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Emerging investigator series: aqueous oxidation of isoprene-derived organic aerosol species as a source of atmospheric formic and acetic acids†

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Atmospheric chemistry models generally assume organic aerosol (OA) to be photochemically inert. Recent mechanisms for the oxidation of biogenic isoprene, a major source of secondary organic aerosol (iSOA), produce excessive OA in the absence of subsequent OA reactivity. At the same time, models underestimate atmospheric concentrations of formic and acetic acids for which OA degradation could provide a source. Here we show that the aqueous photooxidation of an isoprene-derived organosulfate (2-methyltriolsulfate or MTS), an important iSOA component, produces formic and acetic acids in high yields and at timescales competitive with deposition. Experimental data are well fit by a kinetic model in which three sequential oxidation reactions of the isoprene organosulfate produce two molar equivalents of formic acid and one of acetic acid. We incorporate this chemistry and that of 2-methyltetrol, another ubiquitous iSOA component, into the GEOS-Chem global atmospheric chemistry model. Simulations show that photooxidation and subsequent revolatilization of this iSOA may account for up to half of total iSOA loss globally, producing 4 Tg a^{-1} each of formic and acetic acids. This reduces model biases in gas-phase formic acid and total organic aerosol over the Southeast United States in summer by ~30% and 60% respectively. While our study shows the importance of adding iSOA photochemical sinks into atmospheric models, uncertainties remain that warrant further study. In particular, improved understanding of reaction dependencies on particle characteristics and concentrations of particle-phase OH and other oxidants are needed to better simulate the effects of this chemistry on the atmospheric budgets of organic acids and iSOA.

Environmental significance

Atmospheric chemistry models overestimate the production of organic aerosol (OA) from isoprene while simultaneously underestimating gas-phase formic and acetic acid concentrations in isoprene-rich regions. We perform aqueous photooxidation experiments on an isoprene-derived organosulfate – a key isoprene-OA constituent – and show that its oxidation produces formic and acetic acids in high yields on atmospherically relevant timescales. We integrate this chemistry, along with experimental work on other isoprene-derived OA species, into a global atmospheric chemistry model and demonstrate that these reactions improve both model overestimates of isoprene-derived OA and underestimates of gas-phase organic acids. Constraining reactive fates of other OA species and reducing uncertainties in modeling aqueous atmospheric OH will further improve our understanding of global budgets of both OA and organic acids.

Introduction

Formic and acetic acids are ubiquitous and abundant in Earth's troposphere, and play important roles in determining the acidity of cloudwater and rainfall.[1–6] By affecting the acidity of aerosol particles as well, they influence cloud nucleation[7] and the aqueous chemistry of other aerosol constituents.[8,9] Formic and acetic acids are directly emitted to the atmosphere by plants,[10–12] fires,[13–15] soils,[16–18] ice,[19] and anthropogenic activities,[20–23] and can also be produced in the atmosphere

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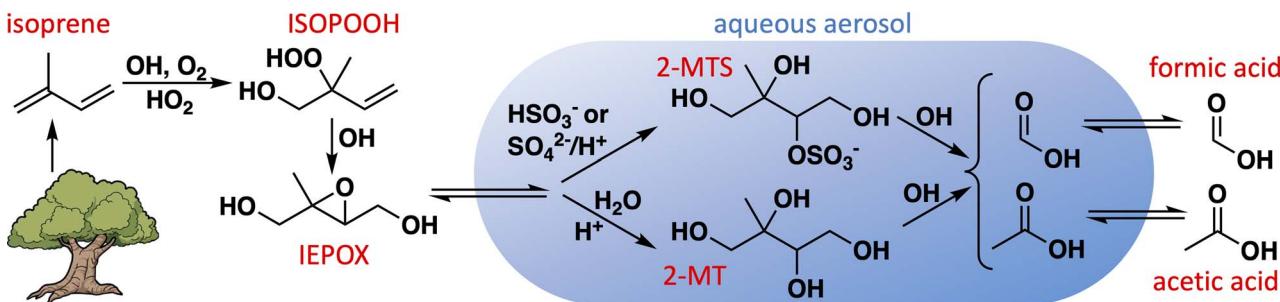
Fig. 1 Route to aqueous production of formic and acetic acids *via* aqueous uptake of isoprene-derived epoxydiol (IEPOX).

through the gas-phase chemistry of biogenic,[5,24–29] pyrogenic,[30–33] and anthropogenic[23,34–36] organic compounds. However, atmospheric modeling studies have repeatedly shown that these known sources are insufficient to account for the observed abundances of formic and acetic acids, especially in areas with strong biogenic influence.[5,37–40]

One route to organic acid formation not generally included in models is the aqueous chemistry of biogenic organic compounds in aerosol and cloudwater (Fig. 1).[41–47] Isoprene is an important primary precursor to biogenic secondary organic aerosol (SOA), especially *via* the intermediate formation of isoprene epoxydiols (IEPOX).[48,49] Experimental evidence has shown that IEPOX efficiently produces 2-methyltetrols (1,2,3,4-tetrahydroxy-2-methylbutane; 2-MT) and their sulfate esters (2-MTS) following its uptake into sulfate-containing deliquesced particles,[48–51] both of which can subsequently form dimers and oligomers.[48,51–54] These compounds have been observed as significant fractions of organic aerosol in biogenically influenced regions around the world.[55–59] Many models treat isoprene-derived SOA (iSOA) as inert, with deposition as its only loss, which may explain why the most up-to-date model mechanisms tend to overestimate its atmospheric abundance.[60,61] Photochemical sinks of aqueous SOA, such as particle-phase oxidation reactions followed by revolatilization of fragmentation products, have been suggested as a possible mechanism to help balance the SOA budget.[62–65] If these fragmentation products include formic and acetic acids, such photochemical sinks of SOA could simultaneously help correct the model underestimates for these acids.

