Chemical reactivity and long-range transport potential of polycyclic aromatic hydrocarbons – a review†

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Cite this: *Chem. Soc. Rev.*, 2013, **42**, 9333

Polycyclic aromatic hydrocarbons (PAHs) are of considerable concern due to their well-recognised toxicity and especially due to the carcinogenic hazard which they present. PAHs are semi-volatile and therefore partition between vapour and condensed phases in the atmosphere and both the vapour and particulate forms undergo chemical reactions. This article briefly reviews the current understanding of vapour–particle partitioning of PAHs and the PAH deposition processes, and in greater detail, their chemical reactions. PAHs are reactive towards a number of atmospheric oxidants, most notably the hydroxyl radical, ozone, the nitrate radical (NO₃) and nitrogen dioxide. Rate coefficient data are reviewed for reactions of lower molecular weight PAH vapour with these species as well as for heterogeneous reactions of higher molecular weight compounds. Whereas the data for reactions of the 2–3-ring PAH vapour are quite extensive and generally consistent, such data are mostly lacking for the 4-ring PAHs and the heterogeneous rate data (5 and more rings), which are dependent on the substrate type and reaction conditions, are less comprehensive. The atmospheric reactions of PAH lead to the formation of oxy and nitro derivatives, reviewed here, too. Finally, the capacity of PAHs for long range transport and the results of numerical model studies are described. Research needs are identified.

Received 26th April 2013

DOI: 10.1039/c3cs60147a

www.rsc.org/csr

Extended summary

Polycyclic aromatic hydrocarbons (PAHs) are a class of organic compounds characterised by the presence of fused aromatic rings. Their ubiquitous presence in the environment is of major concern due to the carcinogenicity of a number of specific PAH compounds (congeners). The European Union has set an indicative limit value of 1 ng m^{−3} of benzo[*a*]pyrene used as a marker for the typically complex mixture of congeners. While elevated concentrations of atmospheric PAHs are encountered in the vicinity of major sources, observations from remote sites indicate long-range transport. In this context, the atmospheric chemistry of PAHs is of particular importance as this has a major influence upon their atmospheric lifetime and hence distribution in the environment.

PAHs are generated mainly as by-products of incomplete combustion processes. Consequently, their sources include domestic burning of coal and wood for heating and cooking; fossil and biomass fuel burning power plants; industrial processes; and road transport. Globally, the combustion of biofuels and wildfires are major sources, while road traffic and specific industries frequently dominate urban emissions.

PAHs range from naphthalene (two aromatic rings) which under ambient conditions exists almost entirely as vapour through to compounds with six or more aromatic rings which partition almost entirely into the particulate phase. The majority of compounds, and especially those with three or four rings, are considered as semi-volatile and such compounds partition between the vapour and particle phases in the atmosphere. These compounds can deposit to surface water and soils where they have a long lifetime but subsequently re-evaporate to the atmosphere. Much of the soil PAH in temperate regions is historic and concentrations in air and soil are relaxing to equilibrium on various time scales. Vapour–particle partitioning of PAH can be quantified through the gas–particle partitioning coefficient. This is influenced by both adsorption and absorption processes and is strongly temperature-dependent; hence a seasonal variation in partitioning is typically observed. Quantitative analysis of the partitioning process suggests that it can be described by the sum of the absorptive process as indicated by the octanol–air partition coefficient and the adsorptive process

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† Electronic supplementary information (ESI) available. See DOI: 10.1039/c3cs60147a

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as described by the soot-air partitioning coefficient. Modes in the mass size distribution for PAH are typically within the ultrafine (less than 0.1 μm) and accumulation (0.1–1 μm) ranges of aerodynamic diameter, but mass size distributions can be transient and measurement is rendered difficult by artefacts in the measurement process caused by PAH semi-volatility. Fast and sensitive *in situ* measurement techniques are not available.

Dry deposition is more effective than wet deposition as a removal process from the atmosphere. Chemical reactions provide the other main sink for atmospheric PAHs. The gas phase reactions of PAHs with the OH radical, the NO_3 radical and ozone have been widely investigated. Available rate coefficient data are most abundant in the case of the OH radical. The established mechanism of PAH reactions with the OH radical involves the formation of a PAH-OH adduct followed by further reaction

with NO_2 or O_3 . The observed reaction products include both ring-retaining nitro-PAHs and quinones, as well as ring-opened products such as phthalic acid, phthalaldehyde and phthalic anhydride. The presence of methyl groups in methyl naphthalenes and methyl phenanthrenes in most cases leads to a modest increase in reactivity relative to the parent PAH. For NO_3 reactions, the predominant reaction pathway involves NO_3 addition followed by reaction with NO_2 leading to nitro-PAH formation. The observed rate coefficients are proportional to the nitrogen dioxide concentration. There have been far fewer studies of the gas phase reactions of PAH with ozone.

The main atmospheric sink for gas phase PAHs appears to be reaction with the OH radical, with rate coefficients for these reactions up to five orders of magnitude greater than for the corresponding reactions with NO_3 for most three to four ring PAHs. While NO_3 reactions appear to be less significant than OH reactions as a PAH degradation process, considerably higher nitro-PAH yields suggest that nighttime reactions of PAHs with NO_3 may be a significant contributor of these compounds in the atmosphere, in addition to daytime OH reactions. Reactions of PAHs with ozone are considered to be of negligible importance in the atmosphere. For many atmospherically relevant substances the available data are insufficient.

Given the predominant association of many PAHs with the particulate phase, heterogeneous reactions of a number of PAHs adsorbed on a number of solid substrates have been extensively studied. Reaction substrates include both carbonaceous aerosol (graphite, diesel exhaust, kerosene flame soot, ethylene flame soot) and mineral particles (silica and MgO). Reactions have been studied for OH, $\text{N}_2\text{O}_5/\text{NO}_3$ and O_3 and rates have been shown to depend not only upon the reactant but also upon the nature of the substrate. Reactions with nitrogen dioxide have also been shown to proceed at a significant rate for some PAHs. For semi-volatile PAHs (mostly three to four ring compounds), comparisons between the rates of homogeneous and

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derivative compounds. This approach involves using atmospheric measurements in combination with chemical kinetics data to assess the contribution of photochemical reactivity of PAHs to observed levels of oxy- and nitro-PAH.

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heterogeneous reactions can be made. In the case of the hydroxyl radical, PAH reactions on carbonaceous particle surfaces are one to three orders of magnitude lower than those derived for gas phase reactions. The presence of a plateau in the experimental decays of PAHs in the reactions of OH indicates that a significant fraction of PAH is unavailable for reaction. However, in the case of ozone, studies of PAH absorbed on graphite and silica substrates show heterogeneous processes to be approximately two orders of magnitude faster than those of the corresponding gas phase reactions.

Estimation of reaction rates in the atmosphere using typical concentrations of atmospheric oxidants indicates that the gas phase reaction with OH remains the dominant loss process for most PAHs studied so far but that heterogeneous reactions with nitrogen dioxide and ozone may be of some importance. However, field measurements, particularly from remote sites, and recent laboratory experiments suggest the stabilisation of particle-associated PAHs in the atmosphere, probably due to incorporation in a particle matrix which limits their accessibility to atmospheric oxidants. A number of ring-opened and ring-retaining products have been identified from these heterogeneous reactions including nitro-PAHs and quinones. In some cases a difference has been noted between heterogeneous and gas-phase reaction products.

Long-range atmospheric transport of PAHs has been studied through both atmospheric measurements and numerical modelling. PAHs show a global distribution with appreciable concentrations observed at sites within the Arctic. Models suggest that atmospheric half-lives of 3–5 ring PAHs are of the order of hours or days, but vary considerably amongst model studies. Since there is significant uncertainty in emissions inventories and the processes determining gas–particle partitioning, and models include different processes (*e.g.* most models neglect revolatilisation from surfaces), model predictions currently have a high degree of uncertainty.

The role of atmospheric reactivity in influencing the observed levels of oxy- and nitro-PAH compounds relative to primary emissions has been investigated using atmospheric measurements. Differences in the reaction mechanism between primary combustion emissions and gas-phase decomposition of PAH have been shown to result in differences in nitro-PAH isomer distributions. Ratios of oxy- and nitro-PAHs to their parent PAHs have also been used to assess the importance of atmospheric reactions in influencing the concentrations of these compounds.

Measurements of temporal (daily, seasonal and diurnal) variations in the concentrations of PAHs and their oxy- and nitro-derivatives have been used to evaluate the reactivity of PAHs and the formation of derivative compounds. Such studies have also been used to determine the atmospheric oxidants having the greatest impact on the atmospheric chemistry of PAH. In the case of low molecular weight compounds reacting predominantly in the vapour phase, some studies have shown a good agreement between atmospheric observations and predictions based upon laboratory kinetic data. A full analysis of atmospheric processing of PAH is limited by the limited current

knowledge of the atmospheric reactivity and reaction products of the nitro- and oxy-derivatives of PAH.

1. Introduction

Polycyclic aromatic hydrocarbons (PAHs) are a group of compounds with proven carcinogenicity. Although the carcinogenic potency varies greatly between congeners,^{1,2} they are generally emitted to the atmosphere as a mixture which presents a substantial toxic hazard. Levels in the atmosphere are almost everywhere dominated by combustion sources, while aluminium and iron and steel production as well as petrogenic sources are significant in some environments.³ The World Health Organisation recommends guidelines in terms of a carcinogenic slope factor⁴ and the European Union indicative limit value is set at 1 ng m^{−3} of benzo[*a*]pyrene.⁵ The United Kingdom has set an air quality standard of 0.25 ng m^{−3} benzo[*a*]pyrene.⁶ Because of the complexity of PAH mixtures, the standards and guidelines are set in terms of the concentration of benzo[*a*]pyrene as this compound typically represents a substantial proportion of the total carcinogenic potential of at least the parent PAH mixture.² In turn, the carcinogenic potential of PAH frequently makes a substantial contribution to the carcinogenicity of urban air.^{7–10}

Although often referred to as persistent organic pollutants (POPs), PAHs are reactive in the atmosphere. In fact, such reactivity presents problems for receptor modelling methods in source attribution studies. Oxidised products are formed, the most notable being oxy-derivatives (mostly quinones) and nitrated compounds. Some such compounds are also present in primary emissions. PAH derivatives have attracted interest because some are very potent mutagens and carcinogens.^{11–13}

A further complexity in studying PAH and their derivatives arises from their semi-volatility. Compounds partition between the particle and vapour phases and hence undergo chemical reactions in both phases. While for the lower molecular weight compounds, reactions in the gas phase are likely to dominate, for compounds with four aromatic rings and higher, reactions in both phases need to be considered. Furthermore, if contact with the ground does not imply immediate reaction, semi-volatility implies the potential of re-volatilisation.

Observations of PAHs at remote sites, *i.e.* far from industrial and transport sources, indicate long-range transport. This is a concern, because of the related hazard for human health and ecosystems. Some PAHs, benzo[*a*]pyrene, benzo[*b*]fluoranthene, benzo[*k*]fluoranthene, and indeno[1,2,3-*cd*]pyrene beside others, are bio-accumulative. Therefore, and because of resistance to degradation PAHs are considered persistent organic pollutants (POPs) by the Convention on Long-range Transboundary Air Pollution and are listed by conventions for UNECE¹⁴ and OSPAR.¹⁵ Monitoring in the atmospheric environment mandated by convention processes is on-going in Europe and North America.^{14,16,17}

Much of the scientific literature is focussed upon the 16 USEPA priority PAHs *i.e.*, acenaphthene (Ace), acenaphthylene (Acy), fluorene (Flu), naphthalene (Nap), anthracene (Ant), fluoranthene (Flt), phenanthrene (Phe), benzo[*a*]anthracene (BaA), benzo[*b*]fluoranthene (BbF), benzo[*k*]fluoranthene (BkF),

Table 1 Physico-chemical properties for 298 K, i.e. vapour pressure of the supercooled liquid p_L (Pa; semivolatility is corresponding to $p_L = 10^{-6}$ – 10^{-2}), Henry coefficient K^H (M atm $^{-1}$), octanol–air partition coefficient K_{oa} , rate coefficients of the gaseous substance with the hydroxyl radical and ozone, and total photochemical residence time τ_{total} of 17 parent PAHs (USEPA 16 priority PAHs + BeP). Estimated values (EPIWIN; USEPA)¹⁸ in lack of experimental data are given in brackets. τ_{total} refers to oxidation by O $_3$ and OH in the gaseous and particulate phases but neglects the possible particle matrix effect

	p_L^a (Pa)	K^H^b (M atm $^{-1}$)	$\log K_{oa}^c$	k_{OH}^g (10 $^{-12}$ cm 3 molec $^{-1}$ s $^{-1}$)	$k_{O_3}^g$ (10 $^{-318}$ cm 3 molec $^{-1}$ s $^{-1}$)	τ_{total}^d (h)
Naphthalene (Nap)	38	2.2	5.19	22	<0.3	13
Acenaphthylene (Acy)	2.6	10.5	6.46	110	550	1
Acenaphthene (Ace)	1.7	7.1	6.44	58	<0.5	5
Fluorene (Fln)	0.54	11.3	6.85	13		21
Phenanthrene (Phe)	0.10	23.5	7.64	31	0.40	9
Anthracene (Ant)	5.9×10^{-2}	20.0	7.70	130		2
Fluoranthene (Flt)	6.8×10^{-3}	76	8.81	11		14–25
Pyrene (Pyr)	4.2×10^{-3}	76	8.86	50		5–6
Benzo(a)anthracene (BaA)	3.8×10^{-4}	159	10.28			3–11
Chrysene (Chr)	1.3×10^{-4}	270	10.30	(50)		9–27
Benzo(b)fluoranthene (BbF)	1.0×10^{-5}	1550	11.34	(16)		34 to >330
Benzo(k)fluoranthene (BkF)	7.8×10^{-6}	1790	11.37	(54)		8–21
Benzo(e)pyrene (BeP)	1.8×10^{-5e}	276 ^e (1240)	(11.35)	(50)		10–15
Benzo(a)pyrene (BaP)	7.9×10^{-6}	1320	11.48	(50)		2–5
Indeno(1,2,3-c,d)pyrene (IPy)	6.6×10^{-7}	2050	12.43	150		6–9
Dibenz(a,h)anthracene (DBA)	9.5×10^{-8e}	120 ^e (2045)	12.59 ^f	(50)		34 to >330
Benzo(g,h,i)perylene (BPe)	4.6×10^{-7}	2410	12.55	5.9		5

^a Ma *et al.*,¹⁹ for temperature dependent data see Tenhulscher *et al.*²⁰ ^b Ma *et al.*,¹⁹ for temperature dependent data see Lei *et al.*²¹ ^c Ma *et al.*,¹⁹ for temperature dependent data see Odabasi *et al.*²² ^d Based on oxidant levels characteristic of the continental background in mid latitudes (10⁶ OH per cm 3 , 50 ppbV O $_3$) and assuming model aerosols being representative for ambient aerosols (silica for unspecific particulate mass, diesel and flame soot for BC). Ranges given reflect uncertainties of kinetic data (see Tables 3, 5 and 6), particulate mass fraction, θ (adopted from a range of field observations during various seasons and in various climates, i.e. Simcik *et al.*,²³ Mandalakis *et al.*,²⁴ Tsapakis and Stephanou,²⁵ Ding *et al.*,²⁶ He and Balasubramanian,²⁷ Lammel *et al.*,¹⁰ Demircioglu *et al.*,²⁸ besides others), and of BC content (adopted from Spindler *et al.*²⁹). ^e Paasivirta *et al.*³⁰ ^f Odabasi *et al.*²²

chrysene (Chr), pyrene (Pyr), benzo[ghi]perylene (BgP), benzo[a]pyrene (BaP), dibenzo[a,h]anthracene (DBaA), and indeno[1,2,3-cd]pyrene (IPy) (Table 1). This list is of largely historic significance and was developed when knowledge of relative toxicity of PAH

congeners was more limited than at present. It does not include some of the more potent carcinogens, but focuses upon those compounds typically present at higher concentrations which are hence measurable in many environmental samples.

In this article, we review current knowledge of PAH and nitro-PAH with a focus upon chemical reactivity in both the gaseous and condensed phases, and where possible elucidate rates and mechanisms. Reactivity is put in the context of exposure of remote atmospheric environments, *i.e.* long-range transport (LRT) potential of these pollutants.

2. Sources

2.1. Primary emissions

Polycyclic aromatic hydrocarbons (PAHs) are by-products of all types of combustion, in particular incomplete. Major anthropogenic emission sources to air are, domestic burning of coal and wood, fossil fuel burning power plants, road transport, wood burning, cooking, and others. For global sources in 2004, Zhang and Tao³¹ reported biofuel (*i.e.* straw, firewood and animal dung, 56.7%) and wildfire (*i.e.* forest and savanna fires, 17.0%) as the main contributions, while traffic, domestic coal combustion and burning of agricultural waste contribute only 4.8%, 3.7% and 2.7%, respectively (referring to the sum of the 16 USEPA priority PAHs). In open fires of agricultural debris, forest or savannah grass, PAHs are not only emitted into air but also into soil. Additionally, petrogenic and biological³² sources may contribute, but anthropogenic sources prevail by far.

The countries ranking highest in emissions of the 16 USEPA priority PAHs in 2004 were China (114 kt year⁻¹), India (90 kt year⁻¹) and the USA (32 kt year⁻¹).³¹ Dominated by the use of biofuels, the PAH pattern may differ considerably among countries. For example, health risk, expressed in BaPeq/sum of 16 PAHs, ranges 0.18–3.58% among countries with highest values expected for countries with high usage of non-anthracite coal or wheat straw.³¹ High spatial resolution emission inventories are available for a number of countries (usually only BaP or few substances, namely BaP, BbF, BkF and InP; *e.g.* EEA,³³ Xu *et al.*³⁴). Emission estimates have also been compiled for whole regions (*e.g.*, Gusev *et al.*³⁵) as well as globally^{31,36} (Table 2). Global anthropogenic emissions of the 16 USEPA priority PAHs in 2004 were close to 4 kg km⁻² year⁻¹ with biofuel (56.7%), wildfire (17.0%) and consumer product usage (6.9%) contributing most.³¹ Power plants, open biomass burning, road transport (mostly diesel), industrial processes, air and sea transport also contribute.

Table 2 Regional and global emission estimates of (a) the 16 USEPA priority PAHs and (b) benzo[a]pyrene (t year⁻¹)

Europe and Russia	USA	China	World	Year	
(a)					
	26 500			1990	USEPA ³⁹
		25 300		2003	Xu <i>et al.</i> ³⁴
51 000	32 000	114 000	520 000	2004	Zhang and Tao ³¹
(b)					
1047			22 436	1990	Münch and Axenfeld ³⁷
590				1995	Pacyna <i>et al.</i> ⁴⁰
264	88	239	3700	1996	Lammel <i>et al.</i> ³⁶
		540		2003	Xu <i>et al.</i> ³⁴
907				2000	Denier van der Gon ⁴¹

For BaP 1990 global anthropogenic emissions, it was estimated that 13% is due to biomass burning and 87% due to fossil fuel combustion.³⁷

Emission inventories are based on emission factors from the various combustion technologies and fuel consumption and transformation data and usually apply for one particular year. Emission inventory determination is very uncertain as substance-specific emission factors (ng g⁻¹ fuel burnt or µg km⁻¹ for transport) determined experimentally are very sensitive to even minor differences in combustion technology, engine operation, fuel composition, *etc.* and, hence, may vary by several orders of magnitude for the same type of emission. Most uncertain are emission factors of open fires (*e.g.* forest fires), because of difficulties in measuring under realistic and representative conditions. Data gaps have been filled assuming a constant ratio of emission factors of PAH and another trace substance (*e.g.* black carbon³⁶). To distribute on a grid, country-based data are often scaled according to population density (under the assumption that emission per capita is constant) and annual data are temporally distributed evenly neglecting seasonal variation of major sources, such as residential heating related fossil fuel burning. More realistic temporal functions have been successfully used in regional PAH modelling.^{35,38}

Known substance patterns of emission sources partly combined with known emission fluxes have been used to quantify the contribution of individual sources to pollution levels at receptor sites using various types of source-receptor modelling.^{42,43} The most simple approach in this context is the comparison of ratios of pairs individual PAH's concentrations at the receptor and source sites. Pairs of PAHs of similar degradability in air have been suggested for this purpose (so-called diagnostic ratios^{44–46}). This approach is not successful due to the variability of sink (degradation, deposition) conditions and incomplete knowledge of chemical kinetic data in both the gaseous and particulate phases (Section 3^{47,48}). The success of multivariate statistical approaches, making use of substance patterns of emission sources which encompass many PAHs (so-called fingerprint methods), is often limited by the similarity of sources in terms of the PAH pattern.⁴⁹ The inclusion of other substances, such as the anhydrous sugar levoglucosan as a marker for biomass combustion⁵⁰ and hopanes as markers for motor vehicle exhausts⁵¹ and coal combustion,⁵² can solve the problem. In the future, this challenge may consistently be covered using inverse modelling (adjoint model). In this context, the ratio of the concentrations of relatively fast degrading benzo[a]pyrene over relatively slow degrading benzo[e]pyrene (BeP) has been suggested as a measure for age of air mass.⁵³

2.2. Re-volatilisation from ground compartments

Most of the 3–4 ring PAHs are semivolatile (*i.e.*, their saturation vapour pressure at 298 K is in the range $p_{\text{sat}} = 10^{-6}$ – 10^{-2} Pa; Table 1). Therefore, these substances are subject to re-volatilisation upon condensation on surfaces and to gas exchange with vegetation,^{54–57} soils^{28,58–62} and water surfaces.^{63–65} Persistence of PAHs in soils is on the order of decades.⁶¹ Therefore, LRT of PAHs may be enhanced by re-evaporation.³⁶ Secondary emissions due to re-volatilisation from surfaces need to be included in modelling of semivolatiles, in particular if the substance had been accumulated

historically in the ground compartments. For example, at the central European background observatory Košetice, Czech Republic, the burden of fluoranthene and pyrene in soil, 121 and 95 ng g⁻¹ (mean of 2006–2008, depth sampled 10 cm) exceeds the burden in air by 4 orders of magnitude (under the assumption of a bulk soil density of 1.5 g cm⁻³ and that the concentrations in ground-level air, 1.83 and 1.19 ng m⁻³, respectively, were representative for the air column within the planetary boundary layer of 1000 m depth) and concentrations in soil and air have been approaching equilibrium during the last decade.⁶⁶ Much of the soil PAH in temperate regions is historic and concentrations in air and soil are relaxing to equilibrium on a multi-year time scale.

The volatilisation flux is described in models based on an empirical parameterization or as given by diffusion. However, not all contributions to effective diffusion are represented in models and additional sources to the atmosphere exist. For example, the possible enhancement of soil–air exchange by co-evaporation with soil water vapour and of air–sea exchange by turbulence in surface water (breaking wave phenomena, bubbles^{67,68}) is usually neglected. Apart from volatilization additional secondary emission paths are possibly relevant, e.g. the formation of aerosol particles from drying sea spray droplets. This is a relevant source as PAHs, like other hydrophobic organics, were found enriched in the sea surface microlayer by up to a factor of 10, sometimes even up to a factor of 100 as compared to the subsurface water.^{69,70} The fraction of the PAH in surface seawater which is partitioning to suspended or sinking particulate matter can volatilize following dissolution in water in relaxation to phase equilibria – or be removed from the surface layer carried by sedimenting particles or by deep water formation. Similarly, the PAH burden stored in soils may only partly be mobile and subject to re-volatilization: in soil, through chemical interaction with fulvic acid or other organic matter^{71–75} and black carbon^{76–78} components PAH may in the long-term turn immobile ('ageing') and more persistent. Sorption strength is dependent on the type of soil organic matter (OM) and sorption history.^{79–81}

Degradation in ground compartments is in models generally assumed to obey first order chemical kinetics without spatial variability in soils or in seawater. However, there is evidence for significant spatial heterogeneity of PAH biodegradability at least in soils.⁸² The rates used for seawater are often merely extrapolated from experimental data obtained in freshwater using a default factor to account for a lower level of bioactivity. The temperature dependence of the degradation reaction coefficient, if not neglected, is usually fixed to some default function, e.g. doubling per 10 K temperature increase.

3. Sinks

3.1. Gas–particle partitioning and deposition

3.1.1. Gas–particle partitioning. The phase partitioning of a semivolatile organic compound, *i*, in aerosols can be quantified by the gas–particle partitioning coefficient K_p (m³ μg⁻¹):

$$K_p = c_{ip}/(c_{ig}c_m) \quad (1)$$

with the concentrations in the gaseous and particulate phases, c_{ig} and c_{ip} , and particulate matter mass concentration c_m . Note that K_p is operationally defined, which means that data are often biased by sampling artefacts.^{83–89} The most relevant contributions under many ambient and sampling conditions are due to oxidative or evaporative losses from particulate matter deposited on filter membranes, which leads to an underestimation of the particulate phase mass fraction.^{86,89}

Adsorption:

If adsorption is dominating the process, then for the particulate phase mass fraction of semivolatile substances, $\theta = c_{ip}/(c_{ig} + c_{ip}) = [1 + 1/(K_p c_{TSP})]^{-1}$, it holds:^{90,91}

$$\theta = c_j S / [c_j S + p_L^0] \quad (2)$$

with constant c_j , aerosol surface S (cm⁻¹), and saturation vapour pressure of the super-cooled liquid p_L^0 (Pa). This relationship is empirically based. c_j should be substance specific, but was hardly ever specified, but rather generally adopted as 17.2 Pa cm. K_p could be expressed as:⁹¹

$$K_p = N_s A_m \text{Te}^{(Q_1 - Q_v)/RT} / (16 p_L^0) \quad (3)$$

with surface concentration of sorption sites N_s , specific surface area of particulate matter $A_m = 10^6$ S cm⁻¹ (cm² μg⁻¹), the enthalpy of desorption from the surface Q_1 and the enthalpy of vapourisation of the super-cooled liquid Q_v . For equilibrium partitioning the slope, m , of plots of the form:

$$\log K_p = m \log p_L^0 + b \quad (4)$$

should be equal to -1 and the intercept, b , for constant temperature should be determined by N_s and A_m . Deviation of m from -1 , however, is inconclusive, as various possible reasons exist which usually cannot be addressed, such as non-equilibrium, temperature change during sampling, weaker interaction between the gas and solid surface, and absorption into OM.^{92,93} Non-equilibrium can be caused by kinetic constraints (introduction of fresh particles or clean air) or by the presence of non-exchangeable compounds on or in the particles.

As $\log p_L^0 \sim 1/T$ (Clausius–Clapeyron equation) eqn (4) is physically equivalent to plots of the form

$$\log K_p = A/T + B \quad (5)$$

which was derived from Langmuir adsorption theory.⁹⁴ For PAH absorption in primary or secondary organic particulate matter K_p would be given as:⁹⁵

$$K_p = K_{OA} f_{OM} MW_{OCT} \zeta_{OCT} / (\rho_{OCT} MW_{OM} \zeta_{OM} 10^{12}) \quad (6)$$

with K_{OA} being the octanol–air partition coefficient (values in Table 1), f_{OM} the mass fraction of OM in particulate matter, MW_{OCT} and MW_{OM} the mean molecular weights of octanol and the OM phase (g mol⁻¹), respectively, ρ_{OCT} the density of octanol (0.820 kg L⁻¹), ζ_{OCT} the activity coefficient of the absorbing compound in octanol, ζ_{OM} the activity coefficient of the compound in the OM phase. With the assumptions that $\zeta_{OCT} = \zeta_{OM}$ and $MW_{OCT}/MW_{OM} = 1$, eqn (6) can be simplified as:

$$\log K_p = m' \log K_{OA} + b' \quad (7)$$

then, and when phase equilibrium is established the slope, m' , of plots of eqn (7) should be equal to +1 and the intercept, b' , equals $\log f_{\text{OM}} - 11.91$.

$\log K_p$ of PAHs based on differentiating sampling of the gaseous and particulate fractions was found to be correlated with $1/T$ (or $\log p_L^0$) in a number of studies.^{92,96–98} Therefore, as a general trend a doubling of θ per ≈ 5 K temperature decrease had been found,⁹⁵ θ being significantly higher in winter,^{10,25,99,100} similar to other semivolatile aerosol constituents. This temperature trend is in accordance with both adsorptive and absorptive mechanisms.⁹⁵ However, K_{OA} was suggested to be a better descriptor for PAH gas–particle partitioning than p_L^0 .¹⁰¹ Measurement of f_{OM} has partly been included, however. Over-prediction of θ by the surface adsorption model had been suggested by other studies at both source and receptor sites.⁶⁵ In these studies, however, S has been estimated and assumed to be constant, rather than measured. Gas–particle partitioning of PAHs is presently understood as having adsorptive and absorptive contributions. In particular, a combination of absorption into OM and adsorption onto soot seems to have a high potential to explain the observations.^{102,103}

With, again, the assumptions $\zeta_{\text{OCT}} = \zeta_{\text{OM}}$ and $\text{MW}_{\text{OCT}}/\text{MW}_{\text{OM}} = 1$ adsorptive and absorptive terms can be combined to eqn (8) (Lohmann and Lammel;¹⁰³ similar equation suggested by Dachs and Eisenreich¹⁰²).

$$K_p = 10^{-12} [f_{\text{OM}}/\rho_{\text{oct}}K_{\text{Oa}} + f_{\text{BC}}/\rho_{\text{BC}}(S_{\text{BC}}/S_{\text{soot}})K_{\text{soot-air}}] \quad (8)$$

with f_{BC} being the mass fraction of soot (black carbon) in particulate matter, $K_{\text{soot-air}}$ the soot–air partitioning coefficient, ρ_{oct} and ρ_{BC} the densities of octanol and soot, respectively, and S_{BC} and S_{soot} available specific surface areas. Diesel soot is taken as representative for soot. A particularly high affinity to particulate matter and to soot was found for the semivolatile PAHs Ant and BaA.^{96,104}

While f_{OM} , likely to influence absorption, and soot, f_{BC} , likely to influence adsorption, had been covered in field studies,^{65,88,105,106} the aerosol surface, S , has only rarely been covered so far,^{10,107} and never together with the other parameters, *i.e.* f_{OM} , f_{BC} , S_{BC} and S_{soot} . Moreover, a more differentiated specification of carbonaceous fractions of PM would be needed in order to advance the understanding of PAH's gas–particle partitioning.

Matrix specific adsorption to and absorption of PAH in particulate matter components can in principle be described accounting for all types of chemical interaction at the molecular level (so-called poly-parameter linear free energy relationships).¹⁰⁸ The substance specific^{109,110} and many relevant aerosol matrix parameters are available to quantify PAH's electron donor and acceptor, and van der Waals interactions (for adsorption) and a cavity formation term accounting for solvation (absorption into the particulate matter organic phase), including their humidity dependence.^{111,112} However, there are still limitations due to data gaps in the experimentally determined substance and matrix parameters.

3.1.1.1. PAH mass size distribution. Upon emission in combustion processes and close to the sources mass size distributions of PAHs are peaking in the ultrafine and accumulation

modes (0.01–0.1 and 0.1–1 μm aerodynamic diameter size ranges, respectively). One or two mass size modes are found at urban and rural sites, with mass median diameters, D_m , almost exclusively in the accumulation mode (0.1–1.0 μm of size).^{113–118} This should be seen in the context of limitations by particle size resolution (number of impactor stages) and analytical sensitivity. During atmospheric transport re-distribution in the aerosol occurs, at least of the semivolatile PAHs, which undergo gas–particle partitioning. Their mass size distribution is determined by condensation to surfaces, which leads to a maximum in the accumulation mode^{113,119} and the number median diameter is in the Aitken mode. This was also confirmed for PAHs.^{113,119–121}

3.1.2. Dry and wet deposition contributions. Wet scavenging of hydrophobic organic substances (low Henry coefficients, K^H , listed in Table 1) is expectedly of low efficiency, unless the substance is particle-associated.¹²² Atmospheric lifetimes with respect to removal by dry and wet deposition at a Central European site were ranging between <2 weeks and >2 weeks, respectively.¹²¹ Correspondingly, for temperate climates dry deposition of PAHs may dominate over wet deposition, although this has not been substantiated yet. Atmospheric lifetime of PAHs towards removal by dry deposition is dominated by particle deposition.^{122,123} However, for individual PAHs the opposite may apply, as was found for anthracene, fluoranthene and pyrene at a suburban site,²⁸ and is obviously related to air–soil exchange and, hence, to the concentration of PAH in soil subject to volatilisation.

Particulate dry deposition fluxes of individual PAHs were measured and mean particle dry deposition velocities ranged between 0.4–10 cm s^{-1} .^{28,64,124,125} In samples of deposited particulate matter 3-ring PAHs were found enriched as compared to air samples, *i.e.* had higher dry deposition velocities than 4- and 5-ring PAHs.²⁸

When there is no exchange of material between the particulate and dissolved phases in rain, the total scavenging ratio, W_t , *i.e.* apparent volume of air washed out by volume of rainwater of a given compound can be expressed by:

$$W_t = 10^3 c_r/c_a = W_g \times (1 - \theta) + W_p \times \theta \quad (9)$$

where c_r and c_a are the rainwater and near-ground air concentrations in units of ng L^{-1} and ng m^{-3} , respectively.

The gaseous PAH scavenging ratio, W_g , *i.e.* the apparent volume of air washed out by volume of rainwater, ranges from 5×10^2 – 5×10^5 in precipitation events.^{63,121,126,127} W_g was found to be in good agreement with predictions,¹²⁶ suggesting that phase equilibria were established.^{128,129} Besides dissolution, adsorption of gaseous molecules onto the surface of rain drops or snow flakes could contribute to the phase transfer process.¹²⁸ The particle scavenging ratio W_p is dependent on particle size and chemical surface properties, rain intensity, raindrop diameter and collision efficiency,¹²⁹ as it is the carrier's (particle) washout efficiency which is determining the process. Typical values of W_p of PAH are mostly in the range $W_p = 5 \times 10^2$ – 10^5 ,^{63,121,124,130,131} and even higher in a tropical climate.¹²⁷ In conclusion, the particle scavenging contribution

to the total scavenging exceeds the gas scavenging contribution. The particle scavenging contribution, $W_p \times \theta/W_t$, accounted for 14–100% in a temperate (Central Europe¹²¹) and 86–100% in a tropical environment (Singapore⁶⁵). More efficient scavenging for semivolatile PAHs than for non-volatile PAHs found in various studies was attributed to a different mass size distribution.¹³⁰

Snow is considered to be a more efficient scavenger for nonpolar organics in air than rain.^{124,132} Efficiency is negatively correlated with the vapour pressure of a substance, indicating that adsorption onto the air–ice interface is the process responsible for vapour scavenging.¹³³

3.1.3. Recommendations for future work. • Artefact-free (*e.g.*, *in situ*) method to determine PAH in the gas and particulate phases at remote sites

- Studies of PAH air–soil exchange
- Improved characterisation of the carbonaceous fraction of PM in field and laboratory studies in order to better account for specific sorption of PAHs to these fractions and, hence, better understand PAH's gas–particle partitioning
- Extension of sorption studies to fill the data gaps describing chemical interaction of PAH at the molecular level with regard to relevant aerosol matrix parameters and their humidity dependence (for poly-parameter linear free energy relationships).

3.2. Chemical reactions

3.2.1. Gas phase PAH reaction chemistry. In the troposphere, lower molecular weight (LMW) PAHs exist predominantly in the gas phase due to their relatively high vapour pressure.^{134–137} As discussed in Section 3.1.2, wet and dry deposition of gas phase PAHs is unlikely to be a significant atmospheric loss process for these compounds.^{121,122,138,139} Additionally, no evidence has been observed for direct photolysis as a sink for gas phase PAHs.¹³⁵

The key sink for PAHs in the atmosphere is chemical reactions with atmospheric oxidants, *i.e.* OH, NO₃ and O₃.^{135,140–142} While it has been suggested that reaction with Cl atoms could also be an important loss process for gas phase PAHs in the marine boundary layer and coastal environments,¹⁴³ there is, as yet, no conclusive evidence of this process being significant in the wider environment. Direct reaction with NO₂ or HNO₃ is also not an important atmospheric loss process.¹⁴⁴

A significant amount of pioneering work investigating the kinetics of atmospheric reactions of PAHs over the past 30 years has been conducted by Atkinson, Arey and co-workers at the University of California, Riverside, USA. Through this research, much insight has been gained into the reaction chemistry (rate coefficients, mechanisms and products) of atmospheric reactions of gas phase 2–4 ring PAHs with OH, NO₃ and O₃ and their relative importance in the removal of PAHs from the atmosphere and conversion to derivative compounds.

Here, we review the reactions of PAHs with OH radicals, NO₃ radicals and O₃ and their transformation into nitro- and oxy-derivative compounds.

Calculated rate coefficients, proposed reaction mechanisms and observed products are discussed. Rate coefficient values and reaction products for the reactions of LWM PAHs with OH,

NO₃ and O₃ have previously been reviewed by Atkinson and Arey.¹³⁵ Since this review was published, additional research has been undertaken to validate or enhance our understanding of the reaction kinetics of these processes. We therefore provide an updated overview and discussion for specific 2–4 ring PAHs.

3.2.1.1. Gas phase reactions of PAHs with OH radicals

3.2.1.1.1. Overview. The hydroxyl radical, OH, plays an essential role in atmospheric behaviour and transformations of organic compounds.^{145–147} For a wide range of organic compounds, reaction with OH represents an important removal mechanism from the atmosphere and is also a potential pathway for the transformation into more toxic compounds.¹⁴⁸

In the troposphere, OH radicals can be formed from a series of reactions initiated by the photolysis of ozone in the presence of water vapour:^{147,149}

Other mechanisms involve the reaction of RO₂ with NO,¹⁴⁵ and photolysis of nitrous acid and formaldehyde.¹⁵⁰

The average 24 h OH concentration has been estimated to be 1.0×10^6 molecules cm^{−3}, but is generally reported as 2.0×10^6 molecules cm^{−3}, a daytime average adjusted for a 12 h average daytime period.^{151–153}

Due to the nature of its formation and reaction chemistry, OH radical concentrations are dependent on temporal and meteorological factors including time of day, season, latitude, cloud cover, and relative humidity^{151,152} as well as concentration of O₃ (a key source compound) and NO₂ (a key sink reactant). For example, the highest OH concentrations are expected in the tropics where humidity and UV light intensity levels will be highest.¹⁵²

3.2.1.1.2. Reaction mechanisms. Despite the considerable amount of research into the reactions of PAHs with OH and calculations of rate coefficient values for these reactions, the precise mechanisms by which these processes occur are not fully understood. Experimental studies have a limited capability to assess these reaction routes due to the lack of analytical methods to detect the radical intermediate species believed to be involved in these processes.¹⁵⁴ Therefore a combination of experimental and theoretical approaches is required.

The potential mechanism(s) for the reaction of OH radicals with aromatic hydrocarbons have been discussed by Atkinson¹⁴⁷ and Atkinson and Arey¹³⁵ and more recently by Atkinson and Arey.¹⁵¹

The reaction of PAHs with OH radicals can be initiated *via* two possible reaction pathways. These are illustrated for the acenaphthene and acenaphthylene in Scheme 1. The proposed pathways involve: (1) OH radical interaction with substituent groups either through H-atom abstraction from C–H groups (Scheme 1a), or, in the case of acenaphthylene which contains

Scheme 1 Mechanism for the reaction of gas-phase PAHs with OH radicals; (a) H-atom abstraction; (b) OH addition to substituent groups; (c) OH addition to the aromatic ring.^{135,147}

an unsaturated cyclopentafused ring, addition to the $>C=C<$ bond of this substituent (Scheme 1b),^{135,153} or (2) OH addition to the aromatic ring, forming an initially energy-rich hydroxycyclohexadienyl-type radical intermediate (hereafter referred to as the PAH-OH adduct), which then further reacts with NO₂ or O₂, to form products or thermally decompose back to reactants^{135,155} (Scheme 1c).

