


 Cite this: *Nanoscale*, 2024, **16**, 13183

The role of fluocerite in the genesis of bastnäsite: mechanistic insights and transformation pathways†

 Luca Terribili, ^{*}a Remi Rateau, ^a Melanie Maddin^a and Juan Diego Rodriguez-Blanco^{a,b}

Fluocerite is a rare earth element (REE) fluoride found as an accessory mineral in magmatic-hydrothermal REE ore deposits, including alkaline complexes and carbonatites, where it is often associated with REE-fluorocarbonates. This study investigates the crystallisation kinetics, mechanisms and energetics of fluocerite (REEF₃) and its role as a precursor phase of bastnäsite, one of the key minerals used for the extraction of REE. Fluocerite was synthesized by reacting pure fluorite (CaF₂) with La-, Ce- and Nd-bearing solutions at temperatures ranging from ambient to low hydrothermal (30–90 °C). The synthetic fluocerites were then placed in contact with Na₂CO₃ solutions at temperatures up to 200 °C. A combined approach using powder X-ray diffraction and scanning electron microscopy with energy dispersive spectroscopy was used to determine the nature and quantify the crystallising solids. Our findings reveal a temperature-dependent fluorite–fluocerite transformation, with completion observed within 1 hour at 90 °C and extending to around 30 days at 30 °C. The rate of crystallisation decreases proportionally with the atomic number of the rare earth elements. On the other hand, the carbonation reaction of fluocerite exhibits a significantly slower rate, by ~3 orders of magnitude, in comparison to the fluorite–fluocerite transformation, regardless of temperature conditions. The synthetic La, Ce and Nd fluocerites transformed into bastnäsite at all temperatures, also forming cerianite (CeO₂) in the Ce-bearing experiments and metastable kozoite, NdCO₃OH_(orth), in the Nd-bearing experiments. Activation energies of fluocerite nucleation increase proportionally with the ionic radii of the REE (81 ± 6 (La); 84 ± 5 (Ce), 96 ± 10 (Nd) kJ mol⁻¹), while the activation energies associated with fluocerite crystallisation are slightly higher for La (90 ± 12 kJ mol⁻¹) and similar for Ce and Nd (76 ± 12 and 72 ± 8 kJ mol⁻¹, respectively).

 Received 12th April 2024,
 Accepted 11th June 2024

DOI: 10.1039/d4nr01614a

rsc.li/nanoscale

Introduction

The peculiar electromagnetic properties of the rare earth elements^{1–5} make them indispensable in high-tech industries and clean energy applications (*i.e.*, from wind turbines to rechargeable electric car batteries).^{6–8} However, REE are considered as critical resources because of the concerns over potential supply disruptions and looming shortages in the near future.⁹

There are a wide range of ore deposits where REE can be found. Among them there are magmatic-hydrothermal ones, including alkaline complexes and carbonatites^{4,10,11} of which an example is the Bayan Obo REE deposit in China, the largest

known in the world.^{3,4,10–12} These kinds of deposits can also constitute ideal sources for REE-enriched hydrothermal fluids, leading to deposits containing both primary igneous and hydrothermal REE mineralization.¹³ Hydrothermal aqueous fluids are of primary importance in the post-magmatic phase formation of carbonatite and alkaline associated REE deposits^{11,12,14,15} mobilising the REE and concentrate them, in some cases even to economic level.^{10–12,16} Nevertheless, REE could require an atypical hydrothermal activity such as, for example, the presence of hydrothermal fluids containing significant amounts of F and Cl.^{10,11,13} The enrichment in F in the hydrothermal fluids in carbonatite and alkaline REE deposits is thought to be produced by fluorite and fluorocarbonates (*i.e.* bastnäsite, synchysite and parisite)¹² which are common and important constituents of these deposits.^{11,12,14,17–22}

One accessory mineral found in magmatic-hydrothermal REE ore deposits is fluocerite, a fluoride mineral with chemical formula REEF₃¹⁶ that can be associated with monazite, xenotime, REE-bearing fluorite, bastnäsite, cerianite, synch-

^aDepartment of Geology, School of Natural Sciences, Trinity College Dublin, College Green, Dublin D02PN40, Ireland. E-mail: terribil@tcd.ie
^biCRAG, Department of Geology, School of Natural Sciences, Trinity College Dublin, College green, Dublin D02PN40, Ireland