We recently showed using nuclear magnetic resonance (NMR) spectroscopy that the aqueous oxidation of 2-MT by OH produces large yields of formic and acetic acids.[66] Here we extend that analysis to a synthetic 2-MTS isomer with similar experimental techniques and use a kinetic model to constrain organic acid yields from its oxidation. We then implement the aqueous reactions of 2-MT and 2-MTS into GEOS-Chem, a global chemical transport model, to assess the contributions of these pathways both to the global iSOA and organic acid budgets.

Experimental

Photochemical experiments

NMR experiments to determine formic and acetic acid yields from the aqueous photochemistry of 2-MTS followed the procedures described by Cope *et al.*[66] for the study of 2-MT. In each of four experiments, 80–400 mM of 2-MTS was mixed with 0.8–4 M H_2O_2 (50% in H_2O, Sigma Aldrich) in D_2O, and the solution was irradiated directly in the quartz NMR tube using a UV-B broadband fluorescent light (peak emission at 310 nm; photon flux shown in Fig. S1†) within a photochemical enclosure. At regular intervals, the tube was removed from the photochemical enclosure and ^{1}H NMR data were collected on a 400 MHz Bruker Avance III HD instrument (Fig. S2†). Data were analyzed using TopSpin, with suppression of the H_2O peak using standard WATERSUP parameters. Control experiments without H_2O_2 showed no significant loss of 2-MTS.

Experiments were performed with two methods to test the dependence of measured organic acid yields on potential gas–liquid partitioning, as in Cope *et al.*[66] In the first, cyclohexane in $CDCl_3$ was placed in a narrow glass capillary, which was flame sealed and dropped into the NMR tube containing 2-MTS and H_2O_2 in D_2O prior to the experiment to act as an internal standard. In the second, cyclohexane in $CDCl_3$ was instead placed in the NMR tube itself, and the 2-MTS reaction mixture was sealed in a glass capillary and dropped into the tube. In the former experiments, the 2-MTS and H_2O_2 solution is exposed to air in the headspace of the NMR tube during the reaction and products can partition between phases, while in the latter experiments the glass capillary contains only an aqueous phase.

2-Methyl-1,2,3-trihydroxy-4-butylsulfate, the isomer of 2-MTS investigated in this work, was synthesized according to previously published methods.[67,68] All reagents were purchased from Sigma Aldrich. Briefly, 2-methylvinyloxirane (95%) was converted to 1,2-dihydroxy isoprene (DHI) by acid-catalyzed hydrolysis, which was then epoxidized with 3-chloroperbenzoic acid (≤77%) in acetonitrile at 0 °C to yield d1-IEPOX. Reaction of d1-IEPOX with tetrabutylammonium hydrogen sulfate (97%) then afforded 2-methyl-1,2,3-trihydroxy-4-butylsulfate, the purity of which (>95%) was confirmed by ^{1}H NMR.

Kinetic modelling

To constrain the organic acid yields and multigenerational aqueous chemistry of 2-MTS, we use a simple kinetic model (Mechanism S1†) run on Matlab (MathWorks, Inc.). The mechanism uses light flux as measured in the photochemical reaction chamber (Fig. S1†) to initiate OH radical production *via* photolysis of H_2O_2, and includes the reactions OH + H_2O_2, HO_2

+ HO$_2$, and OH + HO$_2$.[69] We also include the aqueous reactions of formic and acetic acid with OH,[70] and use the OH + 2-MTS oxidation rate constant ($k_{MTS+OH} = 1.52 \times 10^9$ M^{-1} s^{-1}) recently measured by Abellar et al.[68] for the 2-MTS isomer used here. We initialize the model with aqueous concentrations of 2-MTS and H$_2$O$_2$ used in each experiment and vary the direct organic acid yields from 2-MTS, the number and OH oxidation rates of intermediate stable products of 2-MTS + OH, and the organic acid yields from those stable products to find the best fit to measured organic acid yields.

