} is the relevant gas volume for O_3 adsorption to the interface, the relevant rate for diffusion into the adsorption volume is found by multiplying eqn (5) by $L_{\text{diff}}/L_{\text{ads}}$

$$k_{\text{diff}} = \frac{2 \cdot D_g}{r^2} \cdot \frac{L_{\text{diff}}}{L_{\text{ads}}} = \frac{2 \cdot D_g}{r \cdot L_{\text{ads}}}. \quad (6)$$

This definition of k_{diff} is used in the simulation to supply $O_{3(\text{diff})}$ directly to the interface, providing an upper bound on diffusional transport of $O_{3(\text{g})}$ to the surface. Diffusion is simulated in the kinetic model by including the following step directly in the surface compartment:



In this implementation, the concentration of species $[O_{3(\text{g})}]$ is defined to remain constant, while $[O_{3(\text{diff})}]$ may deviate from $[O_{3(\text{g})}]$ due to competition between steps (A1)–(A3), as well as any chemistry downstream of these steps.

4.2. Aqueous solute surface partitioning scheme

Surface-adsorption of solutes to the air–water interface is treated using a similar Langmuir description where species in solution, after diffusing into the surface compartment, may adsorb to the surface with a rate proportional to the bulk ion concentration and number of available surface sites. Equilibrium surface-adsorbed $[I^-]$ is expressed by a Langmuir isotherm:

$$[I_{(\text{ads})}^-] = \frac{\Gamma_{I^-}^\infty}{\delta} \cdot \frac{K_{\text{eq}}^{I^-} \cdot [I_{(\text{b})}^-]}{1 + K_{\text{eq}}^{I^-} \cdot [I_{(\text{b})}^-]}, \quad (7)$$



where the Langmuir equilibrium coefficient is defined kinetically as,

$$K_{\text{eq}}^{\text{I}^-} = \frac{k_{\text{desolv}}^{\text{I}^-}}{k_{\text{solv}}^{\text{I}^-}}. \quad (8)$$

The isotherm in eqn (7) expresses the surface concentration of iodide in terms of the bulk concentration when the system is at equilibrium (*i.e.*, without reaction). While the adsorption description of iodide follows an analogous scheme to the ozone adsorption expression in step (A1), iodide and ozone are assumed to occupy a distinct set of surface sites. This is not the case for other solutes, where we assume that reaction products and intermediates in the aqueous phase compete for the same set of surface sites as iodide. A maximum surface site concentration $\Gamma_{\text{I}^-}^{\infty}/\delta$ is defined for these solutes. We estimate a maximum concentration $\Gamma_{\text{I}^-}^{\infty}/\delta$ from previous measurements and simulations of iodide at the air–water interface that indicate surface concentrations ranging from ~ 3 M to 10.5 M.^{58–60} In the model we use $\Gamma_{\text{I}^-}^{\infty}/\delta = 10.5$ M as a representative maximum concentration for solutes at the air–water interface, in accordance with measurements of surficial iodide concentrations using nonlinear spectroscopy.⁶⁰

Previous investigations of I^- at the air–water interface have repeatedly found a strong surface preference for iodide relative to other aqueous ions. Such determinations have been made through a variety of techniques including kinetic measurements,^{61,62} photoemission spectroscopy,^{19,63,64} and second-harmonic generation spectroscopy.^{65,66} Measurements of surface-affinity are quantified for iodide through a Gibbs free energy of adsorption to the air–water interface, $\Delta G_{\text{ads}}^{\text{I}^-}$ which can be directly related to a Langmuir equilibrium coefficient using a commonly employed Langmuir adsorption framework:⁶⁵ $K_{\text{eq}}^{\text{I}^-} = e^{-\Delta G_{\text{ads}}^{\text{I}^-}/RT}/C_w$ where C_w is the solvent concentration (assuming water with $C_w \sim 55$ M). Values for $\Delta G_{\text{ads}}^{\text{I}^-}$ have also been computed from ion density profiles obtained by MD simulations of sodium iodide at the air–water interface.^{58,67} Reports of $\Delta G_{\text{ads}}^{\text{I}^-}$ from experiment and simulations range from -0.8 kcal mol $^{-1}$ to -6.2 kcal mol $^{-1}$, demonstrating a strong dependence on solution composition and model framework employed for analysis.^{61–63,65,66,68} This range of $\Delta G_{\text{ads}}^{\text{I}^-}$ corresponds to a large uncertainty in $K_{\text{eq}}^{\text{I}^-}$, ranging from 0.1 M $^{-1}$ to 650 M $^{-1}$. To determine the appropriate $K_{\text{eq}}^{\text{I}^-}$ for the present model, we first performed MD simulations of the initial experimental solution composition and observe the initial surface concentration under these conditions to be $[\text{I}_{\text{ads}}^-] = 780$ mM (see Fig. S5 in SI-3†). Assuming a maximum surface concentration of $\Gamma_{\text{I}^-}^{\infty}/\delta = 10.5$ M, the simulated $[\text{I}_{\text{ads}}^-]$ constrains $K_{\text{eq}}^{\text{I}^-} = 0.32$ M $^{-1}$, a value on the lower side of the reported range in the literature.

Although the surface concentrations in the kinetic model are constrained by the MD simulation results, a number of uncertainties persist due to the complex nature of the air–water interface. The surface selectivity of the halide ions depends strongly on the choice of water model and force field parameters in MD simulations.^{64,69–71} The prediction of our current model

and force field is in agreement with recent observation from experiments and MD simulations.^{60,69,72} One clear deficiency of the current approach is the coarse-grain perspective of the surface, where a single kinetic volume is used to define the entire interface. In reality, small angstrom-scale features in the density profiles of the solutes and solvent at the interface are important for understanding the structure of the interface. Such features can be observed in the free energy and density profiles for aqueous iodide and ozone included in ESI Fig. S4 and S5 in Section SI-3.† Fine structural details of this aqueous system have also been observed using liquid microjet techniques combined with X-ray photoemission spectroscopy.⁷¹ From the perspective of reactivity, however, the interface may in fact be sufficiently represented without such fine details in some instances since the chemical loss of I^- at the interface will be determined most strongly by the overlap region of the density profiles of I^- and O_3 at the interface, where concentrations will remain approximately constant (in the absence of diffusive limitations).

The surface affinities of the reaction products are known with less accuracy than iodide. However, some of the oxidation products in solution have been observed to be depleted from the interface and are suspected to be $\sim 10\text{--}100\times$ less surface active than I^- .^{73,74} Nevertheless, for the sake of simplicity we use the same $K_{\text{eq}}^{\text{I}^-}$ for all solutes in the simulation. Furthermore, we also assign the volatile, neutral products generated in solution (I_2 and HOI) to possess the same partitioning behavior with $K_{\text{eq}}^{\text{I}^-}$. Forward and reverse rates, k_{desolv} and k_{solv} , describing the kinetic components of $K_{\text{eq}}^{\text{I}^-}$ are included for each simulated species in ESI Table S1† and the sensitivity of these rates analyzed in Section SI-5.†

4.3. Product evaporation description

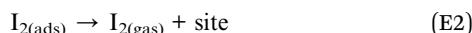
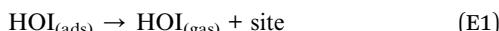
Volatile products generated by oxidation of iodide (*e.g.*, I_2 and HOI) evaporate from the simulated droplet surface through an irreversible step. As with ions in solution, the first step for a volatile species to partition out of the condensed phase is an adsorption step to the air–water interface. Once adsorbed, the species may evaporate from the surface described by a first-order rate. First-order evaporative rate coefficients for each volatile species are derived using the Hertz–Knudsen equation, where the flux of gas from the surface is calculated as,⁷⁵

$$J = \frac{\alpha_e \cdot p_{\text{vap}}}{\sqrt{2\pi mk_B T}},$$

where α_e is the evaporation coefficient (assumed to be 1 for simplicity), m is the molecular weight of the volatile species, k_B is the Boltzmann constant and T is the absolute temperature. The saturation vapor pressure of the species, p_{vap} , is related to its Henry law coefficient and adsorbed surface concentration $p_{\text{vap}} = [\text{I}_{2\text{ads}}]/H_{\text{cc}}^{\text{gb}}$. The surface flux is converted to a rate coefficient by scaling J with the surface area-to volume ratio (SA/V) to account for the overall surface accessibility from the bulk. Individual rate coefficients for volatile species are therefore calculated as,⁷⁶

$$k_{\text{evap}} = \frac{\text{SA}}{V} \cdot \frac{\alpha_e \cdot \left(\frac{1}{H_{\text{cc}}^{\text{gb}}} \right)}{\sqrt{2\pi m k_B T}}.$$

k_{evap} for HOI and I₂, computed from their associated Henry's law constants^{77,78} are used to define evaporative rates for the following steps (E1) & (E2) (and simulation steps S21 and S22 in Table S1†).



4.4. Reaction steps

The chemical mechanism used in the present model relies on previous literature describing the I⁻ + O₃ reaction in both the gas and liquid phase. The complete set of reactive steps and associated rate coefficients are shown in Table S1.† A scheme showing each reaction step is shown in Fig. 5B and C. Below we first review the primary oxidation step of iodide outlined in Fig. 5B, followed by the subsequent reactions involving the primary oxidation products found in Fig. 5C.

4.4.1 Primary oxidation reaction. The initial oxidation step illustrated in Fig. 5B is proposed to proceed through a reaction intermediate IOOO⁻ in analogy to the BrOOO⁻ intermediate produced during bromide ozonation.^{14,31} Although the intermediate IOOO⁻ has not been directly observed, previous experimental and theoretical work proposed its existence at the air-water interface.^{17,22,23,26} The mechanism implemented in the model framework here is adapted from the mechanism for aqueous Br⁻ oxidation by ozone as introduced by Liu *et al.*¹⁴ wherein a stable ozonide adduct is first generated, followed by dissociation to HOI or IO⁻.



The initiation reaction step (R1) involves a fast equilibrium for adduct formation between I⁻ and O₃. Once formed, the IOOO⁻ adduct subsequently reacts with a proton or water to form HOI. The proton-assisted pathway is provided as step (R2) and the water-assisted pathway is defined as a unimolecular decay step (R3), assuming the droplet water content remains constant. While rate coefficients for these individual steps are unknown, we propose that previous measurements of the consumption rate of ozone in bulk solutions^{14,18} provide the rate of adduct formation, *i.e.*, $k_1 = 1.2 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$. For the sake of simplicity, and in the absence of evidence to the contrary, we assume steps (R1)–(R3) have the same rate at the surface and within the bulk. Given the measured gas- and liquid-phase rate coefficients for this reaction are nearly identical,^{12,14} we believe

this is a reasonable assumption. We anticipate the validity of this assumption will be tested by measuring the iodide concentration dependence of the oxidation kinetics, and ideally through future experiments targeting the chemical lifetime of the ozonide intermediate in the bulk and at the surface. We also assume that the proton-assisted decay rate (R2) is diffusion limited, with a rate coefficient of $k_2 = 10^{11} \text{ M}^{-1} \text{ s}^{-1}$, an estimated upper bound for the ion recombination rate in solution as measured previously in the proton-hydroxide recombination reaction.^{79,80} While this assumption likely overestimates the rate of (R2), we similarly suggest that such details could be gathered from future experiments studying the ozonide intermediate directly, particularly in light of HOI and IO⁻ formation kinetics in solution.