The relative importance of the two proposed reaction pathways is dependent on the reaction temperature and pressure conditions. While the OH-addition mechanism is expected to dominate at room temperature, at elevated temperatures, H-atom abstraction from the C-H bonds will become of increasing importance as the PAH-OH adduct will be too thermally unstable and will decompose back to the original reactants.^{147,151} This shift in the reaction mechanism is indicated by the observation of non-exponential OH radical decay at higher temperature ranges.^{147,151}

The dominance of the OH-addition pathway at room temperature has been demonstrated for naphthalene,¹⁵⁶ anthracene^{157,158} and phenanthrene.^{159,160} These studies have generally reported a negative dependence of the rate constant on temperature. Rate coefficient expressions to fit these experimental data, derived from these studies, are shown in Table 3. A lack of a significant 'isotope effect' *i.e.* differences in rate coefficients between unsubstituted PAHs and their deuterated equivalent, as noted by Ananthula *et al.*^{157,160} and Lee *et al.*¹⁵⁹ also indicates a minor contribution of the H-abstraction mechanisms over the temperature ranges studied.

Studies indicate that the OH-addition mechanism is the dominant reaction route up to 773 K for anthracene,¹⁵⁷ 525 K for naphthalene¹⁵⁶ and $\sim$ 380 K for monocyclic aromatic compounds.¹⁶¹ It is therefore suggested that an increase in molecular size (number of aromatic rings) could have a positive effect on the OH-addition rate and/or the thermal stability of the PAH-OH adduct, with 3-ring PAHs being more stable at

higher temperatures than 2-ring structures. Furthermore, Brubaker and Hites¹⁴⁰ investigated the kinetics of the 4-ring PAH fluoranthene and indicated that this compound would form a more stable OH-adduct than 2- or 3-ring structures.

Ananthula *et al.*¹⁶⁰ noted from assessment of previous kinetic studies that the rate of H-atom abstraction at elevated temperatures appears to be largely unaffected by the molecular size. This is presumably due to the dependence of reactivity on the number of possible reaction sites, which does not vary significantly between 1-, 2- and 3-ring species. The reason for the apparent dependence of the OH reaction rate upon molecular size at lower (< 773 K) temperatures is not currently known.

3.2.1.1.3. Reactions of the PAH-OH adducts. Reactions of the PAH-OH adduct with NO₂ are expected based on the observation of nitro-PAH compounds in a wide number of ambient measurements with relative concentrations and isomer profiles consistent with their formation from gas phase reactions (as discussed in detail in Section 5). Indeed, initial investigations into the mechanism(s) of PAH reactions suggested that the PAH-OH adduct will react with NO₂ to form nitro-derivative products^{162–164} and would not react with O₂ to a significant degree.

However, kinetic and product data from experimental studies now suggest that for OH-PAH adducts, the reaction with O₂ dominates over NO₂ reactions under atmospheric conditions.^{151,165,166} It is suggested that the NO₂ and O₂ reactions with the naphthalene-OH adduct may be of equal importance for NO₂ mixing ratios down to 60 ppbV.¹⁶⁶

In addition to laboratory studies, the reactions of naphthalene with OH radicals and subsequent reactions of the Nap-OH adduct have been investigated theoretically in studies by Ricca and Bauschlicher¹⁶⁷ and Qu *et al.*¹⁵⁴ Their investigations of the bond energies of the OH-naphthalene adduct indicate that OH will react with naphthalene at either C₁ or C₂ without an activation energy barrier above the reactants. It was inferred that, after formation of the OH-Nap adduct, decomposition back to reactants is the most favourable pathway.

Qu *et al.*¹⁵⁴ also carried out theoretical study of the mechanism of the subsequent reactions of the Nap-OH adduct with NO₂ and/or O₂ using molecular orbital calculations. The derived pathways for these reactions are summarized in Scheme 2.

It was suggested that H-atom abstraction from the OH-naphthalene adduct would primarily yield the closed-ring 1- and 2-naphthol isomers, from C₁ and C₂ interaction of OH respectively, which would be more stable than open-shell intermediate species^{154,167} (Scheme 2a). It was indicated that the C₁-isomer will be more stable than the C₂-isomer for these hydroxyl products with a ratio of 2 : 1.¹⁶⁷

It was expected that the C₁ and C₂ naphthalene-OH adduct would undergo a competitive decomposition with O₂ and/or NO_x. The N atom of NO₂ can attack the OH-Nap adduct at either the *trans*- or *cis*-position of OH. Attack at the *trans*-position is expected to yield 1- and 2-nitronaphthalene isomers

Table 3 Second-order rate coefficients $k^{(2)}$ for gas-phase reactions of PAH with OH radicals

	k_{OH} ($\text{cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$)	Ref.	T (K)	Notes
Nap	2.4×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR relative to $k(1,2,3\text{-trimethylbenzene}) = 3.27 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$
	2.2×10^{-11}	Atkinson ¹⁴⁷	298	Recommended value based on previous data, overall uncertainty of $\pm 30\%$
	2.3×10^{-11}	Brubaker and Hites ¹⁴⁰	298	Measured over the temperature range 306–366 K
	2.7×10^{-11}	Klamt ¹⁷²	n/a	Theoretical calculation based on a new molecular orbital based estimation method
	2.4×10^{-11}	Biermann <i>et al.</i> ¹⁶⁹	298 ± 1	RR, relative to $k(\text{propene}) = 2.63 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$
	1.9×10^{-11}	Lorenz and Zellner ¹⁵⁶	300	Absolute rate, temperature range 300–873 K, extrapolated using Arrhenius parameter
	2.2×10^{-11}	Klöpffer <i>et al.</i> ¹⁷⁰	300	RR, relative to $k(\text{ethene}) = 8.44 \times 10^{-12} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$
	2.4×10^{-11}	Atkinson <i>et al.</i> ¹⁶⁸	294 ± 1	RR, relative to $k(n\text{-nonane}) = 1.07 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$
1M-Nap	2.6×10^{-11}	Atkinson and Aschmann ¹⁴⁴	295 ± 1	RR, relative to $k(2\text{-methyl-1,3-butadiene}) = 1.02 \times 10^{-10} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$
	4.1×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
	5.3×10^{-11}	Atkinson and Aschmann ¹⁷³	298 ± 2	RR, 2-methyl-1,3-butadiene used as reference compound, $T = 298 \pm 2$
2M-Nap	6.0×10^{-11}	Klamt ¹⁷²	n/a	Theoretical calculation based on a new molecular orbital based estimation method
	4.9×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
	5.2×10^{-11}	Atkinson and Aschmann ¹⁴⁴	295 ± 1	RR, relative to $k(2\text{-methyl-1,3-butadiene}) = 1.02 \times 10^{-10} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$
1E-Nap	5.7×10^{-11}	Klamt ¹⁷²	n/a	Theoretical calculation based on a new molecular orbital based estimation method
	3.6×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
	4.0×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1,2DM-Nap	6.0×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
	2.2×10^{-11}	Banceu <i>et al.</i> ¹⁷⁵	295	RR [relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$]
	7.5×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1,4DM-Nap	5.8×10^{-12}	Klamt ¹⁷²	n/a	Theoretical calculation based on a new molecular orbital based estimation method
	5.8×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1,5DM-Nap	6.0×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
	6.3×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1,7DM-Nap	6.8×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
	6.3×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
2,3DM-Nap	6.2×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
	1.0×10^{-10}	Klamt ¹⁷²	n/a	Theoretical calculation based on a new molecular orbital based estimation method
	7.7×10^{-11}	Atkinson and Aschmann ¹⁴⁴	295 ± 1	RR, relative to $k(2\text{-methyl-1,3-butadiene}) = 1.02 \times 10^{-10} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$
2,6DM-Nap	6.7×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
	6.9×10^{-11}	Phousongphouang and Arey ¹⁷¹	298 ± 2	RR, relative to $k(\text{naphthalene}) = 2.39 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
Ace	8.0×10^{-11}	Reisen and Arey ¹⁵⁵	296	RR [relative to $k(\text{trans-2-butene}) = 6.48 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$]
	5.8×10^{-11}	Brubaker and Hites ¹⁴⁰	298	Measured over the temperature range 325–365 K
	1.0×10^{-10}	Atkinson and Aschmann ¹⁷³	296 ± 1	RR [relative to $k(2,3\text{-dimethyl-2-butene}) = 1.11 \times 10^{-10} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$]
	5.8×10^{-11}	Klöpffer <i>et al.</i> ¹⁷⁰	300	RR [relative to $k(\text{ethene}) = 10^{-12} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$]
	6.4×10^{-11}	Banceu <i>et al.</i> ¹⁷⁵	295	RR [relative to $k(\text{naphthalene}) = 2.2 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$]
	8.0×10^{-11}	Klamt ¹⁷²	n/a	Theoretical calculation based on a new molecular orbital based estimation method

Table 3 (continued)

	k_{OH} (cm ³ molecules ⁻¹ s ⁻¹)	Ref.	T (K)	Notes
Acy	1.2×10^{-10}	Reisen and Arey ¹⁵⁵	296	RR [relative to $k(\text{trans-2-butene}) = 6.48 \times 10^{-11}$ cm ³ molecules ⁻¹ s ⁻¹]
	1.3×10^{-10}	Banceu <i>et al.</i> ¹⁷⁵	295	RR [relative to $k(\text{naphthalene}) = 2.2 \times 10^{-11}$ cm ³ molecules ⁻¹ s ⁻¹]
	1.1×10^{-10}	Atkinson and Aschmann ¹⁷³	296 ± 1	RR [relative to $k(2,3\text{-dimethyl-2-butene}) = 1.11 \times 10^{-10}$ cm ³ molecules ⁻¹ s ⁻¹]
Fln	1.6×10^{-11}	Kwok <i>et al.</i> ¹⁷⁶	297	Placed on an absolute basis by using $k_2(\text{cyclohexane}) = 7.47 \times 10^{-11}$ cm ³ molecules ⁻¹ s ⁻¹
	1.3×10^{-11}	Brubaker and Hites ¹⁴⁰	298	Measured over the temperature range 326–366 K
	9.9×10^{-12}	Klamt ¹⁷²	n/a	Theoretical calculation based on a new molecular orbital based estimation method
	1.3×10^{-11}	Klöpffer <i>et al.</i> ¹⁷⁰	300	RR [relative to $k(\text{ethene}) = 7.47 \times 10^{-12}$ cm ³ molecules ⁻¹ s ⁻¹]
Phe	3.4×10^{-11}	Biermann <i>et al.</i> ¹⁶⁹	298 ± 1	RR [relative to $k(\text{propene}) = 4.85 \times 10^{-12}$ e ^{504/T} cm ³ molecules ⁻¹ s ⁻¹]
	3.1×10^{-11}	Atkinson ¹⁴⁷	298	Recommended value based on previous data, overall uncertainty of ±30%
	2.6×10^{-11}	Klamt ¹⁷²	n/a	Theoretical calculation based on a new molecular orbital based estimation method
	1.6×10^{-11}	Lorenz and Zellner ¹⁵⁶	338	Absolute rate study, measured over a temperature range 338–748 K
	1.3×10^{-11}	Kwok <i>et al.</i> ¹⁷⁶	296	RR [relative to $k(\text{propene}) = 2.66 \times 10^{-11}$ cm ³ molecules ⁻¹ s ⁻¹]
	2.7×10^{-11}	Brubaker and Hites ¹⁴⁰	298	Measured over the temperature range 346–386 K, extrapolated using Arrhenius parameters
	3.2×10^{-11}	Lee <i>et al.</i> ¹⁵⁹	298	Measured over the temperature range 298–386 K, extrapolated using Arrhenius parameters
	$4.98 \pm 2.96 \times 10^{-6} \text{ T}^{-1.97 \pm 0.10}$	Ananthula <i>et al.</i> ¹⁶⁰	373–1000 K	Two-parameter expression to best fit experimental data
	2.9×10^{-11}	Lee <i>et al.</i> ¹⁵⁹	298	Measured over the temperature range 363–403 K, extrapolated using Arrhenius parameters
2M-Phe	6.5×10^{-11}	Lee <i>et al.</i> ¹⁵⁹	298	Measured over the temperature range 338–398 K, extrapolated using Arrhenius parameters
3M-Phe	6.6×10^{-11}	Lee <i>et al.</i> ¹⁵⁹	298	Measured over the temperature range 353–388 K, extrapolated using Arrhenius parameters
9M-Phe	7.6×10^{-11}	Lee <i>et al.</i> ¹⁵⁹	298	Measured over the temperature range 333–373 K, extrapolated using Arrhenius parameters
Ant	1.1×10^{-10}	Biermann <i>et al.</i> ¹⁶⁹	325 ± 1	RR [relative to $k(\text{propene}) = 2.29 \times 10^{-11}$ cm ³ molecules ⁻¹ s ⁻¹]
	1.9×10^{-10}	Brubaker and Hites ¹⁴⁰	298	Measured over the temperature range 346–365 K
	1.3×10^{-11}	Kwok <i>et al.</i> ¹⁷⁶	296	Based on a derived $k(\text{anthracene})/k(\text{phenanthrene})$ value of 1.0 ± 0.5
	2.0×10^{-10}	Klamt ¹⁷²	n/a	Theoretical calculation based on a new molecular orbital based estimation method
	1.3×10^{-10}	Atkinson; ¹⁴⁷ Biermann <i>et al.</i> ¹⁶⁹	298	Recommended value based on previous data, overall uncertainty of ±30%
	$1.12 \times 10^{-10} (\text{T}/298)^{-0.46}$	Goulay <i>et al.</i> ¹⁵⁸	58–470	Two-parameter expression to best fit experimental data
	$8.17 \times 10^{-14} \text{ T}^{-8.3} \text{ e}^{(-3171.71/\text{T})}$	Ananthula <i>et al.</i> ¹⁵⁷	373–923	Modified Arrhenius equation to best fit experimental data
	$2.18 \times 10^{-11} \text{ e}^{(-1734.11/\text{T})}$	Ananthula <i>et al.</i> ¹⁵⁷	999–1200	Modified Arrhenius equation to best fit experimental data
Flt	1.1×10^{-11}	Brubaker and Hites ¹⁴⁰	298	Measured over the temperature range 346–366 K
Pyr	5.0×10^{-11}	Atkinson <i>et al.</i> ¹⁶²	296 ± 2	RR Relative to $k(\text{naphthalene}) = 3.6 \times 10^{-28}$ cm ³ molecules ⁻¹ s ⁻¹
1N-Nap	5.4×10^{-11}	Atkinson ¹⁴⁷	298	Recommended value
2N-Nap	5.6×10^{-11}	Atkinson ¹⁴⁷	298	Recommended value

for the reactions of the C₂ and C₁ adducts respectively (Scheme 2bi), while attack at the *cis*-position would yield 2-hydroxy-1-nitronaphthalene and 1-hydroxy-2-nitronaphthalene (Scheme 2bii). It was indicated that the nitronaphthalene isomers will be the dominant products formed.

O₂ is expected to attack the C₁- and C₂-adducts forming energy-rich intermediates, which will decompose *via* different pathways to yield either 2-formylcinnamaldehyde (Scheme 2ci) or 1- or 2-naphthol isomers (Scheme 2cii). It was also indicated that naphthols can further react with HO₂ which yields

1,4-naphthoquinone (1,4-NQu) (Scheme 2d). It was suggested that 1,4-NQu can also be formed from 1-naphthol by further addition of OH to yield a diol intermediate which can rearrange to form 1,4-diketone followed by loss of H₂.¹⁶⁷ Furthermore, it is indicated that the C₁ and C₂ adducts can react with an NO–O₂ mixture to form an additional route to 2-formylcinnamaldehyde (FCA) (Scheme 2e).

3.2.1.1.4. Rate coefficients. Experimental studies have been carried out to derive second-order rate coefficients for the gas

Scheme 2 Proposed mechanisms for the further reaction of the PAH-OH adduct: (a) H-atom abstraction; (b) reactions with NO₂; (c) reaction with O₂; (d) further reaction of 1-naphthol; (e) reaction with NO/O₂.¹⁵⁴

phase reaction of OH with a number of individual 2–4 ring PAHs, including naphthalene, alkyl naphthalene derivatives, acenaphthene, acenaphthylene, anthracene, phenanthrene, fluorene, fluoranthene and pyrene. Values of the calculated second-order rate coefficient, $k^{(2)}$, derived in these studies and the experimental conditions used are detailed in Table 3.

While for most studies experimental investigation of $k^{(2)}$ values for PAH reaction with OH has been conducted over relatively narrow temperature ranges, Lorenz and Zellner¹⁵⁶ investigated the rate coefficients for the reactions of OH with naphthalene and phenanthrene over a wide range of elevated temperatures (~300–873 K). Similarly Brubaker and Hites¹⁴⁰ measured rate coefficients for a range of temperatures above ambient levels and then extrapolated to environmentally relevant temperatures using the Arrhenius equation.

Atkinson¹⁴⁷ recommended a $k^{(2)}$ value for the reaction of naphthalene with OH based on a unit-weighted least squares analysis of data from previous studies by Atkinson *et al.*,¹⁶⁸ Atkinson and Aschmann,¹⁴⁴ Lorenz and Zellner¹⁵⁶ and Biermann *et al.*¹⁶⁹ The value reported by Klöpffer *et al.*¹⁷⁰ was not used in this calculation. Subsequent experimental studies by Brubaker and Hites¹⁴⁰ and Phoussongphouang and Arey,¹⁷¹ and a theoretical value derived by Klamt *et al.*,¹⁷² using molecular orbital calculations, are shown to be in good agreement with this recommended $k^{(2)}$ value.

$k^{(2)}$ values for a range of alkyl naphthalene derivatives (MNs, DMNs and ENs) with OH have been studied experimentally.^{144,171,173} The $k^{(2)}$ values for 1M-Nap, 2M-Nap and 2,3DM-Nap calculated by Phoussongphouang and Arey¹⁷¹ and Atkinson and Aschmann^{144,173} are shown to be in reasonable agreement,

the small difference in measured rate coefficient values being attributed to differences in product quantification methods.¹⁷¹ These $k^{(2)}$ values are also within a factor 2 of the theoretical value derived by Klamt.¹⁷²

$k^{(2)}$ values derived in different experimental studies for the reaction of acenaphthene with OH are clearly not in agreement. Differences between $k^{(2)}$ reported by Reisen and Arey¹⁵⁵ and Atkinson and Aschmann¹⁷⁴ could be attributed to the use of capillary GC columns in the former and packed column GC techniques in the latter, leading to a more accurate measurement in the later study.¹⁵⁵ The $k^{(2)}$ value reported by Klöpffer *et al.*¹⁷⁰ may be erroneously low as the reference compound, ethene, reacts substantially slower than acenaphthene.¹⁵⁵ The $k^{(2)}$ value reported by Reisen and Arey¹⁵⁵ is very similar to the theoretical value derived by Klamt.¹⁷² There is generally good agreement between the three $k^{(2)}$ values derived for the OH reaction with acenaphthylene.^{155,174,175}

For reactions of anthracene, there is reasonable agreement, between $k^{(2)}$ values calculated theoretically by Klamt¹⁷² and experimentally by Brubaker and Hites.¹⁴⁰ $k^{(2)}$ values measured at 325 K by Biermann *et al.*¹⁶⁹ and Goulay *et al.*¹⁵⁸ were in agreement but both lower than that reported by Brubaker and Hites.¹⁴⁰ Kwok *et al.*¹⁷⁶ reported a much lower $k^{(2)}$ based on a $k(\text{Ant})/k(\text{Phe})$ value of 1.0 ± 0.5 derived from samples taken during the reaction of phenanthrene with OH. It has therefore been suggested that the $k^{(2)}$ value measured by Kwok *et al.*¹⁷⁶ may be erroneously low.

For the reaction of phenanthrene with OH, Atkinson¹⁴⁷ recommended a $k^{(2)}$ value based on previous studies by Lorenz and Zellner¹⁵⁶ and Biermann *et al.*¹⁶⁹ This $k^{(2)}$ value is shown to be in good agreement with the subsequent studies by Brubaker and Hites¹⁴⁰ and Lee *et al.*,¹⁵⁹ both extrapolated to 298 K using the Arrhenius relationship, and slightly higher than the theoretical value derived by Klamt.¹⁷² The $k^{(2)}$ value derived by Kwok *et al.*¹⁷⁶ is a factor of two lower than that of the other rate coefficient values in Table 3 for a similar temperature.

The $k^{(2)}$ values for the reaction of fluorene with OH from three experimental studies^{140,170,177} and a theoretical study¹⁷² do agree well. To our knowledge, the only $k^{(2)}$ measurements for the gas phase reaction of fluoranthene and pyrene with OH have been made by Brubaker and Hites¹⁴⁰ and Atkinson *et al.*,¹⁶³ respectively.

It is clear that $k^{(2)}$ values for gas-phase OH reactions depend strongly on the specific PAH molecular structure. For example, as noted by Brubaker and Hites,¹⁴⁰ Ace has a higher reaction rate than Nap. Both compounds consist of two fused aromatic rings but it is suggested that the reactivity of Ace towards OH is enhanced by H-atom abstraction occurring at the cyclopenta-fused ring.¹⁵⁵ Similarly, Acy is shown to be more reactive than Ace. This has been attributed to the higher reactivity of the unsaturated cyclopentafused ring of Acy with a more rapid OH addition to the >C=C< than H-atom abstraction from the saturated cyclopenta-fused ring.^{155,173,174}

The $k^{(2)}$ values for 1M-Nap and 2M-Nap are shown to be essentially identical suggesting that the mechanism of this reaction with OH will be the same for the two isomers. The presence

Table 4 Rate coefficient ratios for the reaction of methylnaphthalenes and methylphenanthrenes with OH radicals, relative to the reactions of their parent compound

Compound	Relative reactivity ^a ($k_{1(\text{obs})}/k_{\text{Nap}(\text{obs})}$)
Nap	1.00
1M-Nap	1.71
2M-Nap	2.02
1E-Nap	1.52
2E-Nap	1.68
1,2DM-Nap	2.49
1,3DM-Nap	3.13
1,4DM-Nap	2.42
1,5DM-Nap	2.51
1,6DM-Nap	2.65
1,7DM-Nap	2.84
1,8DM-Nap	2.62
2,3DM-Nap	2.57
2,6DM-Nap	2.78
2,7DM-Nap	2.87

Compound	Relative reactivity ^b ($k_{1(\text{obs})}/k_{\text{Phe}(\text{obs})}$)
Phe	1.00
1M-Phe	0.90
2M-Phe	2.02
3M-Phe	2.08
9M-Phe	2.36

^a From Phousongphouang and Arey.¹⁷¹ ^b From Lee *et al.*¹⁵⁹

of alkyl groups has been observed to have an 'activating effect' on OH reactivity for Nap¹⁷¹ and Phe¹⁵⁹ (Table 4). In contrast, Banceu *et al.*¹⁷⁵ reported $k^{(2)}$ values for 1,3DM-Nap and 2,3DM-Nap, which are lower than that of Nap. Furthermore, enhanced reactivity is not observed for 1-methylphenanthrene (1M-Phe). As yet, it is not known why 1M-Phe is not similarly activated.

It is indicated that the reactivity of Ant towards OH is much faster than that of other 3-ring PAHs such as Phe and FlN.

The difference between Ant and FlN reactivity could possibly be explained by differences in the reaction mechanism at C₉ which is the most likely position for OH attack for these compounds.^{178,179} It is suggested, based on reaction product studies (see Section 3.2.1.1.5.), that the H-abstraction mechanism will dominate at the C₉ position for FlN,¹⁷⁸ while the OH addition mechanism will dominate for reactions of Ant,¹⁴⁰ which could explain the enhanced reactivity of Ant relative to FlN.

Ant and Phe have essentially the same 3-ring structure, only differing by the relative position of their aromatic rings. However, Ant appears to be significantly more reactive.

The most reactive sites for Phe are expected to be at the C₁₁ and C₁₂ positions, where net charge is greatest. However, product studies show that reaction occurs predominantly at the C₉ and C₁₀ positions.^{179–182} This suggests that the reactivity at the C₁₁- and C₁₂-positions could be sterically hindered, resulting in lower overall reactivity. Conversely, Ant is expected to have highest net charge at C₉ and C₁₀, which are sterically unhindered. This is confirmed by product studies which indicate reactions occurring at these sites,^{158,180} which possibly explain the relatively enhanced reactivity.

3.2.1.1.5. Reaction products. Products from the reactions of LMW PAHs with OH radicals in the presence of O₂ and NO_x

have been investigated in a number of laboratory studies. More attention in the literature has been paid to the reactions of naphthalene but product investigations have also been carried out for alkyl naphthalenes, acenaphthene, acenaphylene, phenanthrene, anthracene, fluoranthene and pyrene.

3.2.1.1.5.1. Naphthalene reaction products. The gas phase reaction of Nap with OH radicals has been shown to yield both ring-retaining and ring-opened products. Key ring-retaining products identified include naphthols, nitronaphthalenes and naphthoquinones.

1- and 2-naphthol isomers have been identified in product studies^{162,181–183} in yields of approximately 3–7%. The formation of naphthol isomers is expected to result from OH–Nap adduct rearomatization (Scheme 2a) or reaction with O₂ and decomposition of the peroxy radical (Scheme 2cii).

1- and 2N-Nap isomers are also formed in these reactions.^{162,166,181,183,184} This would result from the pathway shown in Scheme 2b. Nishino *et al.*¹⁶⁶ measured the formation yields of 1- and 2N-Nap to be 0.35% and 0.60% respectively. The higher yield observed for reaction occurring at C₁ compared to C₂ was also noted by Wang *et al.*¹⁸⁵ who indicated that 68% of OH addition occurs at C₁. This is in agreement with the theoretical calculations of Qu *et al.*,¹⁵⁴ who calculated that the C₁ adduct lies 10 kcal mol^{−1} lower in energy than the C₂ adduct.

1,2- and 1,4-NQu isomers have also been observed as reaction products in yields of 1–6%.^{181,183,184} This could be expected to occur *via* the reaction in Scheme 2d. An alternative pathway is suggested by Lee and Lane¹⁸¹ and Kautzman *et al.*¹⁸⁴ *via* 1-hydroxynaphthalene-2-one and 1-hydroxynaphthalene-4-one to yield 1,2-NQu and 1,4-NQu respectively (Scheme 3). It has also been suggested that 1,4-NQu can be formed from the photolysis of 1-nitronaphthalene.¹⁶² Other ring retaining products identified from the reaction of Nap with OH include 2,3-epoxy-naphthoquinone, oxygenated indenenes such as inden-1-one and 1,3-indene-dione, benzopyrones and nitro-naphthol isomers.¹⁸¹

Sasaki *et al.*¹⁸² reported that ring-retaining products represented only ~30% of the reaction products from this process. Therefore understanding the ring-opening mechanism is important to fully understand these reactions. In particular, much interest has been paid to the formation and reactivity of the ring-opened species 2-FCA, identified as the main reaction product of naphthalene with OH.^{166,181–186}

The product yield of 2-FCA from this reaction was estimated to be 46–71% by Nishino *et al.*¹⁸⁶ Other studies have reported lower formation yields^{182,183} but these did not take the rapid loss of 2-FCA due to photolysis and reactivity with OH into account.¹⁸⁶ The formation of FCA can be attributed to the mechanism shown Scheme 2ci, involving Nap–OH adduct reaction with O₂ to form the peroxy radical intermediate or 2e involving reaction with O₂ and NO, as predicted by Qu *et al.*¹⁵⁴

Other ring opened products identified include phthalic acid, phthalaldehyde and phthalic anhydride.^{181,182,184} Kautzman *et al.*¹⁸⁴ suggested that these were formed from the further reaction of 2-FCA with OH. An alternative reaction pathway was

also suggested involving reaction with O₂ to yield C₇ and C₉ compounds such as benzoic acid.

Kautzman *et al.*¹⁸⁴ investigated the products from the naphthalene reaction with OH under both high (~80 ppbV) and low NO_x conditions. Two reaction pathways were proposed to yield 2-FCA: (i) a series of reactions initiated by the reaction of the Nap–OH adduct with O₂ to yield a peroxy radical intermediate; (ii) reaction of the Nap–OH adduct involving O₂ and NO. These are summarized in Scheme 4. These mechanisms are in agreement with other product studies^{182,183} and theoretical calculations.¹⁵⁴ It is not currently possible to establish the relative dominance of either of these proposed routes.

The presence of organic peroxides in the product study by Kautzman *et al.*¹⁸⁴ was attributed to an alternative reaction mechanism, initiated similarly by the reaction of the Nap–OH adduct with O₂, but proceeding *via* the formation of a bicyclic peroxy radical (Scheme 5). This mechanism is suggested to lead to the formation of both observed ring-opened species as well as peroxide or epoxide compounds.

Under low NO_x conditions, Kautzman *et al.*¹⁸⁴ indicated that the dominant mechanism for the formation of products from naphthalene will differ from that at higher NO_x levels. It was suggested that the reaction is dominated by the alkoxy radical reaction predominantly with HO₂ to form ring-opened products such as 2-FCA. Under these low NO_x conditions, 1,4-NQu is suggested to result from the reaction pathway detailed in Scheme 2d involving the addition of HO₂ to 1-naphthol followed by reaction with O₂ and two further additions of HO₂.¹⁸⁴

The identified products are broadly consistent with the suggested reaction mechanisms of Qu *et al.*¹⁵⁴ (as detailed in Scheme 2). The product studies detailed above have allowed the identification of >90% of the products formed from the reaction of naphthalene with OH radicals. This has allowed relatively detailed reaction pathways for these processes to be suggested, despite some discrepancies in the literature. However, obtaining a full mass balance for this reaction is difficult due to the secondary reactions of the PAH reaction products.¹⁸³ More work is clearly needed to fully understand these precise reaction mechanisms.

3.2.1.1.5.2. Reactions of other PAHs. Products identified from the reactions of Phe have included ring-opened species such as 2,2'-diformylbiphenyl, 9-fluorenone (9-Flr), 1-,2-,3-,4-, and 9-phenanthrols, 2-,3-,4- and 9-nitrophenanthrenes (N-Phe), 1,4- and 9,10-phenanthraquinones (PQu), dibenzopyranone and 2- and 4-nitrodibenzopyranones, and ring-opened products such as 1,2-naphthalenedicarboxaldehyde and 1,2-naphthalic anhydride.^{180,185,187}

Corresponding to the OH addition mechanism, it was suggested that the Phe–OH adduct would be formed *via* addition of OH to C₁.^{182,187} This can then form the OH–phenanthrene–O₂ peroxy radical, analogous to the reaction of naphthalene (Scheme 2), from which stable products can be formed through reaction with O₂ (see Lee and Lane,¹⁸⁷ for detailed discussion). Phenanthrols could also be formed through H-atom abstraction from the Phe–OH adduct.¹⁸⁷ It was suggested that ring-opened products such as 1,2-naphthalenedicarboxaldehyde and 1,2-naphthalic anhydride

Scheme 3 A proposed alternative reaction pathway for the formation of naphthoquinone isomers.¹⁸³

are formed through cleavage of the central ring between C₉–C₁₀, the most active site for electrophilic attack for Phe.^{178,187}

Wang *et al.*¹⁸⁵ suggested that 2,2'-diformylbiphenyl is a primary product in this reaction, which then reacts further to produce 9-fluorenone and dibenzopyranone.^{178,179,185} The suggested pathway leading to PQu involved initial formation of a keto alcohol at the C₉–C₁₀ bond.^{179,185} The calculated formation yield of 9,10-PQu was ~3%.¹⁸⁵ Combining the measured formation yield with the rate coefficient,¹⁵⁹ the authors calculated a PQu formation rate of 80 pg m⁻³ h⁻¹.

Helmig *et al.*¹⁷⁸ identified products from the reaction of Fln with OH in the presence of NO_x, including hydroxyfluorenes,

nitrofluorenes (N-Fln), 9-fluorenone (9-Flr) and nitrofluorenes. The formation of N-Fln isomers is attributed to the OH-addition mechanism analogous to that discussed for Nap and Phen (Scheme 1c and 2bi). The formation of 9-Flr is expected to proceed *via* H-atom abstraction of the –CH₂– group followed by reaction with O₂.¹⁷⁸ The relatively high fluorenone yield (~9%) compared to nitro-fluorenes (~1%) may indicate a higher importance of the H-atom abstraction mechanism for Fln.

Nitro-PAH isomers have been identified from the gas-phase reactions of Ant, Flth and Pyr with OH.^{163,180,188} As with other studies, nitro-PAH product yields are low: 1-nitroAnt (0.2%),

Scheme 4 Summary of pathways to 2-formylcinnamaldehyde from the naphthalene-OH adduct: (a) *via* reaction with O_2 , (b) *via* reaction with $O_2/NO/RO_2$.

Scheme 5 An alternative reaction scheme for the reaction of naphthalene-PAH adduct proceeding *via* a bicyclic peroxy radical.¹⁸⁴

2-nitroAnt (0.2%), 2-nitroPyr (0.5%) and 4-nitroPyr (0.06%) 2-nitroFlt (3%) 7-nitroFlt (1%) and 8-nitroFlt (0.3%).

It is clear that the formation of the observed products from the gas phase reactions of PAHs with OH radicals is initiated by

the formation of the PAH-OH adduct. The nature and extent of the subsequent reaction pathway will then dictate the precise products formed. Different pathways have been suggested leading to the formation of both ring-retaining and ring-opened products.

The relatively low yield of nitro-PAH products from these processes clearly suggests that the reaction of PAH-OH adducts with NO₂ does not dominate and other mechanisms such as peroxy radical formation and reaction/decomposition are more significant.

In general, formation yields of nitro-PAHs are low, ranging from 0.2 to 5%. It should be noted that, for most PAHs, despite this considerable work, the majority of OH-radical reaction products remain unidentified.¹³⁵ While it can be envisaged that reactions of 3- and 4-ring PAHs will proceed in ways analogous to naphthalene, to fully understand the mechanisms of these processes, the products from these reactions need to be investigated in more detail.

3.2.1.2. Gas phase reactions of PAHs with NO₃ radicals

3.2.1.2.1. Overview. The nitrate radical, NO₃, plays an important role in tropospheric chemistry and is thought to be a key contributor to the oxidising capacity towards VOCs in the lower atmosphere.^{189–191}

This has been demonstrated through investigations of NO₃ reactions with alkanes,^{192–194} organosulphur compounds,¹⁹² aldehydes^{192,195–197} and hydroxy-substituted aromatics.¹⁹⁸ These reactions may have further impacts on tropospheric reactions through the generation of peroxy radicals.^{192,199,200} A detailed overview of NO₃ radical chemistry is provided by Geyer *et al.*^{191,201} but a brief discussion is provided here.

The levels of NO₃ in the troposphere are controlled by the complex interplay of the reactions that lead to its formation or removal. In the troposphere, NO₃ radicals are formed *via* the sequential reactions of NO and NO₂ with O₃.^{202–205}

NO₃ can be removed from the troposphere *via* photolysis by solar radiation.^{204–206}

NO₃ reaction with NO will also lead to its removal from the troposphere.²⁰⁷

NO₃ radicals can also be removed from the troposphere *via* reaction with NO₂ to form N₂O₅, which can subsequently thermally decompose to NO₃.^{202,208}

The reaction rates for this process are shown to be the same for M = O₂ and M = N₂.²⁰⁹ It is also suggested that N₂O₅ can be removed by hydrolysis to form HNO₃, which leads to the removal of NO₃ from the atmosphere.¹⁸⁹

Because of the rapid photolysis of NO₃ and reaction with NO and O₃ (as well as the reaction of NO with O₃) the

concentrations of the NO₃ radical in the ambient troposphere will be low during daytime hours.²¹⁰ Concentrations of NO₃ will therefore be present in the troposphere at significant levels only during evening and night time hours when photolysis is absent and NO levels are low,^{135,151,203} except close to major sources of emissions.

Levels of NO₃ in the troposphere will therefore be subject to substantial temporal and spatial variability,¹⁵¹ ranging from $<5 \times 10^7$ molecules cm⁻³ to 1×10^{10} molecules cm⁻³.^{204,210} Atkinson²¹⁰ has suggested an average 12 h night time NO₃ radical concentration in the lower troposphere of $\sim 5 \times 10^8$ molecules cm⁻³ (~ 20 pptV) over continental areas. This value is expected to be significantly lower in marine environments, with mixing ratios of ~ 0.25 pptV measured at 3 km altitude in Mauna Loa, Hawaii,²¹¹ due to lower NO₂ concentrations.^{194,211}

3.2.1.2.2. NO₃ reaction mechanism. N₂O₅, NO₂ and NO₃ are effectively in equilibrium in the lower troposphere.²⁰³ Initial kinetic studies in mixtures of N₂O₅-NO₃-NO₂ suggested that these reactions are kinetically equivalent to the reaction with N₂O₅ alone.^{162,173,174,212} However, in a study by Atkinson *et al.*,²⁰³ 2,3-dimethyl-2-butene was added to the reaction mixture of N₂O₅-NO₃-NO₂-N₂-O₂ and Nap to scavenge NO₃ radicals. The removal of NO₃ was shown to reduce the amount of naphthalene reacted by a factor of up to >100 indicating that Nap does not undergo elementary reaction with N₂O₅ to a significant degree.

It is therefore suggested that the reaction of PAHs in N₂O₅-NO₃-NO₂ mixtures involves the addition of NO₃, formed from the thermal decomposition of N₂O₅, to the aromatic ring, in a reaction analogous to the reaction of OH radicals with PAH. The reaction scheme for this process is represented in Scheme 6. The reaction proceeds *via* initial addition of NO₃ to the aromatic ring to form a nitratocyclohexadienyl-type radical (6a) (the PAH-NO₃ adduct), which, similarly to the PAH-OH adduct can decompose back to reactants (6b), reacts with NO₂ (6c) or O₂ (6d) to form products^{135,151,210} or can undergo unimolecular decomposition (6e) that would lead to the formation of hydroxy-PAH products and further reaction to form nitro-hydroxy-PAHs.²¹³

NO₃ radicals can also interact with the substituent groups of compounds such as Ace, Acy and M-Nap^{210,214} (Scheme 7). For Ace, this involves H-atom abstraction from the C-H bond of the cyclopenta-fused ring and the formation of HNO₃.^{174,180} For Acy, NO₃ addition to the >C=C< bond of the cyclopenta-fused ring is considered to be the dominant reaction pathway.^{174,180} However, for unsubstituted PAHs, the NO₃-addition pathway is expected to dominate.²¹⁰ It is thought that, for PAHs containing substituent groups, reaction will not result in nitro-PAH formation.^{180,210}

The measured rate coefficient for this reaction will therefore be:

$$k_{\text{obs}} = k_a(k_c[\text{NO}_2] + k_d[\text{O}_2] + k_e)/(k_b + k_c[\text{NO}_2] + k_d[\text{O}_2] + k_e)$$

Assuming, as indicated by experimental data,^{162,203,210,212–215} under conditions used in reaction studies and in the ambient troposphere:

$$k_b \gg (k_c[\text{NO}_2]) \text{ and } k_c[\text{NO}_2] > (k_d[\text{O}_2] + k_e)$$

Scheme 6 Potential pathways for the reaction of PAHs with NO_3 .¹⁵¹

Scheme 7 H-atom abstraction mechanism for the reaction of a methyl-substituted PAH with NO_3 .²¹⁴

this equation simplifies to:

$$k_{\text{obs}} = k_a(k_c[\text{NO}_2])/k_b$$

The observed rate coefficient will therefore be proportional to NO_2 concentration. The decomposition rate coefficient of the Nap-OH adduct was calculated to be $5 \times 10^5 \text{ s}^{-1}$. The PAH- NO_3 adduct is therefore more stable towards decomposition compared to the monocyclic- NO_3 adduct (decomposition rate coefficient of $3 \times 10^8 \text{ s}^{-1}$ at 298 K). However, the NO_3 adduct is estimated to be 46 kJ mol^{-1} less stable towards decomposition than the OH-adduct.²¹⁰

It is also suggested from kinetic data that reaction of the NO_3 -PAH adduct with NO_2 will dominate over the reaction with O_2 under atmospherically relevant conditions. Atkinson *et al.*²¹³ measured the rate coefficient of the NO_3 reaction with Nap as a function of NO_2 and O_2 concentration. An upper limit of $k_d/k_c < 4 \times 10^{-7}$ was obtained at $298 \pm 2 \text{ K}$. Therefore the reaction of the Nap- NO_3 adduct with NO_2 is expected to dominate relative to the reaction with O_2 down to a NO_2 mixing ratio of at least 80 ppbV ($2 \times 10^{12} \text{ molecules cm}^{-3}$) and possibly much lower.¹⁵¹

Qu *et al.*²¹⁶ used high-level molecular orbital theory to predict NO_3 -initiated PAH reactions theoretically. It was shown that the Nap- NO_3 adduct is formed by the attack of the O atom of NO_3 at either the C_1 or C_2 positions. The proposed further

reactions of the PAH- NO_3 adduct are illustrated in Scheme 8. Two energetically favourable primary reaction pathways are proposed for the Nap- NO_3 adduct(s) formed. In the first pathway, loss of NO_2 from the adduct leads to the formation of 1,2-epoxynaphthalene or 2,3-epoxynaphthalene for C_1 and C_2 adducts respectively (Scheme 8a). In the second pathway, the isomerisation of the NO_3 -naphthalene adduct followed by unimolecular decomposition leads to the formation of 2-FCA for the C_1 adduct or 2-acetylcinnamaldehyde for the C_2 adduct (Scheme 8b).