Global modelling

We use the atmospheric chemistry model GEOS-Chem (**http://geos-chem.org**) to simulate the aqueous chemistry of 2-MT and 2-MTS on a global scale. GEOS-Chem is driven by assimilated meteorological observations from the NASA Modern-Era Retrospective analysis for Research and Applications (MERRA-2). We use model version 12.8 (**https://doi.org/10.5281/zenodo.3860693**) with complex SOA chemistry[71] as a base and perform global simulations at 2° × 2.5° horizontal resolution with 47 vertical layers. Regional results reported for the Amazon (53.75 to 76.25°W, 11°S to 3°N), Southeast United States (81.25 to 93.75°W, 31 to 39°N), and Congo Basin (11.25 to 28.75°E, 5°S to 5°N) are from the same global simulations. Transport and convection are calculated at 15 min intervals, including advection based on the TPCORE algorithm,[72] nonlocal boundary layer mixing,[73] and convective mass fluxes from MERRA-2. Emissions and chemistry are calculated at intervals of 30 min. We use the standard Harmonized Emission Component (HEMCO) configuration in GEOS-Chem for emissions,[74] which includes biogenic emissions calculated from the Model of Emissions of Gases and Aerosols from Nature (MEGAN) version 2.1 inventory with isoprene emissions scaled to 535 Tg a^{-1},[75] anthropogenic emissions from the Community Emissions Data System (CEDS) inventory,[76] and biomass burning emissions from the Global Fire Emissions Database (GFED4s).[77] CEDS emissions in Africa are overwritten with the Diffuse and Inefficient Combustion Emissions inventory (DICE-Africa).[78]

GEOS-Chem uses a recently compiled mechanism of isoprene chemistry[61] based on the Reduced Caltech Isoprene Mechanism described by Wennberg et al.[79] The model also includes the reactive uptake of IEPOX into aqueous aerosol as implemented by Marais et al.,[80] with an aqueous-phase first order rate coefficient dependent on particle acidity and on the concentrations of sulfate, bisulfate, and nitrate ions ($k_{aq} = k_1[H^+] + k_2 [SO_4^{2-}][H^+] + k_3[HSO_4{-}] + k_4[NO_3^-][H^+]$).[81,82] In the base model, IEPOX reactive uptake produces a single inert SOA species (SOAIE); here, we separate the reactive uptake into one acid-dependent reaction producing 2-MT ($k_1 = 3.6 \times 10^{-2}$ M^{-1} s^{-1}), one sulfate- and bi-sulfate-dependent reaction producing 2-MTS ($k_2 = 2.0 \times 10^{-4}$ M^{-2} s^{-1}, $k_3 = 7.3 \times 10^{-4}$ M^{-1} s^{-1}), and one nitrate-dependent reaction producing 2-methyltetrol nitrates (2-MTN; $k_4 = 2.0 \times 10^{-4}$ M^{-2} s^{-1}), where the individual rate coefficients $k_{1–4}$ are identical to those used in the base model. We then add 2-MT + OH aqueous chemistry under acidic

conditions from Cope et al.[66] and 2-MTS + OH aqueous chemistry from this study (Mechanism S1†) to the GEOS-Chem mechanism. All formic and acetic acid produced from 2-MT and 2-MTS chemistry is allowed to partition instantaneously back to the gas phase. 2-MTN, which accounts for <1% of IEPOX uptake in our simulations, is still treated as inert. For aqueous OH, we use the parameterization described by Jacob et al.[83] in which $[OH]_{aq} = (1 \times 10^{-19}$ M cm^3 per molecule) $[OH]_g$.

We further update the "base" GEOS-Chem mechanism with additional formic acid sources as described by Chen et al.,[84] building on the prior work of Paulot et al.[5] and Millet et al.[39] These updates include phototautomerization of acetaldehyde,[85] oceanic emissions of alkene precursors to secondary formic acid,[5,39] agricultural and soil emissions of formic acid scaled to ammonia and nitrogen oxide fluxes,[5] and uptake of formic acid on ice and dust.[5] We also include the newly implemented gas-phase chemistry of DHI as described by Bates et al.,[86] which incorporates DHI formation from isoprene hydroxynitrate hydrolysis,[87] and the reaction of isoprene hydroxyhydroperoxide with aqueous bisulfite.[88,89] All results are for annual or seasonal simulations using meteorology from 1 December 2015 to 30 November 2016, following a six-month model initialization.

Results and discussion

Photochemical experiments and kinetic models

Fig. 2 shows the combined results of the aqueous-phase NMR experiments in which 2-MTS was oxidized by OH along with results from the kinetic model. In order to compare results across experiments, the reaction time coordinate was converted
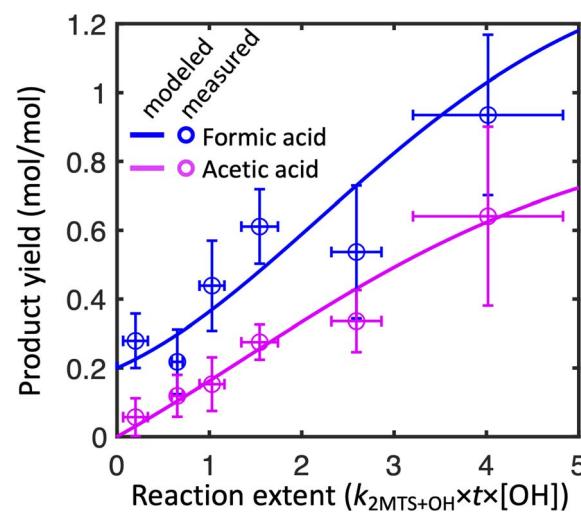
Fig. 2 Modeled and measured yields of formic (blue) and acetic (pink) acids from 2-MTS + OH oxidation in aqueous-phase NMR experiments. Points represent observed product yields from all NMR experiments binned by reaction extent, with vertical and horizontal error bars representing standard deviations of yields and reaction extent respectively among the observations in each bin. Lines represent the simulated yields of formic and acetic acid from the kinetic model.

to reaction extent (equivalent to $k_{\text{MTS+OH}} \times [\text{OH}] \times$ time) by taking the natural log of the inverse of the fraction of 2-MTS remaining in the NMR tube. When 2-MTS concentrations were too low to quantify (<10 mM), reaction extent was extrapolated as simple exponential decay from previous points. The kinetic model could then be fit to the combined data from all experiments. Measurements are binned by reaction extent in Fig. 2 for ease of interpretation; full plots of each experiment are shown in Fig. S3.† The kinetic model satisfactorily captures the acid yield trends and timing in individual experiments (Fig. S3†) in addition to the binned experimental results as a function of reaction extent.