 † Electronic supplementary information (ESI) available. See DOI: <https://doi.org/10.1039/d4nr01614a>


site, samarskite, zircon, thorite, allanite, cerite or gadolinite.^{23,24} Fluocerite is LREE enriched^{12,25,26} and usually displays the following order of preference for the REE: Ce > La > Nd.²⁵ Even if fluocerite is not as common as fluorocarbonates in nature it has been documented in several locations worldwide but always as a minor phase (e.g.,^{12,16,17,19,23,24,26–31}). It has also been suggested that the formation of fluorocarbonates and cerianite (CeO₂) may have involved the carbonation and oxidation of a fluocerite precursor.^{12,16,24,29} This hypothesis could shed light on the difference between natural observations and the results anticipated by experimental studies. Fluocerite is typically found in limited quantities in rare earth element REE-bearing deposits, whereas experimental studies predict its abundant precipitation in F- and REE-enriched hydrothermal fluids.^{12,32,33}

The literature on fluocerite primarily focuses on natural samples, mentioning its presence in specific ore deposits and occasionally delving into its chemistry, crystal orientation, and formation conditions in detail.^{12,23–26,31,34} Only a few experimental studies have dealt with its synthesis^{16,35} or its occurrence as a secondary phase.¹¹ To date, and to the best of our knowledge there is no experimental work addressing the kinetics and mechanisms of fluocerite formation nor investigating its potential role as precursor of REE-fluorocarbonates. The present work will try to fill this knowledge gap and contribute to shed light on the crystallisation mechanisms of REE-fluorocarbonates, which are still unclear.¹⁶ Enhancing our understanding of this topic will contribute to further advancements in comprehending the complex mineralisation processes taking place in REE-bearing deposits such as carbonatites.^{36,37}

We present a comprehensive multi-stage experimental investigation into the kinetics and mechanisms of La-, Ce-, and Nd-bearing fluocerite crystallisation, encompassing reactions involving fluorite and their carbonation processes across a spectrum of temperatures, from ambient to low hydrothermal conditions. Our study does not only elucidate the intricate processes involved in fluocerite formation processes but also underscores the significance of the fluorine ion in the crystallization of bastnäsité.

Materials and methods

In order to study the kinetics and mechanisms which rule the transformation reaction fluorite–fluocerite–bastnäsité two set of experiments were carried out. The first set involved the fluorite–fluocerite transformation while the second focused on the fluocerite–bastnäsité one. The fluorite–fluocerite experiments were carried out at 30, 50, 70, and 90 °C by adding 0.5 g of pure fluorite powder (CaF₂) in a glass bottle reactor to 80 mL of 50 mM La-, Ce- or Nd-bearing aqueous solutions, preheated to specific temperatures. The reactors were then sealed and placed in a preheated oven. All the solutions were prepared using reagent grade REE nitrate hexahydrates (La (NO₃)₃·6H₂O, Ce(NO₃)₃·6H₂O, Nd(NO₃)₃·6H₂O; Sigma-Aldrich 99.99% trace metals basis) and pure deionized (Milli-Q) water.

The fluocerite–bastnäsité experiments were carried out at 50, 90 and 200 °C by adding the synthetic fluocerite powder obtained from the fluorite–fluocerite experiments to a 50 mM Na₂CO₃ solution, prepared using reagent grade chemicals and pure deionized (Milli-Q) water. For the experiments carried out at 50 and 90 °C, 0.25 g of fluocerite were added to a glass bottle reactor with 80 mL of Na₂CO₃-bearing solution and placed in a preheated oven. For the 200 °C experiments, 0.25 g of fluocerite powder was added to 20 mL of CO₃-bearing solution solutions in 20 ml Teflon-lined stainless-steel autoclaves at saturated water vapor pressures and placed in a preheated oven. To monitor the ongoing fluocerite–bastnäsité transformation reaction, 5 mL suspensions were regularly extracted from the reactor using a pipette at specific intervals, in all experiments except the one conducted at 200 °C. For the experiments at 200 °C, multiple repetitions were performed to gather samples at various time intervals. The collected samples were immediately filtered using a vacuum filtration system with 0.2 μm polycarbonate filters and rinsed with isopropanol to prevent potential recrystallisation or formation of other solids from any residual interstitial water and then quickly dried in air.^{38–40} For the 200 °C fluocerite–bastnäsité reaction, the resultant solution was cooled to room temperature and subsequently filtered entirely using the aforementioned method.