It was also indicated that the Nap- NO_3 adduct will also undergo secondary reactions with NO_2 , O_2 or O_2/HO_2 and these reaction pathways will be strongly competitive with the unimolecular decomposition pathway. Mechanisms for the proposed reactions are shown in Scheme 8c-e respectively. The energetically favourable reaction with NO_2 is expected to produce 2-nitronaphthalene and 1-nitronaphthalene isomers and HNO_3 for the reactions of C_1 - and C_2 -adducts respectively, while the reaction with O_2 is expected to produce a nitric ether species. Both of these pathways are expected to proceed *via* the formation and decomposition of an energy-rich intermediate.²¹⁶ The reaction of the Nap- NO_3 adduct with O_2/HO_2 is expected to produce 1,4-NQu. (Scheme 8e).

3.2.1.2.3. Rate coefficients. Relatively fewer rate coefficient studies have been carried out for gas-phase PAH reactions with

Scheme 8 Proposed reaction pathways for the further reaction of the PAH-NO₃ adducts: (a) unimolecular decomposition; (b) isomerisation followed by unimolecular decomposition; (c) reaction with NO₂; (d) reaction with O₂; (e) reaction with O₂-HO₂.²¹⁶

NO₃ than for the corresponding reactions with OH radicals. PAHs with NO₃ radicals, along with the reaction conditions are Measured or derived $k^{(2)}$ values for the gas phase reactions of presented in Table 5.

Table 5 Second-order rate coefficients $k^{(2)}$ for gas-phase reactions of PAH with the NO_3 radical

	$k_{\text{NO}_3} (\text{cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1})$ ($\times [\text{NO}_2]$)	$k_{\text{NO}_3} (\text{s}^{-1})$ $[\text{NO}_2] = 6.91 \times 10^{11}$ molecules cm^{-3} ^a	Ref.	T (K)	Notes
Nap	8.5×10^{-28}	1.1×10^{-16}	Pitts <i>et al.</i> ²¹²	298 ± 2	RR relative to $K_5(\text{NO}_3 + \text{NO}_2 \rightarrow \text{N}_2\text{O}_5) = 3.41 \times 10^{-11} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$
	4.8×10^{-28}	6.2×10^{-17}	Atkinson <i>et al.</i> ¹⁶²	298 ± 2	RR relative to $K_5(\text{NO}_3 + \text{NO}_2 \rightarrow \text{N}_2\text{O}_5) = 3.41 \times 10^{-11} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$
	3.3×10^{-28}	4.3×10^{-17}	Atkinson and Aschmann ¹⁷⁴	296 ± 2	RR relative to $k(\text{propene}) = 9.45 \times 10^{-15} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$
	3.7×10^{-28}	4.7×10^{-17}	Atkinson <i>et al.</i> ¹⁶³	~ 297	RR relative to $k(\text{thiophene}) = 9.93 \times 10^{-14} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, measured over temp. range 272–297 K
	4.2×10^{-28}	5.5×10^{-17}	Atkinson <i>et al.</i> ¹⁶³	~ 297	RR relative to $K_5(\text{NO}_3 + \text{NO}_2 \rightarrow \text{N}_2\text{O}_5) = 1.26 \times 10^{-27} \text{ e}^{11275/T} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, measured over temp. range 272–297 K
1M-Nap	3.6×10^{-28}	4.6×10^{-17}	Atkinson ²¹⁰	298	Recommended value
	8.4×10^{-28}	1.1×10^{-16}	Atkinson and Aschmann ¹⁷³	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.6 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$
	7.0×10^{-28}	9.0×10^{-17}	Atkinson and Aschmann ¹⁷⁴	296 ± 2	RR relative to $k(\text{trans-2-butene}) = 3.89 \times 10^{-13} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$
	7.7×10^{-28}	9.9×10^{-17}	Atkinson ²¹⁰	298	Recommended value
	7.2×10^{-28}	9.2×10^{-17}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
2M-Nap	1.1×10^{-27}	1.4×10^{-16}	Atkinson and Aschmann ¹⁷³	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.6 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$
	1.1×10^{-27}	1.4×10^{-16}	Atkinson and Aschmann ¹⁷⁴	296 ± 2	RR relative to $k(\text{propene}) = 9.45 \times 10^{-15} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$
	1.1×10^{-27}	1.4×10^{-16}	Atkinson ²¹⁰	298	Recommended value
	1.0×10^{-27}	1.3×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1E-Nap	9.8×10^{-28}	1.3×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
2E-Nap	8.0×10^{-28}	1.0×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1,2DM-Nap	6.4×10^{-27}	8.3×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(2,7\text{-DMN}) = 21 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1,3DM-Nap	2.1×10^{-27}	2.7×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1,4DM-Nap	1.3×10^{-27}	1.7×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1,5DM-Nap	1.4×10^{-27}	1.8×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1,6DM-Nap	1.7×10^{-27}	2.1×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1,7DM-Nap	1.4×10^{-27}	1.7×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
1,8DM-Nap	2.1×10^{-26}	2.7×10^{-15}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(2,7\text{-DMN}) = 21 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
2,3DM-Nap	1.5×10^{-28}	1.9×10^{-17}	Atkinson and Aschmann ¹⁷³	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.6 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$
	1.6×10^{-27}	2.1×10^{-16}	Atkinson and Aschmann ¹⁷⁴	296 ± 2	RR relative to $k(\text{propene}) = 9.45 \times 10^{-15} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$
	1.6×10^{-27}	2.0×10^{-16}	Atkinson ²¹⁰	298	Recommended value
	1.5×10^{-27}	2.0×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
2,6DM-Nap	2.1×10^{-27}	2.7×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$, derived from the same work
2,7DM-Nap	2.1×10^{-27}	2.7×10^{-16}	Phouongphouang and Arey ²¹⁴	298 ± 2	RR relative to $k(\text{naphthalene}) = 3.65 \times 10^{-28} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$

Table 5 (continued)

	k_{NO_3} (cm^{-3} molecules $^{-1}$ s $^{-1}$) ($\times [\text{NO}_2]$)	k_{NO_3} (s $^{-1}$) [NO_2] = 6.91×10^{11} molecules cm^{-3} ^a	Ref.	T (K)	Notes
Ace		4.6×10^{-13} ^b	Atkinson and Aschmann ¹⁷⁴	296 $\pm$ 2	RR relative to $k(\text{trans-2-butene}) = 3.89 \times 10^{-13}$ cm^{-3} molecules $^{-1}$ s $^{-1}$
	1.7×10^{-27}	2.1×10^{-16}	Atkinson and Aschmann ¹⁷⁴	296 $\pm$ 2	RR relative to $k(\text{trans-2-butene}) = 3.89 \times 10^{-13}$ cm^{-3} molecules $^{-1}$ s $^{-1}$
Acy		5.5×10^{-12} ^b	Atkinson and Aschmann ¹⁷⁴	296 $\pm$ 2	RR relative to $k(\text{trans-2-butene}) = 3.89 \times 10^{-13}$ cm^{-3} molecules $^{-1}$ s $^{-1}$
Flu		3.5×10^{-14} ^c	Kwok <i>et al.</i> ¹⁷⁷	297 $\pm$ 2	RR relative to $k(1\text{-butene}) = 1.19 \times 10^{-14}$ cm^{-3} molecules $^{-1}$ s $^{-1}$
Phe		1.2×10^{-13} ^d	Kwok <i>et al.</i> ¹⁷⁶	296 $\pm$ 2	RR relative to $k(1\text{-butene}) = 1.19 \times 10^{-14}$ cm^{-3} molecules $^{-1}$ s $^{-1}$
Flt	5.1×10^{-28}	6.6×10^{-17}	Atkinson <i>et al.</i> ¹⁶³	296 $\pm$ 2	RR relative to $k(\text{naphthalene}) = 3.6 \times 10^{-28}$ cm^{-3} molecules $^{-1}$ s $^{-1}$
Pyr	1.6×10^{-27}	2.1×10^{-16}	Atkinson <i>et al.</i> ¹⁶³	296 $\pm$ 2	RR relative to $k(\text{naphthalene}) = 3.6 \times 10^{-28}$ cm^{-3} molecules $^{-1}$ s $^{-1}$
1N-Nap	3.0×10^{-28}	3.9×10^{-17}	Atkinson ²¹⁰	298	Recommended value
2N-Nap	2.7×10^{-28}	3.5×10^{-17}	Atkinson ²¹⁰	298	Recommended value

^a [NO_2] = 6.91×10^{11} molecules cm^{-3} ; annual average, Harwell, UK (2011). ^b [NO_2] = $<1.2 \times 10^{15}$ molecules cm^{-3} . ^c [NO_2] = (7.2–24) $\times 10^{13}$ molecules cm^{-3} . ^d [NO_2] = (4.8–24) $\times 10^{13}$ molecules cm^{-3} .

Atkinson²¹⁰ derived a $k^{(2)}$ value for the reaction of naphthalene with NO_3 based on a unit-weighted least squares analysis of three $k^{(2)}$ values from previous studies.^{163,173,174} It was also shown that the $k^{(2)}$ value for the NO_3 reaction with naphthalene is lower than that of naphthalene- d_8 ($k_{\text{D}}/k_{\text{H}} \sim 1.2$),²⁰³ indicating that there is very little deuterium isotope effect, suggesting that the H-atom abstraction mechanism is of minor importance.

Atkinson²¹⁰ recommended $k^{(2)}$ values for 1M-Nap, 2M-Nap and 2,3DM-Nap using unit-weighted averages of the previous values provided by Atkinson and Aschmann.^{173,174} $k^{(2)}$ values measured by Phouongphouang and Arey²¹⁴ for 1M-Nap, 2M-Nap, and 2,3DM-Naps are in very good agreement with values recommended by Atkinson.²¹⁰ It is indicated that the $k^{(2)}$ for the reaction of naphthalene and alkyl-naphthalenes with NO_3 radicals is linearly dependent on the NO_2 mixing ratio over the range ~ 0.4 –50 ppmV.^{174,210,213} This is therefore consistent with the initial formation of the NO_3 -alkylnaphthalene adduct.

It is indicated that, as for OH radical reactions, alkyl substitution enhances PAH reactivity towards NO_3 following the general order of reactivity DM-Naps > E-Naps > M-Naps > Nap.²¹⁴ The 'enhancement factor' reported for alkylnaphthalene compounds towards NO_3 generally parallels those for the reactions of OH, with relative enhancement ratios (NO_3/OH) of between 1.1 and 2.1 for most compounds. The exceptions are 1,2DM-Nap and 1,8DM-Nap, where the relative enhancement is 7.0 and 22 respectively. While potential explanations, including the steric influence of the methyl group positioning and the possible decrease of ring aromaticity, have been suggested, the precise reason for this enhanced reactivity is not currently known.

Atkinson²¹⁰ noted that the rate coefficient for Ace is similar to that of 2,3DM-Nap, suggesting that the cyclopenta-fused ring in Ace acts like two methyl substituent groups. As mentioned above, Ace is expected to react *via* H-atom abstraction from the $-\text{CH}_2-$ group(s) of the cyclopenta-fused ring in addition to the NO_3 addition mechanism.

To our knowledge, the only $k^{(2)}$ value derived for the gas phase reaction of NO_3 with fluorene was that reported by Kwok *et al.*¹⁷⁷ In this study, no evidence of an NO_2 -dependent $k^{(2)}$ was observed, suggesting that the NO_3 -addition mechanism does not dominate.²¹⁰ Indeed, it was suggested that the reaction proceeds *via* H-atom abstraction from the $-\text{CH}_2-$ group(s). $k^{(2)}$ values for the gas phase reactions of Flt and Pyr with NO_3 radicals have only been studied in one investigation.¹⁶³ For both compounds it was suggested that the reaction proceeds *via* initial addition of NO_3 to the aromatic ring, as detailed above.

NO_3 reactivity with Phe was studied by Kwok *et al.*²¹⁵ A linear relationship between $k^{(2)}$ and NO_2 concentration indicates that the reaction proceeds by the initial addition of NO_3 to the aromatic ring, forming the Phe- NO_3 adduct. Kinetic data²¹⁵ also suggested that an additional mechanism occurs, independent of NO_2 concentration. This may involve NO_3 addition to the C_9 - C_{10} bond²¹⁵ or unimolecular decomposition of the NO_3 -phenanthrene adduct.²¹⁶

The $k^{(2)}$ values in Table 5 indicate that the gas phase reactions of LMW PAHs with NO_3 radicals proceed significantly slower (1–3 orders of magnitude) than the corresponding reactions with OH radicals. This would suggest that OH radical reactions represent a much more important loss process for PAHs than reactions with NO_3 . The relative scarcity of rate coefficient data makes observations regarding the differences in NO_3 reactivity between individual PAHs more difficult. As with the OH radical reaction, it appears that Acy reacts faster than Ace and Nap, which can be attributed to the additional H-atom abstraction from the $>\text{C}=\text{C}<$ bond of the cyclopenta-fused ring. To our knowledge, no rate coefficient data are available for the reaction of NO_3 with Ant so no comparison can be made between this and the other 3 ring compounds.

3.2.1.2.4. Reaction products. Reaction of PAHs through the more dominant NO_3 addition mechanism followed by preferential

addition of NO₂ is expected to yield nitro-PAH isomers, while reactions involving NO₃ interaction with substituent groups through H-atom abstraction, for example Ace and Acy, and Fl_n, are not. As with the reactions involving OH radicals, the NO₃ radical product studies have focused primarily on the reactions of naphthalene with a smaller number of studies investigating other 3–4 ring compounds.

As with OH reactions, product studies show that the NO₃ initiated reaction of Nap forms the ring-retaining products naphthols, nitronaphthalenes and naphthoquinones. 1- and 2-nitronaphthalenes have been identified as reaction products in yields of 15–24% and 7–11% respectively.^{163,173,182,212} The reaction mechanism is consistent with the expected NO₃ addition at the C₁ or C₂ of Nap to form the Nap–OH adduct, followed by addition of NO₂ (Scheme 8c). The formation yield ratio for 1N-Nap/2N-Nap observed in these studies is in good agreement, and is shown to be ~2.2.

Other lower yield products identified¹⁸² include 1,4-NQu (1.9%), 1-hydroxy-2-nitronaphthalene (1.5%) and 2-hydroxy-1-nitronaphthalene (1.3%). The observed products are broadly consistent with the reaction mechanism detailed in Scheme 8. The formation of 1,4-NQu could be explained by pathways involving either (i) reaction of the Nap–NO₃ adduct with O₂/HO₂ (Scheme 8e) or (ii) reaction of naphthols. It should be noted that the epoxynaphthalene and nitric ether species predicted by Qu *et al.*²¹⁶ have not been identified in product studies.

Atkinson *et al.*²¹³ identified naphthols and nitronaphthols. It was shown that the ratio of nitronaphthols/nitronaphthalenes was a factor of 1.6 higher when there was no initial NO₂ addition compared to a reaction where NO₂ was initially present. This could indicate that the unimolecular decomposition of the PAH–NO₃ adduct (e) could be competitive with the assumed dominant reaction with NO₂.

Wang *et al.*²¹⁷ found that (with the exception of 1,4DM-Nap) the yields of EN-Naps and DM-Naps over the range of NO₂ mixing ratios of 0.2–2 ppmV are not linearly dependent upon NO₂ concentration and that the EN-Nap/DMN-Nap profile did not change significantly with varying NO₂ concentration over this range. It was shown that, as the NO₂ concentration increased by a factor of 8, the EN-Nap and DMN-Nap formation yield only increased by a factor of 3. This suggests that the OH-addition mechanism may not dominate for these compounds.

However, Zielinska *et al.*²¹⁸ reported that nitro-PAH products were formed from the gas phase reaction of 1M-Nap and 2M-Nap with NO₃ in yields of ~30%, approximately two orders of magnitude higher than the corresponding reaction initiated by OH. The MNN isomer distribution from the reaction of 1M-Nap was shown to be 1M3N-Nap > 1M5N-Nap > 1M4N-Nap > 1M8N-Nap ~ 1M6N-Nap > 1M7N-Nap > 1M2N-Nap. For the reaction with 2M-Nap the isomer distribution was 2M4N-Nap > 2M1N-Nap ~ 2M5N-Nap > 2M8N-Nap ~ 2M3N-Nap ~ 2M7N-Nap ~ 2M6N-Nap.

Wang *et al.*²¹⁹ found that the isomers with the highest mass formation rates are 2,7DM4NN, 1,2DM4NN, 2,6DM4N-Nap, 2,6DM1N-Nap, 1,6DM8N-Nap, 2,7DM1N-Nap, 1,6DM5N-Nap,

2E4N-Nap. The formation of these products was attributed to the addition of NO₃ to an alpha-carbon next to the –CH₃ group (in the case of the formation of 2,7DM4N-Nap, 2,6DM4N-Nap, 1,6DM8N-Nap, 2E4N-Nap), to an alpha-carbon on the methyl-substituted ring (in the case of 2,7DM1N-Nap, 2,6DM1N-Nap and 1,6DM5N-Nap) or ipso addition at C₁ (in the case of 1,2DM4N-Nap), followed by the addition of NO₂ at the *para*-position and removal of HNO₃. This proposed mechanism is therefore consistent with the NO₃ addition to the aromatic ring. It was also indicated that these reactions also yield ring-opened aromatic carbonyl (MW 170) products and alkynaphthoquinones.²¹⁷

For the gas phase reaction of Acy with NO₃, Atkinson and Aschmann¹⁷⁴ observed negligible levels of N-Acy isomers with the major products observed being oxygenated. Therefore a reaction pathway, possibly involving NO₃ addition to the C=C bond of the cyclopenta-fused ring, is proposed.^{173,174} Arey *et al.*¹⁸⁰ observed the main reaction product of the reaction of Ace and NO₃ to be 4-NAce with a yield of ~40% with smaller amounts of the 3- and 5-isomers (~2%), in agreement with the NO₃ addition mechanism of to the aromatic ring and subsequent reaction with NO₂.

The gas phase reaction of Fl_n with NO₃ was shown to yield 9-fluorenone,¹⁷⁷ which is consistent with the observation that this reaction proceeds *via* initial H-atom abstraction from a –CH₂– group of Fl_n, rather than through initial NO₃ addition to the aromatic ring.

For Ant, Phe and Pyr, nitro-PAH products are noted in low formation yields (<2%).^{163,174,180} The formation of 2-nitroFlt from the reaction of Flt with N₂O₅ in CCl₄ and in the gas phase has been observed with a relatively high yield (24%).^{163,220} The observation of these nitro-PAH products is in agreement with the expected NO₃ addition mechanism with further reaction with NO₂ (Scheme 8c).

Wang *et al.*²²¹ observed the presence of 9-Flr, dibenzopyranone, 9,10-PQu and 2,2'-diformylbiphenyl from the gas phase reaction of Phe with NO₃. The reaction was suggested to proceed analogously to that of the OH-initiated reaction, with NO₃ addition at the C₉–C₁₀-position expected. The calculated formation yield of 9,10-PQu was ~33%. Combining the measured formation yield with the room temperature rate coefficient,¹⁷⁶ the authors calculated a PQu formation rate of 800 pg m^{–3} h^{–1}, suggesting that nighttime NO₃-initiated reaction can contribute significantly to the atmospheric levels of PQu.²²¹

The above studies indicate that for a number of PAHs (*e.g.* naphthalene, alkynaphthalenes, acenaphthene, fluoranthene) the yield of nitro-PAH products from the NO₃-initiated reaction is notably higher than that of the corresponding reaction initiated by OH radicals, suggesting the Nap–NO₃ adduct reacts more preferentially with NO₂ than the Nap–OH adduct. This would suggest that, while NO₃ reactions may represent a less important atmospheric loss process for gas phase PAHs, these reactions are a more important source of nitro-PAH products. This indicates therefore that while during daylight hours the formation of PAH-derivatives will result from OH-initiated reactions, nighttime formation of these compounds from NO₃-initiated reactions will also be significant.

3.2.1.3. *Gas phase reactions of PAHs with O₃*. Ozone (O₃) forms part of a complex reaction cycle with volatile organic compounds, for example alkenes, alkanes, aromatic hydrocarbons, aldehydes and ketones, and other atmospheric gases such as O₂ and NO_x, which will lead to the formation of radical species, which influence the composition and oxidising potential of the atmosphere.^{148,222,223} In addition, direct reaction with ozone can be an important atmospheric loss process for volatile organic chemicals.^{148,223}

The available $k^{(2)}$ values for the reactions of ozone with gas phase PAHs are shown in Table 6. Experimental studies have reported no loss, or very small loss of PAH during exposure studies with relatively high ($\sim 4 \times 10^{13}$ molecules cm⁻³) O₃ concentrations.^{144,168,173,177} These studies therefore report estimated upper limits for $k^{(2)}$ values considerably lower (5–8 orders of magnitude) than the corresponding reactions with OH or NO₃.

Kwok *et al.*¹⁷⁶ did observe the reaction of phenanthrene with $(8.7\text{--}15.6) \times 10^{13}$ molecules cm⁻³ of O₃ and derived a $k^{(2)}$ value for this reaction, which was shown to be of the same order of magnitude as the upper limits calculated for other PAHs in previous studies. These observations therefore indicate that reactions of gas phase PAHs with O₃ will be negligible under atmospheric conditions and for most PAHs the reaction with ozone is too slow to be a significant degradation process or route to the formation of derivative products.

The relatively slow reaction of PAHs indicates that O₃ does not react to any observable extent with the aromatic ring or with saturated side-chain groups. However, it should be noted that the reaction of acenaphthylene with O₃ proceeds significantly (~ 3 orders of magnitude) faster than other PAHs. This is attributed to ozone reacting with the >C=C< bond of the unsaturated cyclopenta-fused ring.^{155,174}

Reisen and Arey¹⁵⁵ conducted a product study for the reaction of acenaphthylene with ozone. API-MS analysis identified a single secondary ozonide product. GC-MS analysis identified six other products: 1-naphthaldehyde, oxaacenaphthylene-2-one, 2-hydroxy-1-naphthaldehyde, 1,8-naphthalic anhydride (NtA) and two dialdehyde isomers (MW 184). These products were attributed to breakdown products of the secondary ozonide. A more comprehensive and quantitative product study is required to fully understand the reaction mechanism of this process.

Wang *et al.*²²¹ observed the presence of 9-fluorenone (9-Fln), dibenzopyranone, 9,10-phenanthrenequinone (PQu) and 2,2'-diformylbiphenyl from the gas phase reaction of phenanthrene with O₃. The authors suggested that the mechanism would involve addition of O at the C₉–C₁₀ bond and rearrangement to a secondary ozonide. The calculated formation yield of 9,10-PQu was $\sim 2\%$. Combining the measured formation yield with the room temperature rate coefficient,¹⁷⁶ the authors calculated a PQu formation rate of $0.2 \text{ pg m}^{-3} \text{ h}^{-1}$, suggesting that reaction with O₃ is not a significant source of PQu in the atmosphere, compared with OH and NO₃ radical induced reactions.

3.2.1.4. Summary and concluding remarks

3.2.1.4.1. *Key conclusions.* • The main sink for gas phase PAHs appears to be reaction with the OH radical, with rate coefficients for these reactions shown to be considerably (up to 5 orders of magnitude) higher than the corresponding reactions with NO₃ for most 3–4 ring PAHs.

• For most PAHs, the dominant gas phase reaction mechanism involves addition of the OH or NO₃ radical to the aromatic ring to form an energy-rich adduct. Further reaction of this adduct with NO₂ will yield nitro-PAH products while adduct reaction with O₂ will form a peroxy-radical intermediate which can further react or decompose *via* a number of different pathways to yield ring-retaining products such as quinones, or a wide range of ring-opened products.

• Product studies suggest, for a number of PAHs, that the reaction is expected to proceed, at least in part due to the H-atom abstraction mechanism from C–H groups from the aromatic ring (fluorene), substituted alkyl groups (methyl- and ethyl-naphthalenes) or radical interaction with the >C=C< bond of the cyclopenta-fused ring (acenaphthylene).

• While NO₃ reactions appear to be less significant than OH reactions as a PAH degradation process, considerably higher nitro-PAH yields suggest that night-time reactions of PAHs with NO₃ may be a significant contributor of these compounds in the atmosphere, in addition to day-time OH reactions.

• Reactions of PAHs with O₃ are considered to be of negligible importance in the atmosphere due to a slow reaction with the aromatic ring. The reaction of acenaphthylene is shown to

Table 6 Rate coefficients reported for gas-phase reactions of PAH with ozone

	k_{O_3} (cm ⁻³ molecules ⁻¹ s ⁻¹)	Ref.	T (K)	Notes
Nap	$< 2.0 \times 10^{-19}$	Atkinson <i>et al.</i> ¹⁶⁸	294 ± 1	Upper limit
	$< 3.0 \times 10^{-19}$	Atkinson and Aschmann ¹⁴⁴	295 ± 1	Upper limit
1M-Nap	$< 1.3 \times 10^{-19}$	Atkinson and Aschmann ¹⁷³	298 ± 2	Upper limit
2M-Nap	$< 3.0 \times 10^{-19}$	Atkinson and Aschmann ¹⁴⁴	295 ± 1	Upper limit
	$< 4.0 \times 10^{-19}$	Atkinson and Aschmann ¹⁷³	295 ± 2	Upper limit
Ace	$< 5.0 \times 10^{-19}$	Atkinson and Aschmann ¹⁷⁴	296 ± 2	Upper limit
Acy	5.5×10^{-16}	Atkinson and Aschmann ¹⁷⁴	296 ± 2	
	1.6×10^{-16}	Reisen and Arey ¹⁵⁵	296 ± 2	RR, relative to $k(2\text{-methyl-2-butadiene}) = 3.96 \times 10^{-16} \text{ cm}^{-3} \text{ molecules}^{-1} \text{ s}^{-1}$
2,3DM-Nap	$< 4.0 \times 10^{-19}$	Atkinson and Aschmann ¹⁴⁴	295 ± 1	Upper limit
	$< 2.0 \times 10^{-19}$	Kwok <i>et al.</i> ¹⁷⁷	297 ± 2	Upper limit
Phe	4.0×10^{-19}	Kwok <i>et al.</i> ¹⁷⁶	296 ± 2	
1N-Nap	$< 6.0 \times 10^{-19}$	Atkinson ²⁰⁷	298 ± 2	Upper limit
2N-Nap	$< 6.0 \times 10^{-19}$	Atkinson ²⁰⁷	298 ± 2	Upper limit

be faster due to interaction of ozone with the >C=C< bond of the cyclopenta-fused ring.

- While reasonable agreement is shown between rate coefficients for the reactions of OH, NO₃ and O₃ derived in different studies (absolute, relative, theoretical) for individual PAHs, discrepancies can often be explained by differences in reaction conditions and quantification methods used.

- There is shown to be considerable variation in reactivity with OH between individual PAH molecules. Reactivity is dependent upon the precise reaction mechanism and can be influenced by steric hindrance of reactive sites and the number and type of substituted groups. Alkyl groups are shown to have an 'activating effect' compared to the reactivity of 2- and 3-ring PAHs.

- The mechanisms of gas phase PAH reactions are still not fully understood. For most PAHs, the majority of reaction products remain unidentified and the precise reaction routes of these processes are still largely unclear. The applicability of observations from experimental studies in the ambient atmosphere is also uncertain (discussed in Section 5).

3.2.1.4.2. Recommendations for future work. The following are areas we suggest for continued research in the area of gas phase PAH reactivity:

- The rates and products of the reactions of Ant with NO₃ and O₃ and of most (predominantly gaseous) 4-ring PAHs with OH, NO₃ and O₃ should be studied.
- A more comprehensive elucidation of the reaction products of the reactions of many PAHs with OH and NO₃ in order to gain a clearer understanding of the reaction mechanisms.
- Verifications of reaction rates based on atmospheric observations.

3.2.2. Heterogeneous reactions of PAHs

3.2.2.1. Overview. Volatile and semi-volatile organic compounds associated with particulate matter can be influenced by heterogeneous photochemical reactions in the atmosphere.^{224,225} Semi-volatile PAHs are shown to have a high affinity for soot and particulate matter.^{96,104} Since ambient air sampling studies have revealed that PAHs and their derivative compounds are found to be associated with atmospheric particles^{137,226–229} and given the wide range of particulate matter (soot, sands, mineral dusts, metal oxides, diesel exhausts, pollens) present in the atmosphere, it is important to understand the extent and nature of heterogeneous photochemical reactions that can lead to removal of PAHs from the atmosphere and/or their conversion to derivative compounds.

In the troposphere, direct photolysis of PAHs associated with solid surfaces is a potentially important process.^{1,230,231} The chemistry of direct photolytic PAH decomposition has been discussed elsewhere for PAHs associated with organic aerosols,^{232–238} inorganic substrates^{239–248} and the air–ice interface.^{249,250} Therefore, photolysis of surface-bound PAH will not be covered in this review. Here, we will discuss the present understanding of particle-associated PAHs with OH radicals, NO₃ radicals, NO₂ and O₃.

Heterogeneous PAH reactivity towards atmospheric oxidants has previously been reviewed by Van Cauwenberghe and Van Vaeck²³⁰

and Van Cauwenberghe,²⁵¹ who provided an overview and extensive literature review of atmospheric reactions of PAH on various surfaces including primary combustion particles and ambient particulate matter with O₃ and NO₂. Finlayson-Pitts and Pitts¹ also provide an in depth discussion of PAH reactions with O₃ and NO₂ on particle surfaces. Here we provide an updated overview of heterogeneous PAH reactivity, incorporating more recent kinetic studies.

3.2.2.2. Heterogeneous PAH reactions with OH radicals

3.2.2.2.1. Overview. Heterogeneous reactions of OH radicals with PAHs adsorbed on different solid surfaces have been reported in a number of experimental studies.^{252–257} $k^{(2)}$ values derived for OH reactions with PAHs associated with different substrate types are presented in Table 7.

Bertram *et al.*²⁵³ studied the heterogeneous loss rate of OH radicals on various organic surfaces, coated on Pyrex tubes, using a flow-tube reactor coupled to a chemical ionization mass spectrometer. The average reaction probabilities ($\gamma_{\text{OH,ave}}$) for these organic surfaces were generally large, (>0.1) for relatively low OH concentrations (1×10^8 molecules cm⁻³) and short exposure time (<5 minutes). For example for soot and solid pyrene surfaces, the $\gamma_{\text{OH,ave}}$ values were 0.88 and 0.32, respectively. It was noted that the $\gamma_{\text{OH,ave}}$ value was lower for an alumina surface than for organic surfaces. These results indicate that uptake of OH radicals to particle surfaces may be rapid and thus have the potential to contribute significantly to oxidative reactions of PAHs on organic surfaces in the atmosphere.

Estève *et al.*²⁵⁴ used a discharge flow technique to study the heterogeneous reaction of OH with Phe adsorbed on a Pyrex surface, in the presence of NO₂ and in the absence of light. They investigated the reaction kinetics by measuring the total loss of OH, followed by laser-induced fluorescence (LIF). The efficiency of the heterogeneous reaction and the change in the surface properties with exposure time were demonstrated. In addition, oxidation products of the heterogeneous reaction of OH with Phe in the presence of NO₂, identified by GC-MS, were noted as mono-hydroxylated isomers of phenanthrene.

In their subsequent investigations, Estève *et al.*^{255,256} studied the heterogeneous reactions of both OH radicals and NO₂ with PAHs associated with carbonaceous (graphite) particles and diesel particulate exhaust. This was achieved using an adapted fast flow reactor and by monitoring the decrease in particle-bound PAH concentration by GC-MS. This therefore allowed the calculation of pseudo-first order rate coefficients of these reactions from the experimental decays of PAH concentrations in relation to OH exposure time.

Rate coefficients derived for PAH reactions with OH on diesel exhaust particles were approximately an order of magnitude lower for most compounds than for reactions of PAHs adsorbed on graphite.²⁵⁶ The matrix effects are detailed below (Section 3.2.2.6). These $k^{(2)}$ values fall approximately within the same order of magnitude for all PAHs investigated in the two studies.

Bedjanian *et al.*²⁵² investigated the reactions of 15 PAHs adsorbed on laboratory generated kerosene soot surfaces using

Table 7 Second-order rate coefficients $k^{(2)}$ for heterogeneous reactions of PAH with OH, NO₂, O₃ and NO₃

PAH	OH reactions			NO ₂ reactions			O ₃ reactions			NO ₃ reactions		
	k_{OH} (cm ³ molec ⁻¹ s ⁻¹)	Ref.	Notes	k_{NO_2} (cm ³ molec ⁻¹ s ⁻¹)	Ref.	Notes	k_{O_3} (cm ³ molec ⁻¹ s ⁻¹)	Ref.	Notes	k_{NO_3} (cm ³ molec ⁻¹ s ⁻¹)	Ref.	Notes
Nap												
Phe	5.0 × 10 ⁻¹²	Estève <i>et al.</i> ²⁵⁵	Graphite particles	2.8 × 10 ⁻¹⁹	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	0.9 × 10 ⁻¹⁸ (1/ <i>C</i> _{O₃} + 10 ⁻¹⁵)	Kahan <i>et al.</i> ²⁹⁶	Octanol			
	3.2 × 10 ⁻¹³	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a	3.4 × 10 ⁻¹⁸	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	2.4 × 10 ⁻¹⁷	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
				3.5 × 10 ⁻¹⁷	Estève <i>et al.</i> ²⁵⁵	Graphite particles	2.3 × 10 ⁻¹⁷	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
				1.1 × 10 ⁻¹⁷	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a	2.3 × 10 ⁻¹⁹ (1/ <i>C</i> _{O₃} + 4.6 × 10 ⁻¹⁶)	Kahan <i>et al.</i> ²⁹⁶	Octanol			
				2.3 × 10 ⁻²¹	Butler and Crossley ²⁷⁰	Ethylene flame soot	1.0 × 10 ⁻¹⁸ (1/ <i>C</i> _{O₃} + 1.4 × 10 ⁻¹⁶)	Kahan <i>et al.</i> ²⁹⁶	Ice			
Ant	4.4 × 10 ⁻¹²	Estève <i>et al.</i> ²⁵⁶	Graphite particles	1.0 × 10 ⁻¹⁶	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	9.8 × 10 ⁻¹⁷	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
				3.4 × 10 ⁻¹⁸	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	1.4 × 10 ⁻¹⁶	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
				6.9 × 10 ⁻¹⁷	Estève <i>et al.</i> ²⁵⁵	Graphite particles	5.1 × 10 ⁻¹⁸ (1/ <i>C</i> _{O₃} + 1.96 × 10 ⁻¹⁵)	Kwamena <i>et al.</i> ²⁹²	Octanol			
				1.0 × 10 ⁻¹⁶	Ma <i>et al.</i> ²⁷⁸	Silica particles	1.3 × 10 ⁻¹⁶ (1/ <i>C</i> _{O₃} + 2.2 × 10 ⁻¹⁵)	Kwamena <i>et al.</i> ³⁰¹	Azelaic acid (wet)			
				5.3 × 10 ⁻¹⁷	Ma <i>et al.</i> ²⁷⁸	MgO particles	1.0 × 10 ⁻¹⁵ (1/ <i>C</i> _{O₃} + 10 ⁻¹³)	Kwamena <i>et al.</i> ³⁰¹	Phenylsiloxane oil			
Flt	1.4 × 10 ⁻¹⁴	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	3.2 × 10 ⁻²¹	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	1.2 × 10 ⁻¹⁸ (1/ <i>C</i> _{O₃} + 1.2 × 10 ⁻¹⁸)	Mmereki <i>et al.</i> ²⁹⁴	Water			
	3.2 × 10 ⁻¹²	Estève <i>et al.</i> ²⁵⁵	Graphite particles	1.0 × 10 ⁻¹⁹	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	1.9 × 10 ⁻¹⁷	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
	3.8 × 10 ⁻¹³	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a	2.9 × 10 ⁻¹⁷	Estève <i>et al.</i> ²⁵⁵	Graphite particles	1.5 × 10 ⁻¹⁷	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
				1.0 × 10 ⁻¹⁷	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a						
				2.5 × 10 ⁻²¹	Butler and Crossley ²⁷⁰	Ethylene flame soot						
Pyr	1.6 × 10 ⁻¹⁴	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	2.0 × 10 ⁻¹⁷	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	2.5 × 10 ⁻¹⁷	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
	3.2 × 10 ⁻¹²	Estève <i>et al.</i> ²⁵⁵	Graphite particles	1.0 × 10 ⁻¹⁹	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	5.9 × 10 ⁻¹⁷	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
	4.1 × 10 ⁻¹³	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a	5.1 × 10 ⁻¹⁷	Estève <i>et al.</i> ²⁵⁵	Graphite particles	9.3 × 10 ⁻¹⁷	Miet <i>et al.</i> ²⁹⁷	Silica particles			
	2.4 × 10 ⁻¹³	Miet <i>et al.</i> ²⁵⁷	Silica particles	1.5 × 10 ⁻¹⁷	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a	2.2 × 10 ⁻¹⁹ (1/ <i>C</i> _{O₃} + 3.1 × 10 ⁻¹⁶)	Kahan <i>et al.</i> ²⁹⁶	Octanol			
				4.8 × 10 ⁻²¹	Butler and Crossley ²⁷⁰	Ethylene flame soot						
				3.2 × 10 ⁻¹⁶	Miet <i>et al.</i> ²⁷⁷	Silica particles						
Chr	9.2 × 10 ⁻¹⁵	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	6.0 × 10 ⁻¹⁹	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	1.5 × 10 ⁻¹⁷	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
	5.0 × 10 ⁻¹²	Estève <i>et al.</i> ²⁵⁵	Graphite particles	1.0 × 10 ⁻¹⁹	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	3.1 × 10 ⁻¹⁷	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
										4.0 × 10 ⁻¹²	Liu <i>et al.</i> ²⁶⁰	Azelaic acid particles

Table 7 (continued)

PAH	OH reactions		NO ₂ reactions		O ₃ reactions		NO ₃ reactions	
	k_{OH} (cm ³ molec ⁻¹ s ⁻¹)	Ref.	Notes	k_{NO_2} (cm ³ molec ⁻¹ s ⁻¹)	Ref.	Notes	k_{NO_3} (cm ³ molec ⁻¹ s ⁻¹)	Notes
BaA	4.4×10^{-13}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a	3.9×10^{-17}	Estève <i>et al.</i> ²⁵⁵	Graphite particles		
				1.0×10^{-17}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a		
				2.6×10^{-21}	Butler and Crossley ²⁷⁰	Ethylene flame soot		
	9.2×10^{-15}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	6.7×10^{-18}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	2.8×10^{-17}	Perraudin <i>et al.</i> ²⁹⁸ Graphite particles
	5.6×10^{-13}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	8.7×10^{-17}	Perraudin <i>et al.</i> ²⁹⁸ Silica particles
BkF	3.2×10^{-13}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a	3.3×10^{-17}	Estève <i>et al.</i> ²⁵⁵	Graphite particles		
				1.3×10^{-17}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a		
				6.2×10^{-21}	Butler and Crossley ²⁷⁰	Ethylene flame soot		
	1.0×10^{-14}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	2.2×10^{-18}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	1.9×10^{-17}	Perraudin <i>et al.</i> ²⁹⁸ Graphite particles
	3.5×10^{-12}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	3.6×10^{-17}	Perraudin <i>et al.</i> ²⁹⁸ Silica particles
BaP				2.5×10^{-17}	Estève <i>et al.</i> ²⁵⁵	Graphite particles		
	1.1×10^{-14}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	9.3×10^{-16}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	5.3×10^{-17}	Perraudin <i>et al.</i> ²⁹⁸ Graphite particles
	4.1×10^{-12}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	1.4×10^{-16}	Perraudin <i>et al.</i> ²⁹⁸ Silica particles
	2.9×10^{-13}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a	7.8×10^{-17}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	2.0×10^{-18} (1/ c_{O_3})	Kahan <i>et al.</i> ²⁹⁶ Octanol
				1.5×10^{-17}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a	$< 3.8 \times 10^{-18}$ (1/ c_{O_3})	Kwamena <i>et al.</i> ^{291,292} NaCl
BeP				1.0×10^{-20}	Butler and Crossley ²⁷⁰	Ethylene flame soot	4.2×10^{-15} (1/ c_{O_3})	Kwamena <i>et al.</i> ²⁹² Soot
							$+ 2.8 \times 10^{-13}$	Pöschl <i>et al.</i> ²⁹⁰
							$+ 1.7 \times 10^{-16}$ (1/ c_{O_3})	Kwamena <i>et al.</i> ²⁹¹ Azelaic acid (wet)
							$+ 1.2 \times 10^{-15}$	Kwamena <i>et al.</i> ²⁹² Silica particles
							$+ 3.0 \times 10^{-16}$ (1/ c_{O_3})	Kwamena <i>et al.</i> ²⁹² Silica particles
Per	1.1×10^{-14}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	2.9×10^{-18}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	1.6×10^{-17}	Perraudin <i>et al.</i> ²⁹⁸ Graphite particles
	4.7×10^{-12}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	2.9×10^{-17}	Perraudin <i>et al.</i> ²⁹⁸ Silica particles
	4.7×10^{-13}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a	3.5×10^{-17}	Estève <i>et al.</i> ²⁵⁵	Graphite particles		
				7.5×10^{-18}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a		
				2.8×10^{-21}	Butler and Crossley ²⁷⁰	Ethylene flame soot		
	5.0×10^{-12}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	1.1×10^{-16}	Estève <i>et al.</i> ²⁵⁵	Graphite particles

Table 7 (continued)

PAH	OH reactions			NO ₂ reactions			NO ₂ reactions			O ₃ reactions			NO ₃ reactions		
	k_{OH} (cm ³ molec ⁻¹ s ⁻¹)	Ref.	Notes	k_{NO_2} (cm ³ molec ⁻¹ s ⁻¹)	Ref.	Notes	k_{NO_2} (cm ³ molec ⁻¹ s ⁻¹)	Ref.	Notes	k_{O_3} (cm ³ molec ⁻¹ s ⁻¹)	Ref.	Notes	k_{NO_3} (cm ³ molec ⁻¹ s ⁻¹)	Ref.	Notes
IPy	3.5×10^{-13}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a	6.2×10^{-18}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	1.9×10^{-17}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles	1.9×10^{-17}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
				1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot			Kerosene flame soot	3.8×10^{-17}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
				7.5×10^{-18}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles SRM 1650a			Diesel exhaust particles SRM 1650a						
BgP	5.9×10^{-12}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	4.7×10^{-17}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles			Silica particles						
				1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot			Kerosene flame soot						
				3.8×10^{-17}	Estève <i>et al.</i> ²⁵⁵	Graphite particles			Graphite particles						
				7.9×10^{-21}	Butler and Crossley ²⁷⁰	Ethylene flame soot			Ethylene flame soot						
BgF	8.4×10^{-15}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot			Kerosene flame soot						
AcP	1.0×10^{-14}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot			Kerosene flame soot						
DBaH	1.6×10^{-14}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot			Kerosene flame soot						
DBaP	1.0×10^{-14}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot			Kerosene flame soot						
BbF	1.2×10^{-14}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot			Kerosene flame soot						
Cor	1.1×10^{-14}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	1.3×10^{-18}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles			Silica particles						
				1.0×10^{-19}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot			Kerosene flame soot						
				2.3×10^{-21}	Butler and Crossley ²⁷⁰	Ethylene flame soot			Ethylene flame soot						
DBaP				1.8×10^{-16}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles			Silica particles	1.3×10^{-16}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
1-NP	1.0×10^{-13}	Miet <i>et al.</i> ²⁵⁷	Silica particles	6.2×10^{-18}	Miet <i>et al.</i> ²⁷⁷	Silica particles			Silica particles	1.3×10^{-16}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
										2.2×10^{-17}	Miet <i>et al.</i> ²⁹⁷	Silica particles	1.3×10^{-12}	Liu <i>et al.</i> ²⁶⁰	Azelaic acid particles

a low-pressure discharge-flow reactor combined with an electron-impact mass spectrometer. An uptake coefficient of ≈ 0.2 was reported under the reaction conditions used. Pseudo- $k^{(1)}$ coefficients calculated in this study generally fell between those derived by Estève *et al.*²⁵⁵ and Estève *et al.*,²⁵⁶ suggesting that the substrate type influences PAH reactivity. However, the OH concentration reported in the study by Bedjanian *et al.*²⁵² was 1–2 orders of magnitude higher than that used in the studies by Estève and co workers^{255,256} leading to much lower $k^{(2)}$ values being derived in this study.