Observed yields (moles of acid in solution per mole of 2-MTS reacted) of formic and acetic acid from 2-MTS oxidation are high, generally exceeding 100% and 30% respectively by the time 2-MTS is below detection limits. Actual yields are higher owing to the simultaneous oxidative losses of formic and acetic acids during the experiment. Yields are not constant with time but rather rise sharply with reaction extent, implying that the acids are produced in multigenerational oxidation processes. Production of formic acid starts promptly upon irradiation of the solution, with an initial yield of 20–35%, while observed acetic acid yields start near zero, suggesting that formic acid is made both in the initial and later oxidation steps while acetic acid is only produced in a multistep reaction mechanism. While the yields and identities of other products in the reaction could not be determined in these experiments, the temporal evolution of the NMR spectra closely resembles that shown by Cope et al. for 2-MT,[66] suggesting that C_4–C_5 hydroxy-carbonyl compounds similar to those identified from 2-MT (but potentially retaining the sulfate group) might serve as intermediates in the aqueous oxidation of 2-MTS.

In experiments in which the 2-MTS and H_2O_2 solution was sealed in the capillary, no remaining 2-MTS was detected after 2–3 hours; in contrast, experiments with the 2-MTS and H_2O_2 solution outside the capillary required up to 8 hours to reach a point at which 50% of the 2-MTS had been oxidized. For experiments in the former category, because the capillary is sealed without headspace, organic acids cannot partition to the gas phase. Other differences between experiments include initial 2-MTS concentrations, distance between the NMR tube and light source, and attenuation of some UV-B light by the glass capillary in the quartz NMR tube. Despite differences in experimental design, no statistically significant difference in yields as a function of effective OH exposure ([OH] × time) was observed between the two experimental methods, so we aggregate results from all experiments together in our analysis.

To best fit experimental yields while accounting for simultaneous oxidative losses of the organic acids, the 2-MTS oxidation mechanism used in the kinetic model (Mechanism S1†) requires overall yields of 2 moles of formic acid and 1 mole of acetic acid per mole of 2-MTS oxidized, implying that four of the five carbon atoms in 2-MTS are converted to small organic acids during its aqueous oxidation. In the best-fit mechanism, these yields are spread over three oxidative steps, which is necessary to account for the prompt production of formic acid and then allow both yields to rise as the reaction progresses. The optimized fit also requires that the second and third oxidation steps
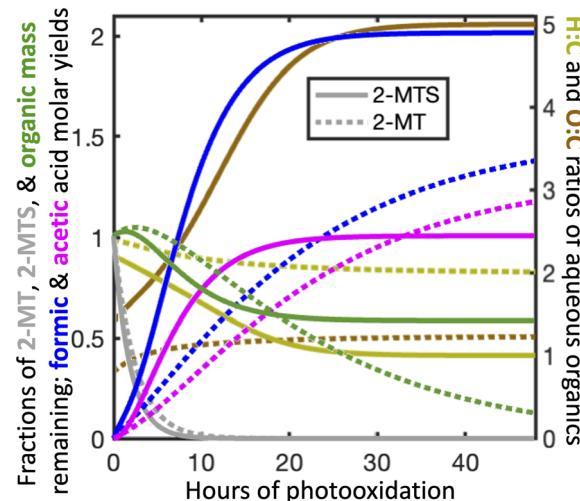
Fig. 3 Temporal evolution of aqueous atmospheric chemistry of the 2-MTS (solid lines) and 2-MT (dashed lines) systems in the optimized mechanism (Mechanism S1†) as implemented in GEOS-Chem for $[\text{OH}]_g = 10^6$ molecule per cm^3.

progress a factor of two more slowly than the initial reaction of 2-MTS with OH. The kinetic mechanism for 2-MTS is similar to that proposed for 2-MT by Cope et al.[66] (1.5 moles of formic acid and 1.25 moles of acetic acid over three oxidative steps per mole of 2-MT reacted), but with a higher formic : acetic acid ratio and faster later-generation oxidation reactions.