The nature of the new crystallised solid was determined with powder X-ray diffraction (XRD). The samples were analysed with a Bruker D5000 powder X-ray diffractometer (Cu Kα radiation, 0.02 per step from 5° to 70° in 2θ at 0.2° min⁻¹) located at TRinity Technology and Enterprise Centre (Dublin). Identification of crystalline phases present in the samples was carried out with a Bruker DIFFRAC.EVA software in combination with the ICDD Powder Data File (PDF-4, The International Centre for Diffraction Data). Pattern-matching refinement, quantification of crystalline phases was carried out with the Rietveld refinement software TOPAS,⁴¹ using the structural files from the American Mineralogist Crystal Structure Database (Fluorite: AMCS D 0011683; Fluocerite-(Ce): AMCS D 0009548; Bastnasite-(Ce): AMCS D 0001553; Cerianite-(Ce): AMCS D 0011686; Kozoite-(Nd): AMCS D 0002475).

Images of the samples were obtained with scanning electron microscopy equipped with energy dispersive spectroscopy (SEM-EDS) to study potential changes in the morphology and size of the crystalline phases. In order to prepare the samples for SEM characterisation, these were placed on mounts and Au–Pd or C coated respectively for imaging or the EDS analyses. SEM-EDS analyses were conducted using a Tescan TIGER MIRA3 FEG-SEM operating under high vacuum conditions and equipped with two Oxford Instruments X-Max 150 mm² energy dispersive spectroscopy (EDS) detectors running Oxford Instruments AZtec analysis software. The analyses were carried out using a beam current of 300 pA and an accelerating voltage of either 5 or 10 kV, for detailed imaging, or 20 kV, for EDS analysis at the iCrag Lab at TRinity College Dublin.

Multiple Point EDS analyses and compositional maps were carried out on the samples to determine the atomic % of La, Ce, Nd, Ca and O present in the crystalline phases. The stan-



standard used for the SEM-EDS analysis was a Co wire with a degree of purity of 99.995%. The particle size distribution in the samples was measured with ImageJ Software.^{42–44}

Quantification of the reaction extents was used to provide information on the rate of crystallisation by using the empirical Avrami equation:⁴⁵

$$\alpha = 1 - \exp(-kt)^n \quad (1)$$

where k is a rate constant, t is time, α is the fraction crystallised, and n is a constant which depends on the transition mechanism. Rewriting the Avrami Equation gives:

$$-\ln(1 - \alpha) = n \ln k + n \ln t \quad (2)$$

The reaction with kinetics that conform to this equation give a straight line when $-\ln(1 - y)$ is plotted against $\ln t$.^{40,46–48} The empirical parameter n value is given by the value of the slope, which can be used to compare reaction mechanism. Parallel lines indicate a constant value of n , suggesting that the reaction mechanism is the same. The intercept on the y axis gives the value of $n \ln k$, by which the k value can be determined.^{40,48}

The apparent activation energy of crystallisation ($E_{a(\text{cryst.})}$) (eqn (3)) and the apparent activation energy for nucleation ($E_{a(\text{nucl.})}$) (eqn (4)) were calculated using the derived induction times (t_0) and rate constants (k) as a function of temperature using an Arrhenius approach.^{49–51} $E_{a(\text{cryst.})}$ is an “apparent” activation energy because it represents contributions from all of the many processes occurring in the reaction (*e.g.*, dissolution, aggregation) and not just the potential barrier for a single reaction.^{49,50,52} The t_0 values were determined through the interpolation of the curve obtained from the % transformation values from fluorite to fluocerite *vs.* time.

$$\ln k = \ln A_{(\text{cryst.})} - (E_{a(\text{cryst.})}/RT) \quad (3)$$

$$\ln t_0 = \ln A_{(\text{nucl.})} - (E_{a(\text{nucl.})}/RT) \quad (4)$$

In the equations A is a pre-exponential factor (s^{-1}), $E_{a(\text{cryst.})}$ is the activation energy for crystallization (kJ mol^{-1}), $E_{a(\text{nucl.})}$ is the activation energy for nucleation (kJ mol^{-1}), R is the gas constant ($8.314 \text{ J mol}^{-1} \text{ K}^{-1}$), and T is temperature (K).