In the above three investigations there is relatively little variation in reactivity between each individual PAH associated with the varying types of organic particulate matter. This was also observed in a laboratory investigation of PAH decay profiles caused by heterogeneous reactions with OH radicals occurring on natural atmospheric particles.²⁵⁸ In each of the three studies detailed in Table 6 the $k^{(2)}$ values for all PAHs studied fall approximately within one order of magnitude. The above studies have demonstrated that rapid reaction with OH can potentially be an important potential pathway for degradation of PAHs associated with carbonaceous particles in the atmosphere.

To our knowledge, the only study to investigate the reaction of PAHs on silica particles was carried out by Miet *et al.*²⁵⁷ The $k^{(2)}$ value derived for this reaction is lower than that of the corresponding reaction on diesel particles. More work is clearly needed to assess these reactions for a wider range of compounds on mineral substrates to compare with other surface reactions.

3.2.2.2.2. Reaction mechanism and products. Ringuet *et al.*²⁵⁸ investigated the reaction products formed from PAHs reactions with OH (+NO₂) occurring on natural atmospheric particles. Formation of NPAH derivative compounds was observed. For example, 1N-Pyr was observed, attributed to the heterogeneous reaction of Pyr as reported in NO₃ radical reaction studies.^{259,260} This apparent discrepancy in radical-induced reaction products between the gas and particle phases has also been reported by Miet *et al.*²⁵⁷ for OH reactions on silica. Furthermore, 2N-Pyr and 4N-Pyr were also reported as reaction products in this study, which were not reported elsewhere and had previously been thought to be solely gas phase reaction products.

The precise mechanism of surface-bound PAHs with OH radicals has not been elucidated and the lack of product evaluation makes it difficult to assess the possible reaction chemistry occurring. The non-linear dependence of the reaction rate on OH concentration²⁵² would indicate that this process is not a simple bi-molecular reaction. While the surface reaction of PAH towards OH is shown to be relatively fast, the presence of the 'plateau' in PAH degradation with exposure time would suggest that gas phase transport of OH to the soot surface is not the rate limiting step in these processes. Clearly more work is required to address these issues and investigate the mechanisms of these reactions in more detail.

3.2.2.3. Heterogeneous PAH reactions with NO₃ radicals

3.2.2.3.1. Overview. Heterogeneous reactions of PAHs with NO₃ radicals have received comparatively little attention in the literature compared with the corresponding reaction in the

gas phase, and the reaction of surface-bound PAHs with OH, NO₂ and O₃. Initial investigations focused on the reactivity of pyrene and perylene^{212,261} adsorbed on glass fibre filters towards N₂O, NO₃ radicals, NO₂ and N₂O₅. It was concluded that NO₃ did not react to a detectable extent with these compounds and that N₂O₅ was the main oxidising species. Kamens *et al.*²⁶² reported that the heterogeneous reaction rate coefficient for the reaction of PAHs with N₂O₅ ($\approx 2.5 \times 10^{14}$ molecules cm⁻³) on wood and diesel soot is of the order of 10^{-18} cm⁻³ molecules⁻¹ s⁻¹.

Recently, however, studies have been conducted that contradict these original assumptions and suggest that NO₃ does play an important role in heterogeneous PAH reactivity. Mak *et al.*²⁶³ studied the reaction of NO₃ with methane and hexane soot and solid pyrene surfaces using a flow tube reactor and reported that the uptake of NO₃ on fresh soot and pyrene surfaces was fast (uptake coefficient >0.1). The authors suggested that NO₃ has the potential to rapidly oxidize soot surfaces and their reaction with PAHs potentially provides an important loss process for these compounds under certain atmospheric conditions.

Similarly, Karagulian and Rossi²⁶⁴ investigated the reaction of soot and NO₃ and also reported rapid uptake (uptake coefficient of 5×10^{-2}). Furthermore, Gross and Bertram²⁶⁵ studied the reactive uptake of NO₃ as well as other oxidants (including NO₂ and O₃) with three solid PAH surfaces. Their results suggested that NO₃ has the potential to be a more important sink for PAHs than reactions with NO₂ and O₃. In an investigation by Kwamena and Abbatt,²⁶⁶ exposing anthracene and pyrene to a mixed flow of NO₃, NO₂ and N₂O₅, it was concluded that the nitro-PAH products formed were as the result of a reaction initiated by NO₃ or N₂O₅ and not through direct reaction with NO₂ or HNO₃.

3.2.2.3.2. Reaction mechanism and products. The first reported quantitative heterogeneous reactivity data for PAH reaction with NO₃ were provided by Liu *et al.*²⁶⁰ Rate coefficients were measured for the reactions of BaA, Chr and Pyr, as well as 1-nitropyrene, adsorbed on azelaic acid particles. The $k^{(2)}$ values measured for these reactions are shown in Table 7. These $k^{(2)}$ values for the NO₃ reactions are of the same order of magnitude or 1–3 orders of magnitude higher than for reactions with OH radicals and 5–7 orders of magnitude higher than the heterogeneous reactions with NO₂ and O₃, depending on the particle surface (discussed in the following sections). This suggests that heterogeneous reactions of PAHs with NO₃ may be very significant. However, more research in this area is required to assess the precise rates and mechanisms for a wider range of surface-bound PAHs.

Product studies for the heterogeneous reaction of PAHs with NO₃ radicals have been conducted for a number of PAHs on azelaic acid particles using Vacuum Ultraviolet Photoionization Aerosol Time of Flight Mass Spectrometry (VUV-ATOFMS).^{259,260} Products identified included 9N-Ant, AQu, 1N-Pyr, 2-, 4- and 9N-Phe isomers, BaA-7,12-dione and 7N-BaA for the reactions of Ant, Pyr, Phe and BaA. The reaction of 1N-Pyr was also shown to produce 1,3-, 1,6-, and 1,8-DN-Pyr isomers.²⁶⁰ While some of the products identified, *e.g.* AQu, are thought to be analogous

to those formed from the gas phase reaction, the formation of 1N-Pyr is in contrast to gas phase reactions, in which 2N-Pyr is the major product.

3.2.2.4. Heterogeneous PAH reactions with NO₂

3.2.2.4.1. Overview. The occurrence of heterogeneous reactions of PAHs with NO₂ was initially demonstrated through experimental studies exposing PAHs on various substrates to NO₂ in the presence of nitric acid, at ambient or elevated levels, resulting in the formation of nitro-PAH products. This has been observed on a number of different surfaces including filters,^{267,268} wood smoke particles,²⁶⁹ soot particles,²⁷⁰ propane flame-generated soot,²⁷¹ coal fly ash,²⁷² silica plates,²⁷³ metal oxides TiO₂, ZnO, Al₂O₃, Fe₂O₃, MgO, CaO, SiO₂ and graphite.²⁷⁴ Details of these studies are summarized by Finlayson-Pitts and Pitts.¹

More recently, full kinetic investigations of this process have been carried out with rate coefficient values for the reaction of NO₂ adsorbed on various particle surfaces calculated.^{255,256,275,276} The $k^{(2)}$ values derived from these studies are presented in Table 7. In general, heterogeneous reactions of PAHs with NO₂ on carbonaceous particles are considerably slower than the corresponding reactions with OH radicals and O₃. Indeed, the reactions of PAHs with OH are 1–6 orders of magnitude faster than the reactions with NO₂, depending on the surface type and specific experimental conditions.

Estève and co-workers investigated the heterogeneous reactions of NO₂ with PAHs associated with graphite²⁵⁵ and diesel exhaust particles.²⁵⁶ It was shown that Pyr and BaP are the most reactive towards NO₂ on diesel exhaust and Ant, Per and BaP are the most reactive on graphite. Nguyen *et al.*²⁷⁵ investigated the reactions of 17 PAHs associated with fresh kerosene flame soot over a temperature range (255–330 K) and the decay of PAHs due to the reaction of NO₂ was considered to be negligible, suggesting that PAH degradation by NO₂ alone is of minor importance in the atmosphere.

Perraudin *et al.*,²⁷⁶ Miet *et al.*²⁷⁷ and Ma *et al.*²⁷⁸ studied the reaction of NO₂ with PAHs adsorbed on silica particles. Ma *et al.*²⁷⁸ also investigated the reaction of gas phase NO₂ with anthracene MgO particles. $k^{(2)}$ values for this reaction were much more variable between individual PAH species than was observed for reactions with OH and O₃ or with NO₂ on carbonaceous substrates and are lower for most compounds by 1–5 orders of magnitude. The matrix effects are detailed below (Section 3.2.2.6).

3.2.2.4.2. Reaction mechanism and products. The formation of nitro-PAH isomers from heterogeneous reactions of PAHs with NO₂¹ would indicate a direct molecular reaction between surface PAH and gaseous NO₂. Consistent with this, 1-nitropyrene has been identified as the only the product from the reaction of pyrene with NO₂ on silica,^{277,279,280} alumina²⁷⁹ and soot⁸⁴ particles. This is in contrast to gas-phase reactions with OH/NO₃ radicals followed by NO₂ where 2-nitropyrene is observed as the main reaction product.¹⁸⁸

Miet *et al.*²⁷⁷ suggested the mechanism for the nitration of pyrene due to the heterogeneous reaction with NO₂ on silica particles as being a simple elementary reaction between the two species, indicated by the linear relationship between pyrene

reactivity and NO₂ concentration. The formation of 1-NPyr is attributed to electrophilic addition of NO₂ at C₁.²⁷⁷ However, Inazu *et al.*²⁷⁴ observed that the main products from the reaction of fluoranthene with NO₂ on various particle types were 1-, 2- and 7N-Flt isomers, indicating that an electrophilic nitration (which would primarily yield 3- or 8-isomers) does not occur.

The heterogeneous reaction of PAHs with NO₂ appears to be significantly enhanced by the presence of gaseous nitric acid.^{267,268,272,281–283} For example, it was shown that the 24 h nitration yield for pyrene adsorbed on filters exposed to 10 ppmV NO₂ rose from 0.02% to 2.85% when traces of HNO₃ were present.²⁶⁸ Grosjean *et al.*²⁸¹ studied the transformation of perylene and BaP deposited on glass and Teflon filters loaded with fly ash and diesel exhaust particles, exposed to 100 ppbV of NO₂. They reported no evidence for chemical transformations in these experiments. This could be attributed to the absence of HNO₃.¹

In the study by Ma *et al.*²⁷⁸ both 9N-Ant and 9,10-AQu were observed as the products from the nitration of anthracene with NO₂ on SiO₂ particles, whereas only 9,10-AQu was observed from the equivalent reaction on MgO particles. The difference between these two processes was attributed by the authors to the formation of HNO₃ on SiO₂ that does not occur on MgO particles.

For the reactions of NO₂ with PAHs on silica particles, it is proposed that HNO₃ is formed on the silica surface, which acts as a catalyst for the process.²⁸⁰ Wang and co-workers proposed that the HNO₃ dissociates producing H⁺ leading to the protonation of NO₂ and N₂O₄ to form HNO₂⁺ and HN₂O₄⁺. This initiates an electrophilic nitration process with reaction occurring at the C₁ position (the carbon with highest electron density), producing intermediate species that further react with NO₂ and/or decompose to form 1N-Pyr (Scheme 9).

It was suggested that HNO₂ gas, which is also formed from NO₂ on the silica surface, also acts catalytically in the heterogeneous reaction of pyrene.²⁸⁴ Furthermore, considering the combined catalytic effect of HNO₃ and HNO₂, a different electrophilic nitration mechanism was proposed (Scheme 10b) in which NO⁺ is formed on the silica surface from the dissociation of HNO₂.²⁸⁵ These species then reacts with NO₂ or N₂O₄ to form NONO²⁺ and NON₂O₄⁺, respectively, which react with surface bound Pyr, analogous to the reaction in Scheme 10a.

3.2.2.5. Heterogeneous PAH reactions with O₃

3.2.2.5.1. Overview. The occurrence of heterogeneous ozonolysis reactions of PAHs was first demonstrated in studies exposing PAHs on solid surfaces to O₃ and observing decay in the PAH profile and/or conversion to oxidised derivative compounds.^{83,231,242,267,282,286–288} These studies indicated that under certain atmospheric conditions, PAHs adsorbed onto surfaces may undergo rapid heterogeneous reaction with gas phase O₃. Detailed review of these studies is provided by Finlayson-Pitts and Pitts.¹

More recently, studies have focussed on full kinetic investigations of the ozonation of PAHs associated with different types of surfaces such as soot particles,^{289,290} organic aerosols and sea salts,^{291,292} air–water interface,^{293–295} organic films,²⁹⁶ mineral surfaces^{297,298} and graphite.²⁹⁸

Scheme 9 Suggested mechanisms for the heterogeneous reaction of pyrene with NO_2 .^{280,284}

Perraudin *et al.*²⁹⁸ studied the reaction of 13 PAHs with O_3 adsorbed on graphite and silica particles. Pseudo- $k^{(1)}$ values were calculated and shown to be proportional to gaseous O_3 concentration. Therefore $k^{(2)}$ values were derived considering the concentration of O_3 (Table 7). The reactivity of PAHs followed the order Chr < BeP < Cor, BgP, Phe, Flt, BkF, Ipy < Pyr < BaA < BaP < Ant < DBaP and Flth < Cor < Phe < BeP < Chr < BkF < Ipy < Pyr < BgP < BaA < DBP < Ant, BaP for graphite particles and silica particles, respectively.

Wu *et al.*²⁹⁹ had previously measured $k^{(2)}$ values for the reaction of O_3 with BaP adsorbed on silica particles of $1.78 \times 10^{-17} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$ at room temperature, about 7 times higher than that measured by Perraudin *et al.*²⁹⁸ This may be explained by the fact that the earlier study was not performed in the dark. This is consistent with a previous study by Katz *et al.*²⁸⁶ where the reactivity of PAHs adsorbed on cellulose was shown to be more rapid under light irradiation conditions.

Miet *et al.*²⁹⁷ investigated the reaction of O_3 with pyrene adsorbed on silica particles. The $k^{(2)}$ value derived from this study was approximately 5 times higher than that derived by Perraudin *et al.*²⁹⁸ despite the reaction conditions (temperature, pressure, particulate concentration) being essentially identical in the two studies. Previously, Kamens *et al.*³⁰⁰ investigated the reactions on wood smoke at low PAH loading with 0.52 ppmV O_3 and reported $k^{(2)}$ values ($k, \text{cm}^3 \text{ molec}^{-1} \text{ s}^{-1}$) of 4.1×10^{-18} (BPy), 7.2×10^{-18} (Chr), 1.0×10^{-18} (BaP), 1.3×10^{-18} (BaA), all roughly one order of magnitude lower than reported by Perraudin *et al.*²⁹⁸ These observations clearly highlight the influence of the specific conditions and nature of the PAH and/or particle surface on the reaction kinetics of these processes.

3.2.2.5.2. Reaction mechanism and products. The precise reaction mechanism(s) and influencing factors governing the heterogeneous reaction of surface-bound PAHs with O_3 have

Scheme 10 Suggested mechanism for the heterogeneous reaction of pyrene with gaseous ozone, forming 1-hydroxypyrene (a), phenanthrene-4,5-dicarboxaldehyde and 4-oxapyren-5-one (b).²⁹⁷

been investigated. A direct (Eley–Rideal) mechanism is possible, whereby a gas phase molecule collides directly with an adsorbed PAH molecule and facilitates a reaction. This would lead to a linear dependence between observed first order rate coefficient and the gaseous ozone concentration.

However, experimental studies suggest that the oxidation of surface-bound PAHs by gas phase ozone will follow Langmuir–Hinshelwood type kinetics.^{266,290–294,301} These authors investigated the kinetics and products of these reactions as a function of ozone concentration and relative humidity (Table 8).

The Langmuir–Hinshelwood mechanism involves one species (in this case the PAH molecule) being strongly adsorbed to the particle surface and the gas species (O_3) being in equilibrium between the gas phase and the particle surface.²⁹¹ The reaction therefore proceeds in two steps: (i) adsorption of the gas molecule,

(ii) surface reaction of PAH and the gas molecule. The reaction rate is therefore dependent on the concentration of both PAH and O_3 and will exhibit a non-linear dependence on gas species concentration.

Earlier studies had indicated a linear relationship between ozone concentration and pseudo-first order reaction rate for ozone mixing ratios of 0.5–0.40 ppmV³⁰² and 0–1.5 ppmV²⁹⁹ on silica. It has been suggested that for these studies, the Langmuir–Hinshelwood mechanism may have been reported if the experiments had been conducted over larger ozone concentration ranges.²⁹¹ A linear dependence of pseudo-first order reactivity on ozone concentration over the concentration range 0.4×10^{14} – 3.3×10^{14} cm³ molecules^{−1} was observed by Perraudin *et al.*²⁹⁸ for a range of PAHs, which may also indicate a simple direct mechanism. However, the suitability of a

Table 8 Comparison of kinetics data for the heterogeneous reactions of BaP and anthracene on different aerosol substrates. Adapted from Kwamena *et al.* (ref. 301) and Mmereki *et al.* (ref. 294)

PAH	Langmuir adsorption equilibrium constant, K_{O_3} ($\text{cm}^3 \times 10^{-15}$)	Pseudo-first order rate coefficient, k ($\text{s}^{-1} \times 10^3$)	Surface	Ref.
BaP	280 ± 20	15 ± 1	Soot aerosols	Pöschl <i>et al.</i> ²⁹⁰
	28	32 ^a	Fused silica	Wu <i>et al.</i> ²⁹⁹
	9.5	32 ^a	Non activated silica	Alebic-Juretic <i>et al.</i> ³⁰²
	2.80 ± 1.4	60 ± 18	Azelaic acid aerosol (72% RH)	Kwamena <i>et al.</i> ²⁹¹
	1.20 ± 0.4	48 ± 8	Azelaic acid aerosol (<1% RH)	
	<0.12 ^b	32 ^a	NaCl	
Ant	100 ± 40	10 ± 3	Phenylsiloxane oil aerosols	Kwamena <i>et al.</i> ³⁰¹
	6.4 ± 1.8	2.8 ± 0.9	Pyrex glass	Kwamena <i>et al.</i> ²⁹²
	2.2 ± 0.9	57 ± 9	Azelaic acid aerosols	Kwamena <i>et al.</i> ³⁰¹
	2.143 ± 0.441	2.55 ± 0.17	Water	Mmereki <i>et al.</i> ²⁹⁴
	1.922 ± 0.478	2.26 ± 0.20	Stearic acid film on water	
	0.681 ± 0.291	1.11 ± 0.12	Octanoic acid ^d	
	0.56	2	Octanol thin film	Kahan <i>et al.</i> ²⁹⁶
	0.508 ± 0.088	2.59 ± 0.14	1-Octanol ^c	Mmereki and Donaldson ²⁹³
	0.118 ± 0.036	0.48 ± 0.07	Hexanoic acid ^e	Mmereki <i>et al.</i> ²⁹⁴

^a k value used is an average from the studies of Pöschl *et al.*²⁹⁰ and Kwamena *et al.*²⁹¹ ^b Calculated assuming $k_1 = 3 \times 10^{-3} \text{ s}^{-1}$, which is the typical error of k_1 obtained for azelaic acid aerosol kinetics under dry conditions. This was taken to be the fastest reaction occurring at the highest O_3 concentrations used (31 ppm).²⁹¹ ^c $2.5 \times 10^{-3} \text{ M}$. ^d $3.79 \times 10^{-3} \text{ M}$. ^e $8 \times 10^{-3} \text{ M}$.

Langmuir–Hinshelwood-type mechanism was not tested in this study, so no definitive conclusion can be drawn.

The observation of Langmuir–Hinshelwood kinetics emphasises the potential importance of the degree of ozone partitioning to the particle surface, as this will have a direct influence on the reaction rate. Therefore, in addition to the pseudo-first order rate coefficient (pseudo- $k^{(1)}$), which describes the surface reaction rate, the Langmuir–Hinshelwood mechanism is also characterized by the ozone partitioning coefficient (K_{O_3}), which describes the degree of partitioning of ozone between the gas phase and the surface (Table 8).

There is notably more variation in K_{O_3} values between these different reaction systems than for the pseudo- $k^{(1)}$ values. This may suggest that differences in heterogeneous reactivity of PAHs with O_3 observed on different substrates is due, to a large extent, to differences in the partitioning rate of O_3 to the surface rather than differences in the surface reaction mechanism(s). However, no direct correlation between K_{O_3} and pseudo- $k^{(1)}$ is found in these data and the lack of $k^{(2)}$ values derived from these studies mean that no solid conclusions can be drawn regarding these reactions.

Differences in K_{O_3} values have been attributed to differences in binding interactions of the initial ozone-substrate complex.²⁹² Therefore, K_{O_3} will be influenced to a large degree by the specific type of particle and/or the coating of additional species (*e.g.* water, organic molecules *etc.*) that will affect the affinity of ozone to the surface and therefore influence PAH reactivity.²⁹¹ It has been suggested, based on adsorption enthalpy values, that ozone will interact more strongly with non-polar surfaces than with polar surfaces.²⁹⁰ This would explain the higher rate of O_3 partitioning to non-polar soot and phenylsiloxane oil aerosols, for example, compared with polar NaCl and azelaic acid aerosols.

A number of studies have identified both ring-opened (*e.g.* dialdehydes, dicarboxylic acids, keto carboxylic acids), and ring-retaining (*e.g.* quinones) products from the heterogeneous

reactions of a number of PAHs with O_3 on particle surfaces.¹ This suggests the significance of both bond-attack (ring-opening) and atom-attack (substitution) mechanisms at the most electrophilic positions.²⁹⁷

Van Cauwenberghe²⁵¹ proposed 2 mechanisms for the heterogeneous reaction of PAHs with O_3 . (1) A one-step electrophilic-nucleophilic attack on olefin bonds with the highest electron density to initially form a primary ozonide. This will decompose to yield ring-opened products such as aldehydes and carboxylic acid compounds. (2) A two-step electrophilic attack that will ultimately lead to the formation of quinones.¹

Miet *et al.*²⁹⁷ tentatively identified several oxidation products formed from the heterogeneous reaction of pyrene with O_3 on silica including 1-hydroxypyrene, phenanthrene-4,5-dicarboxaldehyde and 4-oxapyren-5-one. The proposed mechanisms, suggested by the authors to explain the formation of these products are depicted in Scheme 10. It was suggested that hydroxypyrene would be formed from the reaction of O_3 with the most reactive carbon atom while phenanthrene-4,5-dicarboxaldehyde would be formed by ozonolysis of a C–C bond.²⁹⁷ It was also suggested that 4-oxapyren-5-one would be formed by further reaction of phenanthrene-4,5-dicarboxaldehyde.

Ringuet *et al.*²⁵⁸ investigated PAH decay and PAH derivative formation resulting from heterogeneous reactions of PAHs adsorbed on natural atmospheric particles collected from a high traffic area, exposed to OH and O_3 . BaP was noted as the most reactive PAH studied, in agreement with the findings of Perraudin *et al.*²⁹⁸ for reactions on silica and graphite particles. PAH decay and the formation of PAH ketones and quinones *e.g.* 9-Flr and AQu were observed. However, it was noted that the rate of OPAH formation did not match with the rate of parent PAH degradation in these cases, suggesting that other compounds/species are involved in these reactions.

The formation of NPAH compounds was also noted in these reaction studies, which is surprising as NO_2 was absent in this experiment. Organic nitrogen present was suggested as

the precursor, which may react with O_3 to form a reactive species (e.g. NO_x , NO_2^- , NO_3^-) which then reacts with the PAH upon O_3 attack.

PAH degradation and NPAH formation was also observed for reaction of PAH with a mixture of NO_2 and O_3 . However, it was noted that the formation of these products accounted for only 0.2% and 0.4% of the parent PAH decay for Flt and BaP, respectively. The observation of significant amounts of 2N-Flt and 2N-Pyr from these reactions indicates that these compounds are formed in homogeneous as well as heterogeneous reactions.

The reaction of ozone with anthracene adsorbed onto aqueous surfaces or Pyrex leads to the formation of 9,10-AQ only,^{292,294} while both anthraquinone and anthrone are formed on silica particles.³⁰³

A possible mechanism for the reaction has been proposed, consistent with that described by Bailey.³⁰⁴ This was suggested

to involve two potential pathways (Scheme 11): the first of these involves three consecutive electrophilic attacks by ozone at the C_9 and C_{10} positions of anthracene with the first two ozone attacks being followed by the loss of oxygen and a proton transfer and the third attack followed by loss of oxygen and water.²⁹⁴ The second of these pathways involves the formation of an ozonide due to O_3 addition across the C_9 – C_{10} bond of Ant, which re-arranges to hydroperoxide, which finally dehydrates to form anthraquinone.³⁰³ There is some disagreement between these studies as to which of these reaction pathways dominates. More work is clearly required to establish the mechanism responsible for the formation of quinones from heterogeneous reactions of PAHs with ozone.

The products of the heterogeneous reactions of BaP with ozone were investigated by Letzel *et al.*³⁰⁵ They identified BaP-1,6-dione, BaP-3,6-dione and BaP-6,12-dione as the main reaction products, as previously observed by Pitts *et al.*,²⁶⁷ as well as

Scheme 11 Suggested mechanisms for the heterogeneous reaction of anthracene with O_3 .²⁹²

BaP-4,5-dione and 4-oxa-benzo[*d,e,f*]chrysene-5-one (B(def)C-lactone). The formation of these products would seem to be broadly consistent with the mechanism for the reaction of anthracene (Scheme 11).

The reaction yield of the production of anthraquinone from the reaction of anthracene with O₃ on a pyrex surface was calculated by Kwamena *et al.*²⁹² to be approximately 30% at O₃ concentrations of up to 9.1×10^{15} molecules cm⁻³. However at atmospherically relevant ozone mixing ratios (100 ppbV) the AQ_u formation yield was expected to be very small (<1%), suggesting that this reaction route may be of minor importance in the ambient atmosphere.

3.2.2.6. Factors influencing heterogeneous PAH reactivity

3.2.2.6.1. Substrate type and conditions. It is clear from the discussion above that the chemical composition and specific conditions of the substrate to which PAHs are associated will have a critical role in the heterogeneous reactivity of these compounds towards oxidising species. This 'substrate effect' has the potential to influence the degree of partitioning of the oxidant to the particle surface (as demonstrated for ozone reactions), the relative accessibility of the PAH within the particle matrix, as well as the precise reaction mechanism.

3.2.2.6.1.1. Reaction with NO₂. Ramdahl *et al.*²⁷⁹ showed that PAH reactivity on different surfaces follows the order silica > alumina > charcoal, for six PAHs exposed to 0.5 ppmV of NO₂ in air containing water vapour and traces of HNO₃. It was estimated that PAH reactivity on alumina was 14–24% of that on silica. Similarly, Jäger and Hanus³⁰⁶ observed the order of reactivity for the reaction of several PAHs with 1.3 ppmV of NO₂ to be silica gel > fly ash > aluminium oxide > carbon.

The enhanced reactivity of NO₂ with PAH adsorbed on silica relative to that associated with alumina was attributed to the catalytic effect of HNO₃, as discussed above.^{280,284,285,307} The formation of nitric acid on the acidic silica surface is likely to be greater than that on neutral alumina surfaces which may explain the higher reactivity.²⁷⁹ Similarly, Ma *et al.*²⁷⁸ noted that the reaction rate on SiO₂ was almost two times faster than on MgO. The authors attributed this to the formation of HNO₃, which occurs on the surface of SiO₂ but not on MgO.

Inazu *et al.*²⁷⁴ investigated the reaction of Flt with NO₂ on different particle types and the factors influencing the degradation of Flt and the yield of N-Flt's. The amount of Flt degradation ranged from 20.9% for graphite to 79.5% for TiO₂ and followed the order TiO₂ > ZnO > Al₂O₃ > MgO > CaO > SiO₂ > Fe₂O₃ > graphite. The authors suggested three possible types of matrix effect to explain the different degradation rates and N-Flt yields on different substrate types.

One possible impact is suggested to be the photocatalytic generation of active species such as O₂⁻ and O⁻ on the particle surface.²⁷⁴ This would apply only to TiO₂ and ZnO and Fe₂O₃, which have a band gap energy of less than 4.3 eV, explaining the higher degradation rates of Flt on TiO₂ and ZnO. The reason for Fe₂O₃ degradation being lower than expected is not currently known. While this phenomenon may explain the higher degradation of Flt on these compounds, the overall yield and selectivity of

N-Flt products are lower than for other matrices such as MnO. This was attributed to the photoactivation of the support leading to the formation of oxidising agents as opposed to nitrating agents.

Another effect proposed to influence the Flt reactivity was potential fluoranthene–surface interaction due to the acidity/basicity of the matrix.²⁷⁴ The order of matrix acidity was Al₂O₃ > SiO₂ > TiO₂ > Fe₂O₃ > ZnO > MgO > CaO. This would explain the relatively high degradation rate of Al₂O₃ resulting from cation radical formation. The selectivity of nitration on different supports was observed to be almost opposite to the acidity order. It is therefore suggested that basicity of the support facilitates N-Flt formation and hence explains the higher reaction yield and selectivity on MgO.

The photochemical activation of fluoranthene sorbed to a surface was not considered to be an important factor in these studies.²⁷⁴ It was noted that irradiation and the presence of O₂ are important factors in heterogeneous reactions of PAHs but the specific surface area and particle size have a relatively small influence compared to the surface chemical properties.

3.2.2.6.1.2. Reactions with O₃ and OH. The reactivity of PAHs associated with graphite, diesel exhaust and urban particles towards ozone, OH and NO₂ have been shown to reach a 'plateau', indicating that no further reaction occurs after a certain exposure time.^{255,256,258} These heterogeneous processes proceed *via* a rapid surface reaction but a significant fraction of PAH is unavailable for the reaction due to poor diffusion of the oxidant through and/or inaccessibility of the PAH within the bulk particle. This was studied for the reaction of BaP sorbed to ammonium sulfate particles with O₃.³⁰⁸ Solid EC coating and various liquid organic coatings (when exceeding 10 nm of thickness) prevented the reaction. These results provide a strong indication for matrix effects to enhance the LRTP of PAHs.

Interesting differences were noted between the reactions of O₃ with PAHs for graphite and silica particles: the plateau observed in the case of graphite was relatively similar for the different PAHs studied (40–60%), while the fraction of unreacted PAH was highly variable in the case of silica particles (10–90%) and appeared to be dependent on the reactivity of individual PAH molecules.²⁹⁸

For reactions with both NO₂ and OH degradation on diesel particles is slower or stops earlier than on graphite particles.²⁵⁶ For example the amount of PAH degradation was 7–25% on diesel particles, compared to 10–75% on graphite particles (the same NO₂ concentration). These differences may be related to differences in particle surface area (108 m² g⁻¹ for diesel *vs.* 13.2 m² g⁻¹ for graphite) and the corresponding presence of pores that could house adsorbed pollutant molecules, leading to differences in the relative accessibility of oxidants.²⁵⁶ This difference may be caused by differences in particle generation: in contrast to physisorption of PAHs to (pure) graphite particles, PAHs present in diesel are generated simultaneously and incorporated in the carbonaceous core of the particles.²⁵⁶

It has been noted that the ozonation reactions of PAHs on solid substrates (*e.g.* silica and graphite) are considerably faster

than the reactions on organic films. For example, the rate coefficients for the reaction of ozone with BaP and Anth were shown to be up to a factor of ten higher in the studies conducted by Kwamena and co-workers^{292,301} than those calculated by Kahan *et al.*,²⁹⁶ seemingly contradicting the above suggestion of enhanced affinity of ozone for organic surfaces. However, this observation has been attributed to the potential formation of PAH dimers on the solid surface which may react differently with the gaseous oxidant and thus alter the reaction kinetics.²⁹⁶

Ozone will have a higher affinity for non-polar surfaces, therefore will be more attracted to hydrophobic, organic particles like soot or azelaic acid and will have very low affinity for polar, ionic particles like NaCl.²⁹¹ For example, Kwamena *et al.*²⁹¹ noted that the partitioning efficiency for O₃ followed the order elemental carbon > fixed silica > non-activated silica > solid organic carbon > inorganic salts.

Therefore, the heterogeneous reaction of PAH with ozone could be enhanced by the presence of water. For example, Kwamena *et al.*²⁹¹ observed a partitioning coefficient for ozone to azelaic acid at high RH that was 230% higher than the value that was observed under dry conditions, leading to a pseudo-first order rate coefficient increase for the reaction of anthracene of 25%.²⁹¹ It was suggested that the partitioning of gas phase water to the particle provided a substrate to which O₃ had a greater affinity and hence led to enhanced O₃ partitioning. While the *K*_{O₃} values for these two RH conditions differ by more than a factor 2, the actual pseudo-first order rate coefficient does not vary to such a degree (Table 8).

In contrast, however, in the presence of water both gas phase ozone loss and BaP decay rates were reduced,²⁹⁰ attributed to the competitive adsorption of O₃ and H₂O. Additionally, the reaction kinetics of anthracene reacting with ozone on Pyrex tubes was unaffected by RH conditions.²⁹² Pyrex may be too hydrophobic for water adsorption and, hence, water may not have partitioned to the Pyrex surface during the high RH experiments.²⁹²

The enhanced reaction rate of anthracene with ozone on the octanol-coated aqueous surface compared to the pure water surface^{293,294} can be explained by the fact that octanol is less polar than water and, hence, ozone will have a higher affinity for this surface. A similar explanation was given for the higher reaction rate for naphthalene on water droplets in the presence of fulvic acid.²⁹⁵

3.2.2.6.2. Influence of PAH molecular structure. Where heterogeneous reactions of PAHs with NO₂ have been reported, the relative heterogeneous reactivity of NO₂ with different PAHs has been observed to follow the general order BaP > Pyr > Chr > Flt, Phe.^{255,256,270,279} It is indicated that for the reactions of PAHs with O₃, Anth, BaA, BaP and DBaP are the most reactive compounds and Flt, Chr, BkF and Cor are the least reactive.^{286,288,298}

It has been suggested that the nature of the PAH structure has the potential to influence heterogeneous reactivity of PAHs associated with silica particles, both for reactions with NO₂ and O₃.^{276,298} The values of the derived second-order rate coefficients are considerably more variable between molecules

for the reactions of PAHs with NO₂ (variation of ≈ 5 orders of magnitude) than the reactions with O₃ (variation of ≈ 1 order of magnitude).

The variability of reaction kinetics for different PAH structures observed for silica particles is in contrast to observations of PAHs adsorbed on carbonaceous particles (graphite, diesel exhaust, wood soot) where oxidation rates do not appear to be strongly dependent on the nature of the PAH for OH, NO₂ and O₃ reactions.^{252,255,256,276,298} The derived *k*⁽²⁾ values in these studies fall within the same order of magnitude for all PAH molecules assessed in these studies.