Fig. 3 shows the temporal evolution of aqueous atmospheric 2-MTS and 2-MT chemistry in the optimized mechanism under representative conditions as implemented in GEOS-Chem. To properly track aerosol mass and oxidation state in the model, we assign chemical formulae to the oxidative intermediates of both precursors; these formulae (see Mechanism S1†) are not meant to represent actual species, but rather the approximate elemental ratios of aqueous products remaining at each reaction step. Based on the small fraction of mass lost to formic and acetic acids in the first oxidation step and the likelihood that the reaction adds oxygen to the molecule, we assume that the first-generation intermediates from both 2-MT and 2-MTS are slightly more massive than their parent compounds, resulting in an initial increase in both molar mass and oxidation state. We assume that the final carbon from 2-MTS that does not form formic or acetic acid remains in the particle phase as a C_1 organosulfate, resulting in a high particulate O : C ratio of 5 and a final organic mass 58% as large as the initial mass. In contrast, the remaining carbon from 2-MT is presumed to form CO_2, removing all organic mass from the particle phase.[66] The optimized mechanism is meant as a simple parameterization of the reaction progress for modeling; for example, it cannot rule out the possibility that a larger number of faster oxidation steps are involved in the overall reaction process, and the chemical formulae assigned to intermediates are not constrained by our experimental results.

Global modeling

Fig. 4 shows the change in annually averaged surface IEPOX-derived secondary organic aerosol (IEPOX-SOA) loadings when

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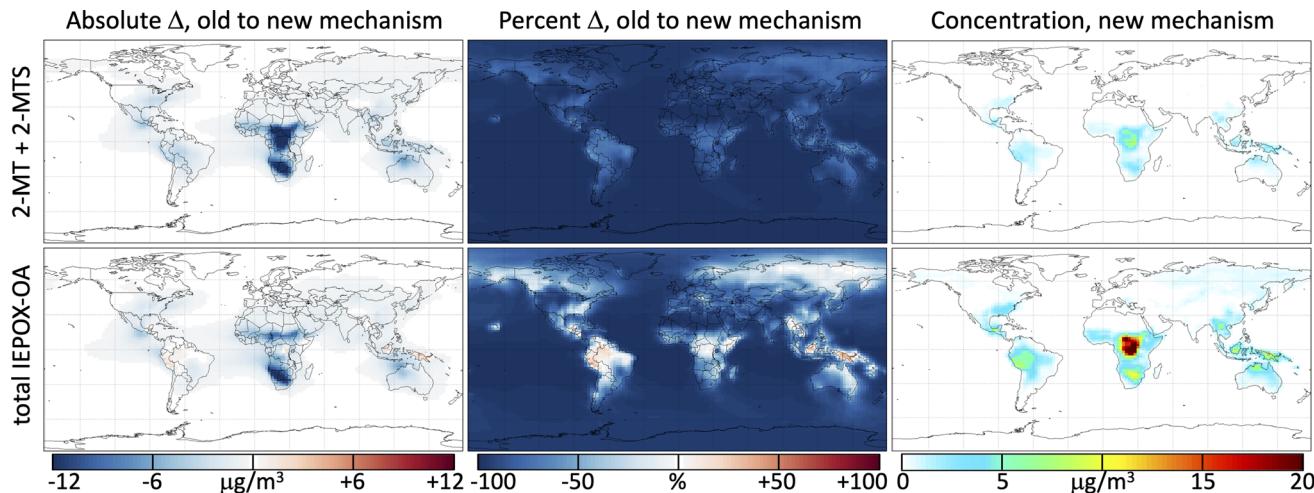
Fig. 4 IEPOX-derived secondary organic aerosol (IEPOX-SOA) and changes in IEPOX-SOA concentrations between the base GEOS-Chem mechanism and the updated mechanism with 2-MT and 2-MTS oxidation. Panels show the absolute change between the mechanisms (left), the percent changes between the mechanisms (center), and the absolute concentrations in the updated mechanism (right) for both 2-MT + 2-MTS concentrations (top) and total IEPOX-SOA concentrations (bottom). All maps show annual averages at 0–1 km altitude.

the reactions from the kinetic model (Mechanism S1†) used to simulate the photochemical experiments of both 2-MT (Cope *et al.*[66]) and 2-MTS (this work) are incorporated into GEOS-Chem. The initial reactions of 2-MT and 2-MTS with OH are fast compared with deposition; as a result, the simulated total tropospheric mass burden of 2-MT + 2-MTS is reduced by 90% when their oxidation by OH is included in GEOS-Chem. Their aqueous-phase oxidation products, however, have longer lifetimes, so the total tropospheric burden of IEPOX-SOA only decreases by 67% with the new mechanism. Global surface values, which generally reflect more recently formed SOA with lower OH exposure, exhibit smaller declines – an 85% total reduction in 2-MT + 2-MTS and 47% total reduction in IEPOX-SOA. Results for the free troposphere (Fig. S4†) are spatially similar to those shown in Fig. 4, but with much lower total burdens and slightly greater fractional decreases in 2-MT, 2-MTS, and IEPOX-SOA due to longer OH exposures. Seasonal variability in these changes, shown in Fig. S5–S10,† predominantly reflects the seasonal patterns of isoprene emissions.