Results

Experimental results revealed that the kinetics of fluorite–fluocerite transformation were dependent on temperature and influenced by the mass of REE. Also, the fluocerite carbonation experiments revealed the formation of newly-formed CO_3 -bearing phases. The combined use of powder X-ray powder diffraction (XRD) and scanning electron microscopy with energy dispersive spectroscopy (SEM-EDS) allowed the identification and quantification of the newly formed phases, which were used to derive the kinetics and mechanistic processes. Selected examples of the Rietveld refinement (Fig. SI-1†) and unit cells parameters from the refinements are presented in the ESI (Tables SI-1, SI-2, SI-3†).

Fluorite–fluocerite transformation

The interaction between the fluorite powder and the preheated 50 mM La, Ce, and Nd-bearing solutions resulted in the crystallisation of fluocerite (PDF 00-002-0529) as the sole product phase, irrespective of the REE used or the temperature. The XRD analyses carried out of the solids, taken at increasing time intervals, revealed a progressive replacement of the original fluorite by fluocerite (Fig. 1).

The analysis of the quantitative data obtained *via* Rietveld refinement highlighted the temperature and the atomic number of REE as crucial factors influencing the kinetics of the fluorite–fluocerite transformation. Experimental observations showed that the transformation rates were notably faster with lower atomic number REE. For instance, experiments conducted with La exhibited the most rapid transformation rates (>175, 25, 2.5, and <1 hours at 30, 50, 70, and 90 °C, respectively). Similar trends were evident in experiments involving Ce and Nd, albeit with longer transformation times compared to La (Fig. 2).

Specifically, transformation reactions at 30 °C required >100, >175, and >720 hours in the La, Ce, and Nd systems, respectively (Fig. 3). Detailed values regarding the extent of transformation reactions and the weight percentage of fluorite *versus* fluocerite are provided in Tables SI-4, SI-5, and SI-6.†

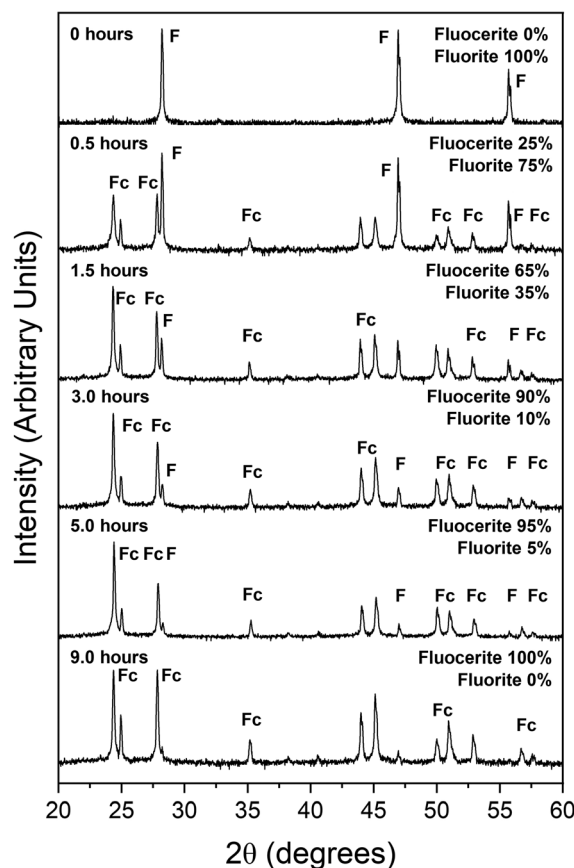


Fig. 1 Powder X-ray diffraction patterns of the Ce 70 °C experiment showing the progressive replacement from fluorite to fluocerite.