Similar rate coefficients of PAHs associated with carbonaceous particles, for example graphite, may be explained by chemical stabilisation of the PAH due to delocalised π -electrons, which is effective regardless of PAH structure.²⁷⁶ The interactions between PAHs and silica are not as strong as those for graphite, which may explain the enhanced differences in heterogeneous reactivity between the individual PAHs.²⁷⁶

Kwamena *et al.*³⁰¹ suggested that the values of the adsorption coefficient (*K*_{O₃}) for the reactions of surface-bound PAH and ozone reflect the partitioning behaviour of ozone to the surface irrespective of the PAH adsorbed onto it. This indicates that, once ozone is adsorbed onto the aerosol surface, there is an inherent barrier to the reaction regardless of the specific PAH that would result in similar rate coefficient values, although there is significant variation between calculated pseudo-first order rate coefficients in the reported studies. This barrier could be related to the level of mobility of the two reactants or to the transformation of ozone into more reactive forms on the substrate surface.

The lack of an observable trend between PAH structure and reactivity for heterogeneous reactions with OH is in contrast to the observations made for the corresponding reactions in the gas phase. In general, it is shown that the reactivity of PAHs with OH in the gas phase depends strongly on the structure of the specific molecule.^{135,140,210} This clearly highlights the importance of the 'stabilising effect' of organic particles on the reactivity of PAHs towards OH radicals. Similar comparisons for NO₂ and O₃ are not possible due to lack of data.

3.2.2.6.3. Influence of PAH surface concentration. It is suggested that particle PAH reactivity with oxidants is influenced by the amount of PAH coated on the particle surface. Perraudin *et al.*²⁹⁸ and Miet *et al.*²⁹⁷ reported that the rate coefficient for the reaction of O₃ with pyrene adsorbed on silica particles can be up to 2 orders of magnitude higher when the particle surface concentration of pyrene was $\approx 50 \mu\text{g g}^{-1}$ compared to that of a surface concentration of $\approx 500 \mu\text{g g}^{-1}$. The effect of this 'bulk shielding' due to multilayer coverage of BaP on soot particles was also shown to reduce the heterogeneous reactivity with ozone.²⁹⁰ This demonstrates that heterogeneous reactivity of PAH with atmospheric oxidants is dependent on both the gaseous concentration of reactant and the particulate loading of PAH.

Perraudin *et al.*²⁹⁸ commented that the enhanced heterogeneous reactivity of PAHs on silica particles at lower PAH particle loadings could partly be explained by the nature of the

surface adsorption sites. They suggested that PAHs may be adsorbed on at least two different types of silica surface site, initially being coated on the more reactive sites, leading to higher reaction rates at lower surface concentrations. Miet *et al.*²⁹⁷ noted that, while reactivity of PAH adsorbed on silica particles towards O₃ is linearly dependent on gaseous O₃ concentration for a constant PAH surface coating, the relationship for a constant O₃ concentration and different PAH particle loading is rather more complex. The authors attributed this to the possibility of PAH being adsorbed on two types of silica surface sites.

Similarly, Alebic-Juretic *et al.*³⁰² indicated that the rate coefficient for the reaction of ozone with PAHs adsorbed on silica gel surfaces is 1.3–2.3 times higher for less than a monolayer coverage of PAH than for a more than one monolayer. Persistence of BaP on soot and ammonium sulfate particles exposed to O₃ was attributed to this ‘bulk shielding’ effect^{290,308} (Section 3.2.2.6.1.2). Therefore, instances of negligible reactivity reported for the reaction of PAHs exposed to ambient O₃ levels^{83,281,309} may be explained by the lack of availability of BaP for reaction at the surface.

Butler and Crossley²⁷⁰ and Nguyen *et al.*²⁷⁵ reported a very low reactivity for PAHs on soot particles towards NO₂, in contrast to the studies of Estève *et al.*^{255,256} for graphite and SRM 1650a diesel exhaust particles. It was noted by Nguyen *et al.*²⁷⁵ that the typical PAH concentrations on SRM 1650a diesel exhaust are of the order of a few tens of µg g⁻¹ while those used by Butler and Crossley²⁷⁰ and Nguyen *et al.*²⁷⁵ are nearly two orders of magnitude higher.

A further explanation for the lower reactivity at higher PAH loading could be the presence of higher molecular weight oxidation products formed due to heterogeneous reactions.^{273,290} These may have a lower volatility than the parent PAH and thus prevent the attacking gaseous oxidant from reaching the PAH that is ‘hidden’ underneath the exposed surface when there is a greater than monolayer coverage.

3.2.2.7. Gas phase reaction vs. heterogeneous reaction. For semivolatile PAHs (mostly 3–4 ring compounds) comparisons for reactions with atmospheric oxidant gases can be carried out between reactions occurring in the gas phase and reactions occurring on various substrate types. This is possible only for the reactions of OH, NO₃ and O₃, as direct reactions of PAHs with gaseous NO₂ are not expected to occur in the gas phase (see Section 3.2.1).

It is suggested that reactivity of PAHs in the gas phase would be significantly greater than when associated with carbonaceous particles,^{255,256} suggesting that the particle surface may have an inhibiting effect on the heterogeneous reaction with OH. It can be seen that $k^{(2)}$ values for PAH reactions with OH on carbonaceous particle surfaces are 1–3 orders of magnitude lower than those derived for gas phase reactions.

The presence of a ‘plateau’ in the experimental decays of PAHs in the reactions of OH indicates that a significant fraction of PAH is unavailable for reaction.^{255,256} The higher level of this plateau observed in the case of diesel particles relative to that

for graphite particles confirms the enhanced inhibiting factor of diesel particles compared to graphite. This ‘stabilising’ effect of particles towards PAH reactivity is also reflected in the relatively similar rate coefficients observed for different PAHs associated with organic particles compared with the apparent PAH-dependent reactivity in the gas phase.

However, it is indicated that the reaction of PAHs with ozone is enhanced on particles in comparison with reactions in the gas phase. For example, Perraudin *et al.*²⁹⁸ studied the reaction of PAHs with ozone adsorbed on graphite and silica particles and reported $k^{(2)}$ values approximately 2 orders of magnitude higher than those calculated for corresponding gas phase reactions (see Tables 2 and 6). Similar comparisons with studies investigating the reaction of ozone with individual PAHs on soot particles,²⁹⁰ Pyrex,²⁹² organic aerosols³⁰¹ and the air–aqueous interface^{293,294} are more difficult to make as second-order rate coefficients were not reported in these studies.

3.2.2.8. Comparison of oxidising species: absolute reaction rates. To assess the relative importance of the second order rate coefficients for the reactions of PAHs in the gas phase (Tables 3, 5 and 6) and associated with particle surfaces (Table 7), absolute reaction rates have been derived using the second order reaction rates derived from the literature and concentrations of atmospheric species typically found in the background atmosphere (Tables 9 and 10).

These tables indicate that the gas phase reaction with OH remains the dominant loss process for PAHs and these reactions are approximately 3 orders of magnitude faster than the corresponding gas phase reactions with NO₃.

It appears that for heterogeneous processes, PAH reactions with NO₂ and O₃ may be more important than second order rate coefficients would suggest. Indeed the absolute rates for PAH reactions with NO₂ and O₃ are shown to be comparable to those of OH reactions, depending on the specific surface type and conditions.

3.2.2.9. Summary and concluding remarks

3.2.2.9.1. Key conclusions. • Heterogeneous reactivity of PAHs with atmospheric oxidants involves complex processes that have been shown to be influenced by a number of factors including the nature (chemical composition, surface area, porosity) of the matrix, presence of other species (*e.g.* nitric acid, water or other organic species), oxidant concentration, PAH molecule involved, and PAH surface concentration.

• The wide diversity in particle properties, including their chemical composition (organic, mineral, biogenic), sources (combustion, erosion, gas phase condensation), origin (natural, anthropogenic), the method of particle formation (temperature, pressure), physical properties (size, porosity, specific surface area), surface coatings (water, nitric acid, organic molecules), means that gaining a full understanding of this reactivity will be extremely difficult as the relative importance of these different factors is highly variable across reaction systems.

• Reaction with OH radicals is the dominant pathway for heterogeneous degradation of PAHs compared with reactions

Table 9 Pseudo-first order reaction rate coefficients $k^{(1)}$ for the gas-phase reactions of PAHs with OH, NO₃ and O₃

	OH		NO ₃		O ₃	
	$k^{(1)a}$ (s ⁻¹)	Ref.	$k^{(1)b}$ (s ⁻¹)	Ref.	$k^{(1)c}$ (s ⁻¹)	Ref.
Nap	4.8×10^{-5} 4.3×10^{-5} 4.6×10^{-5} 5.3×10^{-5} 4.7×10^{-5} 3.7×10^{-5} 4.3×10^{-5} 4.8×10^{-5} 5.2×10^{-5}	Phousongphouang and Arey ¹⁷¹ Atkinson ¹⁴⁷ Brubaker and Hites ¹⁴⁰ Klamt ¹⁷² Biermann <i>et al.</i> ¹⁶⁹ Lorenz and Zellner ¹⁵⁶ Klöpffer <i>et al.</i> ¹⁷⁰ Atkinson <i>et al.</i> ¹⁶⁸ Atkinson and Aschmann ¹⁴⁴	5.5×10^{-8} 3.1×10^{-8} 2.1×10^{-8} 2.4×10^{-8} 2.7×10^{-8} 2.3×10^{-8}	Pitts <i>et al.</i> ²¹² Atkinson <i>et al.</i> ¹⁶² Atkinson and Aschmann ¹⁷⁴ Atkinson <i>et al.</i> ¹⁶³ Atkinson <i>et al.</i> ¹⁶³ Atkinson ²¹⁰	1.4×10^{-7} 2.1×10^{-7}	Atkinson <i>et al.</i> ¹⁶⁸ Atkinson and Aschmann ¹⁴⁴
1M-Nap	8.2×10^{-5} 1.1×10^{-4} 1.2×10^{-4}	Phousongphouang and Arey ¹⁷¹ Atkinson and Aschmann ¹⁷³ Klamt ¹⁷²	5.4×10^{-8} 4.5×10^{-8} 5.0×10^{-8}	Atkinson and Aschmann ¹⁷³ Atkinson and Aschmann ¹⁷⁴ Atkinson ²¹⁰	9.0×10^{-8}	Atkinson and Aschmann ¹⁷³
2M-Nap	9.7×10^{-5} 1.1×10^{-4} 1.2×10^{-4}	Phousongphouang and Arey ¹⁷¹ Atkinson and Aschmann ¹⁴⁴ Klamt ¹⁷²	4.6×10^{-8} 6.9×10^{-8} 7.0×10^{-8} 6.6×10^{-8}	Phousongphouang and Arey ²¹⁴ Atkinson and Aschmann ¹⁷³ Atkinson and Aschmann ¹⁷⁴ Phousongphouang and Arey ²¹⁴	2.1×10^{-7} 2.8×10^{-7}	Atkinson and Aschmann ¹⁴⁴ Atkinson and Aschmann ¹⁷³
1E-Nap	7.3×10^{-5}	Phousongphouang and Arey ¹⁷¹	6.3×10^{-8}	Phousongphouang and Arey ²¹⁴		
2E-Nap	8.0×10^{-5}	Phousongphouang and Arey ¹⁷¹	5.2×10^{-8}	Phousongphouang and Arey ²¹⁴		
1,2DM-Nap	1.2×10^{-4}	Phousongphouang and Arey ¹⁷¹	4.1×10^{-7}	Phousongphouang and Arey ²¹⁴		
1,3DM-Nap	4.4×10^{-5} 1.5×10^{-4}	Banceu <i>et al.</i> ¹⁷⁵ Phousongphouang and Arey ¹⁷¹	1.4×10^{-7}	Phousongphouang and Arey ²¹⁴		
1,4DM-Nap	1.2×10^{-5} 1.2×10^{-4}	Klamt ¹⁷² Phousongphouang and Arey ¹⁷¹	8.4×10^{-8}	Phousongphouang and Arey ²¹⁴		
1,5DM-Nap	1.2×10^{-4}	Phousongphouang and Arey ¹⁷¹	9.1×10^{-8}	Phousongphouang and Arey ²¹⁴		
1,6DM-Nap	1.3×10^{-4}	Phousongphouang and Arey ¹⁷¹	1.1×10^{-7}	Phousongphouang and Arey ²¹⁴		
1,7DM-Nap	1.4×10^{-4}	Phousongphouang and Arey ¹⁷¹	8.7×10^{-8}	Phousongphouang and Arey ²¹⁴		
1,8DM-Nap	1.3×10^{-4}	Phousongphouang and Arey ¹⁷¹	1.4×10^{-6}	Phousongphouang and Arey ²¹⁴		
2,3DM-Nap	1.2×10^{-4} 2.0×10^{-4} 1.5×10^{-4}	Phousongphouang and Arey ¹⁷¹ Banceu <i>et al.</i> ¹⁷⁵ Atkinson and Aschmann ¹⁴⁴	9.5×10^{-9} 1.0×10^{-7} 1.0×10^{-7} 9.8×10^{-8}	Atkinson and Aschmann ¹⁷³ Atkinson and Aschmann ¹⁷⁴ Atkinson ²¹⁰ Phousongphouang and Arey ²¹⁴	2.8×10^{-7}	Atkinson and Aschmann ¹⁴⁴
2,6DM-Nap	1.3×10^{-4}	Phousongphouang and Arey ¹⁷¹	1.4×10^{-7}	Phousongphouang and Arey ²¹⁴		
2,7DM-Nap	1.4×10^{-4}	Phousongphouang and Arey ¹⁷¹	1.4×10^{-7}	Phousongphouang and Arey ²¹⁴		
Ace	1.6×10^{-4} 1.2×10^{-4} 2.1×10^{-4} 1.2×10^{-4} 1.3×10^{-4} 1.6×10^{-4}	Reisen and Arey ¹⁵⁵ Brubaker and Hites ¹⁴⁰ Atkinson and Aschmann ¹⁷⁴ Klöpffer <i>et al.</i> ¹⁷⁰ Banceu <i>et al.</i> ¹⁷⁵ Klamt ¹⁷²	2.3×10^{-4} 1.1×10^{-7}	Atkinson and Aschmann ¹⁷⁴ Atkinson and Aschmann ¹⁷⁴	3.5×10^{-7}	Atkinson and Aschmann ¹⁷⁴
Acy	2.5×10^{-4} 2.6×10^{-4} 2.2×10^{-4}	Reisen and Arey ¹⁵⁵ Banceu <i>et al.</i> ¹⁷⁵ Atkinson and Aschmann ¹⁷⁴	2.7×10^{-3}	Atkinson and Aschmann ¹⁷⁴	3.8×10^{-4} 1.1×10^{-4}	Atkinson and Aschmann ¹⁷⁴ Reisen and Arey ¹⁵⁵
Flu	3.2×10^{-5} 2.6×10^{-5} 2.0×10^{-5} 2.6×10^{-5}	Kwok <i>et al.</i> ¹⁷⁷ Brubaker and Hites ¹⁴⁰ Klamt ¹⁷² Klöpffer <i>et al.</i> ¹⁷⁰	1.8×10^{-5}	Kwok <i>et al.</i> ¹⁷⁷	1.4×10^{-7}	Kwok <i>et al.</i> ¹⁷⁷
Phe	6.8×10^{-5} 6.2×10^{-5} 5.1×10^{-5} 3.1×10^{-5} 2.5×10^{-5} 5.4×10^{-5} 6.4×10^{-5}	Biermann <i>et al.</i> ¹⁶⁹ Atkinson ¹⁴⁷ Klamt ¹⁷² Lorenz and Zellner ¹⁵⁶ Kwok <i>et al.</i> ¹⁷⁶ Brubaker and Hites ¹⁴⁰ Lee <i>et al.</i> ¹⁵⁹	6.0×10^{-5}	Kwok <i>et al.</i> ¹⁷⁶	2.8×10^{-7}	Kwok <i>et al.</i> ¹⁷⁶
1M-Phe	5.8×10^{-5}	Lee <i>et al.</i> ¹⁵⁹				
2M-Phe	1.3×10^{-4}	Lee <i>et al.</i> ¹⁵⁹				
3M-Phe	1.3×10^{-4}	Lee <i>et al.</i> ¹⁵⁹				
9M-Phe	1.5×10^{-4}	Lee <i>et al.</i> ¹⁵⁹				
Ant	2.2×10^{-4} 3.8×10^{-4} 2.6×10^{-5} 4.0×10^{-4} 2.6×10^{-4}	Biermann <i>et al.</i> ¹⁶⁹ Brubaker and Hites ¹⁴⁰ Kwok <i>et al.</i> ¹⁷⁶ Klamt ¹⁷² Atkinson ¹⁴⁷ , Biermann <i>et al.</i> ¹⁶⁹				
Flt	2.2×10^{-5}	Brubaker and Hites ¹⁴⁰	3.3×10^{-8}	Atkinson <i>et al.</i> ¹⁶³		
Pyr	1.0×10^{-4}	Atkinson <i>et al.</i> ¹⁶³	1.0×10^{-7}	Atkinson <i>et al.</i> ¹⁶³		
1N-Nap	1.1×10^{-4}	Atkinson ¹⁴⁷	1.9×10^{-8}	Atkinson ²¹⁰	4.1×10^{-7}	Atkinson ²⁰⁷
2N-Nap	1.1×10^{-4}	Atkinson ¹⁴⁷	1.7×10^{-8}	Atkinson ²¹⁰	4.1×10^{-7}	Atkinson ²⁰⁷

^a [OH] = 2×10^6 molecules cm⁻³; global 12 h average (Atkinson and Arey¹⁵¹). ^b [NO₃] = 5×10^8 molecules cm⁻³; global 12 h average (Atkinson and Arey¹⁵¹). ^c [O₃] = 6.9×10^{11} molecules cm⁻³; 2011 annual average, Harwell, UK.

Table 10 Pseudo-first order reaction rate coefficients $k^{(1)}$ for the heterogeneous reactions of PAHs with OH, NO₂, O₃ and NO₃

PAH	OH reactions			NO ₂ reactions			O ₃ reactions			NO ₃ reactions		
	$k_{OH}^{(1) a}$ (s ⁻¹)	Ref.	Notes	$k_{NO_2}^{(1) b}$ (s ⁻¹)	Ref.	Notes	$k_{O_3}^{(1) c}$ (s ⁻¹)	Ref.	Notes	$k_{NO_3}^{(1) d}$ (s ⁻¹)	Ref.	Notes
Phe	1.0×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	1.9×10^{-7}	Perraudin <i>et al.</i> ²⁷⁶	Perraudin <i>et al.</i> ²⁷⁵	Silica particles	3.1×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles		
	6.5×10^{-7}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles	2.3×10^{-6}	Nguyen <i>et al.</i> ²⁷⁵	Nguyen <i>et al.</i> ²⁵⁵	Kerosene flame soot	3.0×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles		
				2.3×10^{-6}	Estève <i>et al.</i> ²⁵⁵	Estève <i>et al.</i> ²⁵⁵	Graphite particles					
Ant				7.8×10^{-6}	Estève <i>et al.</i> ²⁵⁶	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles					
				1.6×10^{-9}	Butler and Crossley ²⁷⁰	Butler and Crossley ²⁷⁰	Ethylene flame soot					
	8.8×10^{-6}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	6.9×10^{-5}	Perraudin <i>et al.</i> ²⁷⁶	Perraudin <i>et al.</i> ²⁷⁵	Silica particles	1.3×10^{-5}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles		
Flo				2.3×10^{-6}	Nguyen <i>et al.</i> ²⁷⁵	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	1.8×10^{-5}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles		
				4.8×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Estève <i>et al.</i> ²⁵⁵	Graphite particles					
	2.9×10^{-8}	Bedjanian <i>et al.</i> ²⁵²		6.9×10^{-5}	Ma <i>et al.</i> ²⁷⁸	Ma <i>et al.</i> ²⁷⁸	Silica particles					
Pyr				3.7×10^{-5}	Ma <i>et al.</i> ²⁷⁸	Ma <i>et al.</i> ²⁷⁸	MgO particles					
	6.5×10^{-6}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	2.2×10^{-9}	Perraudin <i>et al.</i> ²⁷⁶	Perraudin <i>et al.</i> ²⁷⁵	Silica particles	2.5×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles		
	8.2×10^{-7}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles	6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	1.9×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles		
Chr				6.9×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Estève <i>et al.</i> ²⁵⁵	Graphite particles					
	7.6×10^{-7}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles	2.0×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Estève <i>et al.</i> ²⁵⁵	Graphite particles					
	3.1×10^{-8}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	6.9×10^{-6}	Estève <i>et al.</i> ²⁵⁶	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles	3.2×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles	3.2×10^{-3}	Liu <i>et al.</i> ²⁶⁰
Pyr				1.7×10^{-9}	Butler and Crossley ²⁷⁰	Butler and Crossley ²⁷⁰	Ethylene flame soot					
	6.5×10^{-6}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	1.4×10^{-5}	Perraudin <i>et al.</i> ²⁷⁶	Perraudin <i>et al.</i> ²⁷⁵	Silica particles	7.6×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles		
	8.2×10^{-7}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles	3.5×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Estève <i>et al.</i> ²⁵⁵	Kerosene flame soot	4.2×10^{-5}	Miet <i>et al.</i> ²⁹⁷	Silica particles		
Chr				1.0×10^{-5}	Estève <i>et al.</i> ²⁵⁶	Estève <i>et al.</i> ²⁵⁶	Graphite particles					
	4.8×10^{-7}	Miet <i>et al.</i> ²⁵⁷	Silica particles	1.0×10^{-5}	Estève <i>et al.</i> ²⁵⁶	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles					
				3.3×10^{-9}	Butler and Crossley ²⁷⁰	Butler and Crossley ²⁷⁰	Ethylene flame soot					
Chr				6.4×10^{-5}	Miet <i>et al.</i> ²⁷⁷	Miet <i>et al.</i> ²⁷⁷	Silica particles					
	1.8×10^{-8}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	4.1×10^{-7}	Perraudin <i>et al.</i> ²⁷⁶	Perraudin <i>et al.</i> ²⁷⁵	Silica particles	1.9×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles	2.0×10^{-3}	Liu <i>et al.</i> ²⁶⁰
	1.0×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	4.0×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles		
Chr				2.6×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Estève <i>et al.</i> ²⁵⁵	Graphite particles					
	8.8×10^{-7}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles	6.9×10^{-6}	Estève <i>et al.</i> ²⁵⁶	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles					
				1.8×10^{-9}	Butler and Crossley ²⁷⁰	Butler and Crossley ²⁷⁰	Ethylene flame soot

Table 10 (continued)

OH reactions				NO ₂ reactions		O ₃ reactions		NO ₃ reactions				
PAH	$k_{\text{OH}}^{(1)}$ ^a (s ⁻¹)	Ref.	Notes	$k_{\text{NO}_2}^{(1)}$ ^b (s ⁻¹)	Ref.	Notes	$k_{\text{O}_3}^{(1)}$ ^c (s ⁻¹)	Ref.	Notes	$k_{\text{NO}_3}^{(1)}$ ^d (s ⁻¹)	Ref.	Notes
BaA	1.8×10^{-8}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	4.6×10^{-6}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	3.6×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles	2.2×10^{-3}	Liu <i>et al.</i> ²⁶⁰	Azelaic acid particles
	1.1×10^{-6}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	1.1×10^{-5}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
	6.5×10^{-7}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles	2.2×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Graphite particles						
BbF				8.6×10^{-6}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles						
				4.3×10^{-9}	Butler and Crossley ²⁷⁰	Ethylene flame soot						
	2.1×10^{-8}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	1.5×10^{-6}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	2.5×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
	7.1×10^{-6}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	4.6×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
				1.7×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Graphite particles						
	2.2×10^{-8}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	6.4×10^{-4}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	6.8×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
BaP	8.2×10^{-6}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	1.8×10^{-5}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
	5.9×10^{-7}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles	5.4×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Graphite particles						
				1.0×10^{-5}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles						
BeP				6.9×10^{-9}	Butler and Crossley ²⁷⁰	Ethylene flame soot						
	2.2×10^{-8}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	2.0×10^{-6}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	2.1×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
	9.4×10^{-6}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	3.7×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
	9.4×10^{-7}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles	2.4×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Graphite particles						
				5.2×10^{-6}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles						
				2.0×10^{-9}	Butler and Crossley ²⁷⁰	Ethylene flame soot						
Per	1.0×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	7.3×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Graphite particles						
IPy	7.1×10^{-7}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles	4.3×10^{-6}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	2.5×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
				6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	4.9×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
				5.2×10^{-6}	Estève <i>et al.</i> ²⁵⁶	Diesel exhaust particles						
BgP				3.2×10^{-5}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	2.5×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
	1.2×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Graphite particles	6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot	9.2×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
				2.6×10^{-5}	Estève <i>et al.</i> ²⁵⁵	Graphite particles

Table 10 (continued)

OH reactions			NO ₂ reactions			O ₃ reactions			NO ₃ reactions			
PAH	$k_{\text{OH}}^{(1) \text{ } a}$ (s ⁻¹)	Ref.	Notes	$k_{\text{NO}_2}^{(1) \text{ } b}$ (s ⁻¹)	Ref.	Notes	$k_{\text{O}_3}^{(1) \text{ } c}$ (s ⁻¹)	Ref.	Notes	$k_{\text{NO}_3}^{(1) \text{ } d}$ (s ⁻¹)	Ref.	Notes
BbF	1.7×10^{-8}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	5.5×10^{-9}	Butler and Crossley ²⁷⁰	Ethylene flame soot						
				6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot						
	2.1×10^{-8}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot						
				6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot						
	2.1×10^{-8}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot						
DBaP	2.4×10^{-8}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot						
				6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot						
	2.2×10^{-8}	Bedjanian <i>et al.</i> ²⁵²	Kerosene flame soot	9.2×10^{-7}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	2.5×10^{-6}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
DBaP				6.9×10^{-8}	Nguyen <i>et al.</i> ²⁷⁵	Kerosene flame soot						
				1.6×10^{-9}	Butler and Crossley ²⁷⁰	Ethylene flame soot						
				1.2×10^{-4}	Perraudin <i>et al.</i> ²⁷⁶	Silica particles	1.7×10^{-5}	Perraudin <i>et al.</i> ²⁹⁸	Graphite particles			
							1.7×10^{-5}	Perraudin <i>et al.</i> ²⁹⁸	Silica particles			
1N-Pyr	2.0×10^{-7}	Miet <i>et al.</i> ²⁵⁷	Silica particles	4.3×10^{-6}	Miet <i>et al.</i> ²⁷⁷	Silica particles	2.9×10^{-6}	Miet <i>et al.</i> ²⁹⁷	Silica particles	6.5×10^{-4}	Liu <i>et al.</i> ²⁶⁰	Azelaic acid particles

^a [OH] = 2×10^6 molecules cm⁻³; global 12 h average.^{151 b} [NO₂] = 1.3×10^{11} ; 2011 annual average, Harwell, UK. ^c [O₃] = 6.9×10^{11} molecules cm⁻³; 2011 annual average, Harwell, UK. ^d [NO₃] = 5×10^8 molecules cm⁻³; global 12 h average.¹⁵¹

with NO₂ and O₃ with second order rate coefficients 1–7 orders of magnitude higher on carbonaceous particles.

- Particles are shown to exhibit a potential ‘inhibiting factor’ on the reactivity of PAHs due to slow diffusion of the oxidant or inaccessibility of PAHs in the bulk particle. This may turn PAHs to persistent compounds in air.

- Carbonaceous particles have the potential to stabilize the reactivity of PAH towards OH with slower rates than observed in the gas phase.

- Heterogeneous reactions of PAHs with O₃ are observed to proceed more quickly than the corresponding reactions in the gas phase, with second order rate coefficients ≈ 2 orders of magnitude greater for certain compounds and model matrices studied so far.

- A number of ring-opened and ring-retaining products have been identified from these heterogeneous reactions including nitro-PAHs and quinones. In some cases, a difference has been noted between heterogeneous and gas phase reaction products. For example the observation of 1-nitropyrene from the reaction of particle-bound PAH is in contrast to the production of 2-nitropyrene from gas phase reactions.

- The heterogeneous reaction of PAH with ozone on a range of surfaces has been shown to follow Langmuir–Hinshelwood-type kinetics, where the pseudo-first order reaction rate is characterized by a dependence on gas phase concentration and available reaction sites and reactivity is strongly dependent on the degree of partitioning of ozone to the surface. This is shown to be strongly influenced by the specific nature of the particle surface.

- When atmospheric concentrations of the different oxidising species are considered, the absolute heterogeneous reaction rates of PAHs with OH, NO₂ and O₃ are comparable, depending on the specific particle surface and reaction conditions.

- Heterogeneous reaction chemistry of PAHs is far from being fully understood with knowledge of the reaction mechanisms, products and the factors that influence the reaction kinetics still lacking. Reaction rates for ambient conditions are very difficult to extrapolate, also because many aerosol matrices have not been included in laboratory studies so far.

3.2.2.9.2. Recommendations for future work. The following are areas we suggest for continued research in the area of heterogeneous PAH reactivity:

- Investigation of the mechanisms of the heterogeneous reactions of PAHs with NO₂, OH and NO₃ for various aerosol matrices and conditions.

- More comprehensive elucidation of reaction products of the heterogeneous reactions of PAHs, particularly for OH and NO₃ reactions.

- Investigation of the reactions with OH radicals on silica particles for comparison with other reaction systems.

- Verification of the reactivity and products based on field studies, with particular emphasis on matrix effects.

3.2.3. Atmospheric implications. Photochemical lifetimes in the background atmosphere of gaseous PAHs, τ_{hom} , can be estimated based on the O₃ and OH kinetics (Sections 3.2.1.1 and 3.2.1.3), while the reactivity with oxidized N species will be

largely limited to the high NO_x source areas. Present knowledge suggests τ_{hom} to range between ≈ 1 hour and ≈ 2 days for oxidant concentrations typical for the continental background in mid latitudes. As an implication of the diel variation of radical abundances, longer lifetimes result for night-time emission or transport in high latitudes.

Clearly, heterogeneous reactions (Section 3.2.2) can significantly contribute to the removal of semivolatile and non-volatile PAHs from the atmosphere and to the formation of mutagenic derivative compounds. A quantitative assessment of this contribution has to be considered preliminary and with great caution, because the representativity of model aerosols and laboratory conditions for ambient conditions is insufficient. For example, the matrix of secondary inorganic and organic aerosols has hardly been addressed. Furthermore, the processes ruling gas–particle partitioning are not fully understood, with the consequence that particulate mass fractions, θ , of the semivolatile PAHs in a given aerosol are uncertain, in particular in the cold (free troposphere, Arctic). Assuming representativity of studied model aerosols for ambient aerosols (namely silica for unspecific particulate mass, diesel and flame soot for BC) and a range of θ observed (although in the near ground atmosphere only) the PAHs’ lifetimes towards heterogeneous photochemical degradation, τ_{het} , are estimated to range between 2 hours and >10 days for oxidant and BC concentrations typical for the continental background in mid latitudes. (For values >10 days atmospheric lifetime would be limited by physical sinks, *i.e.* deposition.) Total photochemical residence times, τ_{total} , resulting as $\tau_{\text{total}} = [(1 - \theta)/\tau_{\text{hom}} + \theta/\tau_{\text{het}}]^{-1}$, are listed in Table 9. Obviously, most of the uncertainty of τ_{total} for semivolatile and non-volatile PAHs results from the consideration of heterogeneous reactivity (related to the aerosol characteristics’ variability). It enhances overall reactivity significantly for the semivolatile (partly partitioning to particles) PAHs, *e.g.* Flt and Chr, but less for Pyr (Table 9).

Zimmerman *et al.*³¹⁰ recently observed the formation of nitro-PAH isomers on PM samples exposed to a mixture of N₂O₅–NO₂–NO₃ in a chamber reaction study. The authors attributed this to a heterogeneous reaction with surface-adsorbed N₂O₅ as opposed to reaction with radical NO₃. It was proposed that such heterogeneous reactivity could lead to the formation of nitro-PAH isomers such as 1-NPyr during atmospheric transport of PM. However, the concentrations of NO₂ and N₂O₅ used in this study were factors of >100 higher than observed under ambient conditions so the applicability of this observation in the atmosphere is uncertain.

4. Long-range transport

Slow photochemical degradation (Section 3.2) and limited deposition along transport (Section 3.1.2) permit long-range atmospheric transport of PAHs.

4.1. Observations at remote sites

Evidence of long-range transport of PAH comes from numerous observations in air and precipitation/deposition at remote

terrestrial and marine sites (Table 11). Only a few monitoring stations at remote sites are operating and these are concentrated in Europe and North America. Levels at remote sites are at least one order of magnitude lower than within source areas^{311,312} and some of the (often addressed) 16 USEPA priority PAHs were below the limits of quantification. Elevated PAH concentrations far from combustion sources were observed in long-range transported plumes of open fires, such as from biomass burning in Africa (over the Atlantic³¹³) and boreal forest fires (in high latitudes^{62,314}). Because of photodegradation being more efficient in summer and emissions being higher in winter,

winter levels usually exceed summer levels at receptor sites. An exception, *i.e.* minima in winter, was observed on mountain sites in Oregon, USA³¹⁵ which can be explained by air mass trajectory statistics. Observations at remote sites are usually considered to prove the LRT potential. In some cases, however, the influences of local emissions are difficult to exclude. For example, the Zeppelin observatory, near Ny Alesund, Svalbard, is considered to reflect the remoteness of the Arctic island's environment with usually very low levels observed (*e.g.* 46–75 pg m⁻³ in July and 260–1500 pg m⁻³ in January of recent years for the sum of 15 PAH; Table 11). However, samples collected at the station

Table 11 Arithmetic mean of PAH measurements in air at remote sites, arithmetic mean concentration (sum of gaseous and particulate phase concentrations, pg m⁻³). Sum of 15 PAHs, namely acenaphthylene, acenaphthene, fluorene, phenanthrene, anthracene, fluoranthene, pyrene, benzo[a]anthracene, chrysene, benzo[b]fluoranthene, benzo[k]fluoranthene, benzo[a]pyrene, indeno[123c,d]pyrene, dibenz[a,h]anthracene and benzo[ghi]perylene (*i.e.* 16 USEPA priority PAHs without naphthalene), unless otherwise specified

		<i>n</i>	<i>c</i> _{tot}	Ref.
Continental remote and mountain sites				
Tagish, Yukon, Canada 60°N/134°W	Oct 1993–Apr 1994/May–Sep 1993	7/5	300/61	Halsall <i>et al.</i> ³¹⁸
Redó, Pyrenees (2250 m a.s.l.)	Jun 1997/Feb 1997	7/3	≈ 1650/2750 ^a	Fernández <i>et al.</i> ³¹⁷
Gossenkölle, Alps (2413 m a.s.l.)	Jul 1997/Mar 1997	6/4	≈ 2750/2000 ^a	
Øvre Neådalsvatn, Central Norway (728 m a.s.l.)	1996–1998	8	1900	van Drooge <i>et al.</i> ³²⁵
Izana, Tenerife (2367 m a.s.l.)	2000	12	240	
Skalnaté Pleso, High Tatras (1778 m a.s.l.)	2001–2002	9	5100	
Mt. Bachelor, Oregon, USA (44°N/122°W, 2763 m a.s.l.)	Spring and summer/fall and winter 2004–2006	26/10	≈ 1400/≈ 450	Primbs <i>et al.</i> ³¹⁵
Lake Nam Co, Tibetan Plateau (31°/91°E, 4730 m a.s.l.)	Oct 2006–Feb 2008	15	75–1470 ^b	Xiao <i>et al.</i> ³²⁶
Jungfrauoch, Alps (3573 m a.s.l.)	Jul–Oct 2006	1	1440 ^{c,d}	Halse <i>et al.</i> ^{319e}
Moussala, Balkans (2925 m a.s.l.)		1	500 ^{c,d}	
Chopok, High Tatras (2008 m a.s.l.)		1	1300 ^{c,d}	
Zugspitze, Alps (2650 m a.s.l.)	Jun 2007/Feb 2008	8/14	71/1027	Lammel <i>et al.</i> ³²²
Molopo Reserve, Barberspan (steppe), South Africa (1366 m a.s.l.)	Jan–Jun 2008	12	36–125 ^e	Klánová <i>et al.</i> ³¹²
Tombouctou, (desert) Mali, West Africa (200 m a.s.l.)		6	55–160 ^e	
Mt. Kenya (3650 m a.s.l.)		6	23–36 ^e	
Marine sites				
Finokalia, Crete, Mediterranean 35°N/26°E	Feb 2000–Feb 2002	23	13 500	Tsapakis and Stephanou ²⁵
North Pacific (cruise) 52–62°N	Jul–Sep 2003	10	11 260	Ding <i>et al.</i> ²⁶
Biscay Bay, Atlantic (cruise) 46–49°N	Oct–Nov 2005	3	1370 ^f	Nizzetto <i>et al.</i> ³¹³
Eastern North Atlantic (cruise) 25–31°N		4	230 ^f	
Eastern Equatorial Atlantic (cruise) 0–25°S		13	170 ^f	
Malta, Giordan Lighthouse, Med. 36°N/14°E	Aug–Nov 2006	1	2510 ^{c,d}	Halse <i>et al.</i> ^{319e}
Mace Head, Ireland 53°N/10°W	Jul–Oct 2006	1	5690 ^{c,d}	
Arctic and Antarctic sites				
Barrow, Alaska, US, 71°N	Mar 1979	3	≈ 1190 ^b	Daisey <i>et al.</i> ³¹⁶
Aug 1979		5	157 ^b	
Dye 3, Greenland 65°N/44°W	Winter-spring 1989 and spring-summer 1991	9	18–200 ^f	Jaffrézo <i>et al.</i> ³²¹
Alert, Canadian Archipelago 82°N/62°W	Feb–Apr 1988	10	≈ 290–1100 ^b	Patton <i>et al.</i> ³²⁴
	Oct 1993–Apr 1994/May–Sep 1994	7/5	695/48	Halsall <i>et al.</i> ³¹⁸
	Jan–Dec 1992/1993/1994	52/52/	900/570/370	Hung <i>et al.</i> ³²⁰
	Jan–Dec 1998/1999/2000	52/52/	430/180/150	
		52		
Dunai Island, Eastern Siberia 74°N/124°E	Nov 1993–Mar 1994/May–Aug 1993	4/4	2450/122	Halsall <i>et al.</i> ³¹⁸
Zeppelinfjell, Spitsbergen 79°N12°E	Jul 1994–1996/Jan 1994–1996 ^g	13/13	180–330/2400–6600 ^h	EMEP ¹⁶
	Jul 2008–2010/Jan 2008–2010 ^g	13/13	46–75/260–1500	
Arctic Ocean (cruise) 65–80°N/150–180°W	Jul–Sep 2003	29	3700	Ding <i>et al.</i> ²⁶
Pallas, Finland 68°N/24°E	Jul–Oct 2006	1	600 ^{c,d}	Halse <i>et al.</i> ^{319e}

^a Sum of 24, including naphthalene. ^b Sum of 11. ^c Sum of 8. ^d Passive sampling, equivalent air volume inferred from performance reference compound or following Melymuk *et al.*³²³ ^e Passive sampling, equivalent air volume inferred from side-by-side sampling (Klánová, personal communication). ^f Sum of 9. ^g Total sampling time of 8 d per month. ^h Sum of 12.

may have been influenced by smoldering combustion in underground coal mines in the years 2005–2006 and 2009–2010 and in a mine dump 2006–2008 in the region (located 150–200 km north of the station; Kallenborn, Norwegian Institute for Air Research, personal communication).