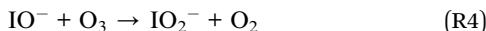
Using the aforementioned values of k_1 and k_2 , coefficients k_{-1} and k_3 are treated as adjustable parameters in the model. A single set of values for k_{-1} ($3.6 \times 10^4 \text{ s}^{-1}$) and k_3 (220 s^{-1}) produce simulation results that agree with the observed droplet kinetics shown in Fig. 2. As presented by Liu *et al.*¹⁴ and shown below in eqn (9), the four rate coefficients k_1 , k_{-1} , k_2 , and k_3 combined with a steady-state approximation for reaction intermediate IOOO⁻ yields an effective bimolecular rate coefficient (k_{obs}) for the overall I⁻ + O₃ reaction,

$$k_{\text{obs}} = \frac{k_1 \left(\frac{k_2}{k_{-1}} \cdot [\text{H}^+] + \frac{k_3}{k_{-1}} \right)}{1 + \frac{k_2}{k_{-1}} \cdot [\text{H}^+] + \frac{k_3}{k_{-1}}} \quad (9)$$

k_{obs} quickly approaches measured rate of $k_{\text{exp}} = 1.2 \times 10^9 \text{ L mol}^{-1} \text{ s}^{-1}$ for neutral to acidic solutions but slows significantly to $\sim 7 \times 10^6 \text{ L mol}^{-1} \text{ s}^{-1}$ under strongly basic conditions. Fig. 6 shows the behavior of k_{obs} from eqn (9) from pH 0 to 13. As discussed further in Section 6, this effective decrease in rate is compounded by the changing overall reactivity of I⁻ in basic solutions, where secondary chemistry involving I⁻ is negligible. We note the experimental rate measured by Liu *et al.*¹⁴ in aqueous solution at pH 6.7 is ~ 50 times larger than k_{obs} predicted from the parametrization shown in Fig. 6. However, this deviation would be expected for bulk measurements of O₃ decay due to IOOO⁻ formation, as the fast initial adduct formation would not be influenced by differences in bulk pH affecting the overall rate of steps (R1)–(R3).

4.4.2 Secondary oxidation reactions. Once formed, the conjugate pair HOI and IO⁻ react further with O₃ and I⁻ yielding different final products depending on pH. For pH values above the pK_a of HOI (pK_a = 10.8), the unimolecular decay pathway of IOOO⁻ shown in (R3) dominates the overall reactivity and the conjugate base IO⁻ is stable in solution. Aqueous IO⁻ can then undergo further oxidation to produce IO₂⁻ and ultimately the iodate anion IO₃⁻. This oxidation process has been studied extensively in wastewater treatment where hypoiodous acid and similar species are used as disinfectants.^{81,82} Mechanistic studies of I⁻ + O₃ in the gas-phase also show IO⁻ as the first oxidation product in the sequential oxidation of I⁻ to IO₃⁻. The sequential addition mechanism of O₃ to the primary oxidation product IO⁻ is implemented as observed by Bhujel *et al.*¹²





Reactions (R4) and (R5) are included directly in the model at both the surface region and the bulk. Here we assume that (R4) and (R5) are irreversible given the relatively slow reverse rate of reaction.¹² As suggested in previous literature, rate coefficients for (R4) and (R5) depend on the phase, where the rate in the gas-phase appears collision limited but the analogous liquid phase rate is $\sim 1000\times$ slower. In the current model, we assume reactions (R4) and (R5) at the droplet surface resemble the gas-phase reactivity more than the liquid phase, and therefore use an approximate diffusion-limited reaction rate coefficient $k \sim 10^{10} \text{ M}^{-1} \text{ s}^{-1}$ for steps (R4) and (R5) at the surface. The slower liquid-phase reaction rates $k \sim 10^6 \text{ M}^{-1} \text{ s}^{-1}$ are used in the bulk phase and are based on past measurements in aqueous solution.⁸¹

For solutions below $\text{pH} = 10.8$, the primary oxidation product is HOI. Here, the proton-assisted decay of IOOO^- to HOI contributes to the overall loss of the intermediate and quickly becomes the dominate pathway below $\text{pH} 10$. To fully capture the reaction mechanism under neutral to acidic conditions, it is important to account for HOI/ IO^- speciation and associated reactions with I^- , OH^- , and I_2 . The fate of the HOI/ IO^- conjugate pair and their reactivities with iodine in solution have long been studied using spectroscopic techniques in an attempt to unravel this complex mechanism relevant to oxidation processes in disinfection, nuclear chemistry, and halogen chemistry in biological systems.⁸³⁻⁸⁸ Here, we include the steps in the model that are most relevant for the fate of I^- and the pH dependent reactivity in solution. This is summarized in the mechanistic overview in Fig. 5C. The most elementary of these reactions is simply the conjugate acid-base conversion



As introduced above for reaction (R2), we assume that recombination reactions occur at the proton neutralization rate^{79,80} in solution $k_{\text{neutral}} = 10^{11} \text{ M}^{-1} \text{ s}^{-1}$. This forward rate of (R6) then constrains the reverse dissociation rate of HOI using the $\text{pK}_a = 10.8$. Solution pH is then simulated in the model using the fixed dissociation rate coefficient for HOI along with the pseudo-first order reaction rate (*i.e.*, $k'_{\text{IO}} = [\text{H}^+] \times k_{\text{neutral}}$).

Having defined solution pH through the speciation of HOI, subsequent chemistry with I^- is addressed. We note that while the $\text{O}_3 + \text{HOI}$ reaction has been previously observed,⁸¹ this reaction is not included in the model since this reaction rate is 10^5 times slower than the $\text{O}_3 + \text{I}^-$ reaction and is significantly slower than other sinks for HOI. In the presence of I^- , HOI is quickly converted to I_2 under acidic and neutral pH.⁸⁶ Production of I_2 has been reported to occur through two distinct mechanisms that dominate at different pH values, both included in the model as (R7)–(R9). The first pathway involves the formation of reactive intermediate I_2OH^- (R7) which dissociates to I_2 through the proton-assisted reaction in (R8). The second formation pathways shown in (R9) involves the direct elimination of OH^- from HOI by I^- ^{89,90}



The forward and reverse rate coefficients for (R7)–(R9) with literature ref. 82, 89 and 90 are included in Table S1.† While step (R7) is independent of solution pH, reactions (R8) and (R9) both depend directly on pH explicitly through $[\text{H}^+]$ and $[\text{OH}^-]$. As in the association between IO^- and H^+ , pseudo-first order rate coefficients are calculated using $[\text{H}^+]$ and $[\text{OH}^-]$, as to avoid the computational cost of simulating these species directly. Lastly, once I_2 is produced in solution, the triiodide anion⁹¹ is formed by,



Simulation results, however, predict I_3^- concentrations under all pH conditions are negligible due to the rapid evaporative loss of I_2 , which is a dominant sink for iodine in small volumes (*i.e.*, droplets).

The complete set of reactions ((R1)–(R10)) as shown in Fig. 5B and C, is included in both the surface and bulk compartments of the kinetic model. As mentioned above, since surface partitioning behaviors of most species involved are unknown, we generally assume the surface partitioning of all solutes match the iodide anion, likely an overestimation of surface activity for most species due to the high surface propensity of I^- . Model sensitivity to absolute k_{desolv} and k_{solv} values is revisited in ESI Section SI-5† for the iodide ion only. We note that the sensitivity of these rates for other solutes is negligible due to the large proportion of bulk chemistry that drives the observed pH sensitivity. All rate coefficients for (R1)–(R10) are also assumed to be equal between the surface and bulk compartments, except for the difference in IO^- and IO_2^- reactivity mentioned above and noted in Table S1.†

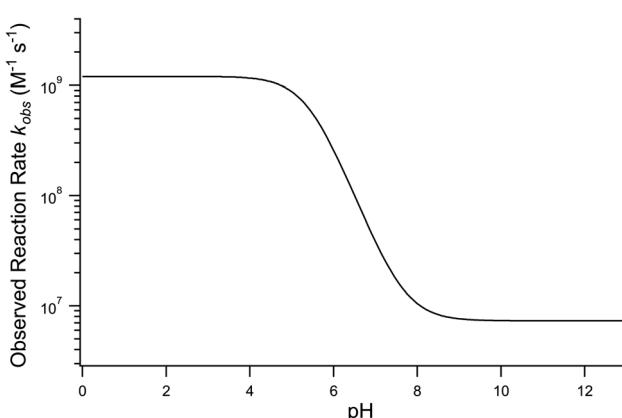


Fig. 6 Calculated pH dependence of the primary oxidation rate from combined reaction steps (R1)–(R3). Observed rate k_{obs} is calculated from eqn (9), following the framework of Liu *et al.*¹⁴ used in the bulk aqueous $\text{Br}^- + \text{O}_3$ reaction system.



5. Model results

The simulation results are compared to experiments in Fig. 2–4 (Section 2). Overall, the simulations capture the observed trends in the iodide-decay kinetics and how they depend upon $[O_3]$ and pH. As discussed further below, the initial rate scaling with $[O_3]$ can be understood from transport limitations and surface concentrations, while the pH dependence arises from a surface-rate pH dependence (*i.e.*, (R1)–(R3)) and the bulk reaction mechanism (*i.e.*, (R6)–(R9)). Although the simulations generally recreate the shape of iodide decay kinetics, the observed experimental decay shapes are not fully replicated. Under basic conditions where $[O_3] < 1$ ppm, simulations predict kinetics that appear more exponential in time, with experiments trending closer to a bi-exponential or linear time-dependence, albeit with larger experimental uncertainty. Although we cannot fully rationalize the origin of these deviations, the differences are relatively small, and the simulated trends generally agree with experiment. Some factors contributing to the shape of time-dependent kinetics will be examined in Section 6, with some potential reasons identified for the disagreement between model and experiment.

6. Analysis and discussion

Here we analyze the major features observed in the experimental and simulated kinetics to provide greater insight into the underlying chemical and physical processes governing this heterogenous reaction. First, in Section 6.1, the pH dependence of the reaction is examined in more detail, with attention given to identifying which chemical steps appear to determine the overall reaction rate and products observed. Section 6.2 examines how adsorbed ozone concentrations evolve during the reaction, where an approximate kinetic derivation is used to predict steady-state surface concentrations of ozone during reaction. Results from this derivation are then used in Section 6.3 to understand the behavior of the overall uptake coefficient for O_3 across a range of reactant concentrations. Results from experiment, simulation, and closed form expressions are compared with past measurements of this system to provide broader context to the current findings.

6.1. pH Dependent reaction kinetics and product distributions

As shown in Fig. 4, the experimental and simulated kinetics exhibit a strong pH dependence. In the model, this pH dependence has two distinct origins; the first is the pH rate dependence in the primary oxidation step ((R1)–(R3)) and the second is the network of secondary iodide loss channels involving H^+ and OH^- ((R6)–(R9)). Model results shown in Section SI-6† demonstrate that the secondary chemistry of iodide alone cannot achieve the model/experiment agreement shown in Fig. 2–4. As summarized in Fig. S10,† this suggests an additional mechanism is responsible for the kinetics observed when varying pH. We propose this additional pH dependence results

from the proton-assisted decay of an ozonide intermediate $IOOO^-$ as was first proposed (and later observed) for the analogous Br^- system.^{14,31} To the best of our knowledge, no direct evidence of the $IOOO^-$ intermediate has been reported, likely due to a short chemical lifetime. However, according to the parameterization of steps (R1)–(R3) introduced above, such an intermediate may be relatively long lived under strongly basic solutions (with lifetimes on the order of 10–100 μ s). On the other hand, as the surface lifetime of $IOOO^-$ and any degree of surface-stabilization is unknown, the overall lifetime may be significantly smaller if bulk solvation occurs on a timescale much faster than chemical decomposition. We also acknowledge that the bulk mechanism for the primary oxidation step may deviate significantly from the surface mechanism, and the kinetics of any intermediates involved may result in a different overall pH dependence. Recent theoretical work investigating the energetics of possible ozonide intermediates for aqueous iodide oxidation suggests a series of possible conformations of the intermediate.²⁶ The evolution of such conformations and the kinetics involved may again depend greatly on the presence of the interface and local pH.