Regional variability in the effects of the added 2-MT and 2-MTS chemistry reveals the extent of oxidation within particles, *i.e.*, the number of oxidative generations away from the C_5 parent compounds. The slightly more massive first-generation oxidative intermediates (see Fig. 3 and Mechanism S1†) can lead to a slight increase in total IEPOX-SOA mass relative to the base mechanism in locations close to IEPOX source regions. For example, in the Amazon, where OH is low due to titration by biogenic hydrocarbons, annual average surface IEPOX-SOA mass increases by 11% in simulations with the new 2-MT and 2-MTS chemistry. The second and third oxidation steps of 2-MT and 2-MTS lead to rapid organic aerosol mass loss and production of formic and acetic acids, which return to the gas phase. Therefore, in other isoprene-rich regions like the Congo basin and Southeast United States, where higher OH concentrations mean that more IEPOX-SOA oxidation is accomplished

close to the source, annual average surface IEPOX-SOA mass decreases by 5% (Congo Basin) and 24% (Southeast US) in simulations with the new mechanism. Over the Southeast United States in the summer, this reduces the simulated boundary layer organic aerosol burden from 7.6 to 6.5 μg m^{-3}, improving the GEOS-Chem bias relative to atmospheric observations of OA by 60%.[71] Even sharper relative decreases in IEPOX-SOA burdens are evident in regions far from IEPOX-SOA sources such as over the oceans, after multiple generations of photooxidation have occurred, where positive model biases have also been noted.[71,90]

As shown in Fig. 5, this added photochemical sink of 2-MT and 2-MTS in GEOS-Chem drastically changes the simulated IEPOX-SOA budget. Where the model previously predicted that all particle-phase 2-MT and 2-MTS (of which 16 Tg of carbon mass is produced annually) deposited unreacted, the updated model suggests that 92% of 2-MT and 2-MTS instead react with OH, forming more oxidized organic aerosol components and fragmenting to produce formic and acetic acids. These oxidized intermediates continue to react in the aqueous phase, and as a result, half of all carbon that enters the particle phase as IEPOX eventually becomes formic or acetic acids. The reactive flux through these oxidation pathways can also act as a major local sink of OH; the updated simulation shows decreased boundary layer OH of up to 15% compared to the base simulation over major isoprene-emitting regions (Fig. S11†).

Fig. 6 shows the global distribution of formic acid produced by 2-MT and 2-MTS oxidation and the contribution of this source to the total atmospheric burden. As implemented in GEOS-Chem, the oxidation of IEPOX-SOA produces 14 Tg a^{-1} of formic acid – approximately equivalent to direct emissions from fires, anthropogenic sources, and plants combined (see Table S1†) – increasing the global atmospheric source (emissions plus secondary production) by 18%. This in turn increases the tropospheric formic acid burden by 15%. These changes are

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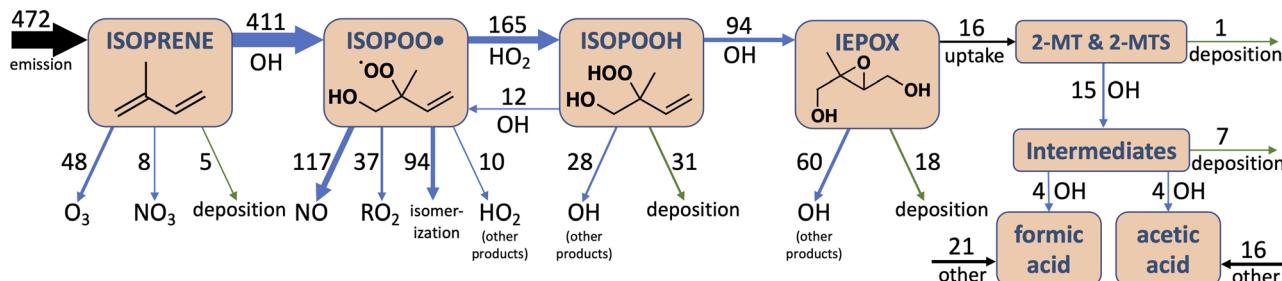
Fig. 5 Schematic of isoprene oxidation to formic and acetic acids *via* IEPOX-derived organic aerosol (OA). Blue arrows denote chemical reactions, green arrows deposition, and black arrows other processes. Numbers denote troposphere-wide annual average carbon fluxes through each pathway in TgC a^{-1} from a GEOS-Chem simulation with the updated 2-MT and 2-MTS oxidation chemistry. The small direct formic and acetic acid yields from 2-MT & 2-MTS + OH are not shown. In the base GEOS-Chem mechanism, the deposition arrow from 2-MT and 2-MTS has a flux of 16 TgC a^{-1}, and the pathways extending downward from the 2-MT & 2-MTS box have no flux.
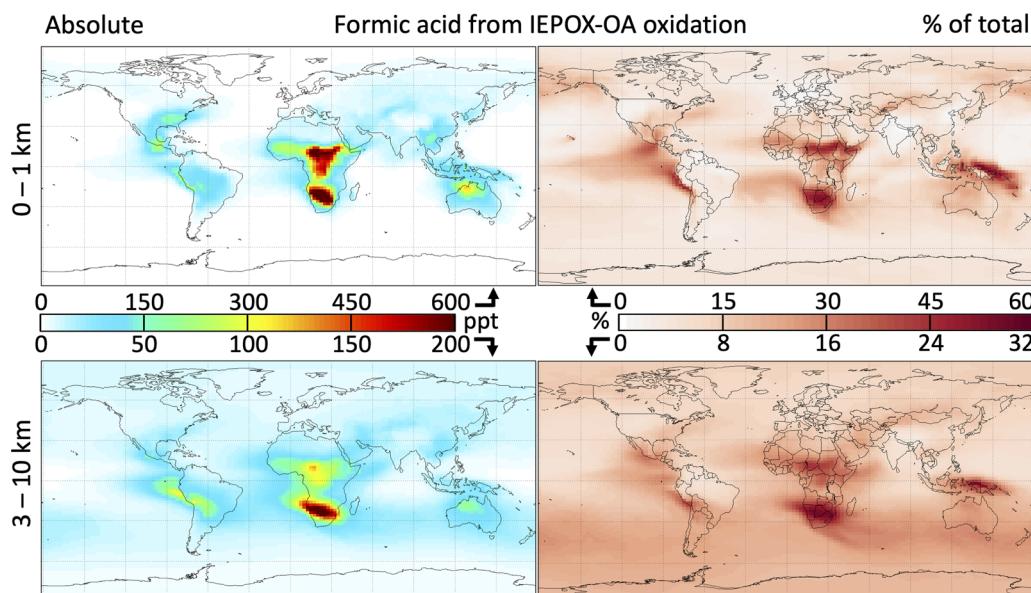
Fig. 6 Formic acid from the multi-generational oxidation of 2-MT and 2-MTS. Panels show the absolute mixing ratios of formic acid produced from IEPOX-derived organic aerosol (IEPOX-OA) oxidation (left) and the fraction of the total formic acid mixing ratios that this IEPOX-OA-derived formic acid comprises (right) at altitudes of 0–1 km (top) and 3–10 km (bottom). All maps show annual averages from a GEOS-Chem simulation with the updated 2-MT and 2-MTS oxidation mechanism.