Snow (firn) and ice core studies allow the assessment of temporal deposition trends. PAH deposition at high mountain sites in Europe and North America was traced back until 1700 and 1875, respectively.^{327,328} Levels in Europe increased with fossil fuel use (coal), significantly by the year 1900.^{327,329} PAH accumulates in high altitude soils and biota.^{330,331} Contamination in the Arctic is comparable with high mountain sites in Europe: PAH levels in the range of 36–660 ng L⁻¹ were reported from an ice core in the Canadian Archipelago attributed to the years 1963–1993³³² and 0.6–237 ng kg⁻¹ in an ice core from Greenland attributed to 1987–1990³³³ (Table 11). In the Arctic, meridional transports carrying PAHs are common in particular in the so-called haze season, *i.e.* spring-time.^{318,334}

No continuous monitoring of PAHs in Antarctic air has been conducted so far. Deposited amounts are lower than in the Arctic, it seems. Snow contamination by PAH on the Ekström shelf ice, 2002–2005, was completely explained by emissions within the region (ships, research stations), whereas no influence of intercontinental transport was suggested.³³⁵ Mostly due to the atmospheric dynamics of the region (polar vortex in the Southern Ocean), intercontinental atmospheric transport to the Antarctic continent is much less likely than in the Arctic. Levels in snow are 18–99 ng L⁻¹ in surface and deep snow for 1987–1991 on the Antarctic Peninsula (Terra Nova Bay, 74–75°S, 163–164°E³³⁶) and 26–197 ng L⁻¹ on the Ekström Shelf Ice in the Weddell Sea.³³⁵

PAH observations in the remote marine environment are very rare. Ship-based measurements are often unreliable because of blank problems related to fuel emissions from the ship.³³⁷ Measurements in seawater demonstrate the deposition of PAH from combustion sources to the open ocean.³³⁷

4.2. Model results

The transport of fluorene, phenanthrene, fluoranthene and benzo[*a*]pyrene from the UK to northern Russia has been studied using a Lagrangian box model.³³⁸ Regional (North America) and trans-Pacific transport pathways of PAHs were documented based on back trajectories.^{314,315,339–341} The atmospheric fate of PAHs was studied for the Los Angeles Basin using another multimedia model.³⁴² The potential influence of seasonally dependent environmental fate processes on the observed seasonality of concentrations of six PAHs in air was investigated for the UK using a multimedia box model (climatological atmospheric transports³⁴³). The regional and hemispheric transport and fate of benzo[*a*]pyrene was studied by atmospheric and multicompartment models with a focus on Europe^{35,344–348} and East Asia.^{349,350} Global distribution and fate of benzo[*a*]pyrene and 1–2 semivolatile PAHs were studied using an atmosphere general circulation model based multicompartment model^{36,351} and using an assimilated meteorology driven atmospheric model.³⁵² Due to lack of consolidated kinetic data the heterogeneous reactivity of PAHs (*i.e.* basically total reactivity in the

case of BaP) has been neglected so far in most model studies (assumption of zero degradability in the particulate phase).

Conventional atmospheric chemistry models neglect re-volatilisation from surfaces. As benzo[*a*]pyrene volatility is low (vapour pressure, $p = 7.3 \times 10^{-7}$ Pa at 298 K) its tendency to re-volatilise upon deposition to surfaces is limited, but still secondary emissions are estimated to contribute 9% to total emissions on the European continent (2007 emissions³⁴⁵). However, parameterisations of air–soil exchange which are in use in multicompartment models are not well validated. Multicompartment models so far have been neglecting PAH cycling in the cryosphere (sea ice, land ice, temporary snow cover).

The LRT potential of PAHs has been quantified.^{36,345,351,352} For benzo[*a*]pyrene, benzo[*b*]fluoranthene, benzo[*k*]fluoranthene, and indeno[1,2,3-*cd*]pyrene it was characterized to be regional and close to the one of polybrominated diphenylethers; *i.e.*, areas of high emission are still reflected in the deposition pattern (*e.g.* Gusev *et al.*³⁵). However, a large part of PAH transport occurs in the free troposphere, above the planetary boundary layer.^{350,351} Atmospheric half-lives of 3–5 ring PAHs are of the order of hours or days and do vary among model studies considerably. This reflects uncertain or unknown rate coefficients and the uncertainty with regard to air–soil exchange and processes which determine partitioning, as well as different methods of representation of these processes. Accordingly, model estimates for the total environmental residence time, which weighs all the compartmental residence times including in soil and water, where degradability is slow, do vary largely across models, namely 2–66 months for benzo[*a*]pyrene.³⁵³ Assuming similar degradability in the particulate and gas phase would lead to unrealistically short lifetimes.^{36,352} Hence, modelling provides evidence that degradation kinetics in the particulate phase must be significantly slower than in the gas phase.

Gas–particle partitioning strongly influences the atmospheric cycling, the total environmental fate including compartmental distributions and the LRT potential. Using a global model including a dynamic aerosol sub-model with the components OM and soot, the comparison of various gas–particle partitioning models suggests that for BaP, Ant and Flt both adsorption (preferentially to soot) and absorption (into OM) processes are contributing, and that total environmental burdens would be significantly lower assuming adsorption alone.³⁶ Most of the total environmental burden of BaP, Ant and Flt, namely 80–90%, is found in soil (including vegetation) after a few years of continuous emission into the model world atmosphere, while the atmospheric burdens of these substances account for 2–4% and oceanic burdens for 5–12%.

The effective meridional spreading (MS) and the spatial range (SR), measures for the tendency of distributions to extend across latitudes, of benzo[*a*]pyrene were quantified to 1000 km in air by a zonal multimedia fate model but and 10 000 km by the general circulation model based multicompartment model.³⁵³ Multimedia fate models underestimate the transport in the free troposphere, because temporal and spatial variability of wind is not represented and, hence, should underestimate the LRT potential. Multimedia models have been used in the context of

substance screening for POPs criteria and accounting for the multicompartimental cycling of PAH on various spatial scales (e.g. Hauck *et al.*³⁵⁴). The LRT potential of semivolatile PAHs is higher than that of benzo[*a*]pyrene and other 5-ring PAHs. It is enhanced by re-volatilisation. Re-volatilisation was estimated to account for >10% of the Ant and Flt global total emissions throughout the year.³⁶ The numerical values of the various LRT metrics in use strongly depend on the evaluation methodology and on model design.

As a number of PAHs, e.g. Pyr (Section 3), are converted in photochemistry into long-lived derivatives which undergo LRT^{1,355} and, eventually, are ubiquitously distributed, the assessment of PAHs' LRT potential and environmental risk should also include these derivatives. This aspect has been addressed in a box model study only.³⁴² As PAHs can enter the food web, integrated modelling is needed to assess the exposure of the ecosystems^{356,357} and, hence, uptake along food chains.

4.3. Recommendations for future work

The assessment of LRTP relies on remote observations and related model results, *i.e.* residence times and LRTP.

- Monitoring at remote sites of continents other than Europe and North America
- Validation of air–soil exchange parameterisations used in models in order to better account for the LRT potential being enhanced by secondary emission (re-volatilisation)
- Extend assessment of pollutants' LRT potential to derivatives of PAHs, in particular nitro-PAHs
- Include PAHs in food web modelling

5. Observational evidence for atmospheric reactions of PAHs in the atmosphere and formation of nitro- and oxy-derivatives

Levels of oxy- and nitro-PAH compounds in the ambient atmosphere, as well as the extent and nature of their direct emission sources have been widely investigated in the last 30 years. Tables S1 and S2 of the ESI† provide an overview of specific oxy- and nitro-PAH compounds respectively, identified in atmospheric samples, in studies of different combustion emission sources, and in product studies of laboratory reaction experiments. A review of environmental occurrence of oxy-PAH in the atmosphere was recently provided by Walgraeve *et al.*,³⁵⁸ which discussed the levels of these compounds in the atmosphere as well as their seasonality, partitioning behaviour and environmental fate. To our knowledge, no such review exists for nitro-PAH compounds. It should be noted that Tables S1 and S2 (ESI†) do not provide details of measured concentrations. For details of this, the reader is directed to the specific references provided or the review by Walgraeve *et al.*³⁵⁸

As discussed in detail in Section 3, laboratory experiments strongly indicate that both gas phase and heterogeneous reactions of PAHs can occur under environmentally relevant conditions in the atmosphere and can contribute to the observed levels of these compounds. The potentially mutagenic nature of

these reaction products is a concern from a public health perspective. However, the key questions remain: (i) to what extent do these reactions occur in the ambient atmosphere? (ii) What is the relative contribution of these reactions to observed levels of PAH-derivative compounds in the atmosphere compared with primary combustion emissions?

A number of ambient sampling approaches have been used to address these questions, in addition to or in conjunction with, laboratory reaction kinetics data. These sampling studies have been discussed previously.^{135,151} Here we provide an updated review of the current evidence of PAH reactivity and the formation of nitro- and oxy-PAHs in the atmosphere. This includes a discussion of differences in isomer profiles, temporal concentration variations, air-mass trajectory studies, theoretical calculations and the impact of source reduction measures.

5.1. Isomer profiles

As discussed in Section 3, the position on the aromatic ring where oxidation or nitration of a PAH compound occurs will determine the specific isomer of the oxy- or nitro-PAH derivative formed. Differences in the reaction mechanism during combustion processes and atmospheric photodecomposition processes can therefore produce distinct isomer product distributions. A key observation that can lead to an assessment of the relative contribution of direct emissions compared to atmospheric formation is therefore comparison of the isomer profiles of specific oxy- and nitro-PAH compounds observed in ambient sampling studies.

It has been noted that the key formation mechanism during combustion processes (see Table S2, ESI†) which produce nitro-PAH isomers is electrophilic nitration which creates products distinct from the isomers formed from radical-initiated reactions.³⁵⁹ For example, electrophilic nitration products for the reactions of pyrene and fluoranthene are noted as 1N-Pyr and 3N-Flt, while gas phase reactions under simulated atmospheric conditions are shown to form mainly 2N-Pyr and 2N-Flt.¹³⁵

As shown in Table S2 (ESI†), 2N-Flt and 2N-Pyr have been observed in ambient air samples in many different locations around the world but are generally not observed in direct emissions. 2N-Flt has been observed in a direct industrial emission from carbon electrode manufacture³⁶⁰ but this is not considered to be a major contributor on a large scale.³⁶¹ Zhu *et al.*³⁶² also identified 2N-Flt in diesel vehicle emissions but with an emission rate significantly lower (~0.2%) than that of 1N-Pyr. The isomers 2N-Flt and 2N-Pyr are therefore commonly used as markers for gas phase reaction products. Conversely, 1N-Pyr has been identified in a number of incomplete emission sources including diesel- and gasoline-fuelled vehicles, coal-fired power plants and aluminium smelting (see Table S2, ESI†) but is not observed as a gas phase reaction product, making it suitable as a marker for direct emissions.

2N-Flt and 2N-Pyr have been reported to be present in the atmosphere at higher levels than 1N-Pyr.^{360,361,363–372} The observation of 2N-Flt and 2N-Pyr associated with PM in ambient studies at these levels is therefore indicative of the occurrence of OH and/or NO₃ radical-induced reactions in the ambient atmosphere.

Furthermore, measured concentrations of 2N-Flt and 2N-Pyr are generally consistent with those predicted based on measured rate coefficients and formation yields.^{188,204,220,373}

Due to the distinct origins of the 2N-Flt and 1N-Pyr isomers, the 2N-Flt/1N-Pyr ratio can be used to assess the relative contribution from atmospheric reactions (OH and NO₃) compared to that of direct emissions.^{355,364,365,372,374,375} It is suggested that, assuming comparable levels of Flt and Pyr are present and the same dispersion and photolytic loss rates of the two isomers, a 2N-Flt/1N-Pyr ratio of >5 indicates the dominance of atmospheric reactions while a ratio of <5 indicates the dominance of direct combustion emissions.^{355,376}

It is indicated from laboratory studies that 2N-Flt can be formed *via* both OH and NO₃ initiated reactions,^{163,188} while 2N-Pyr will be formed only by OH-initiated reactions.^{163,220} The ratio between these two isomers can therefore be used as an indicator for the relative importance of OH (daytime) and NO₃ (night time) reaction pathways.^{364,374,377} A value of between 5 and 10 would indicate the dominance of the OH reaction pathway, while a value of above 100 would suggest the enhanced importance of NO₃-initiated reactions.³⁷⁶

Values of the 2N-Flt/1N-Pyr and 2N-Flt/2N-Pyr ratios measured in a number of different locations and conditions are presented in Tables 12 and 13 respectively. It can be seen that the values of both isomer ratios are highly variable between sites and under

different specific conditions. A 2N-Flt/1N-Pyr value of >5 is more commonly observed,^{355,364,372,378} indicative of atmospheric formation dominating in these studies. Locations with lower 2N-Flt/1N-Pyr ratios are generally areas with a dominant specific source, such as heavily trafficked areas^{367,374,379} and large urban centres.^{364,371,380} It has been noted that the 2N-Flt/1N-Pyr ratio is generally higher at suburban sites relative to their proximate urban site.^{364,374,375} This has been attributed to a longer exposure of the air mass to photochemical oxidants.³⁵⁵

The relatively low (<10) 2N-Flt/2N-Pyr ratios in most studies, are indicative of daytime OH-initiated reactions dominating over NO₃-initiated reactions, as indicated in laboratory reaction studies. Higher 2N-Flt/2N-Pyr ratios have been noted in rural areas compared to urban areas,^{134,376} suggesting increased importance of NO₃ reactions, which may be attributed to lack of fresh inputs of NO.^{364,376} It may be expected that, since NO₃ reactions occur almost entirely at night and OH reaction will occur during the day, there may be a distinct change in 2N-Flt/2N-Pyr over a diurnal cycle. However, Feilberg *et al.*³⁷⁴ noted that there was only a small increase in the 2N-Flt/2N-Pyr value during night time sampling (see Table 13). The authors therefore suggest that NO₃-initiated reactions play a minor role in the formation of 2N-Flt in these night time samples and attributed the presence of 2N-Flt to transport and dilution of OH-initiated reaction products.

Table 12 Summary of 2-NF/1-NP concentration ratios from ambient measurements

Location	Details	2NF/1NP	Ref.
Marseilles area, France	Urban and suburban	<5	Albinet <i>et al.</i> ¹³⁴
Marseilles area, France	Rural	>10	Albinet <i>et al.</i> ¹³⁴
Alpine Valley locations, South France	Mean summer value (one location)	>20	Albinet <i>et al.</i> ³⁷⁶
Alpine Valley locations, South France	Mean winter value (all locations)	<10	Albinet <i>et al.</i> ³⁷⁶
Baltimore, USA	Urban, winter	1–3	Bamford and Baker ³⁶⁴
Baltimore, USA	Urban, summer	6–24	Bamford and Baker ³⁶⁴
Baltimore, USA	Suburban	1–10	Bamford and Baker ³⁶⁴
Baltimore, USA	Urban	8–30	Bamford and Baker ³⁶⁴
Barcelona, Spain	Residential area	4	Bayona <i>et al.</i> ³⁸¹
Milan, Italy	Residential area	6.1	Cecinato <i>et al.</i> ³⁸²
Rome, Italy	Residential area	1.4	Cecinato <i>et al.</i> ³⁸²
Columbus, USA	Residential area	2.5	Chuang <i>et al.</i> ³⁸³
Rome, Italy	Urban	6.7	Ciccioli <i>et al.</i> ³⁵⁵
Milan, Italy	Urban	5.2	Ciccioli <i>et al.</i> ³⁵⁵
Naples, Italy	Residential area	1	Ciccioli <i>et al.</i> ³⁵⁵
Montelibretti, Italy	Suburban	9	Ciccioli <i>et al.</i> ³⁵⁵
Madrid, Spain	Suburban	7	Ciccioli <i>et al.</i> ³⁵⁵
C.Porziano, Italy	Suburban	12	Ciccioli <i>et al.</i> ³⁵⁵
Birmingham, UK	Roadway tunnel	2.5	Dimashki <i>et al.</i> ³⁷⁹
Ho Chi Minh City, Vietnam	Urban	21	Hien <i>et al.</i> ³⁶⁷
Ho Chi Minh City, Vietnam	Traffic site	2.7	Hien <i>et al.</i> ³⁶⁷
Copenhagen, Denmark	Traffic site	0.72	Feilberg <i>et al.</i> ³⁷⁴
Tokyo, Japan	Urban (summer)	8.9	Kojima <i>et al.</i> ³⁸⁷
Tokyo, Japan	Urban (winter)	5.4	Kojima <i>et al.</i> ³⁸⁷
Kanazawa, Japan	Urban	1.8	Murahashi and Hayakawa ³⁸⁰
Athens, Greece	Urban	2.1	Marino <i>et al.</i> ³⁷⁵
Riverside, USA	Urban background	8.75	Pitts <i>et al.</i> ³⁶⁹
Los Angeles, USA	Urban	3.9	Reisen and Arey ³⁷¹
Claremont, USA	Urban background	7.8	Ramdahl <i>et al.</i> ³⁷⁰
St Louis, USA	SRM (1648)	3.5	Ramdahl <i>et al.</i> ³⁷⁰
Washington DC, USA	SRM (1649)	3	Ramdahl <i>et al.</i> ³⁷⁰
Aurskog, Norway	Rural residential	3.7	Ramdahl <i>et al.</i> ³⁷⁰
Beijing, China	2008 Olympic Games	25–46	Wang <i>et al.</i> ³⁹⁷
Houston, USA	Suburban	4.2	Wilson <i>et al.</i> ³⁸⁴
Claremont, USA	Urban	21	Zielinska <i>et al.</i> ²¹⁸

Table 13 Summary of 2-NF/2-NP concentration ratios from ambient measurements

Location	Details	2NF/2NP	Ref.
Marseilles area, France	Rural	3.7	Albinet <i>et al.</i> ¹³⁴
Alpine Valley locations, France	Mean summer value (one location)	< 60	Albinet <i>et al.</i> ³⁷⁶
Alpine Valley locations, France	Mean winter value (all locations)	< 10	Albinet <i>et al.</i> ³⁷⁶
Baltimore, USA	Urban	5–57	Bamford and Baker ³⁶⁴
Baltimore, USA	Suburban	7–60	Bamford and Baker ³⁶⁴
Barcelona, Spain	Residential area	6	Bayona <i>et al.</i> ³⁸¹
Rome, Italy	Residential area	2.2	Cecinato <i>et al.</i> ³⁸²
Milan, Italy	Residential area	4.6	Cecinato <i>et al.</i> ³⁸²
Naples, Italy	Residential area	1.7	Ciccioli <i>et al.</i> ³⁵⁵
Montelibretti, Italy	Suburban	4.5	Ciccioli <i>et al.</i> ³⁵⁵
Madrid, Spain	Suburban	3.5	Ciccioli <i>et al.</i> ³⁵⁵
C.Porziano, Italy	Suburban	6	Ciccioli <i>et al.</i> ³⁵⁵
Copenhagen, Denmark	Urban and suburban	< 10	Feilberg <i>et al.</i> ³⁷⁴
Copenhagen, Denmark	Urban and suburban	14.2	Feilberg <i>et al.</i> ³⁷⁴
Athens, Greece	Urban	1.9	Marino <i>et al.</i> ³⁷⁵
Riverside, USA	Ambient POM	23.3	Pitts <i>et al.</i> ³⁶⁹
Claremont, USA	Urban background	35	Ramdahl <i>et al.</i> ³⁷⁰
St Louis, USA	SRM (1648)	9.3	Ramdahl <i>et al.</i> ³⁷⁰
Washington DC, USA	SRM (1649)	12	Ramdahl <i>et al.</i> ³⁷⁰
Aurskog, Norway	Rural residential	3.3	Ramdahl <i>et al.</i> ³⁷⁰
Los Angeles and Riverside, USA	Winter	16 ± 7	Reisen and Arey ³⁷¹
Los Angeles and Riverside, USA	Summer	>35	Reisen and Arey ³⁷¹
Finokalia, Crete	Mean value from a diurnal study, marine background location	3.5	Tsapakis and Stephanou ³⁷⁷
Beijing, China	2008 Olympic Games	3.4–4.8	Wang <i>et al.</i> ³⁹⁷

It should be noted that the 2N-Flt/1N-Pyr and 2N-Flt/2-NP ratios reflect simply the relative levels of these isomers in the atmosphere and while they can be used as a reasonable marker for OH and/or NO₃ initiated reactions in the atmosphere, it should be considered that these ratios can be influenced by a number of other factors which may alter the ratio. For example, the relative input and removal rates of particulate matter may be of importance. Similarly, meteorological factors such as changes in mixing height and levels of solar irradiation can affect the levels of these isomers by influencing the degree of their dispersion and photolytic loss respectively.³⁶⁶ The use of these ratios should therefore be used with caution when assessing the relative importance of OH and/or NO₃-induced PAH reactivity, particularly over a relatively short sampling period. (Tables 12 and 13)^{377,379–384}

The occurrence of radical-induced reactions has also been assessed by comparing the isomer distribution of naphthalene derivatives from ambient air samples with that of experimental reaction products. For example, it has been noted in several studies that the isomer profiles of N-Naps, MN-Naps and EN-Naps + DMN-Naps observed in ambient air are reasonably consistent with their formation from radical induced atmospheric reactions.^{180,218,361,371,378,385}

As discussed by Atkinson and Arey (2007), MN-Naps profiles from a laboratory study of the reaction of 1- and 2M-Naps with OH radicals^{180,371} strongly resemble that of an ambient air sample taken during the morning hours in Mexico City, suggesting that atmospheric reactions have a strong influence of the levels and isomer profile of these compounds in urban air. Similarly, night-time air samples provide evidence for the occurrence of NO₃-induced reactions in the atmosphere. The profile of MN-Naps from a chamber reaction of NO₃ radicals produces a different profile to that of the OH-induced reaction

and closely resembles the profile of a night time air sample from a receptor site downwind of urban Los Angeles (see Atkinson and Arey¹⁵¹ and Reisen and Arey³⁷¹ for a more detailed discussion of these observations).

Based on similar observations of alkylnitronaphthalene isomer profiles, Wang *et al.*²¹⁹ suggested specific isomers or isomer ratios as indicators for the occurrence of OH and/or NO₃ reactions. For example, it was suggested that the isomers 2E1N-Nap, 1,7DM8N-Nap, 2,7DM1N-Nap, 1,6DM5N-Nap, 1,3DM4N-Nap and 1,2DM4N-Nap are indicative of NO₃-induced formation, while low levels of these compounds in comparison with 1,7DM5N-Nap suggest the dominance of OH-induced reactions.²¹⁹ Furthermore, relatively high ratios of 2,7DM4N-Nap/1,7DM5N-Nap and of 2M4N-Nap/1M5N-Nap are suggested to be indicative of the predominance of NO₃ reaction chemistry relative to OH reactions.²¹⁹

Reisen and Arey³⁷¹ measured concentrations of PAHs and nitro-PAHs in California and noted that the formation ratios (nitro-PAH/PAH) for nitronaphthalenes and alkyl-substituted nitronaphthalenes followed the order N-Naps < MNNs < DMN-Naps + EN-Naps. These formation ratios were suggested to be due to atmospheric reactions of PAHs with OH radicals and were shown to be a factor ~ 10 higher in summer compared to winter, which was attributed to enhanced photochemical activity during summertime. The observed trend in nitro-PAH/PAH ratios for these compounds is consistent with the observed 'activating effect' of alkyl groups on OH and NO₃ reactivity as described by Phousongphouang and Arey^{171,214} (see Section 3.2.1.1.4), providing further evidence for the occurrence of atmospheric reactions, similar to those observed in laboratory reactions.

As discussed by Kameda,³⁸⁶ the gas phase reaction of triphenylene with OH and NO₃ has been shown to form the nitro-PAH derivatives 1-nitrotriphenylene (1N-TPh) and 2-nitrophenylene (2N-TPh). These reactions preferentially form 2N-TPh. Kameda³⁸⁶ noted that

the ratio of 2N-TPh/1N-TPh observed in samples of airborne particles was >1.55 , similar to that observed from gas phase OH- and NO₃-initiated reactions (2N-TPh/1N-TPh = 1.22 to >1.5). The 2N-TPh/1N-TPh ratio observed in diesel exhaust particulate samples was shown to be much lower (2N-TPh/1N-TPh = 0.37), suggesting that radical-induced reactions contribute to the observed levels of these compounds, particularly 2N-TPh and that direct combustion emissions predominantly contribute to the levels of 1N-TPh.

Studies have also used the ratio of oxy- or nitro-PAH to their 'parent' PAH to assess the importance of atmospheric reactions on the levels of these compounds. For example, in the review by Walgraeve *et al.*,³⁵⁸ the ratio of oxy-PAH/'parent' PAH was assessed for ambient sampling studies in the literature. It was shown that during winter, 50% of these ratios was between 0.006 and 0.16 (19 studies³⁵⁸). In summer this ratio was reported to be about 20 times higher than the winter samples, with 50% of ratios between 0.54 and 3.6 (48 studies³⁵⁸). This suggests the increased importance of photochemical activity during summer leading to higher levels of oxy-PAHs, thus providing evidence for the importance of atmospheric reactions.

5.2. Temporal concentration patterns

Temporal (daily, seasonal and diurnal) variations in the concentrations of atmospheric pollutants will be influenced by a number of factors including the variation in specific combustion sources, mixing height level, photolytic loss rates and gas-particle partitioning. Temporal changes in the levels of oxy- and nitro-PAH compounds have been investigated to assess the impacts of these variations and can be viewed with reference to the relative contribution of photochemical reactions to the observed levels of these compounds.

Kojima *et al.*³⁸⁷ investigated daily and seasonal variations in concentrations of nitro- and oxy-PAHs associated with airborne PM in Tokyo, Japan. They observed that summer/winter concentration ratios were higher for 2N-Flt (0.36) than for 1N-Pyr (0.19–0.27), consistent with 2N-Flt being a secondary reaction product and 1N-Pyr being a primary emission. In contrast, the summer/winter concentration ratios for oxy-PAHs were the same as or lower than that of PAHs, contrary to the expected lower concentrations resulting from the contribution of secondary reactions. The authors suggested that this discrepancy may result from the increased partitioning of semi-volatile oxy-PAHs from the particulate phase to the gas phase in summer resulting in relatively low concentrations associated with the PM samples, as previously observed by Albinet *et al.*³⁸⁸ Kojima *et al.*³⁸⁷ observed strong correlations between oxy-PAH concentrations and primary emissions such as PAHs and CO in winter, while a much weaker correlation was noted in summer. This may indicate the relative dominance of direct emissions of oxy-PAHs in winter and a greater importance of atmospheric reactivity in summer.

Reisen and Arey³⁷¹ noted that levels of PAHs in the Los Angeles area were elevated in winter relative to summer which can be attributed to lower mixing heights resulting in lower dispersion rates^{364,371} or less efficient photolytic loss.^{246,389} In contrast, nitro-PAHs at a site ~60 km downwind of Los Angeles

had concentrations generally higher in summer, indicating the greater importance of photochemical reactions in summer, when levels of atmospheric oxidants (OH, NO₃, O₃) are higher,^{178,389} therefore indicating enhanced secondary input of these compounds during summer. A similar observation was made by Eiguren-Fernandez *et al.*³⁹⁰ for PAH-quinone compounds.

2N-Flt/1N-Pyr have been shown to be lower during winter, especially in urban areas,³⁷⁶ indicating that atmospheric reactions are less important in winter relative to direct emissions. Similarly, in winter, the 2N-Flt/2N-Pyr ratio has been shown to be higher than in summer,^{364,371,374} suggesting that NO₃-initiated reaction may play a more important role during winter. These two observations may be explained by lower intensity of photochemical activity in winter, leading to lower levels of OH radicals in the atmosphere during winter. Winter conditions may however be more favourable to the formation of the NO₃ radical, and to lesser photolytic losses of this species.

Variations in pollutant concentration over a 24 h period can be indicative of the nature and extent of their input and output mechanisms to the atmosphere. This is particularly relevant to the case of PAH-derivative compounds. While the diurnal variation of parent PAHs is likely to reflect direct source emission signals as modified by meteorology, the levels of nitro- and oxy-PAHs, which may result in part from atmospheric reactions, may display different patterns, possibly reflecting the relative balance of primary and secondary inputs. Furthermore, the nature of the atmospheric reactions of PAHs in the atmosphere, with OH reactions occurring during daylight and NO₃ reactions occurring at night, will lead to potentially distinct diurnal patterns in the levels of nitro- and oxy-PAHs.

Since OH radicals are only present during daylight hours night time/daytime concentration ratios provide an assessment of unreacted *versus* reacted air masses, and the relative values of these ratios should correspond to the relative reactivity of the individual compounds.

Arey *et al.*³⁹¹ measured 12 h 'daytime' and 'night time' concentrations of LMW PAHs during a photochemical air pollution episode in the Los Angeles Basin. The night time/daytime concentration ratios derived were shown to correlate with the OH radical reaction rate coefficients, for example with the largest ratios being observed for PAHs that are most reactive towards OH radicals. This indicates the occurrence of OH-induced reactions occurring in the atmosphere at similar rates to those predicted in laboratory studies. Similarly, Phousongphouang and Arey¹⁷¹ compared night time/daytime concentration ratios of alkyl-nitronaphthalenes measured in ambient samples to their calculated rate coefficients for the gas phase reaction of OH radicals with alkyl-naphthalenes and observed a reasonable linear correlation with OH rate coefficients, supporting the expectation that OH reactions are dominating the chemistry of these compounds in the atmosphere.

Reisen and Arey³⁷¹ investigated diurnal variations of PAH and nitro-PAH levels to assess atmospheric reactivity by sampling over four time periods (morning, day, evening and night) in California, USA. It was shown that, in summer, PAH concentrations clearly decrease during the day and evening periods.

This could be indicative of PAH loss due to atmospheric reactions with OH and the more reactive compounds were shown to decrease the most.³⁷¹ This 'reactive loss' was shown to be more significant in summer compared to winter.

Tsapakis and Stephanou³⁷⁷ studied the diurnal pattern of nitro-PAH as well as PAH, OH radicals and O₃ at a background site in the eastern Mediterranean. The authors suggested that the diurnal pattern of PAHs was dominated by the input of local sources (volatilization from the sea in this case). 2N-Pyr and 2N-Flt were the most abundant nitro-PAHs identified and displayed a well-defined diurnal pattern, with a concentration maximum occurring at midday followed by a rapid decrease. This pattern closely matched the diurnal variation of OH radicals, suggesting that OH radical-initiated reactions are the dominant processes controlling the levels of these nitro-PAHs in the atmosphere. The subsequent rapid decrease in nitro-PAH concentration after the midday maximum was attributed to photolytic loss of nitro-PAH compounds.

Hien *et al.*³⁶⁷ measured concentrations of 1N-Pyr and 2N-Flt in Ho Chi Minh City, Vietnam during both daytime and night time hours in residential and high traffic locations. They observed lower levels of both compounds during daylight hours in the residential area. The higher levels of 1N-Pyr during the night were attributed to photodecomposition of 1N-Pyr during the day, while higher levels of 2N-Flt during the night were suggested to be due to atmospheric formation from NO₃-induced reactions. In contrast, levels of 2N-Flt at the high traffic location were higher during the daytime, suggesting the importance of OH-radical induced reactions at these locations.

5.3. Quantitative predictions and atmospheric trajectory studies

A number of studies have attempted to make quantitative assessments of PAH reactivity in the atmosphere. These investigations can simply involve relating observed concentrations from sampling studies to laboratory reaction kinetics data and either comparing observed levels to those predicted theoretically or using measured concentration values to calculate a formation yield that can be related to observed laboratory studies. A number of studies have attempted to quantify the relative level of PAH-derivative compounds in the atmosphere originating from secondary input from photochemical reactions. This has been attempted by carrying out a sampling study along a specific air-mass trajectory with increasing distance from the source region towards a receptor site.

Arey *et al.*³⁹² initially predicted concentrations of 1/2N-Nap, 2N-Flt and 2N-Pyr, based on calculations using laboratory-derived OH reaction rate coefficients and nitro-PAH formation yields, measured parent PAH concentrations and incorporating photolytic loss of nitro-PAHs.¹⁸⁰ It was shown that there was a remarkably good agreement between the calculated values and those measured in the ambient atmosphere, suggesting the dominant input of these compounds from atmospheric reactions of parent PAHs compared to direct emissions.

More recently, as discussed by Atkinson and Arey¹⁵¹ (and references therein), based on observed levels of 1- and 2-NNs in

ambient sampling studies, the formation yields of the OH + naphthalene reaction were estimated using the [nitro-nap]/[nap] ratios, OH reaction rate coefficient and photolysis rate coefficient values from the literature and estimated average OH radical concentrations. It was shown that the calculated values for the formation yields of 1N-Nap and 2N-Nap were reasonably consistent with experimental data,^{182,363} providing further evidence that PAH reactions with OH radicals and further reaction with NO₂ occur significantly in the ambient atmosphere.

The predicted PQu formation rates calculated by Wang *et al.*²²¹ (see Section 3) are significant, compared to observed levels in ambient sampling studies (see Table S1, ESI[†]), suggesting that atmospheric formation may be a significant contributor to the observed levels in the atmosphere. Eiguren-Fernandez *et al.*³⁹⁰ investigated the changes in PQ concentration with decreasing distance from the highly polluted Los Angeles basin, following a prevailing wind trajectory towards a more remote downwind receptor site. A significant increase in PQu concentration was observed as the air parcel moved along this trajectory, suggesting the occurrence of photochemical reactions contributing to the observed levels of PQu.

Eiguren-Fernandez *et al.*³⁹⁰ investigated the correlations of PQu concentrations to those of phenanthrene in the particulate phase, Phe(p) and the vapour phase, Phe(v). It was shown that in the source region, there was a good correlation between PQu and Phe(p), indicative of direct emissions dominating, while at the downwind locations, there was a good correlation between PQu and Phe(v), suggesting the dominance of atmospheric reactions as the source of PQu in these locations. Eiguren-Fernandez *et al.*³⁹⁰ also estimated the percentage contribution of atmospheric transformations to the observed levels of PQu at the downwind site, based upon ratios of benzo[ghi]perylene (a marker for vehicular emissions) to PQu at the source region and downwind sites (see Eiguren-Fernandez *et al.*³⁹⁰ for details). It was estimated that ~90% of PQu measured in the Los Angeles basin was of secondary origin from photochemical reactions during atmospheric transport.

Kojima *et al.*³⁹³ measured the concentrations of PAHs, oxy-PAHs and nitro-PAHs at a road-side location (A) and at sites 90 m (B) and 170 m (C) away from the roadside. Concentrations of PAHs and 1N-Pyr followed the order A > B > C (A:C ratio = 0.20–0.53 and 0.20–0.26 respectively) indicating that direct input from the traffic source and dilution with increasing distance from the road controlled their concentration. Oxy-PAH concentrations also decreased with increasing distance from the roadside, although not as significantly as PAHs and 1N-Pyr (A:C ratio = 0.42–0.69). This would suggest that their atmospheric concentrations are influenced by traffic emissions but also by an additional background source affecting all sites similarly.

In contrast, the concentration of 2N-Flt was approximately the same at all three sites (A:C ratio = 0.88–1.03), indicating the occurrence of input from atmospheric reactions. The authors attempted to quantify the contribution of atmospheric formation to the levels of oxy-PAHs, modifying the calculations used by Eiguren-Fernandez *et al.*³⁹⁰ and based on the oxy-PAH and PAH concentrations at sites A and C. Their results indicated

that percentage contribution of atmospheric formation at the downwind site ranged from 9% to 72% depending on the compound and season, with a higher secondary input contribution generally noted in summer. It was also noted that the percentage of secondary input followed the order 1,8-NtA > 9-Flr and 9,10-Aqu > BA with values of >50% for 1,8-NtA, 9-Flr and 9,10-Aqu in both summer and winter.

These findings therefore indicate that a considerable fraction of oxy-PAHs detected in atmospheric particulate matter in urban areas will originate from atmospheric reactions of primary PAH. However, the authors note that these calculations are sometimes inadequate and may not be applicable under all environmental conditions. The amount of oxy-PAH in the atmosphere and the relative importance of atmospheric reactions *versus* primary emissions will also be dependent on a number of other factors including the emission rates of PAH and oxy-PAH, meteorological factors such as mixing layer height, wind speed and direction, which influence dilution levels, and the degree of gas-particle partitioning of both PAH and oxy-PAH.³⁹³

5.4. Emission reduction related studies

Urban air quality has been and will increasingly be deteriorating on most continents³⁹⁴ and related limit or recommended values are often exceeded, even today. For example, in Europe, this has involved setting country-specific emission mitigation targets such as those set by the National Emission Ceiling Directive (2001/81/EC).³⁹⁵

The implementation of strict measures in certain areas to reduce emissions, either on a long-term or short-term period, provides an opportunity to investigate the impacts of these measures on the levels of oxy- and nitro-PAH compounds and to investigate the relative contribution of primary and secondary sources. To date, however, few investigations have addressed this area of research.

Kojima *et al.*³⁸⁷ investigated the impact of government regulation regarding emissions from diesel vehicles on the levels of PAHs and nitro-PAHs, in Tokyo. The emission reduction measures in this case involved a 2003 regulation, enforcing the use of lower-emission vehicles or installation of PM reduction devices. Comparison of ambient PM samples taken in 1997–1998 and 2006–2007 revealed a significant decrease in the average annual concentrations of PAHs and 1-NP over this period, attributed to the reduction in primary vehicular emissions. In contrast, no decrease in concentration was noted for 2N-Flt and 2N-Pyr suggesting the importance of other inputs such as atmospheric reactions to influence the levels of these compounds.

During the 2008 Olympic Games, efforts were made to improve the air quality in Beijing by implementing restrictions on direct emissions from sources such as vehicular traffic, construction sites and coal-fired power plants.^{396,397} This resulted in significant decrease in the levels of primary traffic-related pollutants such as CO and NO₂ during this source reduction period.³⁹⁶ Reductions in concentrations were also noted for a number of nitro-PAH and oxy-PAH compounds ranging from 15.1% to 56.6% and 24.8% to 46.6% respectively.³⁹⁷

The mean 2N-Flt/1N-Pyr value of 25–46 measured over the source control and non-source control periods indicates the dominance of photochemical reactions over this time. However, this ratio was shown to decrease during the source control period, which was attributed to the prevailing meteorological conditions and the decreased traffic emissions during this time.³⁹⁷ The relatively low 2N-Flt/2N-Pyr ratio of 3.4–4.8 calculated over this period suggested the dominance of daytime OH-initiated reactions.

It is evident that air monitoring studies in urban areas carried out before and after the implementation of pollution control measures have the potential to generate interesting observations regarding the differences in primary and secondary inputs of PAH-derivative compounds. More extensive work in this area in different urban locations and for a wider range of compounds is clearly needed.

5.5. Conclusions and recommendations for future work

Field data suggest that gas phase reactions of PAHs with OH (and NO₃) radicals can contribute significantly to the observed levels of nitro- and oxy-PAHs – although for many atmospherically relevant parent PAHs this chemistry is still unknown. The mechanisms should be explored in order to identify precursors and sources, needed for emission mitigation.

(i) *Compounds with multiple origins.* Assessment of the relative primary and secondary inputs of PAH-derivatives in a large number of the studies discussed in this section has focussed on marker compounds for either direct emissions (*e.g.* 1N-Pyr) or atmospheric reactions (*e.g.* 2N-Flt, 2N-Pyr). While this provides a tentative analysis for the relative contributions of OH- and/or NO₃-initiated reactions, this approach will lead to definitive conclusion regarding source apportionment for only these specific compounds.

As can be seen in Tables S1 and S2 (ESI[†]), a large number of oxy- and nitro-PAH compounds identified in ambient air can result from both direct emissions and atmospheric reactions. This makes the task of assessing and potentially of quantifying the relative importance of their sources considerably more difficult. Among the wide range of PAH-derivative compounds in the atmosphere, sources in that sense have only been investigated for PQu and NM-Naps.

(ii) *The role of heterogeneous PAH reactions.* Heterogeneous PAH reactions play an important role in determining the levels of PAH derivative compounds in the atmosphere, but so far the kinetics data and reaction products have been elucidated with the focus on gas phase processes. Heterogeneous PAH reactions are influenced by a wide range of factors in addition to those associated with gas phase systems. Furthermore, partitioning between gas and particulate phases of both parent PAHs and derivatives is subject to ambient conditions (temperature, PM composition) and its determination is subject to artefacts (due to heterogeneous PAH chemistry on sampling media; Section 3.1.1).

To assess exposure towards these pollutants, there is also a need to study the chemical sinks, *i.e.* kinetics and reaction products of the PAH derivatives in the atmosphere.