Across the range of pH studied, a clear difference is observed in the overall product distribution predicted through the kinetic model. While the soluble, terminal product IO_3^- can be observed directly (Fig. S3†), volatile products such as I_2 and HOI go undetected in the experiments. ESI Section SI-7† provides the modeled product distribution across the pH range analyzed for an example ozone concentration of $[O_{3(g)}] = 820$ ppb. As expected, for simulations below the pK_a of HOI, the dominant product is I_2 , with a small fraction of HOI emitted. For strongly basic solution, IO_3^- dominates the product distribution. As observed in the iodate formation kinetics shown in SI-2,† the overall fraction of IO_3^- appears to decrease with decreasing $[O_3]$ below 1 ppm. This trend is not observed in the simulations (Fig. S3†), and the origin of the changing mechanism is currently unknown.

6.2. Surface-adsorbed ozone concentrations

The kinetic simulations indicate that surface-adsorbed ozone concentrations fall below the concentration predicted by Henry's law for a range of solution pH. Fig. 7 shows surface-adsorbed $[O_{3(ads)}]$ for three solution pH values at early times in the simulation. For the simulation at pH 13, $[O_{3(ads)}]$ rapidly approaches its Henry's law concentration at the surface, achieving a steady-state concentration that is ~50% of the expected Henry's law concentration. In the pH 8 simulation, however, $[O_{3(ads)}]$ finds its steady-state value that is ~8× smaller. In the acidic extreme at pH 3, $[O_{3(ads)}]$ is ~100× depleted from the Henry's law concentration, indicating that the primary oxidation step for $I^- + O_3$ is substantially limited by O_3 transport to the air–water interface. This transport limitation results from the diffusive rate to the surface being slower than the reactive loss of $O_{3(ads)}$, ultimately depleting both $O_{3(ads)}$ and $O_{3(diss)}$ in the model results. We expect more generally that surface reactions will be limited by ozone



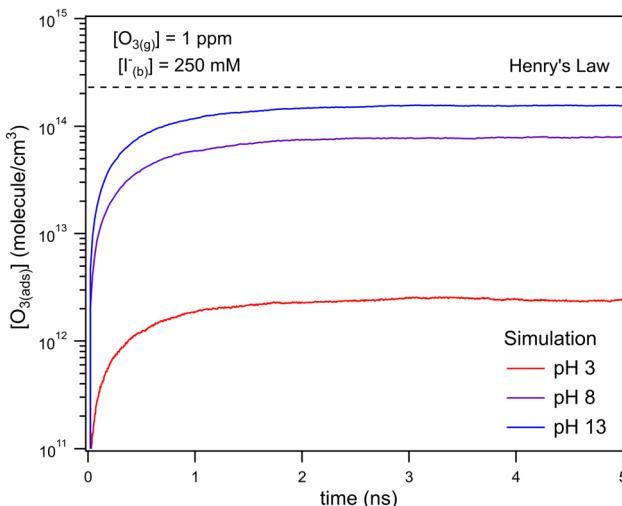


Fig. 7 Simulated $[O_3(\text{ads})]$ for the first 5 ns of reaction with $[O_3(\text{g})] = 1 \text{ ppm}$ and bulk iodide $[I_{(\text{b})}^-] = 250 \text{ mM}$. Adsorbed O_3 is seen reaching a steady state concentration after ~ 1 ns. Simulation results demonstrate $[O_3(\text{ads})]$ is $\sim 100\times$ smaller than the predicted Henry's Law concentration for a pH 3 solution due to the fast chemical loss rate on the surface.

diffusion for cases when $k'_{\text{rxn}} > k_{\text{diff}}$ where the pseudo-first order reaction rate is $k'_{\text{rxn}} = k_{\text{obs}} \cdot [I_{(\text{ads})}^-]$. This concept can be concisely expressed using the second Damköhler number, defined in this system as the ratio of the surface reaction rate to the characteristic gas-phase diffusion rate: $Da_{\text{II}} = k'_{\text{rxn}}/k_{\text{diff}}$. Using this notation, the oxidation reaction under conditions with $Da_{\text{II}} < 1$ will be limited by the surface reaction itself, whereas $Da_{\text{II}} > 1$ denotes the rate will be limited by gas-phase diffusion and will display surface depletion of O_3 . As expressed in eqn (6), k_{diff} depends on both the droplet radius and the equilibrium partitioning of O_3 which is captured in the model by use of L_{ads} . While we reserve a more thorough investigation of the droplet-size and molecular specificity of k_{diff} to future

as a vertical line in Fig. 8 indicates the point at which $k'_{\text{rxn}} = k_{\text{diff}}$, when the reactive loss rate of O_3 on the surface equals the rate of O_3 diffusion into the adsorption length. For ease of example, the pH dependent mechanism and adduct formation is not considered in Fig. 8 and a simple second order rate coefficient of $k_{\text{rxn}} = 1.2 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$ is used for the $I^- + O_3$ reaction. As all experiments performed in the present work utilize $[I^-] = 250 \text{ mM}$, Fig. 8 suggests for $k_{\text{obs}} \sim 10^9 \text{ M}^{-1} \text{ s}^{-1}$, $[O_3(\text{ads})]$ is greatly depleted from its Henry's Law concentration. Since the reaction under this condition is limited by O_3 transport, the kinetics scale as first-order in $[O_3]$ as demonstrated through the high linearity of initial rates *vs.* $[O_3]$ reported in Fig. 3.

The surface depletion of O_3 can also be computed analytically using a simple steady-state analysis of $[O_3(\text{ads})]$ and $[O_{(\text{diff})}]$ at early reaction times. As shown in Fig. 7, $[O_3(\text{ads})]$ approaches steady state after approximately 1 ns. Assuming that $[I_{(\text{ads})}^-]$ is equal to the Langmuir equilibrium concentration, the steady-state expressions for $[O_3(\text{ads})]$ and $[O_{(\text{diff})}]$ are

$$\begin{aligned} \frac{d[O_3(\text{ads})]}{dt} &= 0 \\ &= \sigma \cdot k_{\text{ads}} \cdot [O_3(\text{diff})] \cdot [\text{site}_{O_3}] \\ &\quad - k_{\text{rxn}} \cdot [O_3(\text{ads})] \cdot \frac{\Gamma_{I^-}^\infty}{\delta} \cdot \frac{K_{\text{eq}}^{I^-} \cdot [I_{(\text{b})}^-]}{1 + K_{\text{eq}}^{I^-} \cdot [I_{(\text{b})}^-]} \\ &\quad - k_{\text{des}} \cdot [O_3(\text{ads})] \end{aligned} \quad (10\text{a})$$

$$\begin{aligned} \frac{d[O_3(\text{diff})]}{dt} &= 0 \\ &= k_{\text{diff}} \cdot [O_3(\text{gas})] + k_{\text{des}} \cdot [O_3(\text{ads})] - k_{\text{diff}} \cdot [O_3(\text{diff})] \\ &\quad - \sigma \cdot k_{\text{ads}} \cdot [O_3(\text{diff})] \cdot [\text{site}_{O_3}] \end{aligned} \quad (10\text{b})$$

Rearranging eqn (10a) and (10b) while substituting for $[O_3(\text{diff})]$ produces an expression for $[O_3(\text{ads})]$ as a function of $[I_{(\text{b})}^-]$:

$$[O_3(\text{ads})] = \frac{\sigma \cdot k_{\text{ads}} \cdot k_{\text{diff}} \cdot [O_3(\text{gas})] \cdot [\text{site}_{O_3}]}{k_{\text{des}} \cdot k_{\text{diff}} + k_{\text{rxn}} \cdot \frac{\Gamma_{I^-}^\infty}{\delta} \cdot \frac{K_{\text{eq}}^{I^-} \cdot [I_{(\text{b})}^-]}{1 + K_{\text{eq}}^{I^-} \cdot [I_{(\text{b})}^-]} \cdot (k_{\text{diff}} + \sigma \cdot k_{\text{ads}} \cdot [\text{site}_{O_3}])} \quad (11)$$

work, the MD results in SI-3 Section B† show that the characteristic rate of surface equilibration ($\sim k_{\text{des}}$) is faster than k_{diff} for the particular droplet radius studied—indicating that transport limits of O_3 do not result from kinetic limitations. However, this is not generally the case for smaller droplet sizes, where kinetic limitations at the interface may become dominant.

Fig. 8 shows that the degree of ozone depletion at the interface is also dependent on the bulk $[I_{(\text{b})}^-]$. Simulations in this case are initialized for each given $[I_{(\text{b})}^-]$ by first computing the corresponding equilibrium $[I_{(\text{ads})}^-]$ as dictated by the Langmuir isotherm (eqn (7)). The simulation is then run until a steady state $[O_3(\text{ads})]$ is observed. The inflection point noted

Eqn (11) accounts for the surface depletion of O_3 by the surface reaction only and does not include solvation dynamics of O_3 , which quickly become insignificant when surface O_3 becomes depleted by reaction. Furthermore, we highlight that eqn (11) assumes adsorbed I^- equilibrates on a timescale much faster than the O_3 depletion dynamics such that $[I_{(\text{ads})}^-]$ is equal to the Langmuir equilibrium value calculated from eqn (7). Relaxing this assumption leads to a case where both surface reactants may be depleted. Results from eqn (11) as a function of bulk iodide concentration $[I_{(\text{b})}^-]$ are compared with simulation results in Fig. 8 for $[O_3(\text{g})] = 1 \text{ ppm}$. For simplicity, the results from eqn (11) in Fig. 8 utilize a simple bimolecular $I^- + O_3$ reaction rate of $k_{\text{rxn}} = 1.2 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$.



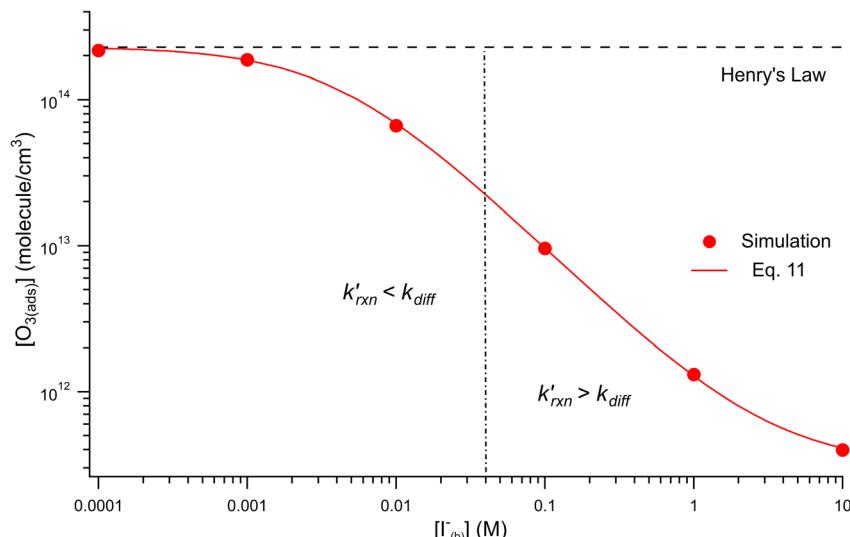


Fig. 8 Steady state $[O_{3(\text{ads})}]$ dependence on $[I_{(\text{b})}^-]$ from μM to M concentrations for $[O_{3(\text{g})}] = 1 \text{ ppm}$. Points denote simulation results from early steady-state analysis of ozone. The red curve shows steady-state expression from eqn (11). Results are compared to the Henry's Law concentration of $[O_{3(\text{ads})}]$. The dashed vertical line indicates the $[I_{(\text{b})}^-]$ where the chemical loss rate of O_3 at the surface ($k'_{\text{rxn}} = [I_{(\text{ads})}^-] \times k_{\text{rxn}}$) is equal to the simulated diffusion rate of O_3 (i.e., k_{diff}).