spatially heterogeneous, with the largest absolute and fractional contributions of formic acid from IEPOX-SOA chemistry occurring over or downwind of isoprene source regions. Regionally, annual average surface formic acid from IEPOX-SOA chemistry can exceed 500 ppt, contributing over 50% to the local burden.

Because the majority of acid production takes place only after multiple aqueous oxidation steps, source areas with low OH such as the Amazon see minimal formic acid production from this mechanism; instead, most acid production occurs downwind. In contrast, the Southeast United States and Congo Basin sustain sufficiently high OH for 2-MT and 2-MTS to undergo multiple oxidation steps, resulting in high local formic acid production. Over the Southeast United States, where Millet *et al.*[39] showed that simulated near-surface formic acid is underestimated in GEOS-Chem by 1.5 ppb in the summer, the

chemistry of 2-MT and 2-MTS produces an additional 500 pptv. IEPOX-SOA oxidation contributes to forested tropical Africa exhibiting some of the highest total formic acid mixing ratios in the world (Fig. S12†) sustained year-round (Fig. S13†), consistent with hotspots of both isoprene and formic acid observed by satellite,[91,92] although previous comparisons between models and satellite observations of formaldehyde have concluded that MEGAN overestimates central African isoprene emissions,[93] and that total African isoprene emissions should be reduced by 26%.[94]

Acetic acid production from IEPOX-SOA oxidation is approximately concurrent with that of formic acid; its distribution (Fig. S14–S16†) therefore looks nearly identical to that shown in Fig. 5. Although 2-MT and 2-MTS oxidation produces slightly less acetic acid (9.5 Tg a^{-1}) than formic acid, this makes up a larger fraction of the total atmospheric acetic acid source.

The effect of this chemistry on the tropospheric acetic acid burden – a 19% increase – is therefore larger than the effect for formic acid, although the additional 300 pptv of acetic acid simulated from this chemistry over the Southeast United States in the summer is insufficient to make up the 1.1 ppb shortfall previously identified in GEOS-Chem.[39]

The modeling results described above rely on a number of assumptions. First, we assume that the distribution of 2-MTS isomers present in the atmosphere reacts similarly to the one synthesized and studied here, which is likely only a minor isomer in the ambient distribution.[52] Second, our experiments were performed in dilute aqueous solution, whereas the atmospheric aerosols in which these reactions occur frequently have high ionic strength and likely contain many other possible coreactants. Third, we assume that 2-MT and 2-MTS and their IEPOX precursors do not undergo other reactions in aerosol, such as dimerization and oligomerization, that are not included in GEOS-Chem, although these processes have been shown to occur.[51–54] While such oligomers could likely fragment upon OH-initiated oxidation as readily as their C_5 precursors, they may produce formic and acetic acids in different yields or timescales. Further mechanistic insights into the chemical pathways to formic and acetic acid formation in particle liquid water will be helpful for modeling.

A large source of uncertainty in these modeling results is the representation of particle-phase OH. GEOS-Chem simulates aqueous OH in aerosols as though it exists in rapid equilibrium with the gas phase, using a parameterization developed for cloud water.[83] Although this is certainly an oversimplification, aqueous atmospheric OH is notoriously difficult both to measure and to accurately model.[95,96] Ambient particles are generally highly concentrated solutions containing many other species with which OH can react, in addition to numerous aqueous-phase sources of OH,[97–100] none of which are explicitly represented here. While aqueous OH concentrations in GEOS-Chem ($\sim 10^{-13}$ M) are comparable to those in some other models,[101–105] few measurements exist for comparison, and only for marine aerosol[106,107] and larger droplets.[95] Improving model representation of aerosol-phase OH will be critical for accurately simulating the fate of IEPOX-SOA in the atmosphere.