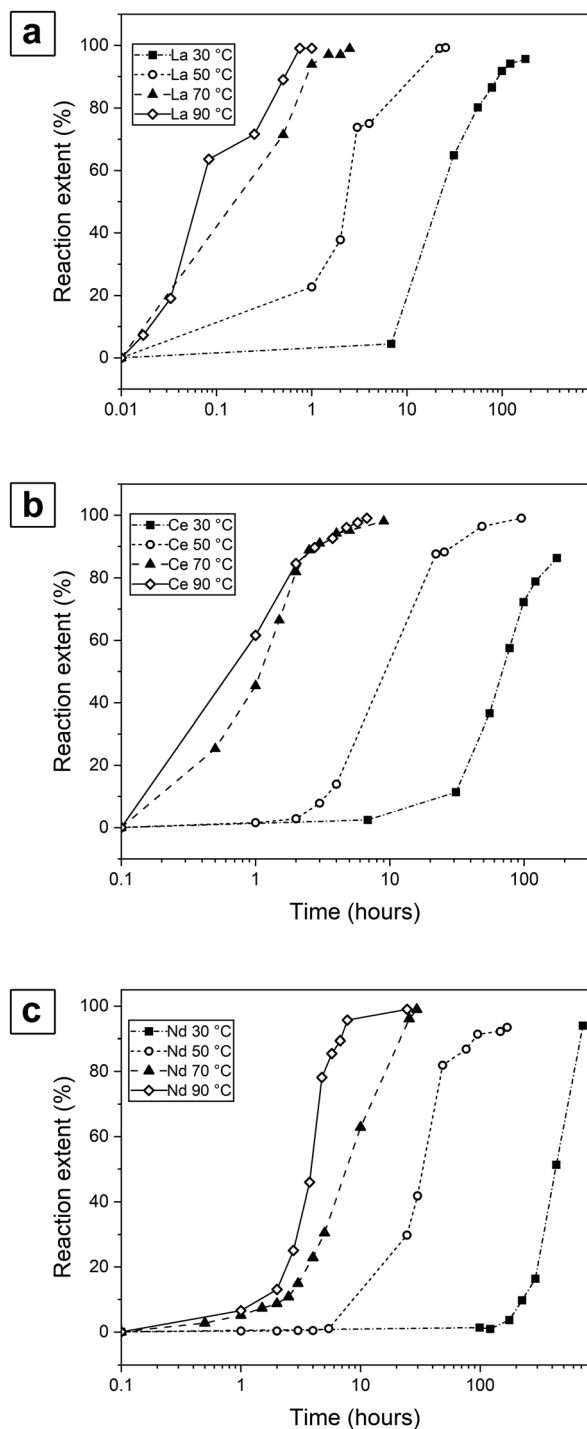


Fig. 2 Crystallisation reaction curves of fluorite for La (a), Ce (b), Nd (c) experiments carried out at temperatures from 30 to 90 °C.

SEM imaging revealed that the crystalline morphology of fluorite was found to be influenced by the presence of different REE in the starting solution and by variations in reaction temperatures, consistently resulting in flat hexagonal crystals in all experiments (Fig. SI-2†). The presence of different REE affected the crystalline dimensions, including diameter and thickness (Fig. 4a–c). It was then possible to recognise a

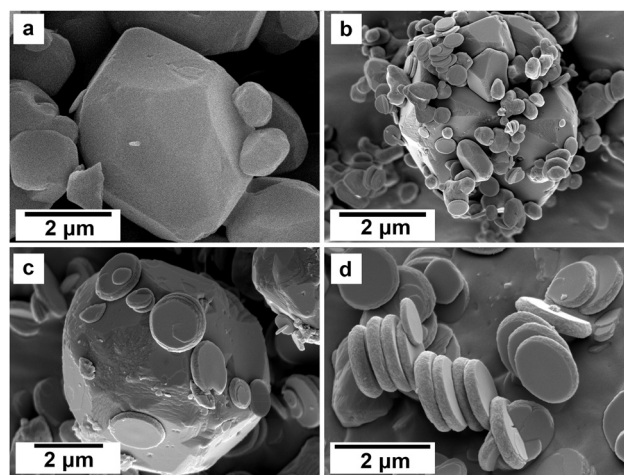


Fig. 3 SEM micrographs showing the progressive replacement of fluorite (a) by La-fluocerite (b, c and d) at 30 °C after increasing time. (a) Pure fluorite at the onset of the reaction, (b) La-fluocerite growing on the surface of a fluorite crystals after 1 day, (c) after 4 days and (d) after 5 days.