Output from eqn (11) and the simulation predict the same surface behavior for adsorbed O_3 , where reaction depletion becomes significant for bulk $[I_{(\text{b})}^-] > 1 \text{ mM}$. Only for $[I_{(\text{b})}^-] < 1 \text{ mM}$ does adsorbed ozone remain at its Henry's Law value. At these lower iodide concentrations, the multiphase kinetics are driven by bulk-reaction dynamics and should be accurately predicted using the closed formed analytical expressions derived in Wilson *et al.*²⁹

6.3. Uptake coefficients

In this section we compute uptake coefficients from the experimental kinetics shown in Fig. 3 and introduce an expression for computing uptake while accounting for surface depletion of O_3 . This approach builds on the recent work by Wilson *et al.*^{28,29} where uptake expressions were derived to describe reaction conditions when O_3 is depleted from the bulk solution, but not depleted from the microdroplet surface.

6.3.1 Experimental uptake calculation. Reactive uptake coefficients for O_3 are computed from the experimental kinetics shown in Fig. 3 using the initial observed reaction rate, k_{init} , to compute γ_{exp} ,⁹²

$$\gamma_{\text{exp}} = \frac{4 \cdot r \cdot S_{\text{rxn}} \cdot k_{\text{init}} \cdot [I_{(\text{b})}^-]_0}{3 \cdot [O_{3(\text{g})}] \cdot c}, \quad (12)$$

where r is the droplet radius, c is the mean molecular speed of O_3 in the gas phase, and S_{rxn} is an additional reaction-stoichiometry factor. The S_{rxn} factor is included in eqn (12) since the experiments yield k_{init} for the decay of iodide, whereas the expression for γ_{exp} refers explicitly to the reactive loss of O_3 . Secondary chemistry of I^- is coupled to solution pH, and therefore, measurements of k_{init} reflect chemical loss of I^- in addition to the primary $I^- + O_3$ reaction. Following the stoichiometry of reaction, however, allows uptake to be

compared across pH using S_{rxn} to properly compute γ_{exp} . Simulation results predict molecular iodine I_2 as the major product for pH 3 and 8 (demonstrated in SI-7†), indicating that for every $I^- + O_3$ reaction an additional I^- is consumed. This 1:2 ozone-to-iodide reaction equivalence is accounted for when calculating γ_{exp} by simply using $S_{\text{rxn}} = 0.5$. At pH 13, however, the only relevant loss channel for I^- is O_3 and as such, $S_{\text{rxn}} = 1$. We note that only conditions where $[O_{3(\text{g})}] > 1 \text{ ppm}$ were used in calculating uptake for pH 13 conditions given the apparent mechanistic complexity for sub-ppm concentrations of O_3 . Fig. 9 shows average γ_{exp} as a function of pH resulting from eqn (12), where a modest increase in γ_{exp} from $\sim 4 \times 10^{-4}$ at pH 13 to 2×10^{-3} at pH 3 is observed. Experimental results in Fig. 9 are compared to the values obtained from the analytical expression introduced in the following section.

6.3.2 Analytic expression for O_3 uptake. To compute reactive uptake in this system directly, we begin with the expressions for surface and bulk reactive uptake of O_3 in microdroplets as recently introduced by Wilson *et al.*:²⁹

$$\gamma_s = \frac{4 \cdot r \cdot k_{\text{rxn}} \cdot [O_{3(\text{ads})}] \cdot [I_{(\text{ads})}^-]}{3 \cdot c \cdot [O_{3(\text{g})}]} \cdot \frac{(r^3 - (r - \delta)^3)}{r^3}, \quad (13)$$

$$\gamma_b = \frac{4 \cdot r \cdot k_{\text{rxn}} \cdot [I_{(\text{b})}^-]}{3 \cdot c} \cdot \left[\frac{k_{\text{transport}} \cdot H_{\text{cc}}^{\text{gb}}}{k_{\text{rxn}} \cdot [I_{(\text{b})}^-] + k_{\text{transport}}} \right]. \quad (14)$$

The radial term in eqn (13) accounts for the relative scaling of surface to bulk volume with changing radius r and surface thickness δ . In eqn (14), $k_{\text{transport}}$ is a term introduced by Wilson *et al.*²⁹ that describes an overall ozone equilibration rate into the bulk solution that includes liquid diffusion and kinetic contributions. While previous work utilized the equilibrium



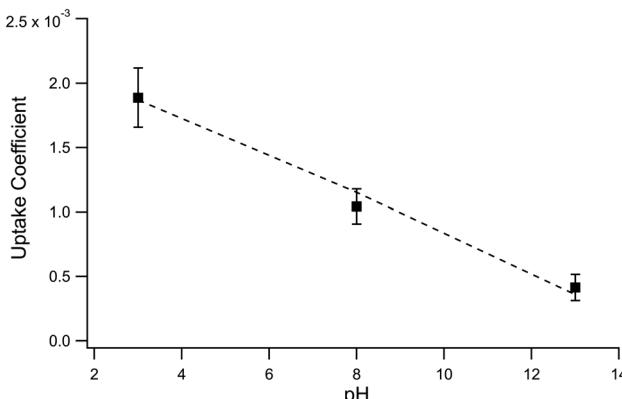


Fig. 9 Experimental reactive uptake coefficients for O_3 calculated from eqn (12) (points) compared with the analytical expression for surface uptake coefficients using eqn (13) (dashed line). Error for uptake coefficients reflects the standard deviation of uptake across different $[\text{O}_3]$ for each pH.

assumption $[\text{O}_{3(\text{ads})}] = H_{\text{cc}}^{\text{gb}} \times [\text{O}_{3(\text{g})}]$, the current investigation with iodide has shown $[\text{O}_{3(\text{ads})}]$ resides far below this equilibrium value. Therefore, eqn (11) which accounts for reactive depletion of $[\text{O}_{3(\text{ads})}]$ is used in eqn (13) rather than assuming the Henry's Law concentration.

The surface uptake coefficient calculated from eqn (13) is compared to experimental uptake coefficients in Fig. 9. The pH dependence of γ_s is captured by using the pH-scaling of k_{obs} discussed in Section 4.4 above and provided in eqn (9). As shown in Fig. 9, agreement is observed between experiment and the uptake expression in eqn (13), reinforcing the conclusion that the surface reaction dominates iodide loss under our experimental conditions. The agreement between experiment and uptake expressions further motivates the use of eqn (13) and (14) to predict uptake across a much larger concentration range. Fig. 10 shows uptake predictions from eqn (13) and (14) as well as the sum total $\gamma_{\text{total}} = \gamma_s + \gamma_b$ for a constant $[\text{O}_{3(\text{g})}] = 1 \text{ ppm}$ and droplet radius $r = 24 \mu\text{m}$. Average uptake from the experiments at $[\text{I}_{(\text{b})}^-] = 250 \text{ mM}$ and pH 3 are included for comparison in Fig. 10, as well as two limiting cases from resistor-based models discussed in more detail below.

For very low iodide concentrations shown in Fig. 10, below 100 nM, the uptake coefficient is dominated by the bulk reaction and is limited by the oxidation reaction rate in the bulk. At approximately $[\text{I}_{(\text{b})}^-] = 100 \text{ nM}$, uptake of O_3 by a bulk reaction starts becoming limited due to depletion of O_3 in the droplet interior and the surface reaction starts to dominate the reactive uptake. For increasing iodide concentrations, the surface reaction totally governs the uptake of O_3 . When iodide concentrations increase above 1 mM, results from Fig. 8 suggest $[\text{O}_{3(\text{ads})}]$ becomes greatly depleted due to gas-phase diffusion limitations. This depletion effect is similarly observed in Fig. 10 in the limiting of surface uptake to values $\sim 10^{-3}$.

Fig. 10 also compares the uptake coefficient expressions with commonly used resistor model limits shown with dashed lines. For very low $[\text{I}_{(\text{b})}^-]$, computed γ_{total} values agree with the dilute limit or "phase-mixed" case introduced by Schwartz *et al.*⁹³

$$\gamma_{\text{dilute}} = \frac{4 \cdot r \cdot H_{\text{cc}}^{\text{gb}}}{3 \cdot c} \cdot k_{\text{rxn}} \cdot [\text{I}_{(\text{b})}^-]. \quad (15)$$

As bulk iodide concentrations increase and γ_s dominates γ_{total} , results approach the "uptake controlled by fast reaction" case from Worsnop *et al.*⁵¹ or the "diffusion limit" case by Schwartz *et al.*⁹³

$$\gamma_{\text{diff}} = \frac{4 \cdot H_{\text{cc}}^{\text{gb}}}{c} \cdot \sqrt{D_{\text{O}_{3(\text{b})}} \cdot k_{\text{rxn}} \cdot [\text{I}_{(\text{b})}^-]} \cdot \left[\coth\left(\frac{r}{L_{\text{rxn}}}\right) - \frac{L_{\text{rxn}}}{r} \right], \quad (16)$$

where the reaction-diffusion length L_{rxn} is defined as $\sqrt{D_{\text{O}_{3(\text{b})}} / k_{\text{rxn}} \cdot [\text{I}_{(\text{b})}^-]}$. Deviation from the diffusion limit becomes pronounced for intermediate $[\text{I}_{(\text{b})}^-]$ as the surface first becomes enriched with $[\text{I}_{(\text{ads})}^-]$ relative to the bulk concentration, followed by the depletion of $[\text{O}_{3(\text{ads})}]$ for $[\text{I}_{(\text{b})}^-] > 1 \text{ mM}$.

6.4. Atmospheric implications & literature comparison

Here, we comment on the relevance of our findings to uptake of O_3 by seawater concentrations of I^- in droplets, and subsequently discuss our findings in context of previous literature results. The expressions for uptake in eqn (13) and (14) allow for the prediction of uptake in systems such as seawater and sea spray aerosol (SSA) interacting with the marine boundary layer, where near-surface seawater iodide concentrations range from $\sim 10\text{--}200 \text{ nM}$ (ref. 94) and typical ozone concentrations in the marine boundary layer are $\sim 20 \text{ ppb}$.^{95\text{--}98} For these concentrations, uptake of O_3 by a droplet with radius $r = 24 \mu\text{m}$ is on the order of 5×10^{-7} with bulk chemistry dominating the overall uptake. As previously reported, salt concentrations in SSA can increase dramatically upon emission from the sea surface, with I^- becoming potentially enhanced by factors up to ~ 1000 (ref. 99\text{--}101) and, in some measurements, making up a third of the total soluble iodine content of aerosol with particle radius $r < 500 \text{ nm}$.¹⁰² Additionally, rapid acidification of SSA¹⁰³ may contribute to an apparent acceleration of iodide oxidation by ozone relative to the pH of seawater ($\sim \text{pH } 8$). Under these concentrated and acidic conditions where iodide concentrations could approach $\sim 100 \mu\text{M}$, the expression obtained for γ_{total} predicts the uptake of ozone is driven almost exclusively by the surface reaction. However, we acknowledge that the current model construction employs a coarse-grain depiction of the droplet surface and bulk regions and therefore potentially overlooks important "sub-surface" dynamics which would affect the predicted transition in bulk *vs.* surface uptake. For reference we include the range of iodide concentrations encountered in seawater and SSA in the highlighted box in Fig. 10, where the transition from bulk to surface chemistry is observed.

While the model predictions presented in Fig. 10 suggest both surface and bulk contributions to the uptake of O_3 under relevant salt concentrations, we emphasize that results in Fig. 10 have resulted from evaluating eqn (13) and (14) for the droplet radius $r = 24 \mu\text{m}$ investigated in the experiments. For SSA and marine aerosol a broad range of sizes are encountered, where typical aerosols have radii $r < 10 \mu\text{m}$ and submicron particles make up the vast majority of aerosol by total number.