Abbreviations

Ace	Acenaphthene
Acy	Acenaphthylene
AcP	Acepyrene
Ant	Anthracene
AQu	Anthraquinone
BaA	Benzo[a]anthracene
BaP	Benzo[a]pyrene
BbF	Benzo[b]fluoranthene
BeP	Benzo[e]pyrene
BgF	Benzo[ghi]fluoranthene
BgP	Benzo[ghi]perylene
BkF	Benzo[k]fluoranthene
Chr	Chrysene
Cor	Coronene
DBA	Dibenzo[a,h]anthracene
DBaP	Dibenzo[a,e]pyrene
DBaP	Dibenzo[a,h]pyrene
DBaP	Dibenzo[a,l]pyrene
DM-Nap	Dimethylnaphthalene
DM-Phe	Dimethylphenanthrene
E-Nap	Ethyl naphthalene
Flt	Fluoranthene
Fln	Fluorene
Flr	Fluorenone
FCA	Formycinmaldehyde
IPy	Indeno[1,2,3-cd]pyrene
M-Nap	Methylnaphthalene
M-Phe	Methylphenanthrene
NQu	Naphthoquinone
NtA	Naphthalic anhydride
N-Nap	Nitronaphthalene
Phe	Phenanthrene
PQu	Phenanthrequinone
Per	Perylene
Pyr	Pyrene
TPh	Triphenylene

References

- 1 B. J. Finlayson-Pitts and J. N. Pitts, *Chemistry of the upper and lower atmosphere*, Academic Press, San Diego, USA, 2000.
- 2 J. M. Delgado-Saborit, C. Stark and R. M. Harrison, *Environ. Int.*, 2011, **37**, 383–392.
- 3 M. J. Wenborn, P. J. Coleman, N. R. Passant, E. Lymberidi, J. Sully and R. A. Weir, *AEAT-3512/REMC/20459131/Issue 1*, 1999.
- 4 WHO, *Air quality guidelines, Global update 2005*, World Health Organisation, Bonn, 2006.
- 5 European Union, *Directive 2004/107/EC of the European Parliament and of the Council of 15 December 2004 23 relating to arsenic, cadmium, mercury, nickel and polycyclic aromatic hydrocarbons in ambient air*, Official Journal 24 of the European Union, L23, 3–16, 2005.
- 6 EPAQS, *A recommendation for a united kingdom air quality standard for polycyclic aromatic hydrocarbons*, Department of the Environment, 1999.
- 7 R. M. Harrison, D. J. T. Smith and A. J. Kibble, *Occup. Environ. Med.*, 2004, **61**, 799–805.
- 8 A. J. Cohen, *Environ. Health Perspect.*, 2000, **108**, 743–750.
- 9 K. B. Okona-Mensah, J. Battershill, A. Boobis and R. Fielder, *Food Chem. Toxicol.*, 2005, **43**, 1103–1116.
- 10 G. Lammel, J. Novák, L. Landlová, A. Dvorská, J. Klánová, P. Čupr, J. Kohoutek, E. Reimer and L. Škrdlíková, in *Urban Airborne Particulate Matter: Origins, Chemistry, Fate and Health Impacts*, ed. F. Zereini and C. L. S. Wiseman, Springer, Berlin, 2010, pp. 39–62.
- 11 T. Enya, H. Suzuki, T. Watanabe, T. Hirayama and Y. Hisamatsu, *Environ. Sci. Technol.*, 1997, **31**, 2772–2776.
- 12 T. Nielsen, A. Feilberg and M. L. Binderup, *Environ. Sci. Pollut. Res.*, 1999, **6**, 133–137.
- 13 A. Feilberg, T. Nielsen, M. L. Binderup, H. Skov and M. W. B. Poulsen, *Atmos. Environ.*, 2002, **36**, 4617–4625.
- 14 UNECE, *United Nations Economic Commission for Europe, Convention on Long-range Transboundary Air Pollution, Protocol on Persistent Organic Pollutants*, Aarhus, Denmark, 1998.
- 15 OSPAR, *Oslo Paris Commission, Quality status report 2000*, Chemistry, London, ch. 4, 2000, p. 108.
- 16 EMEP, *Measurement data online, European Monitoring and Evaluation Programme*, Geneva, 43, <http://www.nilu.no/projects/ccc/emepdata.html>, 2011.
- 17 IADN, *Integrated Atmospheric Deposition Network*, http://www.ec.gc.ca/rs-23_mn/default.asp?lang=En&n=BFE9D3A3-1, accessed 2.2.2012, 2011.
- 18 USEPA, *2007 US Environmental Protection Agency. EPI Suite v4.0, Exposure assessment tools and models*, <http://www.epa.gov/opt/exposure/pubs/episuite>, 2007.
- 19 Y.-G. Ma, Y. D. Lei, H. Xiao, F. Wania and W.-H. Wang, *J. Chem. Eng. Data*, 2010, **55**, 819–825.
- 20 T. E. M. ten Hulscher, L. E. van der Velde and W. A. Bruggeman, *Environ. Toxicol. Chem.*, 1992, **11**, 1595–1603.
- 21 Y. D. Lei, R. Chankalal, A. Chan and F. Wania, *J. Chem. Eng. Data*, 2002, **47**, 801–806.
- 22 M. Odabasi, E. Cetin and A. Sofuoglu, *Atmos. Environ.*, 2006, **40**, 6615–6625.
- 23 M. F. Simcik, T. P. Franz, H. X. Zhang and S. J. Eisenreich, *Environ. Sci. Technol.*, 1998, **32**, 251–257.
- 24 M. Mandalakis, M. Tsapakis, A. Tsoga and E. G. Stephanou, *Atmos. Environ.*, 2002, **36**, 4023–4035.
- 25 M. Tsapakis and E. G. Stephanou, *Environ. Pollut.*, 2005, **133**, 147–156.
- 26 X. Ding, X.-M. Wang, Z.-Q. Xie, C.-H. Xiang, B.-X. Mai, L.-G. Sun, M. Zheng, G.-Y. Sheng, J.-M. Fu and U. Poeschl, *Atmos. Environ.*, 2007, **41**, 2061–2072.
- 27 J. He and R. Balasubramanian, *Atmos. Environ.*, 2009, **43**, 4375–4383.
- 28 E. Demircioglu, A. Sofuoglu and M. Odabasi, *J. Hazard. Mater.*, 2011, **186**, 328–335.

- 29 G. Spindler, T. Gnauk, A. Gruener, Y. Iinuma, K. Mueller, S. Scheinhardt and H. Herrmann, *J. Atmos. Chem.*, 2012, **69**, 127–157.
- 30 J. Paasivirta, S. Sinkkonen, P. Mikkelsen, T. Rantio and F. Wania, *Chemosphere*, 1999, **39**, 811–832.
- 31 Y. Zhang and S. Tao, *Atmos. Environ.*, 2009, **43**, 812–819.
- 32 T. Gocht, J. A. C. Barth, M. Epp, M. Jochmann, M. Blessing, T. C. Schmidt and P. Grathwohl, *Appl. Geochem.*, 2007, **22**, 2652–2663.
- 33 EEA, *EMEP/EEA air pollutant emission inventory guidebook*, European Environment Agency, Technical 28 Report No. 9/2009, Copenhagen, 2009.
- 34 S. S. Xu, W. X. Liu and S. Tao, *Environ. Sci. Technol.*, 2006, **40**, 702–708.
- 35 A. Gusev, S. Dutchak, O. Rozovskaya, V. Shatalov, V. Sokoviyh, N. Vulykh, W. Aas and K. Breivik, *Persistent Organic Pollutants in the Environment. EMEP Status Report 3/2011*, Meteorological Synthesizing 38 Centre - East, Moscow, Russian Federation, 2011, p. 108.
- 36 G. Lammel, A. M. Sehili, T. C. Bond, J. Feichter and H. Grassl, *Chemosphere*, 2009, **76**, 98–106.
- 37 J. Münch, F. Axenfeld, *Historic emission data base of selected organic pollutants (POPs) in Europe (1970–42 95)-PCDD/F, B(a)P, PCB, g-HCH, HCB, DDE/DDT. Report to the European Commission (DG XII)*, 43 Environment and Climate Project No. ENV-CT96-0214, 1999.
- 38 I. Bewersdorff, A. Aulinger, V. Matthias and M. Quante, *Meteorol. Z.*, 2009, **18**, 41–53.
- 39 USEPA, U.S. Environmental Protection Agency. *Locating and estimating air emissions from sources of 2 polycyclic organic matter. EPA-454/R-98-014. Office of Air Quality Planning and Standards*, Research Triangle 3 Park, NC, USA, 1998.
- 40 J. M. Pacyna, K. Breivik, J. Munch and J. Fudala, *Atmos. Environ.*, 2003, **37**, S119–S131.
- 41 H. Denier van der Gon, M. van het Bolscher, A. Visschedijk and P. Zandveld, *Atmos. Environ.*, 2007, **41**, 9245–9261.
- 42 A. L. Robinson, R. Subramanian, N. M. Donahue, A. Bernardo-Bricker and W. F. Rogge, *Environ. Sci. Technol.*, 2006, **40**, 7803–7810.
- 43 E. Galarneau, *Atmos. Environ.*, 2008, **42**, 8139–8149.
- 44 M. Blumer, *Sci. Am.*, 1976, **234**, 35–45.
- 45 B. R. T. Simoneit, *Int. J. Environ. Anal. Chem.*, 1985, **22**, 203–233.
- 46 M. B. Yunker, R. W. Macdonald, R. Vingarzan, R. H. Mitchell, D. Goyette and S. Sylvestre, *Org. Geochem.*, 2002, **33**, 489–515.
- 47 A. Dvorská, G. Lammel and J. Klánová, *Atmos. Environ.*, 2011, **45**, 420–427.
- 48 M. Tobiszewski and J. Namiesnik, *Environ. Pollut.*, 2012, **162**, 110–119.
- 49 A. Dvorská, K. Komprdová, G. Lammel, J. Klánová and H. Plachá, *Atmos. Environ.*, 2012, **46**, 147–154.
- 50 B. R. T. Simoneit, J. J. Schauer, C. G. Nolte, D. R. Oros, V. O. Elias, M. P. Fraser, W. F. Rogge and G. R. Cass, *Atmos. Environ.*, 1999, **33**, 173–182.
- 51 J. J. Schauer, W. F. Rogge, L. M. Hildemann, M. A. Mazurek, G. R. Cass and B. R. T. Simoneit, *Atmos. Environ.*, 1996, **30**, 3837–3855.
- 52 D. R. Oros and B. R. T. Simoneit, *Fuel*, 2000, **79**, 515–536.
- 53 T. Nielsen, *Atmos. Environ.*, 1988, **22**, 2249–2254.
- 54 F. Böhme, K. Welsch-Pausch and M. S. McLachlan, *Environ. Sci. Technol.*, 1999, **33**, 1805–1813.
- 55 M. S. McLachlan, *Environ. Sci. Technol.*, 1999, **33**, 1799–1804.
- 56 S. D. Choi, R. M. Staebler, H. Li, Y. Su, B. Gevaio, T. Harner and F. Wania, *Atmos. Chem. Phys.*, 2008, **8**, 4105–4113.
- 57 A. D. St-Amand, P. M. Mayer and J. M. Blais, *Atmos. Environ.*, 2009, **43**, 4283–4288.
- 58 I. T. Cousins and K. C. Jones, *Environ. Pollut.*, 1998, **102**, 105–118.
- 59 W. Wilcke, *J. Plant Nutr. Soil Sci.*, 2000, **163**, 229–248.
- 60 A. J. Sweetman, M. Dalla Valle, K. Prevedouros and K. C. Jones, *Chemosphere*, 2005, **60**, 959–972.
- 61 G. L. Daly, Y. D. Lei, L. E. Castillo, D. C. G. Muir and F. Wania, *Atmos. Environ.*, 2007, **41**, 7339–7350.
- 62 R. Wang, S. Tao, B. Wang, Y. Yang, C. Lang, Y. Zhang, J. Hu, J. Ma and H. Hung, *Environ. Sci. Technol.*, 2010, **44**, 1017–1022.
- 63 R. M. Dickhut and K. E. Gustafson, *Environ. Sci. Technol.*, 1995, **29**, 1518–1525.
- 64 M. Odabasi, A. Sofuoglu, N. Vardar, Y. Tasdemir and T. M. Holsen, *Environ. Sci. Technol.*, 1999, **33**, 426–434.
- 65 R. Balasubramanian and J. He, *Fate and transfer of persistent organic pollutants in a multi-media environment*, in *Urban Airborne Particulate Matter: Origins, Chemistry, Fate and Health Impacts*, ed. F. Zereini and C. L. S. Wiseman, Springer, Heidelberg, 2010, pp. 277–307.
- 66 J. Komprda, K. Kubosová, A. Dvorská, M. Scheringer, J. Klánová and I. Holoubek, *J. Environ. Monit.*, 2009, **11**, 269–276.
- 67 D. M. Farmer, C. L. McNeil and B. D. Johnson, *Nature*, 1993, **361**, 620–623.
- 68 J. DeWulf, H. Van Langenhove and B. Heireman, *Water Res.*, 1998, **32**, 2106–2112.
- 69 L. Lim, O. Wurl, S. Karuppiyah and J. P. Obbard, *Mar. Pollut. Bull.*, 2007, **54**, 1212–1219.
- 70 C. Guitart, N. Garcia-Flor, J. C. Miquel, S. W. Fowler and J. Albaiges, *J. Marine Syst.*, 2010, **79**, 210–217.
- 71 A. Eschenbach, M. Kastner, R. Bierl, G. Schaefer and B. Mahro, *Chemosphere*, 1994, **28**, 683–692.
- 72 P. B. Hatzinger and M. Alexander, *Environ. Sci. Technol.*, 1995, **29**, 537–545.
- 73 J. J. H. Haftka, H. A. J. Govers and J. R. Parsons, *Environ. Sci. Pollut. Res.*, 2010, **17**, 1070–1079.
- 74 P. Lehnik-Habrink, S. Hein, T. Win, W. Bremser and I. Nehls, *J. Soils Sediments*, 2010, **10**, 1487–1498.
- 75 Y. Yang, N. Zhang, M. Xue and S. Tao, *Environ. Pollut.*, 2010, **158**, 2170–2174.
- 76 R. M. Allen-King, P. Grathwohl and W. P. Ball, *Adv. Water Resour.*, 2002, **25**, 985–1016.
- 77 S. Ribes, B. van Drooge, J. Dachs, O. Gustafsson and J. O. Grimalt, *Environ. Sci. Technol.*, 2003, **37**, 2675–2680.
- 78 M. Obst, P. Grathwohl, A. Kappler, O. Eibl, N. Peranio and T. Gocht, *Environ. Sci. Technol.*, 2011, **45**, 7314–7322.

- 79 S. Endo, P. Grathwohl, S. B. Haderlein and T. C. Schmidt, *Environ. Sci. Technol.*, 2009, **43**, 3094–3100.
- 80 B. Pan, S. Tao, D. Wu, D. Zhang, H. Peng and B. Xing, *Chemosphere*, 2011, **84**, 1578–1583.
- 81 G. E. Schaumann, *J. Plant Nutr. Soil Sci.*, 2006, **169**, 145–156.
- 82 Y. Yang, N. Zhang, M. Xue, S. T. Lu and S. Tao, *Environ. Pollut.*, 2011, **159**, 591–595.
- 83 E. Brorstrom, P. Grennfelt and A. Lindskog, *Atmos. Environ.*, 1983, **17**, 601–605.
- 84 M. Carrara, J. C. Wolf and R. Niessner, *Atmos. Environ.*, 2010, **44**, 3878–3885.
- 85 M. Goriaux, B. Jourdain, B. Temime, J. L. Besombes, N. Marchand, A. Albinet, E. Leoz-Garziandia and H. Wortham, *Environ. Sci. Technol.*, 2006, **40**, 6398–6404.
- 86 I. G. Kavouras, J. Lawrence, P. Koutrakis, E. G. Stephanou and P. Oyola, *Atmos. Environ.*, 1999, **33**, 4977–4986.
- 87 M. Possanzini, A. Febo and A. Liberti, *Atmos. Environ.*, 1983, **17**, 2605–2610.
- 88 C. Schauer, R. Niessner and U. Poschl, *Environ. Sci. Technol.*, 2003, **37**, 2861–2868.
- 89 M. Tsapakis and E. G. Stephanou, *Atmos. Environ.*, 2003, **37**, 4935–4944.
- 90 C. E. Junge, Basic considerations about trace constituents in the atmosphere as related to the fate of global 16 pollutants, in *Fate of Pollutants in the Air and Water Environments, Part I*, ed. I. H. Suffet, Wiley, New York, 1977, vol. 17, pp. 7–26.
- 91 J. F. Pankow, *Atmos. Environ.*, 1987, **21**, 2275–2283.
- 92 J. F. Pankow and T. F. Bidleman, *Atmos. Environ., Part A*, 1992, **26**, 1071–1080.
- 93 K. U. Goss and R. P. Schwarzenbach, *Environ. Sci. Technol.*, 1998, **32**, 2025–2032.
- 94 H. Yamasaki, K. Kuwata and H. Miyamoto, *Environ. Sci. Technol.*, 1982, **16**, 189–194.
- 95 J. F. Pankow, L. M. Isabelle, D. A. Buchholz, W. T. Luo and B. D. Reeves, *Environ. Sci. Technol.*, 1994, **28**, 363–363.
- 96 E. Galarneau, T. F. Bidleman and P. Blanchard, *Atmos. Environ.*, 2006, **40**, 182–197.
- 97 K. E. Gustafson and R. M. Dickhut, *Environ. Sci. Technol.*, 1997, **31**, 140–147.
- 98 V. Subramanyam, K. T. Valsaraj, L. J. Thibodeaux and D. D. Reible, *Atmos. Environ.*, 1994, **28**, 3083–3091.
- 99 S. Ozcan and M. E. Aydin, *Atmos. Res.*, 2009, **93**, 715–722.
- 100 K. Ravindra, L. Bencs, E. Wauters, J. de Hoog, F. Deutsch, E. Roekens, N. Bleux, P. Berghmans and R. Van Grieken, *Atmos. Environ.*, 2006, **40**, 771–785.
- 101 A. Finizio, D. Mackay, T. Bidleman and T. Harner, *Atmos. Environ.*, 1997, **31**, 2289–2296.
- 102 J. Dachs and S. J. Eisenreich, *Environ. Sci. Technol.*, 2000, **34**, 3690–3697.
- 103 R. Lohmann and G. Lammel, *Environ. Sci. Technol.*, 2004, **38**, 3793–3803.
- 104 M. T. O. Jonker and A. A. Koelmans, *Environ. Sci. Technol.*, 2002, **36**, 3725–3734.
- 105 B. T. Mader and J. F. Pankow, *Environ. Sci. Technol.*, 2002, **36**, 5218–5228.
- 106 S. Pongpiachan, K. Thamanu, K. F. Ho, S. C. Lee and P. Sompongchaiyakul, *Southeast Asian J. Trop. Med. Public Health*, 2009, **40**, 1377–1394.
- 107 G. Lammel, J. Klánová, P. Ilić, J. Kohoutek, B. Gasić, I. Kovacic and L. Škrdlíková, *Atmos. Environ.*, 2010, **44**, 5022–5027.
- 108 K. U. Goss and R. P. Schwarzenbach, *Environ. Sci. Technol.*, 2001, **35**, 1–9.
- 109 M. H. Abraham, *J. Chromatogr.*, 1993, **644**, 95–139.
- 110 S. N. Atapattu and C. F. Poole, *J. Chromatogr., A*, 2008, **1195**, 136–145.
- 111 C. M. Roth, K. U. Goss and R. P. Schwarzenbach, *Environ. Sci. Technol.*, 2005, **39**, 6638–6643.
- 112 K. U. Goss, *Crit. Rev. Environ. Sci. Technol.*, 2004, **34**, 339–389.
- 113 J. A. L. Cancio, A. V. Castellano, S. S. Martin and J. F. S. Rodriguez, *Water, Air, Soil Pollut.*, 2004, **154**, 127–138.
- 114 L. P. Chrysikou, P. G. Gemenetzi and C. A. Samara, *Atmos. Environ.*, 2009, **43**, 290–300.
- 115 G. Kiss, Z. Varga-Puchony and Z. Hlavaj, Distribution of polycyclic aromatic hydrocarbons on atmospheric aerosol particles of different size, in *Nucleation and Atmospheric Aerosols*, ed. M. Kulmala and P. O. Wagner, Elsevier, Kidlington, 1996, pp. 501–503.
- 116 J. Schnelle, T. Jansch, K. Wolf, I. Gebefugi and A. Kettrup, *Chemosphere*, 1995, **31**, 3119–3127.
- 117 C. Venkataraman and S. K. Friedlander, *Environ. Sci. Technol.*, 1994, **28**, 563–572.
- 118 Y. Kawanaka, Y. Tsuchiya, S.-J. Yun and K. Sakamoto, *Environ. Sci. Technol.*, 2009, **43**, 6851–6856.
- 119 Y. Kameda, J. Shirai, T. Komai, J. Nakanishi and S. Masunaga, *Sci. Total Environ.*, 2005, **340**, 71–80.
- 120 A. H. Miguel, T. W. Kirchstetter, R. A. Harley and S. V. Hering, *Environ. Sci. Technol.*, 1998, **32**, 450–455.
- 121 L. Škrdlíková, L. Landlova, J. Klanova and G. Lammel, *Atmos. Environ.*, 2011, **45**, 4305–4312.
- 122 T. F. Bidleman, *Environ. Sci. Technol.*, 1988, **22**, 361–367.
- 123 T. Gocht, O. Klemm and P. Grathwohl, *Atmos. Environ.*, 2007, **41**, 1315–1327.
- 124 T. P. Franz and S. J. Eisenreich, *Environ. Sci. Technol.*, 1998, **32**, 1771–1778.
- 125 Y. Tasdemir and F. Esen, *Atmos. Environ.*, 2007, **41**, 1288–1301.
- 126 M. P. Ligocki, C. Leuenberger and J. F. Pankow, *Atmos. Environ.*, 1985, **19**, 1609–1617.
- 127 S. K. Sahu, G. G. Pandit and S. Sadasivan, *Sci. Total Environ.*, 2004, **318**, 245–249.
- 128 M. F. Simcik, *Atmos. Environ.*, 2004, **38**, 491–501.
- 129 W. G. N. Slinn, L. Hasse, B. B. Hicks, A. W. Hogan, D. Lal, P. S. Liss, K. O. Munnich, G. A. Sehmel and O. Vittori, *Atmos. Environ.*, 1978, **12**, 2055–2087.
- 130 M. P. Ligocki, C. Leuenberger and J. F. Pankow, *Atmos. Environ.*, 1985, **19**, 1619–1626.
- 131 J. H. Offenberg and J. E. Baker, *Environ. Sci. Technol.*, 2002, **36**, 3763–3771.
- 132 J. T. Hoff, F. Wania, D. Mackay and R. Gillham, *Environ. Sci. Technol.*, 1995, **29**, 1982–1989.

- 133 F. Wania, D. Mackay and J. T. Hoff, *Environ. Sci. Technol.*, 1999, **33**, 195–197.
- 134 A. Albinet, E. Leoz-Garziandia, H. Budzinski and E. Villenave, *Sci. Total Environ.*, 2007, **384**, 280–292.
- 135 R. Atkinson and J. Arey, *Environ. Health Perspect.*, 1994, **102**, 117–126.
- 136 F. Wania and D. Mackay, *Environ. Sci. Technol.*, 1996, **30**, A390–A396.
- 137 S. C. Zhang, W. Zhang, K. Y. Wang, Y. T. Shen, L. W. Hu and X. J. Wang, *Environ. Monit. Assess.*, 2009, **151**, 197–207.
- 138 A. Birgul, Y. Tasdemir and S. S. Cindoruk, *Atmos. Res.*, 2011, **101**, 341–353.
- 139 E. Terzi and C. Samara, *Atmos. Environ.*, 2005, **39**, 6261–6270.
- 140 W. W. Brubaker and R. A. Hites, *J. Phys. Chem. A*, 1998, **102**, 915–921.
- 141 J. N. Pitts, *Atmos. Environ.*, 1987, **21**, 2531–2547.
- 142 J. Sasaki, J. Arey and W. P. Harger, *Environ. Sci. Technol.*, 1995, **29**, 1324–1335.
- 143 L. Wang, J. Arey and R. Atkinson, *Environ. Sci. Technol.*, 2005, **39**, 5302–5310.
- 144 R. Atkinson and S. M. Aschmann, *Int. J. Chem. Kinet.*, 1986, **18**, 569–573.
- 145 F. L. Eisele, G. H. Mount, D. Tanner, A. Jefferson, R. Shetter, J. W. Harder and E. J. Williams, *J. Geophys. Res.: Atmos.*, 1997, **102**, 6457–6465.
- 146 R. G. Prinn, J. Huang, R. F. Weiss, D. M. Cunnold, P. J. Fraser, P. G. Simmonds, A. McCulloch, P. A. L. Bernard, C. Harth, P. Salameh, S. O'Doherty, R. H. J. Wang, L. Porter and B. R. Miller, *Science*, 2001, **292**, 1882–1888.
- 147 R. Atkinson, *J. Phys. Chem. Ref. Data, Monogr.*, 1989, 1–246.
- 148 R. Atkinson and J. Arey, *Chem. Rev.*, 2003, **103**, 4605–4638.
- 149 M. Krol, P. J. van Leeuwen and J. Lelieveld, *J. Geophys. Res.: Atmos.*, 1998, **103**, 10697–10711.
- 150 J. H. Seinfeld and S. N. Pandis, *Atmospheric Chemistry and Physics*, John Wiley & Sons, Inc., 1998, ISBN 0-471-17815-2.
- 151 R. Atkinson and J. Arey, *Polycyclic Aromat. Compd.*, 2007, **27**, 15–40.
- 152 M. Krol and J. Lelieveld, *J. Geophys. Res.: Atmos.*, 2003, **108**, 4125.
- 153 R. G. Prinn, J. Huang, R. F. Weiss, D. M. Cunnold, P. J. Fraser, P. G. Simmonds, A. McCulloch, C. Harth, S. Reimann, P. Salameh, S. O'Doherty, R. H. J. Wang, L. W. Porter, B. R. Miller and P. B. Krummel, *Geophys. Res. Lett.*, 2005, **32**, L07809.
- 154 X. Qu, Q. Zhang and W. Wang, *Chem. Phys. Lett.*, 2006, **429**, 77–85.
- 155 F. Reisen and J. Arey, *Environ. Sci. Technol.*, 2002, **36**, 4302–4311.
- 156 K. Lorenz and R. Zellner, *Ber. Bunsenges.*, 1983, **87**, 629–636.
- 157 R. Ananthula, T. Yamada and P. H. Taylor, *J. Phys. Chem. A*, 2006, **110**, 3559–3566.
- 158 F. Goulay, C. Rebrion-Rowe, J. L. Le Garrec, S. D. Le Picard, A. Canosa and B. R. Rowe, *J. Chem. Phys.*, 2005, **122**, 104308.
- 159 W. Lee, P. S. Stevens and R. A. Hites, *J. Phys. Chem. A*, 2003, **107**, 6603–6608.
- 160 R. Ananthula, T. Yamada and P. H. Taylor, *Int. J. Chem. Kinet.*, 2007, **39**, 629–637.
- 161 R. A. Perry, R. Atkinson and J. N. Pitts, *J. Phys. Chem.*, 1977, **81**, 296–304.
- 162 R. Atkinson, J. Arey, B. Zielinska and S. M. Aschmann, *Environ. Sci. Technol.*, 1987, **21**, 1014–1022.
- 163 R. Atkinson, J. Arey, B. Zielinska and S. M. Aschmann, *Int. J. Chem. Kinet.*, 1990, **22**, 999–1014.
- 164 R. Atkinson, S. M. Aschmann, J. Arey, B. Zielinska and D. Schuetzle, *Atmos. Environ.*, 1989, **23**, 2679–2690.
- 165 R. Koch, R. Knispel, M. Elend, M. Siese and C. Zetzsch, *Atmos. Chem. Phys.*, 2007, **7**, 2057–2071.
- 166 N. Nishino, R. Atkinson and J. Arey, *Environ. Sci. Technol.*, 2008, **42**, 9203–9209.
- 167 A. Ricca and C. W. Bauschlicher, *Chem. Phys. Lett.*, 2000, **328**, 396–402.
- 168 R. Atkinson, S. M. Aschmann and J. N. Pitts, *Environ. Sci. Technol.*, 1984, **18**, 110–113.
- 169 H. W. Biermann, H. Mac Leod, R. Atkinson, A. M. Winer and J. N. Pitts, *Environ. Sci. Technol.*, 1985, **19**, 244–248.
- 170 W. Klöpffer, R. Frank, E. G. Kohl and F. Haag, *Chem.-Ztg.*, 1986, **110**, 57–61.
- 171 P. T. Phousongphouang and J. Arey, *Environ. Sci. Technol.*, 2002, **36**, 1947–1952.
- 172 A. Klamt, *Chemosphere*, 1993, **26**, 1273–1289.
- 173 R. Atkinson and S. M. Aschmann, *Atmos. Environ.*, 1987, **21**, 2323–2326.
- 174 R. Atkinson and S. M. Aschmann, *Int. J. Chem. Kinet.*, 1988, **20**, 513–539.
- 175 C. E. Banceu, C. Mihele, D. A. Lane and N. J. Bunce, *Polycyclic Aromat. Compd.*, 2001, **18**, 415–425.
- 176 E. S. C. Kwok, W. P. Harger, J. Arey and R. Atkinson, *Environ. Sci. Technol.*, 1994, **28**, 521–527.
- 177 E. S. C. Kwok, R. Atkinson and J. Arey, *Int. J. Chem. Kinet.*, 1997, **29**, 299–309.
- 178 D. Helmig, J. Arey, R. Atkinson, W. P. Harger and P. A. McElroy, *Atmos. Environ., Part A*, 1992, **26**, 1735–1745.
- 179 D. Helmig and W. P. Harger, *Sci. Total Environ.*, 1994, **148**, 11–21.
- 180 J. Arey, B. Zielinska, R. Atkinson and S. M. Aschmann, *Int. J. Chem. Kinet.*, 1989, **21**, 775–799.
- 181 J. Y. Lee and D. A. Lane, *Atmos. Environ.*, 2009, **43**, 4886–4893.
- 182 J. Sasaki, S. M. Aschmann, E. S. C. Kwok, R. Atkinson and J. Arey, *Environ. Sci. Technol.*, 1997, **31**, 3173–3179.
- 183 N. J. Bunce, L. Liu, J. Zhu and D. A. Lane, *Environ. Sci. Technol.*, 1997, **31**, 2252–2259.
- 184 K. E. Kautzman, J. D. Surratt, M. N. Chan, A. W. H. Chan, S. P. Hersey, P. S. Chhabra, N. F. Dalleska, P. O. Wennberg, R. C. Flagan and J. H. Seinfeld, *J. Phys. Chem. A*, 2010, **114**, 913–934.
- 185 L. Wang, R. Atkinson and J. Arey, *Environ. Sci. Technol.*, 2007, **41**, 2803–2810.
- 186 N. Nishino, J. Arey and R. Atkinson, *Environ. Sci. Technol.*, 2009, **43**, 1349–1353.

- 187 J. Lee and D. A. Lane, *Atmos. Environ.*, 2010, **44**, 2469–2477.
- 188 J. Arey, B. Zielinska, R. Atkinson, A. M. Winer, T. Ramdahl and J. N. Pitts, *Atmos. Environ.*, 1986, **20**, 2339–2345.
- 189 D. Asaf, D. Pedersen, V. Matveev, M. Peleg, C. Kern, J. Zingler, U. Platt and M. Luria, *Environ. Sci. Technol.*, 2009, **43**, 9117–9123.
- 190 S. S. Brown, W. P. Dube, J. Peischl, T. B. Ryerson, E. Atlas, C. Warneke, J. A. de Gouw, S. T. Hekkert, C. A. Brock, F. Flocke, M. Trainer, D. D. Parrish, F. C. Feshenfeld and A. R. Ravishankara, *J. Geophys. Res.: Atmos.*, 2011, **116**, D24305.
- 191 A. Geyer, B. Alicke, S. Konrad, T. Schmitz, J. Stutz and U. Platt, *J. Geophys. Res.: Atmos.*, 2001, **106**, 8013–8025.
- 192 U. Platt, G. Lebras, G. Poulet, J. P. Burrows and G. Moortgat, *Nature*, 1990, **348**, 147–149.
- 193 J. M. Roberts, C. J. Hahn, F. C. Feshenfeld, J. M. Warnock, D. L. Albritton and R. E. Sievers, *Environ. Sci. Technol.*, 1985, **19**, 364–369.
- 194 A. M. Winer, R. Atkinson and J. N. Pitts, *Science*, 1984, **224**, 156–159.
- 195 C. A. Cantrell, J. A. Davidson, K. L. Busarow and J. G. Calvert, *J. Geophys. Res.: Atmos.*, 1986, **91**, 5347–5353.
- 196 C. A. Cantrell, W. R. Stockwell, L. G. Anderson, K. L. Busarow, D. Perner, A. Schmeltekopf, J. G. Calvert and H. S. Johnston, *J. Phys. Chem.*, 1985, **89**, 139–146.
- 197 W. R. Stockwell and J. G. Calvert, *J. Geophys. Res., C: Oceans Atmos.*, 1983, **88**, 6673–6682.
- 198 W. P. L. Carter, A. M. Winer and J. N. Pitts, *Environ. Sci. Technol.*, 1981, **15**, 829–831.
- 199 N. Carslaw, L. J. Carpenter, J. M. C. Plane, B. J. Allan, R. A. Burgess, K. C. Clemitshaw, H. Coe and S. A. Penkett, *J. Geophys. Res.: Atmos.*, 1997, **102**, 18917–18933.
- 200 D. Mihelcic, D. Klemp, P. Musgen, H. W. Patz and A. Volzthomas, *J. Atmos. Chem.*, 1993, **16**, 313–335.
- 201 A. Geyer, R. Ackermann, R. Dubois, B. Lohrmann, T. Muller and U. Platt, *Atmos. Environ.*, 2001, **35**, 3619–3631.
- 202 R. Atkinson and A. C. Lloyd, *J. Phys. Chem. Ref. Data*, 1984, **13**, 315–444.
- 203 R. Atkinson, E. C. Tuazon and J. Arey, *Int. J. Chem. Kinet.*, 1990, **22**, 1071–1082.
- 204 R. Atkinson, A. M. Winer and J. N. J. Pitts, *Atmos. Environ.*, 1986, **20**, 331–340.
- 205 R. A. Graham and H. S. Johnston, *J. Phys. Chem.*, 1978, **82**, 254–268.
- 206 F. Magnotta and H. S. Johnston, *Geophys. Res. Lett.*, 1980, **7**, 769–772.
- 207 R. Atkinson, *J. Phys. Chem. Ref. Data, Monogr.*, 1994, **2**, 216.
- 208 M. W. Malko and J. Troe, *Int. J. Chem. Kinet.*, 1982, **14**, 399–416.
- 209 T. J. Wallington, R. Atkinson, A. M. Winer and J. N. Pitts, *Int. J. Chem. Kinet.*, 1987, **19**, 243–249.
- 210 R. Atkinson, *J. Phys. Chem. Ref. Data*, 1991, **20**, 459–507.
- 211 J. F. Noxon, *J. Geophys. Res., C: Oceans Atmos.*, 1983, **88**, 1017–1021.
- 212 J. N. Pitts, B. Zielinska, J. A. Sweetman, R. Atkinson and A. M. Winer, *Atmos. Environ.*, 1985, **19**, 911–915.
- 213 R. Atkinson, E. C. Tuazon, I. Bridier and J. Arey, *Int. J. Chem. Kinet.*, 1994, **26**, 605–614.
- 214 P. T. Phousongphouang and J. Arey, *Environ. Sci. Technol.*, 2003, **37**, 308–313.
- 215 E. S. C. Kwok, R. Atkinson and J. Arey, *Int. J. Chem. Kinet.*, 1994, **26**, 511–525.
- 216 X. G. Qu, Q. Z. Zhang and W. X. Wang, *Chem. Phys. Lett.*, 2006, **432**, 40–49.
- 217 L. Wang, S. M. Aschmann, R. Atkinson and J. Arey, *Environ. Sci. Technol.*, 2009, **43**, 2766–2772.
- 218 B. Zielinska, J. Arey, R. Atkinson and P. A. McElroy, *Environ. Sci. Technol.*, 1989, **23**, 723–729.
- 219 L. Wang, R. Atkinson and J. Arey, *Environ. Sci. Technol.*, 2010, **44**, 2981–2987.
- 220 B. Zielinska, J. Arey, R. Atkinson, T. Ramdahl, A. M. Winer and J. N. Pitts, *J. Am. Chem. Soc.*, 1986, **108**, 4126–4132.
- 221 L. Wang, R. Atkinson and J. Arey, *Atmos. Environ.*, 2007, **41**, 2025–2035.
- 222 M. S. Alam, M. Camredon, A. R. Rickard, T. Carr, K. P. Wyche, K. E. Hornsby, P. S. Monks and W. J. Bloss, *Phys. Chem. Chem. Phys.*, 2011, **13**, 11002–11015.
- 223 R. Atkinson, *Atmos. Environ.*, 2000, **34**, 2063–2101.
- 224 U. Schurath and K. H. Naumann, *Pure Appl. Chem.*, 1998, **70**, 1353–1361.
- 225 D. Vione, V. Maurino, C. Minero, E. Pelizzetti, M. A. J. Harrison, R. I. Olariu and C. Arsene, *Chem. Soc. Rev.*, 2006, **35**, 441–453.
- 226 G. Andreou and S. Rapsomanikis, *J. Hazard. Mater.*, 2009, **172**, 363–373.
- 227 P. Di Filippo, C. Riccardi, D. Pomata and F. Buiarelli, *Atmos. Environ.*, 2010, **44**, 2742–2749.
- 228 Z. P. Gu, J. L. Feng, W. L. Han, L. Li, M. H. Wu, J. M. Fu and G. Y. Sheng, *J. Environ. Sci.*, 2010, **22**, 389–396.
- 229 S. Harrad and L. Laurie, *J. Environ. Monit.*, 2005, **7**, 722–727.
- 230 K. van Cauwenberghe and L. van Vaeck, *Mutat. Res.*, 1983, **116**, 1–20.
- 231 L. van Vaeck and K. van Cauwenberghe, *Atmos. Environ.*, 1984, **18**, 323–328.
- 232 T. D. Behymer and R. A. Hites, *Environ. Sci. Technol.*, 1988, **22**, 1311–1319.
- 233 M. Jang and S. R. McDow, *Environ. Sci. Technol.*, 1995, **29**, 2654–2660.
- 234 M. Jang and S. R. McDow, *Environ. Sci. Technol.*, 1997, **31**, 1046–1053.
- 235 R. M. Kamens, Z. Guo, J. N. Fulcher and D. A. Bell, *Environ. Sci. Technol.*, 1988, **22**, 103–108.
- 236 S. R. McDow, M. S. Jang, Y. Hong and R. M. Kamens, *J. Geophys. Res.: Atmos.*, 1996, **101**, 19593–19600.
- 237 S. R. McDow, Q. R. Sun, M. Vartiainen, Y. S. Hong, Y. L. Yao, T. Fister, R. Q. Yao and R. M. Kamens, *Environ. Sci. Technol.*, 1994, **28**, 2147–2153.
- 238 S. R. McDow, M. Vartiainen, Q. R. Sun, Y. S. Hong, Y. L. Yao and R. M. Kamens, *Atmos. Environ.*, 1995, **29**, 791–797.
- 239 T. D. Behymer and R. A. Hites, *Environ. Sci. Technol.*, 1985, **19**, 1004–1006.