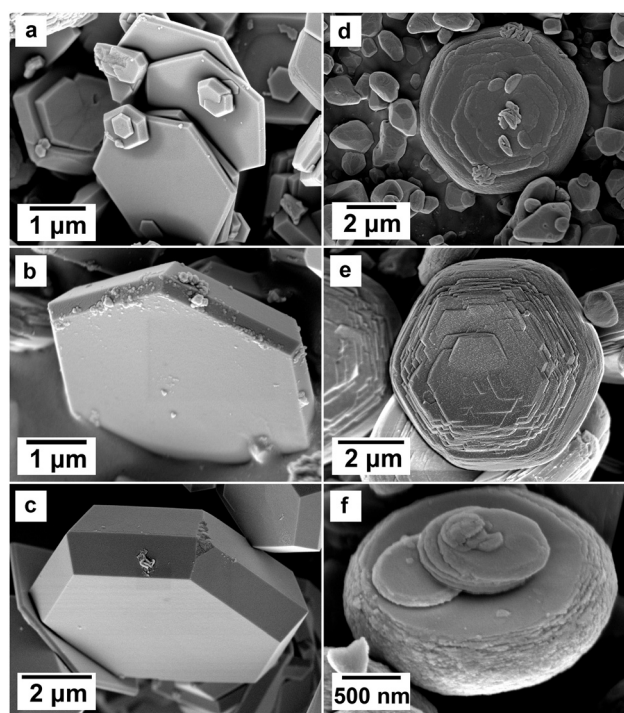


Fig. 4 La- (a), Ce- (b) and, Nd-fluocerite (c) crystals obtained from experiments at 90 °C showing the influence of different REE on the crystal morphology. (d and e) Sub-hexagonal growth imperfections on the (001) face of fluorite-Nd crystals crystallised at 70 °C. (f) Internal arrangement of the growing fluorite crystals consisting of layered nanodomains with thickness <50 nm.

general trend in which fluorite dimensions increased with the increasing REE ionic radius (see Table 1).

Fluocerite-La consisted of very thin (<500 nm) crystals with average sizes of 1.5 μm (Fig. 4a). Conversely, the largest fluo-



Table 1 Average diameters and thickness of REE-fluocerite crystals obtained at different temperature and collected at the end of the crystallisation reactions

T (°C)	La		Ce		Nd	
	Diameter (μm)	Thickness (μm)	Diameter (μm)	Thickness (μm)	Diameter (μm)	Thickness (μm)
30	0.75 ± 0.13	0.14 ± 0.04	1.22 ± 0.33	0.23 ± 0.05	2.44 ± 0.40	0.84 ± 0.22
50	1.49 ± 0.35	0.25 ± 0.07	0.80 ± 0.14	0.13 ± 0.04	2.12 ± 0.54	0.63 ± 0.18
70	0.70 ± 0.13	0.08 ± 0.02	1.26 ± 0.27	0.17 ± 0.06	2.56 ± 0.88	0.83 ± 0.37
90	1.64 ± 0.75	0.46 ± 0.18	2.70 ± 0.94	1.45 ± 0.68	4.45 ± 0.53	1.30 ± 0.30

cerite-Nd crystals crystallised as hexagonal prisms with sizes reaching 6 μm and thicknesses of 2–3 μm (Fig. 4c). Also, at the earliest stages of the crystallisation reactions and at the lower temperatures, fluocerite crystals showed a rounded and sub hexagonal morphology which became more euhedral with time and/or temperature (Fig. SI-3†). At 30 °C, crystals in all REE systems showed a rounded shape throughout the reaction. At 50 °C, this shape persisted for up to 1 day for La, 2 days for Ce, and 4 days for Nd. At 70 °C, the duration shortened to 4 hours for La, 5 hours for Ce, and 10 hours for Nd. At the highest temperature (90 °C) all fluocerites crystals appeared to be flat euhedral hexagons. Sometimes minor imperfections were visible on the (001) face of the crystals, consisting of growth steps of sub-hexagonal morphology. These defects were more common in the Nd system (Fig. 4d and e). Along the lateral faces, at temperature <90 °C it was also possible to see the internal layered arrangement of the growing hexagonal crystals, consisting of nanodomains (<100 nm) stacked on top of each other (Fig. 4f).