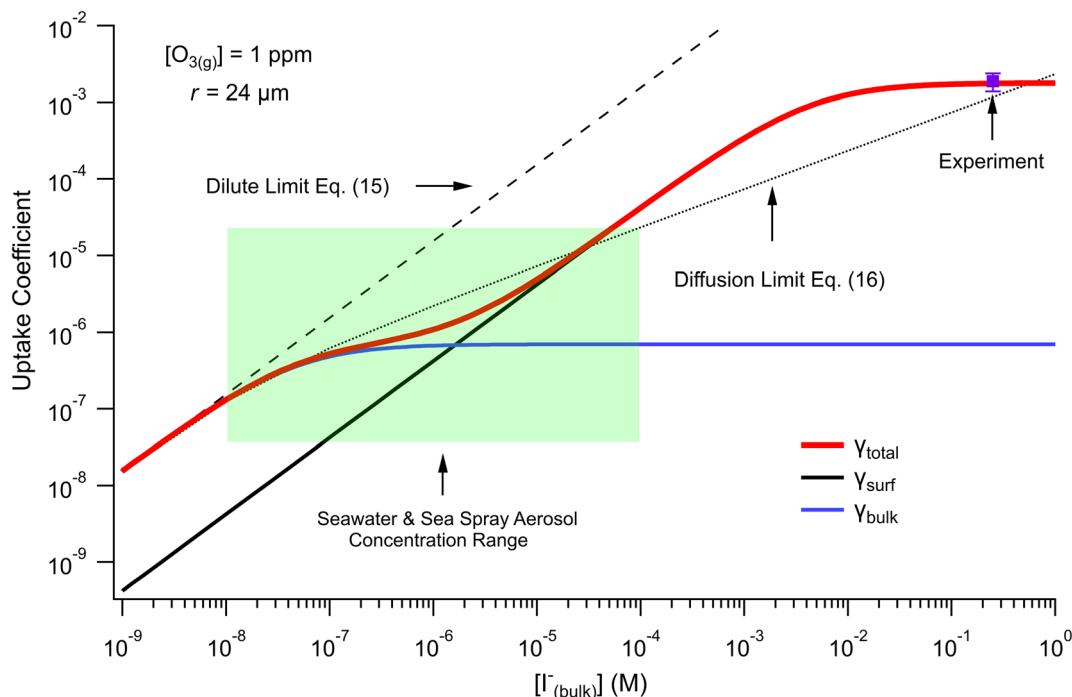


Fig. 10 Uptake coefficients computed from eqn (13) and (14) using $[O_{3(g)}] = 1 \text{ ppm}$ and droplet radius $r = 24 \mu\text{m}$. Total uptake is the sum of surface and bulk components. Results from the analytical approach are compared with two limiting cases from resistor-based uptake modeling shown as dashed and dotted lines. Experimental uptake at pH 8 is compared with the models.

For such small sizes, we expect the surface contribution to the overall uptake of O_3 by I^- to become even more significant due to the inverse radial scaling of the surface to volume ratio. On the other hand, eqn (13) and (14) predict the uptake of ozone by a macro-scale system such as the sea surface would be dominated by bulk chemistry using the same radial scaling argument. However, correctly calculating uptake for macro-scale systems utilizing the current model would require detailed knowledge of the system dimensions and the governing gas-phase mass transport steps for O_3 across large systems such as the marine boundary layer. As such, we reserve a more complete discussion of the size-dependence of γ_{total} under relevant concentrations to future work.

Uptake coefficients calculated from the analysis above and observed in experiments can also be compared to previous laboratory measurements of uptake of O_3 on aqueous iodide solutions and aerosol. One of the most comparable previous reports by Magi *et al.*¹⁰ measures uptake of O_3 using a droplet-train apparatus, with droplet sizes $\sim 100 \mu\text{m}$ and NaI concentrations ranging from 0.5 to 3 M. The reported uptake coefficient from this work under conditions 293 K and $[I^-] = 0.5 \text{ M}$ is $\gamma = 5 \times 10^{-3}$, only about 2–3 times greater than the maximum uptake measured in this work. However, as the droplet-train apparatus operates at reduced pressure (10–30 torr), deviations from uptake measured at ambient pressure may be expected. Additional relevant experiments performed by Ammann and coworkers^{20,21} use submicron sodium iodide aerosol to measure the uptake of ozone. The main difference between the conditions of these experiments and ours is particle size, as the present experiments all utilize much

larger microdroplets. Nevertheless, measurements on submicron aerosol composed of iodide/chloride mixtures by Rouvière *et al.*²⁰ yield uptake coefficients of $\gamma \sim 4.4 \times 10^{-3}$, which differ by a factor of ~ 2 from the largest uptake coefficients observed in our experiments. This reference also reports a relatively weak concentration dependence, where increasing $[I^-]$ from 0.9 M to 7.3 M increases the ozone reactive uptake coefficient by ~ 2.5 . This may point to a similar surface ozone depletion effect observed for sub-micron aerosol, although a more in-depth analysis of the experimental conditions used is needed to draw any specific conclusions from this past work.

Lastly, we note that while the conditions studied and modeled in the current work generally do not accurately represent conditions found in real-world systems, our current approach aims to identify and analyze the molecular-scale processes governing reactivity in simple multiphase systems. A deep understanding of these simple systems, in turn, builds the foundation for investigating the complex systems encountered in the environment.

7. Conclusion

Kinetics of iodide oxidation by ozone in single microdroplets are reported and analyzed with a recently developed multiphase model that describes individual surface and bulk processes occurring during the course of reaction. A kinetic model is constructed using literature references, basic assumptions about rate coefficients, and is further constrained using molecular simulations of O_3 and iodide at the air–water

interface. Use of the model provides insight into important surface mechanisms that govern the overall reactivity in microdroplets and provide a benchmark for developing more general expressions for O_3 uptake in aerosol. Experimental results in conjunction with modeling work suggest the $I^- + O_3$ reaction under the conditions studied occurs exclusively at the air–water interface. Nevertheless, we note that since the model employed discretizes the entire droplet into two compartments, a detailed account of sub-surface diffusion and reaction processes may be lacking which could be important for a clear accounting of surface *vs.* bulk reactivity. We anticipate future work will clarify these subsurface processes to fully understand the transition between surface and bulk reactivity.

A strong dependence on the consumption rate of I^- with bulk solution pH is observed, as previously reported for bulk solution. However, the modeled pH-dependent chemistry involving I^- in solution cannot fully explain the observed kinetics. To explain the observed pH dependence of the $I^- + O_3$ reaction we invoke an $IOOO^-$ intermediate, which is analogous to reaction intermediates found in the mechanism governing bromide oxidation by ozone at the air–water interface. A kinetic analysis also demonstrates a first-order dependence on $[O_3]_{\text{gas}}$ for all pH conditions. Simulation and analytical results suggest this reaction order originates from the depletion of ozone at the droplet surface during reaction for neutral to acidic conditions, and the linear scaling of surface adsorbed O_3 for mild to extremely basic conditions. Surface depletion of reagents appears to be a general feature of fast surface-reactions and may be relevant across scales for many systems involving uptake of trace gases. A kinetic expression for steady-state ozone concentrations on the droplet surface, eqn (11), is derived and subsequently used in a recently developed set of expressions, eqn (13) and (14), for predicting uptake coefficients in droplets due to both surface and bulk reactions. Surface uptake coefficients calculated using eqn (11) for surface-adsorbed O_3 indicate a large sensitivity to reactive depletion of species at the air–water interface, warranting the further study of the conditions under which reactants become depleted at interfaces.

Total uptake coefficients predicted by the analytical expressions obtained are analyzed across a range of concentrations and compared to the experimental uptake values of $\gamma \sim 10^{-3}$ and previous aerosol uptake experiments from the literature. The applicability of the present findings to reaction in seawater and sea spray aerosol is considered, with a brief discussion of the major findings and limitations of the modeling approach. Future work aims to expand the current approach to understand mass transport limitations in heterogeneous chemistry more broadly while investigating which key parameters control uptake coefficients and how these processes are related to reaction mechanisms at the droplet surface, where a number of factors such as partial solvation, pH, and reaction depletion may significantly alter expected uptake.

Data availability

Data are available from the corresponding author upon request.

Author contributions

A. M. P. and K. R. W. conceived of experiments and the kinetic model. A. M. P. performed experiments, analyzed data, and executed kinetic simulations. G. V. B. advised in construction of open-port sampling interface. K. P. and D. T. L. performed molecular simulations, and wrote description and analysis of results. A. M. P. and K. R. W. wrote manuscript with contributions from all authors.

Conflicts of interest

There are no conflicts to declare.

Acknowledgements

This work was supported by the Condensed Phase and Interfacial Molecular Science Program (CPIMS), in the Chemical Sciences Geosciences and Biosciences Division of the Office of Basic Energy Sciences of the U.S. Department of Energy under Contract No. DE-AC02-05CH11231. We thank Prof. Megan Willis (Colorado State University), Prof. Phill Geissler (1974–2022, UC Berkeley), Dr Amr Dodin (UC Berkeley), Liron Cohen (UC Berkeley/LBNL), Dr Shane Devlin (UNLV/LBNL), and Prof. Rich Saykally (UC Berkeley/LBNL) for helpful discussions.

References

- 1 A. Saiz-Lopez, J. M. C. Plane, A. R. Baker, L. J. Carpenter, R. von Glasow, J. C. Gómez Martín, G. McFiggans and R. W. Saunders, Atmospheric Chemistry of Iodine, *Chem. Rev.*, 2012, 112(3), 1773–1804, DOI: [10.1021/cr200029u](https://doi.org/10.1021/cr200029u).
- 2 L. J. Carpenter, Iodine in the Marine Boundary Layer, *Chem. Rev.*, 2003, 103(12), 4953–4962, DOI: [10.1021/cr0206465](https://doi.org/10.1021/cr0206465).
- 3 L. J. Carpenter, S. M. MacDonald, M. D. Shaw, R. Kumar, R. W. Saunders, R. Parthipan, J. Wilson and J. M. C. Plane, Atmospheric Iodine Levels Influenced by Sea Surface Emissions of Inorganic Iodine, *Nat. Geosci.*, 2013, 6(2), 108–111, DOI: [10.1038/ngeo1687](https://doi.org/10.1038/ngeo1687).
- 4 L. J. Carpenter, R. J. Chance, T. Sherwen, T. J. Adams, S. M. Ball, M. J. Evans, H. Hepach, L. D. J. Hollis, C. Hughes, T. D. Jickells, A. Mahajan, D. P. Stevens, L. Tiné and M. R. Wadley, Marine Iodine Emissions in a Changing World, *Proc. R. Soc. A*, 2021, 477(2247), 20200824, DOI: [10.1098/RSPA.2020.0824](https://doi.org/10.1098/RSPA.2020.0824).
- 5 J. C. Gómez Martín, T. R. Lewis, M. A. Blitz, J. M. C. Plane, M. Kumar, J. S. Francisco and A. Saiz-Lopez, A Gas-to-Particle Conversion Mechanism Helps to Explain Atmospheric Particle Formation through Clustering of Iodine Oxides, *Nat. Commun.*, 2020, 11(1), 4521, DOI: [10.1038/s41467-020-18252-8](https://doi.org/10.1038/s41467-020-18252-8).
- 6 Y. J. Tham, X.-C. He, Q. Li, C. A. Cuevas, J. Shen, J. Kalliokoski, C. Yan, S. Iyer, T. Lehmusjärvi, S. Jang, R. C. Thakur, L. Beck, D. Kemppainen, M. Olin, N. Sarnela, J. Mikkilä, J. Hakala, M. Marbouti, L. Yao, H. Li, W. Huang, Y. Wang, D. Wimmer, Q. Zha, J. Virkanen, T. G. Spain, S. O'Doherty, T. Jokinen,

F. Bianchi, T. Petäjä, D. R. Worsnop, R. L. Mauldin, J. Ovadnevaite, D. Ceburnis, N. M. Maier, M. Kulmala, C. O'Dowd, M. D. Maso, A. Saiz-Lopez and M. Sipilä, Direct Field Evidence of Autocatalytic Iodine Release from Atmospheric Aerosol, *Proc. Natl. Acad. Sci. U. S. A.*, 2021, **118**(4), 2021, DOI: [10.1073/PNAS.2009951118](https://doi.org/10.1073/PNAS.2009951118).