- 240 R. Dabestani, K. J. Ellis and M. E. Sigman, *J. Photochem. Photobiol., A*, 1995, **86**, 231–239.
- 241 R. T. Dabestani, J. T. Barbas, K. J. Ellis and M. E. Sigman, *Abs. Am. Chem. Soc.*, 1995, **210**, 148-ENVR.
- 242 D. A. Lane and M. Katz, in *Fate of Pollutants in the Air and Water Environments, Part 2*, ed. A. Suffet, Wiley-Interscience, New York, 1977, pp. 133–154.
- 243 M. E. Sigman, R. Arce, C. Rayes, J. T. Barbas and R. Dabestani, *Abs. Am. Chem. Soc.*, 1997, **213**, 220-ENVR.
- 244 J. T. Barbas, R. Dabestani and M. E. Sigman, *J. Photochem. Photobiol., A*, 1994, **80**, 103–111.
- 245 J. T. Barbas, M. E. Sigman, A. C. Buchanan and E. A. Chevis, *Photochem. Photobiol.*, 1993, **58**, 155–158.
- 246 J. T. Barbas, M. E. Sigman and R. Dabestani, *Environ. Sci. Technol.*, 1996, **30**, 1776–1780.
- 247 R. A. Yokley, A. A. Garrison, E. L. Wehry and G. Mamantov, *Environ. Sci. Technol.*, 1986, **20**, 86–90.
- 248 W. A. Korfmacher, D. F. S. Natusch, D. R. Taylor, G. Mamantov and E. L. Wehry, *Science*, 1980, **207**, 763–765.
- 249 T. F. Kahan and D. J. Donaldson, *J. Phys. Chem. A*, 2007, **111**, 1277–1285.
- 250 T. F. Kahan, R. Zhao, K. B. Jumaa and D. J. Donaldson, *Environ. Sci. Technol.*, 2010, **44**, 1302–1306.
- 251 K. A. van Cauwenberghe, in *Handbook of Polycyclic Aromatic Hydrocarbons*, ed. A. Bjørseth and T. Ramdahl, Dekker, New York, 1985, pp. 351–384.
- 252 Y. Bedjanian, M. L. Nguyen and G. Le Bras, *Atmos. Environ.*, 2010, **44**, 1754–1760.
- 253 A. K. Bertram, A. V. Ivanov, M. Hunter, L. T. Molina and M. J. Molina, *J. Phys. Chem. A*, 2001, **105**, 9415–9421.
- 254 W. Estève, H. Budzinski and E. Villenave, *Polycyclic Aromat. Compd.*, 2003, **23**, 441–456.
- 255 W. Estève, H. Budzinski and E. Villenave, *Atmos. Environ.*, 2004, **38**, 6063–6072.
- 256 W. Estève, H. Budzinski and E. Villenave, *Atmos. Environ.*, 2006, **40**, 201–211.
- 257 K. Miet, H. Budzinski and E. Villenave, *Polycyclic Aromat. Compd.*, 2009, **29**, 267–281.
- 258 J. Ringuet, A. Albinet, E. Leoz-Garziandia, H. Budzinski and E. Villenave, *Atmos. Environ.*, 2012, **61**, 15–22.
- 259 Y. Zhang, B. Yang, J. Gan, C. Liu, X. Shu and J. Shu, *Atmos. Environ.*, 2011, **45**, 2515–2521.
- 260 C. Liu, P. Zhang, B. Yang, Y. Wang and J. Shu, *Environ. Sci. Technol.*, 2012, **46**, 7575–7580.
- 261 J. N. Pitts, J. A. Sweetman, B. Zielinska, R. Atkinson, A. M. Winer and W. P. Harger, *Environ. Sci. Technol.*, 1985, **19**, 1115–1121.
- 262 R. M. Kamens, J. Guo, Z. Guo and S. R. McDow, *Atmos. Environ., Part A*, 1990, **24**, 1161–1173.
- 263 J. Mak, S. Gross and A. K. Bertram, *Geophys. Res. Lett.*, 2007, **34**, L10804.
- 264 F. Karagulian and M. J. Rossi, *J. Phys. Chem. A*, 2007, **111**, 1914–1926.
- 265 S. Gross and A. K. Bertram, *J. Phys. Chem. A*, 2008, **112**, 3104–3113.
- 266 N. O. A. Kwamena and J. P. D. Abbatt, *Atmos. Environ.*, 2008, **42**, 8309–8314.
- 267 J. N. Pitts, K. A. van Cauwenberghe, D. Grosjean, J. P. Schmid, D. R. Fitz, W. L. Belser, G. B. Knudson and P. M. Hynds, *Science*, 1978, **202**, 515–519.
- 268 H. Tokiwa, R. Nakagawa, K. Morita and Y. Ohnishi, *Mutat. Res.*, 1981, **85**, 195–205.
- 269 Z. Guo and R. M. Kamens, *J. Atmos. Chem.*, 1991, **12**, 137–151.
- 270 J. D. Butler and P. Crossley, *Atmos. Environ.*, 1981, **15**, 91–94.
- 271 E. Brorstrom-Lunden and A. Lindskog, *Environ. Sci. Technol.*, 1985, **19**, 313–316.
- 272 M. M. Hughes, D. F. S. Natusch, D. R. Taylor and M. V. Zeller, in *Polynuclear Aromatic Hydrocarbons: Chemistry and Biological Effects*, ed. A. Bjørseth and A. J. Dennis, Battelle Press, Columbus, 1980.
- 273 C. H. Wu and H. Niki, *Environ. Sci. Technol.*, 1985, **19**, 1089–1094.
- 274 K. Inazu, T. Kobayashi and Y. Hisamatsu, *Chemosphere*, 1997, **35**, 607–622.
- 275 M. L. Nguyen, Y. Bedjanian and A. Guilloteau, *J. Atmos. Chem.*, 2009, **62**, 139–150.
- 276 E. Perraudin, H. Budzinski and E. Villenave, *Atmos. Environ.*, 2005, **39**, 6557–6567.
- 277 K. Miet, K. Le Menach, P. M. Flaud, H. Budzinski and E. Villenave, *Atmos. Environ.*, 2009, **43**, 837–843.
- 278 J. Ma, Y. Liu and H. He, *Atmos. Environ.*, 2011, **45**, 917–924.
- 279 T. Ramdahl, A. Bjørseth, D. M. Lokensgard and J. N. Pitts, *Chemosphere*, 1984, **13**, 527–534.
- 280 H. M. Wang, K. Hasegawa and S. Kagaya, *Chemosphere*, 1999, **39**, 1923–1936.
- 281 D. Grosjean, *Atmos. Environ.*, 1983, **17**, 2112–2114.
- 282 A. Lindskog, E. Brorström-Lundén and A. Sjödin, *Environ. Int.*, 1985, **11**, 125–130.
- 283 R. A. Yokley, A. A. Garrison, G. Mamantov and E. L. Wehry, *Chemosphere*, 1985, **14**, 1771–1778.
- 284 H. M. Wang, K. Hasegawa and S. Kagaya, *Chemosphere*, 2000, **41**, 1479–1484.
- 285 H. M. Wang, K. Hasegawa and S. Kagaya, *J. Environ. Sci. Health, Part A: Toxic/Hazard. Subst. Environ. Eng.*, 2000, **35**, 765–773.
- 286 M. Katz, C. Chan, H. Tosine and T. Sakuma, in *Polynuclear Aromatic Hydrocarbons*, ed. P. W. Jones and P. Leber, Ann Arbor Science Publishers, Ann Arbor, MI, 1979, pp. 171–189.
- 287 J. N. Pitts, D. M. Lokensgard, P. S. Ripley, K. A. van Cauwenberghe, L. van Vaecck, S. D. Shaffer, A. J. Thill and W. L. Belser, *Science*, 1980, **210**, 1347–1349.
- 288 J. N. Pitts, H. R. Paur, B. Zielinska, J. Arey, A. M. Winer, T. Ramdahl and V. Mejia, *Chemosphere*, 1986, **15**, 675–685.
- 289 S. Lelievre, Y. Bedjanian, N. Pouvesle, J. L. Delfau, C. Vovelle and G. Le Bras, *Phys. Chem. Chem. Phys.*, 2004, **6**, 1181–1191.
- 290 U. Pöschl, T. Letzel, C. Schauer and R. Niessner, *J. Phys. Chem. A*, 2001, **105**, 4029–4041.
- 291 N. O. A. Kwamena, J. A. Thornton and J. P. D. Abbatt, *J. Phys. Chem. A*, 2004, **108**, 11626–11634.

- 292 N. O. A. Kwamena, M. E. Earp, C. J. Young and J. P. D. Abbatt, *J. Phys. Chem. A*, 2006, **110**, 3638–3646.
- 293 B. T. Mmereki and D. J. Donaldson, *J. Phys. Chem. A*, 2003, **107**, 11038–11042.
- 294 B. T. Mmereki, D. J. Donaldson, J. B. Gilman, T. L. Eliason and V. Vaida, *Atmos. Environ.*, 2004, **38**, 6091–6103.
- 295 S. Raja and K. T. Valsaraj, *J. Air Waste Manage. Assoc.*, 2005, **55**, 1345–1355.
- 296 T. F. Kahan, N. O. A. Kwamena and D. J. Donaldson, *Atmos. Environ.*, 2006, **40**, 3448–3459.
- 297 K. Miet, K. Le Menach, P. M. Flaud, H. Budzinski and E. Villenave, *Atmos. Environ.*, 2009, **43**, 3699–3707.
- 298 E. Perraudin, H. Budzinski and E. Villenave, *J. Atmos. Chem.*, 2007, **56**, 57–82.
- 299 C. H. Wu, I. Salmeen and H. Niki, *Environ. Sci. Technol.*, 1984, **18**, 603–607.
- 300 R. M. Kamens, J. M. Perry, D. A. Saucy, D. A. Bell, D. L. Newton and B. Brand, *Environ. Int.*, 1985, **11**, 131–136.
- 301 N. O. A. Kwamena, M. G. Staikova, D. J. Donaldson, I. J. George and J. P. D. Abbatt, *J. Phys. Chem. A*, 2007, **111**, 11050–11058.
- 302 A. Alebic-Juretic, T. Cvitas and L. Klasinc, *Environ. Sci. Technol.*, 1990, **24**, 62–66.
- 303 E. Perraudin, H. Budzinski and E. Villenave, *Atmos. Environ.*, 2007, **41**, 6005–6017.
- 304 P. S. Bailey, *Ozonation in Organic Chemistry: Vol. II, Nonolefinic Compounds*, Academic Press, New York, 1982, pp. 82–84.
- 305 T. Letzel, E. Rosenberg, R. Wissiack, M. Grasserbauer and R. Niessner, *J. Chromatogr. A*, 1999, **855**, 501–514.
- 306 J. Jäger and V. Hanus, *J. Hyg., Epidemiol., Microbiol., Immunol.*, 1980, **24**, 1–15.
- 307 H. Y. Chen, Y. G. Teng and J. S. Wang, *Sci. Total Environ.*, 2012, **414**, 293–300.
- 308 S. Zhou, A. K. Y. Lee, R. D. McWhinney and J. P. D. Abbatt, *J. Phys. Chem. A*, 2012, **116**, 7050–7056.
- 309 R. W. Coutant, L. Brown, J. C. Chuang, R. M. Riggan and R. G. Lewis, *Atmos. Environ.*, 1988, **22**, 403–409.
- 310 K. Zimmermann, N. Jariyasopit, S. L. Massey Simonich, S. Tao, R. Atkinson and J. Arey, *Environ. Sci. Technol.*, 2013, **47**, 8434–8442.
- 311 F. M. Jaward, N. J. Farrar, T. Harner, A. J. Sweetman and K. C. Jones, *Environ. Toxicol. Chem.*, 2004, **23**, 1355–1364.
- 312 J. Klánová, P. Čupr, I. Holoubek, J. Borůvková, P. Příbylová, R. Kareš, T. Tomšej and T. Ocelka, *J. Environ. Monit.*, 2009, **11**, 1952–1963.
- 313 L. Nizzetto, R. Lohmann, R. Gioia, A. Jahnke, C. Temme, J. Dachs, P. Herckes, A. Di Guardo and K. C. Jones, *Environ. Sci. Technol.*, 2008, **42**, 1580–1585.
- 314 S. A. Genualdi, R. K. Killin, J. Woods, G. Wilson, D. Schmedding and S. L. M. Simonich, *Environ. Sci. Technol.*, 2009, **43**, 1061–1066.
- 315 T. Primbs, A. Piekarz, G. Wilson, D. Schmedding, C. Higginbotham, J. Field and S. M. Simonich, *Environ. Sci. Technol.*, 2008, **42**, 6385–6391.
- 316 J. M. Daisey, R. J. McCaffrey and R. A. Gallagher, *Atmos. Environ.*, 1981, **15**, 1353–1363.
- 317 P. Fernández, J. O. Grimalt and R. M. Vilanova, *Environ. Sci. Technol.*, 2002, **36**, 1162–1168.
- 318 C. J. Halsall, L. A. Barrie, P. Fellin, D. C. G. Muir, B. N. Billeck, A. R. White, L. Lockhart, F. Y. Rovinsky, E. Y. Kononov and B. Pastukhov, *Environ. Sci. Technol.*, 1997, **31**, 3593–3599.
- 319 A. K. Halse, M. Schlabach, S. Eckhardt, A. Sweetman, K. C. Jones and K. Breivik, *Atmos. Chem. Phys.*, 2011, **11**, 1549–1564.
- 320 H. Hung, P. Blanchard, C. J. Halsall, T. F. Bidleman, G. A. Stern, P. Fellin, D. C. G. Muir, L. A. Barrie, L. M. Jantunen, P. A. Helm, J. Ma and A. Konoplev, *Sci. Total Environ.*, 2005, **342**, 119–144.
- 321 J. L. Jaffrézo, P. Masclet, M. P. Clain, H. Wortham, S. Beyne and H. Cachier, *Atmos. Environ., Part A*, 1993, **27**, 2781–2785.
- 322 G. Lammel, J. Klánová, J. Kohoutek, R. Prokeš, L. Ries and A. Stohl, *Environ. Pollut.*, 2009, **157**, 3264–3271.
- 323 L. Melymuk, M. Robson, P. A. Helm and M. L. Diamond, *Atmos. Environ.*, 2011, **45**, 1867–1875.
- 324 G. W. Patton, M. D. Walla, T. F. Bidleman and L. A. Barrie, *J. Geophys. Res.: Atmos.*, 1991, **96**, 10867–10877.
- 325 B. L. van Drooge, P. Fernandez, J. O. Grimalt, E. Stuchlik, C. J. Torres Garcia and E. Cuevas, *Environ. Sci. Pollut. Res.*, 2010, **17**, 1207–1216.
- 326 H. Xiao, S. Kang, Q. Zhang, W. Han, M. Loewen, F. Wong, H. Hung, Y. D. Lei and F. Wania, *J. Geophys. Res.: Atmos.*, 2010, **115**, D16310.
- 327 J. Gabrieli, P. Vallenga, G. Cozzi, P. Gabrielli, A. Gambaro, M. Sigl, F. Decet, M. Schwikowski, H. Gäggeler, C. Boutron, P. Cescon and C. Barbante, *Environ. Sci. Technol.*, 2010, **44**, 3260–3266.
- 328 S. Usenko, D. H. Landers, P. G. Appleby and S. L. Simonich, *Environ. Sci. Technol.*, 2007, **41**, 7235–7241.
- 329 T. Gocht, K. M. Moldenhauer and W. Puttmann, *Appl. Geochem.*, 2001, **16**, 1707–1721.
- 330 L. K. Ackerman, A. R. Schwindt, S. L. M. Simonich, D. C. Koch, T. F. Blett, C. B. Schreck, M. L. Kent and D. H. Landers, *Environ. Sci. Technol.*, 2008, **42**, 2334–2341.
- 331 X. Wang, T. Yao, Z. Cong, X. Yan, S. Kang and Y. Zhang, *Chin. Sci. Bull.*, 2007, **52**, 1405–1413.
- 332 A. J. Peters, D. J. Gregor, C. F. Teixeira, N. P. Jones and C. Spencer, *Sci. Total Environ.*, 1995, **160–161**, 167–179.
- 333 J. L. Jaffrézo, M. P. Clain and P. Masclet, *Atmos. Environ.*, 1994, **28**, 1139–1145.
- 334 V. Hoyau, J. L. Jaffrézo, P. H. Garrigues, M. P. Clain and P. Masclet, *Polycyclic Aromat. Compd.*, 1996, **8**, 35–44.
- 335 P. Kukučka, G. Lammel, A. Dvorská, J. Klánová, A. Möller and E. Fries, *Environ. Chem.*, 2010, **7**, 504–513.
- 336 P. G. Desideri, L. Lepri, L. Checchini and D. Santianni, *Int. J. Environ. Anal. Chem.*, 1994, **55**, 33–46.
- 337 R. Lohmann, R. Gioia, K. C. Jones, L. Nizzetto, C. Temme, Z. Xie, D. Schulz-Bull, I. Hand, D. M. Griffin, E. Morgan and L. Jantunen, *Environ. Sci. Technol.*, 2009, **43**, 5633–5639.
- 338 C. J. Halsall, A. J. Sweetman, L. A. Barrie and K. C. Jones, *Atmos. Environ.*, 2001, **35**, 255–267.

- 339 W. D. Hafner and R. A. Hites, *Environ. Sci. Technol.*, 2003, **37**, 3764–3773.
- 340 C. Lang, S. Tao, W. Liu, Y. Zhang and S. Simonich, *Environ. Sci. Technol.*, 2008, **42**, 5196–5201.
- 341 U. M. Sofowote, H. Hung, A. K. Rastogi, J. N. Westgate, P. F. Deluca, Y. Su and B. E. McCarry, *Atmos. Environ.*, 2011, **45**, 967–976.
- 342 D. Yaffe, Y. Cohen, J. Arey and A. J. Grosovsky, *Risk Anal.*, 2001, **21**, 275–294.
- 343 K. Prevedouros, E. Brorström-Lundén, C. J. Halsall, K. C. Jones, R. G. M. Lee and A. J. Sweetman, *Environ. Pollut.*, 2004, **128**, 17–27.
- 344 A. Aulinger, V. Matthias and M. Quante, *J. Appl. Meteor. Climatol.*, 2007, **46**, 1718–1730.
- 345 A. Gusev, O. Rozovskaya, V. Shatalov, V. Sokovyh, W. Aas and K. Breivik, *Persistent Organic Pollutants in the Environment. EMEP Status Report 3/2009, Meteorological Synthesizing Centre – East*, Moscow, Russian Federation, 2009, p. 84.
- 346 V. Matthias, A. Aulinger and M. Quante, *Atmos. Environ.*, 2009, **43**, 4078–4086.
- 347 V. Shatalov, A. Gusev, S. Dutchak, O. Rozovskaya, V. Sokovykh, N. Vulykh, W. Aas and K. Breivik, *Persistent organic pollutants in the environment. EMEP Status Report 3/2010*, Geneva, 2010, p. 100.
- 348 C. Silibello, G. Calori, M. P. Costa, M. G. Dirodi, M. Mircea, P. Radice, L. Vitali and G. Zanini, *Atmos. Pollut. Res.*, 2012, **3**, 399–407.
- 349 Y. Zhang, H. Shen, S. Tao and J. Ma, *Atmos. Environ.*, 2011, **45**, 2820–2827.
- 350 Y. Zhang, S. Tao, J. Ma and S. Simonich, *Atmos. Chem. Phys.*, 2011, **11**, 11993–12006.
- 351 A. M. Sehili and G. Lammel, *Atmos. Environ.*, 2007, **41**, 8301–8315.
- 352 C. L. Friedman and N. E. Selin, *Environ. Sci. Technol.*, 2012, **46**, 9501–9510.
- 353 G. Lammel, W. Klöpffer, V. S. Semeena, E. Schmidt and A. Leip, *Environ. Sci. Pollut. Res.*, 2007, **14**, 153–165.
- 354 M. Hauck, M. A. J. Huijbregts, J. M. Armitage, I. T. Cousins, A. M. J. Ragas and D. van de Meent, *Chemosphere*, 2008, **72**, 959–967.
- 355 P. Ciccioli, A. Cecinato, E. Brancaleoni, M. Frattoni, P. Zacchei, A. H. Miguel and P. D. Vasconcellos, *J. Geophys. Res.: Atmos.*, 1996, **101**, 19567–19581.
- 356 E. Alonso, N. Tapie, H. Budzinski, K. Lemenach, L. Peluhet and J. V. Tarazona, *Environ. Sci. Pollut. Res.*, 2008, **15**, 31–40.
- 357 D. Marinov, S. Dueri, I. Puillat, R. Carafa, E. Jurado, N. Berrojalbiz, J. Dachs and J. M. Zaldivar, *Mar. Pollut. Bull.*, 2009, **58**, 1554–1561.
- 358 C. Walgraeve, K. Demeestere, J. Dewulf, R. Zimmermann and H. Van Langenhove, *Atmos. Environ.*, 2010, **44**, 1831–1846.
- 359 T. Nielsen, *Environ. Sci. Technol.*, 1984, **18**, 157–163.
- 360 A. Liberti and P. Ciccioli, *HRC & CC, J. High Resolut. Chromatogr. Chromatogr. Commun.*, 1986, **9**, 492–501.
- 361 J. Arey, B. Zielinska, R. Atkinson and A. M. Winer, *Atmos. Environ.*, 1987, **21**, 1437–1444.
- 362 J. P. Zhu, X. L. Cao, R. Pigeon and K. Mitchell, *J. Air Waste Manage. Assoc.*, 2003, **53**, 67–76.
- 363 R. Atkinson, J. Arey, B. Zielinska, J. N. Pitts and A. M. Winer, *Atmos. Environ.*, 1987, **21**, 2261–2262.
- 364 H. A. Bamford and J. E. Baker, *Atmos. Environ.*, 2003, **37**, 2077–2091.
- 365 P. Ciccioli, A. Cecinato, E. Brancaleoni, R. Draisci and A. Liberti, *Aerosol Sci. Technol.*, 1989, **10**, 296–310.
- 366 P. Ciccioli, A. Cecinato, R. Cabella, E. Brancaleoni and P. Buttini, *Atmos. Environ., Part A*, 1993, **27**, 1261–1270.
- 367 T. T. Hien, L. T. Thanh, T. Kameda, N. Takenaka and H. Bandow, *Atmos. Environ.*, 2007, **41**, 7715–7725.
- 368 T. Nielsen, B. Seitz and T. Ramdahl, *Atmos. Environ.*, 1984, **18**, 2159–2165.
- 369 J. N. Pitts, J. A. Sweetman, B. Zielinska, A. M. Winer and R. Atkinson, *Atmos. Environ.*, 1985, **19**, 1601–1608.
- 370 T. Ramdahl, B. Zielinska, J. Arey, R. Atkinson, A. M. Winer and J. N. Pitts, *Nature*, 1986, **321**, 425–427.
- 371 F. Reisen and J. Arey, *Environ. Sci. Technol.*, 2005, **39**, 64–73.
- 372 W. Wang, N. Jariyasopit, J. Schrlau, Y. Jia, S. Tao, T.-W. Yu, R. H. Dashwood, W. Zhang, X. Wang and S. L. M. Simonich, *Environ. Sci. Technol.*, 2011, **45**, 6887–6895.
- 373 J. A. Sweetman, B. Zielinska, R. Atkinson, T. Ramdahl, A. M. Winer and J. N. Pitts, *Atmos. Environ.*, 1986, **20**, 235–238.
- 374 A. Feilberg, M. W. B. Poulsen, T. Nielsen and H. Skov, *Atmos. Environ.*, 2001, **35**, 353–366.
- 375 F. Marino, A. Cecinato and P. A. Siskos, *Chemosphere*, 2000, **40**, 533–537.
- 376 A. Albinet, E. Leoz-Garziandia, H. Budzinski, E. Villenave and J. L. Jaffrezo, *Atmos. Environ.*, 2008, **42**, 43–54.
- 377 M. Tsapakis and E. G. Stephanou, *Environ. Sci. Technol.*, 2007, **41**, 8011–8017.
- 378 F. Reisen, S. Wheeler and J. Arey, *Atmos. Environ.*, 2003, **37**, 3653–3657.
- 379 M. Dimashki, S. Harrad and R. M. Harrison, *Atmos. Environ.*, 2000, **34**, 2459–2469.
- 380 T. Murahashi and K. Hayakawa, *Anal. Chim. Acta*, 1997, **343**, 251–257.
- 381 J. M. Bayona, M. Casellas, P. Fernandez, A. M. Solanas and J. Albaiges, *Chemosphere*, 1994, **29**, 441–450.
- 382 A. Cecinato, *J. Sep. Sci.*, 2003, **26**, 402–408.
- 383 J. C. Chuang, G. A. Mack, M. R. Kuhlman and N. K. Wilson, *Atmos. Environ., Part B*, 1991, **25**, 369–380.
- 384 N. K. Wilson and T. R. McCurdy, *Atmos. Environ.*, 1995, **29**, 2575–2584.
- 385 P. Gupta, W. P. Harger and J. Arey, *Atmos. Environ.*, 1996, **30**, 3157–3166.
- 386 T. Kameda, *J. Health Sci.*, 2011, **57**, 504–511.
- 387 Y. Kojima, K. Inazu, Y. Hisamatsu, H. Okochi, T. Baba and T. Nagoya, *Atmos. Environ.*, 2010, **44**, 2873–2880.
- 388 A. Albinet, E. Leoz-Garziandia, H. Budzinski, E. Villenave and J. L. Jaffrezo, *Atmos. Environ.*, 2008, **42**, 55–64.
- 389 R. M. Kamens, H. Karam, J. H. Guo, J. M. Perry and L. Stockburger, *Environ. Sci. Technol.*, 1989, **23**, 801–806.

- 390 A. Eiguren-Fernandez, A. H. Miguel, R. Lu, K. Purvis, B. Grant, P. Mayo, E. Di Stefano, A. K. Cho and J. Froines, *Atmos. Environ.*, 2008, **42**, 2312–2319.
- 391 J. Arey, R. Atkinson, B. Zielinska and P. A. McElroy, *Environ. Sci. Technol.*, 1989, **23**, 321–327.
- 392 J. Arey, R. Atkinson, S. M. Aschmann and D. Schuetzle, *Polycyclic Aromat. Compd.*, 1990, **1**, 33–50.
- 393 Y. Kojima, K. Inazu, Y. Hisamatsu, H. Okochi, T. Baba and T. Nagoya, *Asian J. Atmos. Environ.*, 2010, **4**, 1–8.
- 394 A. Pozzer, P. Zimmermann, U. M. Doering, J. van Aardenne, H. Tost, F. Dentener, G. Janssens-Maenhout and J. Lelieveld, *Atmos. Chem. Phys.*, 2012, **12**, 6915–6937.
- 395 M. Giannouli, E.-A. Kalognomou, G. Mellios, N. Moussiopoulos, Z. Samaras and J. Fiala, *Atmos. Environ.*, 2011, **45**, 4753–4762.
- 396 H. Cai and S. Xie, *Sci. Total Environ.*, 2011, **409**, 1935–1948.
- 397 W. Wang, S. Simonich, B. Giri, Y. Chang, Y. Zhang, Y. Jia, S. Tao, R. Wang, B. Wang, W. Li, J. Cao and X. Lu, *Sci. Total Environ.*, 2011, **409**, 2942–2950.
- 398 Ref. 398–472 are cited in the ESI†; J. O. Allen, N. M. Dookeran, K. Taghizadeh, A. L. Lafleur, K. A. Smith and A. F. Sarofim, *Environ. Sci. Technol.*, 1997, **31**, 2064–2070.
- 399 C. A. Alves and C. A. Pio, *J. Braz. Chem. Soc.*, 2005, **16**, 1017–1029.
- 400 P. Castells, F. J. Santos and M. T. Gaiceran, *J. Chromatogr., A*, 2003, **1010**, 141–151.
- 401 A. K. Cho, E. Di Stefano, Y. You, C. E. Rodriguez, D. A. Schmitz, Y. Kumagai, A. H. Miguel, A. Eiguren-Fernandez, T. Kobayashi, E. Avol and J. R. Froines, *Aerosol Sci. Technol.*, 2004, **38**, 68–81.
- 402 M. Y. Chung, R. A. Lazaro, D. Lim, J. Jackson, J. Lyon, D. Rendulic and A. S. Hasson, *Environ. Sci. Technol.*, 2006, **40**, 4880–4886.
- 403 O. Delhomme, M. Millet and P. Herckes, *Talanta*, 2008, **74**, 703–710.
- 404 A. Eiguren-Fernandez, A. H. Miguel, E. Di Stefano, D. A. Schmitz, A. K. Cho, S. Thuraiatnam, E. L. Avol and J. R. Froines, *Aerosol Sci. Technol.*, 2008, **42**, 854–861.
- 405 N. Kishikawa, M. Nakao, Y. Ohba, K. Nakashima and N. Kuroda, *Chemosphere*, 2006, **64**, 834–838.
- 406 R. Koeber, J. M. Bayona and R. Niessner, *Environ. Sci. Technol.*, 1999, **33**, 1552–1558.
- 407 J. Konig, E. Balfanz, W. Funcke and T. Romanowski, *Anal. Chem.*, 1983, **55**, 599–603.
- 408 J. Lintelmann, K. Fischer, E. Karg and A. Schroppel, *Anal. Bioanal. Chem.*, 2005, **381**, 508–519.
- 409 J. Lintelmann, K. Fischer and G. Matuschek, *J. Chromatogr., A*, 2006, **1133**, 241–247.
- 410 Y. Liu, M. Sklorz, J. Schnelle-Kreis, J. Orasche, T. Ferge, A. Kettrup and R. Zimmermann, *Chemosphere*, 2006, **62**, 1889–1898.
- 411 J. D. McDonald, B. Zielinska, J. C. Sagebiel and M. R. McDaniel, *Aerosol Sci. Technol.*, 2002, **36**, 1033–1044.
- 412 C. Neususs, M. Pelzing, A. Plewka and H. Herrmann, *J. Geophys. Res.: Atmos.*, 2000, **105**, 4513–4527.
- 413 S. Nicol, J. Dugay and M. C. Hennion, *Chromatographia*, 2001, **53**, S464–S469.
- 414 E.-J. Park, D.-S. Kim and K. Park, *Environ. Monit. Assess.*, 2008, **137**, 441–449.
- 415 N. Re-Poppi and M. Santiago-Silva, *Atmos. Environ.*, 2005, **39**, 2839–2850.
- 416 E. Schnelle-Kreis, M. Sklorz, A. Peters, J. Cyrus and R. Zimmermann, *Atmos. Environ.*, 2005, **39**, 7702–7714.
- 417 J. Schnelle-Kreis, I. Gebefugi, G. Welzl, T. Jaensch and A. Kettrup, *Atmos. Environ.*, 2001, **35**, S71–S81.
- 418 M. D. Sienra, *Atmos. Environ.*, 2006, **40**, 2374–2384.
- 419 B. R. T. Simoneit, X. Bi, D. R. Oros, P. M. Medeiros, G. Sheng and J. Fu, *Environ. Sci. Technol.*, 2007, **41**, 7294–7302.
- 420 M. Sklorz, J.-J. Briede, J. Schnelle-Kreis, Y. Liu, J. Cyrus, T. M. de Kok and R. Zimmermann, *J. Toxicol. Environ. Health, Part A*, 2007, **70**, 1866–1869.
- 421 M. Strandell, S. Zakrisson, T. Alsberg, R. Westerholm, L. Winquist and U. Rannug, *Environ. Health Perspect.*, 1994, **102**, 85–92.
- 422 M. Tsapakis, E. Lagoudaki, E. G. Stephanou, I. G. Kavouras, P. Koutrakis, P. Oyola and D. von Baer, *Atmos. Environ.*, 2002, **36**, 3851–3863.
- 423 A. Valavanidis, K. Fiotakis, T. Vlahogianni, V. Papadimitriou and V. Pantikaki, *Environ. Chem.*, 2006, **3**, 118–123.
- 424 N. Yassaa, B. Y. Meklati, A. Cecinato and F. Marino, *Environ. Sci. Technol.*, 2001, **35**, 306–311.
- 425 T. Kameda, T. Goto, A. Toriba, N. Tang and K. Hayakawa, *J. Chromatogr., A*, 2009, **1216**, 6758–6761.
- 426 M. P. Ligocki and J. F. Pankow, *Environ. Sci. Technol.*, 1989, **23**, 75–83.
- 427 P. M. Fine, G. R. Cass and B. R. T. Simoneit, *Environ. Sci. Technol.*, 2002, **36**, 1442–1451.
- 428 E. M. Fitzpatrick, A. B. Ross, J. Bates, G. Andrews, J. M. Jones, J. Aspden, H. Phylaktou, M. Pourkashanian and A. Williams, *Process Saf. Environ. Prot.*, 2007, **85**, 430–440.
- 429 M. P. Fraser, G. R. Cass and B. R. T. Simoneit, *Environ. Sci. Technol.*, 1998, **32**, 2051–2060.
- 430 M. D. Hays, P. M. Fine, C. D. Geron, M. J. Kleeman and B. K. Gullett, *Atmos. Environ.*, 2005, **39**, 6747–6764.
- 431 Y. Iinuma, E. Brüggemann, T. Gnauk, K. Müller, M. O. Andreae, G. Helas, R. Parmar and H. Herrmann, *J. Geophys. Res.: Atmos.*, 2007, **112**, D08209.
- 432 C. A. Jakober, S. G. Riddle, M. A. Robert, H. Destailats, M. J. Charles, P. G. Green and M. J. Kleeman, *Environ. Sci. Technol.*, 2007, **41**, 4548–4554.
- 433 C. A. Jakober, M. A. Robert, S. G. Riddle, H. Destailats, M. J. Charles, P. G. Green and M. J. Kleeman, *Environ. Sci. Technol.*, 2008, **42**, 4697–4703.
- 434 J. Oda, I. Maeda, T. Mori, A. Yasuhara and Y. Saito, *Environ. Technol.*, 1998, **19**, 961–976.
- 435 J. Oda, S. Nomura, A. Yasuhara and T. Shibamoto, *Atmos. Environ.*, 2001, **35**, 4819–4827.
- 436 X. Rao, R. Kobayashi, R. White-Morris, R. Spaulding, P. Frazey and M. J. Charles, *J. AOAC Int.*, 2001, **84**, 699–705.
- 437 N. Re-Poppi and M. R. Santiago-Silva, *Chromatographia*, 2002, **55**, 475–481.
- 438 W. F. Rogge, L. M. Hildemann, M. A. Mazurek, G. R. Cass and B. R. T. Simoneit, *Environ. Sci. Technol.*, 1993, **27**, 2736–2744.

- 439 W. F. Rogge, L. M. Hildemann, M. A. Mazurek, G. R. Cass and B. R. T. Simoneit, *Environ. Sci. Technol.*, 1993, **27**, 1892–1904.
- 440 W. F. Rogge, L. M. Hildemann, M. A. Mazurek, G. R. Cass and B. R. T. Simoneit, *Environ. Sci. Technol.*, 1993, **27**, 636–651.
- 441 W. F. Rogge, L. M. Hildemann, M. A. Mazurek, G. R. Cass and B. R. T. Simoneit, *Environ. Sci. Technol.*, 1997, **31**, 2731–2737.
- 442 W. F. Rogge, L. M. Hildemann, M. A. Mazurek, G. R. Cass and B. R. T. Simoneit, *Environ. Sci. Technol.*, 1998, **32**, 13–22.
- 443 G. Shen, S. Wei, Y. Zhang, R. Wang, B. Wang, W. Li, H. Shen, Y. Huang, Y. Chen, H. Chen, W. Wei and S. Tao, *Atmos. Environ.*, 2012, **60**, 234–237.
- 444 S. Sidhu, B. Gullett, R. Striebich, J. Klosterman, J. Contreras and M. de Vito, *Atmos. Environ.*, 2005, **39**, 801–811.
- 445 M. Stefanova, S. P. Marinov, A. M. Mastral, M. S. Callen and T. Garcia, *Fuel Process. Technol.*, 2002, **77**, 89–94.
- 446 B. Zielinska, J. Sagebiel, J. D. McDonald, K. Whitney and D. R. Lawson, *J. Air Waste Manage. Assoc.*, 2004, **54**, 1138–1150.
- 447 H. A. Bamford, D. Z. Bezabeh, M. M. Schantz, S. A. Wise and J. E. Baker, *Chemosphere*, 2003, **50**, 575–587.
- 448 K. Hayakawa, T. Murahashi, M. Butoh and M. Miyazaki, *Environ. Sci. Technol.*, 1995, **29**, 928–932.
- 449 H. Kakimoto, M. Kitamura, Y. Matsumoto, S. Sakai, F. Kanoh, T. Murahashi, K. Akutsu, R. Kizu and K. Hayakawa, *J. Health Sci.*, 2000, **46**, 5–15.
- 450 H. Kakimoto, H. Yokoe, Y. Matsumoto, S. Sakai, F. Kanoh, T. Murahashi, K. Akutsu, A. Toriba, R. Kizu and K. Hayakawa, *J. Health Sci.*, 2001, **47**, 385–393.
- 451 L. C. Marr, K. Dzepina, J. L. Jimenez, F. Reisen, H. L. Bethel, J. Arey, J. S. Gaffney, N. A. Marley, L. T. Molina and M. J. Molina, *Atmos. Chem. Phys.*, 2006, **6**, 1733–1745.
- 452 H. F. Nassar, N. Tang, T. Kameda, A. Toriba, M. I. Khoder and K. Hayakawa, *Atmos. Environ.*, 2011, **45**, 7352–7359.
- 453 M. D. Sienra and N. G. Rosazza, *Atmos. Res.*, 2006, **81**, 265–276.
- 454 N. Tang, T. Hattori, R. Taga, K. Igarashi, X. Y. Yang, K. Tamura, H. Kakimoto, V. F. Mishukov, A. Toriba, R. Kizu and K. Hayakawa, *Atmos. Environ.*, 2005, **39**, 5817–5826.
- 455 N. Tang, M. Tabata, V. F. Mishukov, V. Sergineko, A. Toriba, R. Kizu and K. Hayakawa, *J. Health Sci.*, 2002, **48**, 30–36.
- 456 B. L. Valle-Hernandez, V. Mugica-Alvarez, E. Salinas-Talavera, O. Amador-Munoz, M. A. Murillo-Tovar, R. Villalobos-Pietrini and A. De Vizcaya-Ruiz, *Sci. Total Environ.*, 2010, **408**, 5429–5438.
- 457 B. Zielinska, J. Arey, R. Atkinson and A. M. Winer, *Atmos. Environ.*, 1989, **23**, 223–229.
- 458 J. C. Ball and W. C. Young, *Environ. Sci. Technol.*, 1992, **26**, 2181–2186.
- 459 R. M. Campbell and M. L. Lee, *Anal. Chem.*, 1984, **56**, 1026–1030.
- 460 W. M. Draper, *Chemosphere*, 1986, **15**, 437–447.
- 461 T. L. Gibson, *Atmos. Environ.*, 1982, **16**, 2037–2040.
- 462 K. Hayakawa, M. Butoh, Y. Hirabayashi and M. Miyazaki, *Jpn. J. Toxicol. Environ. Health*, 1994, **40**, 20–25.
- 463 K. Hayakawa, M. Butoh and M. Miyazaki, *Anal. Chim. Acta*, 1992, **266**, 251–256.
- 464 IARC, *International Agency for Research on Cancer, Monographs on the Evaluation of Carcinogenic Risks to Humans, Volume 46: Diesel and Gasoline Engine Exhausts and Some Nitroarenes*, <http://monographs.iarc.fr/ENG/Monographs/vol46/volume46.pdf>, 1989.
- 465 G. Karavalakis, S. Stournas and E. Bakeas, *Sci. Total Environ.*, 2009, **407**, 3338–3346.
- 466 T. Murahashi, R. Kizu, H. Kakimoto, A. Toriba and K. Hayakawa, *J. Health Sci.*, 1999, **45**, 244–250.
- 467 M. C. Paputa-Peck, R. S. Marano, D. Schuetzle, T. L. Riley, C. V. Hampton, T. J. Prater, L. M. Skewes, T. E. Jensen, P. H. Ruehle, L. C. Bosch and W. P. Duncan, *Anal. Chem.*, 1983, **55**, 1946–1954.
- 468 S. M. Rappaport, Z. L. Jin and X. B. Xu, *J. Chromatogr.*, 1982, **240**, 145–154.
- 469 M. A. Ratcliff, A. J. Dane, A. Williams, J. Ireland, J. Luecke, R. L. McCormick and K. J. Voorhees, *Environ. Sci. Technol.*, 2010, **44**, 8343–8349.
- 470 D. Schuetzle, *Environ. Health Perspect.*, 1983, **47**, 65–80.
- 471 D. Schuetzle, T. L. Riley, T. J. Prater, T. M. Harvey and D. F. Hunt, *Anal. Chem.*, 1982, **54**, 265–271.
- 472 R. Westerholm and K. E. Egeback, *Environ. Health Perspect.*, 1994, **102**, 13–23.