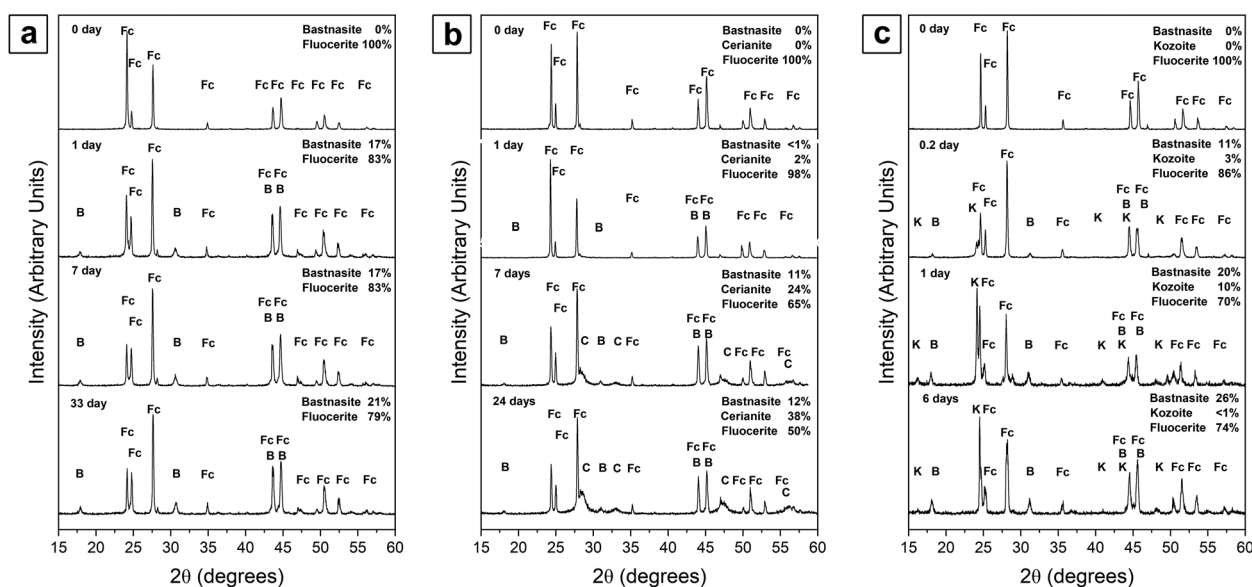
Results from energy dispersive spectroscopy (EDS) microanalysis on selected fluocerite crystals revealed the presence of REE (La, Ce, Nd) and F (Table SI-7†), supporting the XRD data.

In addition, they revealed the presence of trace amount of Ca quantifiable in <6 atomic% or below the detection limit.

Fluocerite–bastnäsäite transformation

Experimental results revealed that the transformation reactions involving synthetic La-, Ce- and Nd-fluocerite powders with a preheated 50 mM Na₂CO₃ resulted into the crystallisation of bastnäsäite (PDF 00-002-1433) (Table SI-8†) as primary phase at all the temperatures tested (50, 90 and 200 °C) and regardless the REE in the system. Cerianite (CeO₂; PDF 00-043-1002) was also found in the Ce system whereas kozoite (PDF 00-027-1296) in the Nd system, along with bastnäsäite.

The analysis of the quantitative data obtained by Rietveld refinement showed that the mineralogy of the final products of the fluocerite carbonation experiments were strongly REE dependent. In the La system the primary and only phase present was bastnäsäite whereas in Ce and Nd system bastnäsäite was always coexisting with cerianite and kozoite, respectively (Fig. 5a, b and c). The only experiments that showed a different trend were the Ce and Nd ones, in which the primary phases were cerianite and kozoite with bastnäsäite only occurring as a minor phase (see Tables SI-10 and SI-11†).

**Fig. 5** Powder X-ray diffraction patterns showing the transformation of the La- (a), Ce- (b) and Nd-fluocerite (c) to REE-fluorocarbonates, at 90 °C (La), 70 °C (Ce), and 200 °C (Nd).

Irrespective of the particular REE employed and the temperature conditions, none of the reactions led to a complete transformation of fluocerite into the newly formed phases (Fig. 5a, b and c). Following the initial stages of the reactions, data indicated a deceleration in the transformation process from fluocerite to the CO₃-bearing phases and cerianite, with the weight percentage of the involved phases exhibiting stability and minor fluctuations over extended periods. There were two exceptions to this general trend: first, the experiment at 50 °C in the La system, in which fluocerite was progressively replaced by bastnäsite during all the duration of the experiment (see Table SI-9†). Second, the Ce system, which showed a different general trend, with a progressive increase in the weight % of cerianite at all tested temperatures.