7 J. C. Gómez Martín, T. R. Lewis, A. D. James, A. Saiz-Lopez and J. M. C. Plane, Insights into the Chemistry of Iodine New Particle Formation: The Role of Iodine Oxides and the Source of Iodic Acid, *J. Am. Chem. Soc.*, 2022, **144**(21), 9240–9253, DOI: [10.1021/jacs.1c12957](https://doi.org/10.1021/jacs.1c12957).

8 L. Ganzeveld, D. Helmig, C. W. Fairall, J. Hare and A. Pozzer, Atmosphere-Ocean Ozone Exchange: A Global Modeling Study of Biogeochemical, Atmospheric, and Waterside Turbulence Dependencies, *Global Biogeochem. Cycles*, 2009, **23**(4), GB4021, DOI: [10.1029/2008GB003301](https://doi.org/10.1029/2008GB003301).

9 J. A. Garland, A. W. Elzerman and S. A. Penkett, The Mechanism for Dry Deposition of Ozone to Seawater Surfaces, *J. Geophys. Res.: Oceans*, 1980, **85**(C12), 7488–7492, DOI: [10.1029/JC085IC12P07488](https://doi.org/10.1029/JC085IC12P07488).

10 L. Magi, F. Schweitzer, C. Pallares, S. Cherif, P. Mirabel and C. George, Investigation of the Uptake Rate of Ozone and Methyl Hydroperoxide by Water Surfaces, *J. Phys. Chem. A*, 1997, **101**(27), 4943–4949, DOI: [10.1021/jp970646m](https://doi.org/10.1021/jp970646m).

11 J. H. Hu, Q. Shi, P. Davidovits, D. R. Worsnop, M. S. Zahniser and C. E. Kolb, Reactive Uptake of $\text{Cl}_2(\text{g})$ and $\text{Br}_2(\text{g})$ by Aqueous Surfaces as a Function of Br^- and I^- Ion Concentration: The Effect of Chemical Reaction at the Interface, *J. Phys. Chem.*, 1995, **99**(21), 8768–8776, DOI: [10.1021/j100021a050](https://doi.org/10.1021/j100021a050).

12 M. Bhujel, D. L. Marshall, T. A. Maccarone, I. B. McKinnon, J. A. Trevitt, G. d Silva, J. S. Blanksby and J. B. L. Poad, Gas Phase Reactions of Iodide and Bromide Anions with Ozone: Evidence for Stepwise and Reversible Reactions, *Phys. Chem. Chem. Phys.*, 2020, **22**(18), 9982–9989, DOI: [10.1039/D0CP01498B](https://doi.org/10.1039/D0CP01498B).

13 R. Teiwes, J. Elm, K. Handrup, E. P. Jensen, M. Bilde and H. B. Pedersen, Atmospheric Chemistry of Iodine Anions: Elementary Reactions of I^- , IO^- , and IO_2^- with Ozone Studied in the Gas-Phase at 300 K Using an Ion Trap, *Phys. Chem. Chem. Phys.*, 2018, **20**(45), 28606–28615, DOI: [10.1039/C8CP05721D](https://doi.org/10.1039/C8CP05721D).

14 Q. Liu, L. M. Schurter, C. E. Muller, S. Aloisio, J. S. Francisco and D. W. Margerum, Kinetics and Mechanisms of Aqueous Ozone Reactions with Bromide, Sulfite, Hydrogen Sulfite, Iodide, and Nitrite Ions, *Inorg. Chem.*, 2001, **40**(17), 4436–4442, DOI: [10.1021/ic000919j](https://doi.org/10.1021/ic000919j).

15 S. R. Schneider, P. S. J. Lakey, M. Shiraiwa and J. P. D. Abbatt, Reactive Uptake of Ozone to Simulated Seawater: Evidence for Iodide Depletion, *J. Phys. Chem. A*, 2020, **124**(47), 9844–9853, DOI: [10.1021/acs.jpca.0c08917](https://doi.org/10.1021/acs.jpca.0c08917).

16 S. R. Schneider, P. S. J. Lakey, M. Shiraiwa and J. P. D. Abbatt, Iodine Emission from the Reactive Uptake of Ozone to Simulated Seawater, *Environ. Sci.: Processes Impacts*, 2023, **25**(2), 254–263, DOI: [10.1039/D2EM00111J](https://doi.org/10.1039/D2EM00111J).

17 Y. Sakamoto, A. Yabushita, M. Kawasaki and S. Enami, Direct Emission of I_2 Molecule and IO Radical from the Heterogeneous Reactions of Gaseous Ozone with Aqueous Potassium Iodide Solution, *J. Phys. Chem. A*, 2009, **113**(27), 7707–7713, DOI: [10.1021/jp903486u](https://doi.org/10.1021/jp903486u).

18 J. Hoigné, H. Bader, W. R. Haag and J. Staehelin, Rate Constants of Reactions of Ozone with Organic and Inorganic Compounds in Water—III. Inorganic Compounds and Radicals, *Water Res.*, 1985, **19**(8), 993–1004, DOI: [10.1016/0043-1354\(85\)90368-9](https://doi.org/10.1016/0043-1354(85)90368-9).

19 E. Woods, C. A. Konys and S. R. Rossi, Photoemission of Iodide from Aqueous Aerosol Particle Surfaces, *J. Phys. Chem. A*, 2019, **123**(13), 2901–2907, DOI: [10.1021/acs.jpca.8b12323](https://doi.org/10.1021/acs.jpca.8b12323).

20 A. Rouvière, Y. Sosedova and M. Ammann, Uptake of Ozone to Deliquesced KI and Mixed KI/NaCl Aerosol Particles, *J. Phys. Chem. A*, 2010, **114**(26), 7085–7093, DOI: [10.1021/JP103257D](https://doi.org/10.1021/JP103257D).

21 A. Rouvière and M. Ammann, The Effect of Fatty Acid Surfactants on the Uptake of Ozone to Aqueous Halogenide Particles, *Atmos. Chem. Phys.*, 2010, **10**(23), 11489–11500, DOI: [10.5194/acp-10-11489-2010](https://doi.org/10.5194/acp-10-11489-2010).

22 R. Teiwes, J. Elm, M. Bilde and H. B. Pedersen, The Reaction of Hydrated Iodide $\text{I}(\text{H}_2\text{O})^-$ with Ozone: A New Route to IO_2^- Products, *Phys. Chem. Chem. Phys.*, 2019, **21**(32), 17546–17554, DOI: [10.1039/C9CP01734H](https://doi.org/10.1039/C9CP01734H).

23 H. B. Pedersen, J. Elm, C. H. Frederiksen, S. P. S. Jessen, R. Teiwes and M. Bilde, The Reaction of Isotope-Substituted Hydrated Iodide $\text{I}(\text{H}_2\text{O})^-$ with Ozone: The Reactive Influence of the Solvent Water Molecule, *Phys. Chem. Chem. Phys.*, 2020, **22**(34), 19080–19088, DOI: [10.1039/D0CP03219K](https://doi.org/10.1039/D0CP03219K).

24 G. C. Moreno, O. Gálvez, V. L.-A. Moreno, E. María Espildora-García and M. Teresa Baeza-Romero, A Revisit of the Interaction of Gaseous Ozone with Aqueous Iodide. Estimating the Contributions of the Surface and Bulk Reactions, *Phys. Chem. Chem. Phys.*, 2018, **20**(43), 27571–27584, DOI: [10.1039/C8CP04394A](https://doi.org/10.1039/C8CP04394A).

25 C. Moreno, M.-T. Baeza-Romero, M. Sanz, Ó. Gálvez, V. L. Arza, C. J. Ianni and E. Espildora, Iodide Conversion to Iodate in Aqueous and Solid Aerosols Exposed to Ozone, *Phys. Chem. Chem. Phys.*, 2020, **22**(10), 5625–5637, DOI: [10.1039/C9CP05601G](https://doi.org/10.1039/C9CP05601G).

26 Ó. Gálvez, M. T. Baeza-Romero, M. Sanz and L. F. Pacios, A Theoretical Study on the Reaction of Ozone with Aqueous Iodide, *Phys. Chem. Chem. Phys.*, 2016, **18**(11), 7651–7660, DOI: [10.1039/C5CP06440F](https://doi.org/10.1039/C5CP06440F).

27 C. Moreno and M. T. Baeza-Romero, A Kinetic Model for Ozone Uptake by Solutions and Aqueous Particles Containing I^- and Br^- , Including Seawater and Sea-Salt Aerosol, *Phys. Chem. Chem. Phys.*, 2019, **21**(36), 19835–19856, DOI: [10.1039/C9CP03430G](https://doi.org/10.1039/C9CP03430G).

28 M. D. Willis and K. R. Wilson, Coupled Interfacial and Bulk Kinetics Govern the Timescales of Multiphase Ozonolysis Reactions, *J. Phys. Chem. A*, 2022, **126**(30), 4991–5010, DOI: [10.1021/acs.jpca.2c03059](https://doi.org/10.1021/acs.jpca.2c03059).

29 K. R. Wilson, A. M. Prophet and M. D. Willis, A Kinetic Model for Predicting Trace Gas Uptake and Reaction, *J.*



Phys. Chem. A, 2022, **126**(40), 7291–7308, DOI: [10.1021/acs.jpca.2c03559](https://doi.org/10.1021/acs.jpca.2c03559).

30 R. Kaur Kohli, G. J. Van Berkel and J. F. Davies, An Open Port Sampling Interface for the Chemical Characterization of Levitated Microparticles, *Anal. Chem.*, 2022, **94**(8), 3441–3445, DOI: [10.1021/acs.analchem.1c05550](https://doi.org/10.1021/acs.analchem.1c05550).

31 L. Artiglia, J. Edebeli, F. Orlando, S. Chen, M.-T. Lee, P. Corral Arroyo, A. Gilgen, T. Bartels-Rausch, A. Kleibert, M. Vazdar, M. Andres Carignano, J. S. Francisco, P. B. Shepson, I. Gladich and M. Ammann, A Surface-Stabilized Ozonide Triggers Bromide Oxidation at the Aqueous Solution-Vapour Interface, *Nat. Commun.*, 2017, **8**(1), 700, DOI: [10.1038/s41467-017-00823-x](https://doi.org/10.1038/s41467-017-00823-x).

32 M. D. Willis, G. Rovelli and K. R. Wilson, Combining Mass Spectrometry of Picoliter Samples with a Multicompartment Electrodynamic Trap for Probing the Chemistry of Droplet Arrays, *Anal. Chem.*, 2020, **92**(17), 11943–11952, DOI: [10.1021/ACS.ANALCHEM.0C02343/SUPPL_FILE/ACOC02343_SI_001.PDF](https://doi.org/10.1021/ACS.ANALCHEM.0C02343).

33 M. I. Jacobs, J. F. Davies, L. Lee, R. D. Davis, F. Houle and K. R. Wilson, Exploring Chemistry in Microcompartments Using Guided Droplet Collisions in a Branched Quadrupole Trap Coupled to a Single Droplet, Paper Spray Mass Spectrometer, *Anal. Chem.*, 2017, **89**(22), 12511–12519, DOI: [10.1021/acs.analchem.7b03704](https://doi.org/10.1021/acs.analchem.7b03704).

34 J. F. Davies, Mass, Charge, and Radius of Droplets in a Linear Quadrupole Electrodynamic Balance, *Aerosol Sci. Technol.*, 2019, **53**(3), 309–320, DOI: [10.1080/02786826.2018.1559921](https://doi.org/10.1080/02786826.2018.1559921).

35 G. J. Van Berkel and V. Kertesz, An Open Port Sampling Interface for Liquid Introduction Atmospheric Pressure Ionization Mass Spectrometry, *Rapid Commun. Mass Spectrom.*, 2015, **29**(19), 1749–1756, DOI: [10.1002/rem.7274](https://doi.org/10.1002/rem.7274).