In addition to IEPOX-SOA, terpene- and aromatic-derived SOA have been shown to lose mass when exposed to OH or light on atmospherically relevant timescales.[64,108,109] While the gas-phase products of these reactions are not characterized, their incorporation into GEOS-Chem would further accelerate losses of organic aerosol and may lead to OA underestimates in parts of the world, which would need to be balanced by larger sources. Indeed, the source strength of IEPOX-SOA was tuned in GEOS-Chem (*via* the effective Henry's Law constant, H*) to match summertime IEPOX-SOA observations from aerosol mass spectrometry (AMS) in the Southeast United States,[80,110] meaning that the model likely underestimates IEPOX-SOA there with the new loss pathways added. Retuning the model's H* to accelerate IEPOX-SOA formation would counterbalance the oxidative loss, and would further increase formic and acetic production from IEPOX-SOA oxidation; however, more recent analyses suggest that IEPOX-SOA mass loadings derived from AMS measurements may be overestimated in the Southeast United States,[111–113] so we do not increase the IEPOX-SOA source here. Future incorporation of oxidative OA losses into models should also reassess OA sources to ensure that burdens remain well represented.

This study, with its noted assumptions, shows that adding photochemical sinks to SOA species in atmospheric models may have significant impacts on atmospheric composition. Although the cumulative effects of our assumptions on the atmospheric IEPOX-SOA oxidation rate and the magnitude of acid production cannot be straightforwardly determined, the results presented here provide a starting point from which to consider the potential impacts of aqueous-phase SOA photo-oxidation on the organic aerosol and organic acid budgets. In addition to this homogeneous (one phase) aging process, heterogeneous aging of IEPOX-SOA – the direct reaction of SOA with gas-phase OH – may also contribute. Studies of heterogeneous IEPOX-SOA aging using oxidation flow reactors (OFRs) measured much longer lifetimes ($\sim$2 weeks) of SOA against oxidative mass loss.[114–116] However, such methods require their own assumptions to generalize from laboratory to ambient conditions; for example, the OH reactive-diffusive distance in particles is small,[117] and the high OH and short residence time in OFRs can make surface reactions self-limiting when the particle-phase organics cannot rapidly reach the surface,[118,119] an effect exacerbated when the particles are highly organic.[120,121] The actual atmospheric reactivity of IEPOX-SOA will depend on both gas- and particle-phase sources and sinks of OH, and on the aerosol phase, viscosity, water content, and chemical composition. Further studies of aerosol oxidative aging in particle water at different particle phases and viscosities, and with aqueous OH sources and conditions relevant to the atmosphere, would facilitate an improved extrapolation from the lab to the ambient and more accurately inform future modeling studies.

Conclusions

We have shown using a synthetic standard in NMR-monitored reaction experiments that the aqueous-phase OH-initiated oxidation of 2-MTS, a ubiquitous product of IEPOX uptake into aerosol, rapidly produces formic and acetic acids in high yields in a multistep process. When incorporated into a global atmospheric model, simulations suggest that this oxidation (along with the analogous chemistry of 2-MT) is the dominant loss process of 2-MTS and 2-MT, reducing their combined tropospheric burden by 90%. These reactions may produce more oxidized aqueous intermediates that can locally increase SOA mass loadings, most notably over the Amazon, but ultimately lead to a 67% reduction in the tropospheric burden of total IEPOX-derived aerosol. The mass is removed from particles *via* conversion to gas-phase formic and acetic acids, increasing the tropospheric burdens of these species by 14.8% and 19.4% respectively. Aqueous-phase oxidation of IEPOX-SOA can also contribute significantly to local reactivity and depletion of OH, particularly over African forests.

While the modeling herein relies on many assumptions that introduce uncertainty to our results, we have shown that OH-initiated oxidation of aqueous IEPOX-SOA could substantially alter the global budgets of both organic aerosol and organic acids, and have provided a benchmark for quantifying these effects. Furthermore, the direction and magnitude of the changes to aerosol and acid burdens we simulate are in line with what previous studies suggest is needed to remedy model biases. Recent modeling studies suggest that incorporating all known sources of SOA from isoprene would lead to a severe overestimate of organic aerosol and that additional losses like those proposed here are needed to balance the SOA budget.[36,60] Our simulated 1.1 µg m^{-3} reduction in IEPOX-SOA over the Southeast United States in summer due to 2-MT and 2-MTS chemistry would reduce the current model-measurement bias in GEOS-Chem by 60%.[71] The increased tropospheric formic and acetic acid burdens from IEPOX-SOA oxidation also serve to reduce model-measurement biases; while the additional 500 ppt of formic acid and 300 ppt of acetic acid we simulate from 2-MT and 2-MTS chemistry over the Southeast United States in summer only makes up ~30% of the model shortfall,[39] similar oxidative chemistry of other particle-phase organics could further contribute. Extending the methods herein to other common aqueous aerosol constituents, improving model representation of aerosol-phase OH, and minimizing the other uncertainties from assumptions used in simulating gas–aerosol interactions will further improve our ability to simulate the burdens and budgets of organic aerosol and organic acids in the atmosphere.

Author contributions

KHB prepared and performed GEOS-Chem simulations with assistance from DJJ, XC, and DBM; JDC performed photooxidation experiments; KHB, JDC, and TBN analysed experimental data; KHB prepared the manuscript with assistance from all authors.

Conflicts of interest

The authors declare no conflicts of interest.

Acknowledgements

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