36 C. Liu, G. J. Van Berkel, P. Kovarik, J. B. Perot, V. Inguva and T. R. Covey, Fluid Dynamics of the Open Port Interface for High-Speed Nanoliter Volume Sampling Mass Spectrometry, *Anal. Chem.*, 2021, **93**(24), 8559–8567, DOI: [10.1021/acs.analchem.1c01312](https://doi.org/10.1021/acs.analchem.1c01312).

37 O. S. Ovchinnikova, D. Bhandari, M. Lorenz and G. J. Van Berkel, Transmission Geometry Laser Ablation into a Non-Contact Liquid Vortex Capture Probe for Mass Spectrometry Imaging, *Rapid Commun. Mass Spectrom.*, 2014, **28**(15), 1665–1673, DOI: [10.1002/rem.6946](https://doi.org/10.1002/rem.6946).

38 G. J. Van Berkel, V. Kertesz and H. Boeltz, Immediate Drop on Demand Technology (I-DOT) Coupled with Mass Spectrometry via an Open Port Sampling Interface, *Bioanalysis*, 2017, **9**(21), 1667–1679, DOI: [10.4155/bio-2017-0104](https://doi.org/10.4155/bio-2017-0104).

39 J. F. Cahill, J. Riba and V. Kertesz, Rapid, Untargeted Chemical Profiling of Single Cells in Their Native Environment, *Anal. Chem.*, 2019, **91**(9), 6118–6126, DOI: [10.1021/acs.analchem.9b00680](https://doi.org/10.1021/acs.analchem.9b00680).

40 R. Kaur Kohli, R. S. Reynolds, K. R. Wilson and J. F. Davies, Exploring the Influence of Particle Phase in the Ozonolysis of Oleic and Elaidic Acid, *Aerosol Sci. Technol.*, 2023, 1–18, DOI: [10.1080/02786826.2023.2226183](https://doi.org/10.1080/02786826.2023.2226183).

41 G. Lamoureux, E. Harder, I. V. Vorobyov, B. Roux and A. D. MacKerell Jr, A Polarizable Model of Water for Molecular Dynamics Simulations of Biomolecules, *Chem. Phys. Lett.*, 2006, **418**(1–3), 245–249.

42 W. D. Hinsberg and F. A. Houle, *Kinetoscope*, <https://hinsberg.net/kinetoscope/>, accessed 2022-02-15.

43 N. Heine, F. A. Houle and K. R. Wilson, Connecting the Elementary Reaction Pathways of Criegee Intermediates to the Chemical Erosion of Squalene Interfaces during Ozonolysis, *Environ. Sci. Technol.*, 2017, **51**(23), 13740–13748, DOI: [10.1021/acs.est.7b04197](https://doi.org/10.1021/acs.est.7b04197).

44 F. A. Houle, A. A. Wiegel and K. R. Wilson, Predicting Aerosol Reactivity Across Scales: From the Laboratory to the Atmosphere, *Environ. Sci. Technol.*, 2018, **52**(23), 13774–13781, DOI: [10.1021/acs.est.8b04688](https://doi.org/10.1021/acs.est.8b04688).

45 K. R. Wilson, A. M. Prophet, G. Rovelli, M. D. Willis, R. J. Rapf and M. I. Jacobs, A Kinetic Description of How Interfaces Accelerate Reactions in Micro-Compartments, *Chem. Sci.*, 2020, **11**(32), 8533–8545, DOI: [10.1039/DOSC03189E](https://doi.org/10.1039/DOSC03189E).

46 R. Vácha, P. Slavíček, M. Mucha, B. J. Finlayson-Pitts and P. Jungwirth, Adsorption of Atmospherically Relevant Gases at the Air/Water Interface: Free Energy Profiles of Aqueous Solvation of N₂, O₂, O₃, OH, H₂O, HO₂, and H₂O₂, *J. Phys. Chem. A*, 2004, **108**(52), 11573–11579, DOI: [10.1021/jp046268k](https://doi.org/10.1021/jp046268k).

47 R. Sander, Compilation of Henry's Law Constants (Version 4.0) for Water as Solvent, *Atmos. Chem. Phys.*, 2015, **15**(8), 4399–4981, DOI: [10.5194/acp-15-4399-2015](https://doi.org/10.5194/acp-15-4399-2015).

48 N. M. Panich, B. G. Ershov, A. F. Seliverstov and A. G. Basiev, Ozone Solubility in Concentrated Aqueous Solutions of Salts, *Russ. J. Appl. Chem.*, 2007, **80**(11), 1812–1815, DOI: [10.1134/S107042707207110067](https://doi.org/10.1134/S107042707207110067).

49 I. Langmuir, The Adsorption of Gases on Plane Surfaces of Glass, Mica and Platinum, *J. Am. Chem. Soc.*, 1918, **40**(9), 1361–1403, DOI: [10.1021/ja02242a004](https://doi.org/10.1021/ja02242a004).

50 J. Vieceli, M. Roeselová, N. Potter, L. X. Dang, B. C. Garrett and D. J. Tobias, Molecular Dynamics Simulations of Atmospheric Oxidants at the Air–Water Interface: Solvation and Accommodation of OH and O₃, *J. Phys. Chem. B*, 2005, **109**(33), 15876–15892, DOI: [10.1021/jp051361+](https://doi.org/10.1021/jp051361+).

51 D. R. Worsnop, J. W. Morris, Q. Shi, P. Davidovits and C. E. Kolb, A Chemical Kinetic Model for Reactive Transformations of Aerosol Particles, *Geophys. Res. Lett.*, 2002, **29**(20), 57, DOI: [10.1029/2002GL015542](https://doi.org/10.1029/2002GL015542).

52 G. M. Nathanson, P. Davidovits, D. R. Worsnop and C. E. Kolb, Dynamics and Kinetics at the Gas–Liquid Interface, *J. Phys. Chem.*, 1996, **100**(31), 13007–13020, DOI: [10.1021/jp953548e](https://doi.org/10.1021/jp953548e).

53 M. E. Saecker and G. M. Nathanson, Collisions of Protic and Aprotic Gases with Hydrogen Bonding and Hydrocarbon Liquids, *J. Chem. Phys.*, 1993, **99**(9), 7056–7075, DOI: [10.1063/1.465425](https://doi.org/10.1063/1.465425).

54 C. F. Clement, M. Kulmala and T. Vesala, Theoretical Consideration on Sticking Probabilities, *J. Aerosol Sci.*, 1996, **27**(6), 869–882, DOI: [10.1016/0021-8502\(96\)00032-8](https://doi.org/10.1016/0021-8502(96)00032-8).



55 S. E. Schwartz and J. E. Freiberg, Mass-Transport Limitation to the Rate of Reaction of Gases in Liquid Droplets: Application to Oxidation of SO₂ in Aqueous Solutions, *Atmos. Environ.*, 1981, **15**(7), 1129–1144, DOI: [10.1016/0004-6981\(81\)90303-6](https://doi.org/10.1016/0004-6981(81)90303-6).

56 J. H. Seinfeld and S. N. Pandis, Mass Transfer Aspects of Atmospheric Chemistry, in *Atmospheric Chemistry and Physics: From Air Pollution to Climate Change*, John Wiley & Sons, Inc., 1998.

57 F. Jin, R. Balasubramaniam and K. J. Stebe, Surfactant Adsorption to Spherical Particles: The Intrinsic Length Scale Governing the Shift from Diffusion to Kinetic-Controlled Mass Transfer, *J. Adhes.*, 2004, **80**(9), 773–796, DOI: [10.1080/00218460490480770](https://doi.org/10.1080/00218460490480770).

58 P. Jungwirth and D. J. Tobias, Ions at the Air/Water Interface, *J. Phys. Chem. B*, 2002, **106**(25), 6361–6373, DOI: [10.1021/jp020242g](https://doi.org/10.1021/jp020242g).

59 A. C. Stern, M. D. Baer, C. J. Mundy and D. J. Tobias, Thermodynamics of Iodide Adsorption at the Instantaneous Air-Water Interface, *J. Chem. Phys.*, 2013, **138**(11), 114709, DOI: [10.1063/1.4794688](https://doi.org/10.1063/1.4794688).

60 L. Piatkowski, Z. Zhang, E. H. G. Backus, H. J. Bakker and M. Bonn, Extreme Surface Propensity of Halide Ions in Water, *Nat. Commun.*, 2014, **5**(1), 4083, DOI: [10.1038/ncomms5083](https://doi.org/10.1038/ncomms5083).

61 D. I. Reeser and D. J. Donaldson, Influence of Water Surface Properties on the Heterogeneous Reaction between O₃(g) and I(Aq)–, *Atmos. Environ.*, 2011, **45**(34), 6116–6120, DOI: [10.1016/J.ATMOSENV.2011.08.042](https://doi.org/10.1016/J.ATMOSENV.2011.08.042).

62 S. N. Wren and D. J. Donaldson, Glancing-Angle Raman Spectroscopic Probe for Reaction Kinetics at Water Surfaces, *Phys. Chem. Chem. Phys.*, 2010, **12**(11), 2648–2654, DOI: [10.1039/B9222254E](https://doi.org/10.1039/B9222254E).

63 R. Weber, B. Winter, P. M. Schmidt, W. Widdra, I. V. Hertel, M. Dittmar and M. Faubel, Photoemission from Aqueous Alkali-Metal-Iodide Salt Solutions Using EUV Synchrotron Radiation, *J. Phys. Chem. B*, 2004, **108**(15), 4729–4736, DOI: [10.1021/jp030776x](https://doi.org/10.1021/jp030776x).

64 B. Credidio, M. Pugini, S. Malerz, F. Trinter, U. Hergenhahn, I. Wilkinson, S. Thürmer and B. Winter, Quantitative Electronic Structure and Work-Function Changes of Liquid Water Induced by Solute, *Phys. Chem. Chem. Phys.*, 2022, **24**(3), 1310–1325, DOI: [10.1039/D1CP03165A](https://doi.org/10.1039/D1CP03165A).

65 P. B. Petersen and R. J. Saykally, Probing the Interfacial Structure of Aqueous Electrolytes with Femtosecond Second Harmonic Generation Spectroscopy, *J. Phys. Chem. B*, 2006, **110**(29), 14060–14073, DOI: [10.1021/jp0601825](https://doi.org/10.1021/jp0601825).

66 P. B. Petersen, J. C. Johnson, K. P. Knutson and R. J. Saykally, Direct Experimental Validation of the Jones-Ray Effect, *Chem. Phys. Lett.*, 2004, **397**(1–3), 46–50, DOI: [10.1016/j.cplett.2004.08.048](https://doi.org/10.1016/j.cplett.2004.08.048).

67 L. X. Dang, Computational Study of Ion Binding to the Liquid Interface of Water, *J. Phys. Chem. B*, 2002, **106**(40), 10388–10394, DOI: [10.1021/jp021871t](https://doi.org/10.1021/jp021871t).

68 Y. Liu, Is the Free Energy Change of Adsorption Correctly Calculated?, *J. Chem. Eng. Data*, 2009, **54**(7), 1981–1985, DOI: [10.1021/je800661q](https://doi.org/10.1021/je800661q).

69 G. Olivieri, K. M. Parry, R. D'Auria, D. J. Tobias and M. A. Brown, Specific Anion Effects on Na⁺ Adsorption at the Aqueous Solution-Air Interface: MD Simulations, SESSA Calculations, and Photoelectron Spectroscopy Experiments, *J. Phys. Chem. B*, 2018, **122**(2), 910–918, DOI: [10.1021/acs.jpcb.7b06981](https://doi.org/10.1021/acs.jpcb.7b06981).

70 R. D'Auria and D. J. Tobias, Relation between Surface Tension and Ion Adsorption at the Air-Water Interface: A Molecular Dynamics Simulation Study, *J. Phys. Chem. A*, 2009, **113**(26), 7286–7293, DOI: [10.1021/jp810488p](https://doi.org/10.1021/jp810488p).

71 N. Ottosson, M. Faubel, S. E. Bradforth, P. Jungwirth and B. Winter, Photoelectron Spectroscopy of Liquid Water and Aqueous Solution: Electron Effective Attenuation Lengths and Emission-Angle Anisotropy, *J. Electron Spectrosc. Relat. Phenom.*, 2010, **177**(2), 60–70, DOI: [10.1016/j.elspec.2009.08.007](https://doi.org/10.1016/j.elspec.2009.08.007).

72 S. Ou, Y. Hu, S. Patel and H. Wan, Spherical Monovalent Ions at Aqueous Liquid–Vapor Interfaces: Interfacial Stability and Induced Interface Fluctuations, *J. Phys. Chem. B*, 2013, **117**(39), 11732–11742, DOI: [10.1021/jp406001b](https://doi.org/10.1021/jp406001b).

73 S. Saha, S. Roy, P. Mathi and J. A. Mondal, Polyatomic Iodine Species at the Air-Water Interface and Its Relevance to Atmospheric Iodine Chemistry: An HD-VSFG and Raman-MCR Study, *J. Phys. Chem. A*, 2019, **123**(13), 2924–2934, DOI: [10.1021/acs.jpca.9b00828](https://doi.org/10.1021/acs.jpca.9b00828).

74 S. Roy and J. A. Mondal, “Breaking” and “Making” of Water Structure at the Air/Water-Electrolyte (NaXO₃; X = Cl, Br, I) Interface, *J. Phys. Chem. Lett.*, 2021, **12**(7), 1955–1960, DOI: [10.1021/acs.jpclett.0c03827](https://doi.org/10.1021/acs.jpclett.0c03827).

75 T. Huthwelker and T. Peter, Analytical Description of Gas Transport across an Interface with Coupled Diffusion in Two Phases, *J. Chem. Phys.*, 1996, **105**(4), 1661–1667, DOI: [10.1063/1.472025](https://doi.org/10.1063/1.472025).

76 A. A. Wiegel, M. J. Liu, W. D. Hinsberg, K. R. Wilson and F. A. Houle, Diffusive Confinement of Free Radical Intermediates in the OH Radical Oxidation of Semisolid Aerosols, *Phys. Chem. Chem. Phys.*, 2017, **19**(9), 6814–6830, DOI: [10.1039/C7CP00696A](https://doi.org/10.1039/C7CP00696A).

77 I. Sanemasa, T. Kobayashi, C. Y. Piao and T. Deguchi, Equilibrium Solubilities of Iodine Vapor in Water, *Bull. Chem. Soc. Jpn.*, 1984, **57**(5), 1352–1357, DOI: [10.1246/bcsj.57.1352](https://doi.org/10.1246/bcsj.57.1352).

78 C.-C. Lin, Volatility of Iodine in Dilute Aqueous Solutions, *J. Inorg. Nucl. Chem.*, 1981, **43**(12), 3229–3238, DOI: [10.1016/0022-1902\(81\)80094-2](https://doi.org/10.1016/0022-1902(81)80094-2).

79 G. C. Barker and D. C. Sammon, Kinetics of Proton-Transfer Reactions, *Nature*, 1967, **213**(5071), 65–66, DOI: [10.1038/213065b0](https://doi.org/10.1038/213065b0).

80 M. Eigen and L. De Maeyer, Untersuchungen über die Kinetik der Neutralisation. I, Z. Elektrochem. Ber. Bunsenges. Phys. Chem., 1955, **59**(10), 986–993, DOI: [10.1002/bbpc.19550591020](https://doi.org/10.1002/bbpc.19550591020).



81 Y. Bichsel and U. von Gunten, Oxidation of Iodide and Hypoiodous Acid in the Disinfection of Natural Waters, *Environ. Sci. Technol.*, 1999, **33**(22), 4040–4045, DOI: [10.1021/es990336c](https://doi.org/10.1021/es990336c).

82 Y. Bichsel, Behavior of Iodine Species in Oxidative Processes during Drinking Water Treatment. *Doctoral thesis*, ETH Zurich, 2000, DOI: [10.3929/ethz-a-003861380](https://doi.org/10.3929/ethz-a-003861380).

83 T. L. Allen and R. M. Keefer, The Formation of Hypoiodous Acid and Hydrated Iodine Cation by the Hydrolysis of Iodine, *J. Am. Chem. Soc.*, 1955, **77**(11), 2957–2960, DOI: [10.1021/ja01616a008](https://doi.org/10.1021/ja01616a008).

84 C. H. Li and C. F. White, Kinetics of Hypoiodite Decomposition, *J. Am. Chem. Soc.*, 1943, **65**(3), 335–339, DOI: [10.1021/ja01243a011](https://doi.org/10.1021/ja01243a011).

85 G. V. Buxton and Q. G. Mulazzani, On the Hydrolysis of Iodine in Alkaline Solution: A Radiation Chemical Study, *Radiat. Phys. Chem.*, 2007, **76**(6), 932–940, DOI: [10.1016/j.radphyschem.2006.06.009](https://doi.org/10.1016/j.radphyschem.2006.06.009).

86 M. Eigen and K. Kustin, The Kinetics of Halogen Hydrolysis, *J. Am. Chem. Soc.*, 1962, **84**(8), 1355–1361, DOI: [10.1021/ja00867a005](https://doi.org/10.1021/ja00867a005).

87 J. C. Wren, J. Paquette, S. Sunder and B. L. Ford, Iodine Chemistry in the +1 Oxidation State. II. A Raman and Uv-Visible Spectroscopic Study of the Disproportionation of Hypoiodite in Basic Solutions, *Can. J. Chem.*, 1986, **64**(12), 2284–2296, DOI: [10.1139/v86-375](https://doi.org/10.1139/v86-375).

88 J. Paquette, J. C. Wren, S. Sunder and B. L. Ford, *The Disproportionation of Iodine (I)*, United Kingdom, 1986, pp. 29–45, 0-7058-1073-9.

89 K. Sebők-Nagy and T. Körtvélyesi, Kinetics and Mechanism of the Hydrolytic Disproportionation of Iodine, *Int. J. Chem. Kinet.*, 2004, **36**(11), 596–602, DOI: [10.1002/kin.20033](https://doi.org/10.1002/kin.20033).

90 I. Lengyel, I. R. Epstein and K. Kustin, Kinetics of Iodine Hydrolysis, *Inorg. Chem.*, 1993, **32**(25), 5880–5882, DOI: [10.1021/ic00077a036](https://doi.org/10.1021/ic00077a036).

91 D. A. Palmer, R. W. Ramette and R. E. Mesmer, Triiodide Ion Formation Equilibrium and Activity Coefficients in Aqueous Solution, *J. Solution Chem.*, 1984, **13**(9), 673–683, DOI: [10.1007/BF00650374](https://doi.org/10.1007/BF00650374).

92 J. D. Smith, J. H. Kroll, C. D. Cappa, D. L. Che, C. L. Liu, M. Ahmed, S. R. Leone, D. R. Worsnop and K. R. Wilson, The Heterogeneous Reaction of Hydroxyl Radicals with Sub-Micron Squalane Particles: A Model System for Understanding the Oxidative Aging of Ambient Aerosols, *Atmos. Chem. Phys.*, 2009, **9**(9), 3209–3222, DOI: [10.5194/acp-9-3209-2009](https://doi.org/10.5194/acp-9-3209-2009).

93 S. E. Schwartz, *Mass-Transport Considerations Pertinent to Aqueous Phase Reactions of Gases in Liquid-Water Clouds*, Springer Berlin Heidelberg, Berlin, Heidelberg, 1986, pp. 415–471.

94 R. Chance, A. R. Baker, L. Carpenter and T. D. Jickells, The Distribution of Iodide at the Sea Surface, *Environ. Sci.: Processes Impacts*, 2014, **16**(8), 1841–1859, DOI: [10.1039/C4EM00139](https://doi.org/10.1039/C4EM00139).

95 J. E. Johnson, R. H. Gammon, J. Larsen, T. S. Bates, S. J. Oltmans and J. C. Farmer, Ozone in the Marine Boundary Layer over the Pacific and Indian Oceans: Latitudinal Gradients and Diurnal Cycles, *J. Geophys. Res.: Atmos.*, 1990, **95**(D8), 11847–11856, DOI: [10.1029/JD095iD08p11847](https://doi.org/10.1029/JD095iD08p11847).

96 H. B. Singh, G. L. Gregory, B. Anderson, E. Browell, G. W. Sachse, D. D. Davis, J. Crawford, J. D. Bradshaw, R. Talbot, D. R. Blake, D. Thornton, R. Newell and J. Merrill, Low Ozone in the Marine Boundary Layer of the Tropical Pacific Ocean: Photochemical Loss, Chlorine Atoms, and Entrainment, *J. Geophys. Res.: Atmos.*, 1996, **101**(D1), 1907–1917, DOI: [10.1029/95JD01028](https://doi.org/10.1029/95JD01028).

97 R. R. Dickerson, K. P. Rhoads, T. P. Carsey, S. J. Oltmans, J. P. Burrows and P. J. Crutzen, Ozone in the Remote Marine Boundary Layer: A Possible Role for Halogens, *J. Geophys. Res.: Atmos.*, 1999, **104**(D17), 21385–21395, DOI: [10.1029/1999JD900023](https://doi.org/10.1029/1999JD900023).

98 K. Watanabe, Y. Nojiri and S. Kariya, Measurements of Ozone Concentrations on a Commercial Vessel in the Marine Boundary Layer over the Northern North Pacific Ocean, *J. Geophys. Res.: Atmos.*, 2005, **110**, D11310, DOI: [10.1029/2004JD005514](https://doi.org/10.1029/2004JD005514).

99 F. Y. B. Seto and R. A. Duce, A Laboratory Study of Iodine Enrichment on Atmospheric Sea-Salt Particles Produced by Bubbles, *J. Geophys. Res.*, 1972, **77**(27), 5339–5349, DOI: [10.1029/JC077i027p05339](https://doi.org/10.1029/JC077i027p05339).

100 R. A. Duce and E. J. Hoffman, Chemical Fractionation at the Air/Sea Interface, *Annu. Rev. Earth Planet. Sci.*, 1976, **4**(1), 187–228, DOI: [10.1146/annurev.ea.04.050176.001155](https://doi.org/10.1146/annurev.ea.04.050176.001155).

101 B. S. Gilfedder, S. C. Lai, M. Petri, H. Biester and T. Hoffmann, Iodine Speciation in Rain, Snow and Aerosols, *Atmos. Chem. Phys.*, 2008, **8**(20), 6069–6084, DOI: [10.5194/acp-8-6069-2008](https://doi.org/10.5194/acp-8-6069-2008).

102 J. C. Gómez Martín, A. Saiz-Lopez, C. A. Cuevas, A. R. Baker and R. P. Fernández, On the Speciation of Iodine in Marine Aerosol, *J. Geophys. Res.: Atmos.*, 2022, **127**(4), e2021JD036081, DOI: [10.1029/2021JD036081](https://doi.org/10.1029/2021JD036081).

103 K. J. Angle, D. R. Crocker, R. M. C. Simpson, K. J. Mayer, L. A. Garofalo, A. N. Moore, S. L. Mora Garcia, V. W. Or, S. Srinivasan, M. Farhan, J. S. Sauer, C. Lee, M. A. Pothier, D. K. Farmer, T. R. Martz, T. H. Bertram, C. D. Cappa, K. A. Prather and V. H. Grassian, Acidity across the Interface from the Ocean Surface to Sea Spray Aerosol, *Proc. Natl. Acad. Sci. U. S. A.*, 2021, **118**(2), e2018397118, DOI: [10.1073/pnas.2018397118](https://doi.org/10.1073/pnas.2018397118).