The SEM images taken at different times after the beginning of the interaction between synthetic fluocerites and CO₃-bearing solution revealed that nm-sized rare earth carbonates (bastnäsite and kozoite) started to grow on the surfaces of fluocerites and quickly covered the host crystals forming a crust. The same phenomenon was observed during the formation of cerianite in the Ce system.

SEM imaging revealed that the REE-bearing carbonates showed very similar morphology. Bastnäsite and kozoite appeared like individual nm-sized (~50 nm) crystals fully covering the surfaces of the fluocerite host (Fig. 6a–f).

In particular, bastnäsite crystallised as nanodomains with a columnar prismatic arrangement. At higher temperatures and

after longer reaction times, these bastnäsite nanocrystals increased in size and became more euhedral, with a pseudo columnar morphology (Fig. 7 and 8).

Overall, the bastnäsite nanostructures consisted of numerous small prisms that exhibited a reduction in thickness as they crystallised closer to the (001) faces of the fluocerite substrate (Fig. 8b–e), sometimes showing spindle-shaped morphologies while other ones (Fig. 8f) consisted of numerous

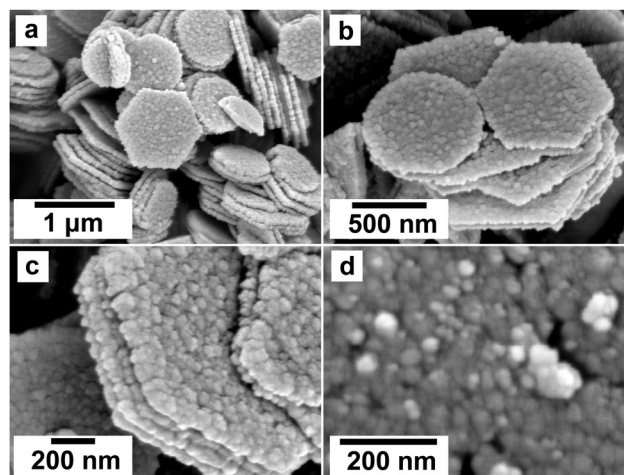


Fig. 7 (a and b) SEM micrographs showing a general view of La-fluocerite crystals replaced by bastnäsite at 90 °C. (c and d) Closer views of nano-bastnäsite overgrowths.

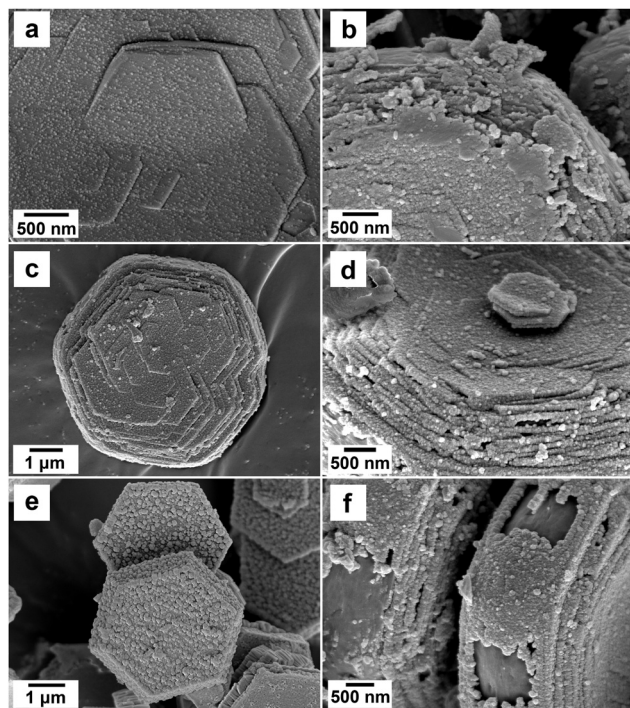


Fig. 6 SEM micrographs showing the progressive replacement of the surface of fluocerite by nanocrystalline REE-carbonates (a–f). Images (b), (d) and (f) show details of the newly formed nanocrystals covering the fluocerite host phase.

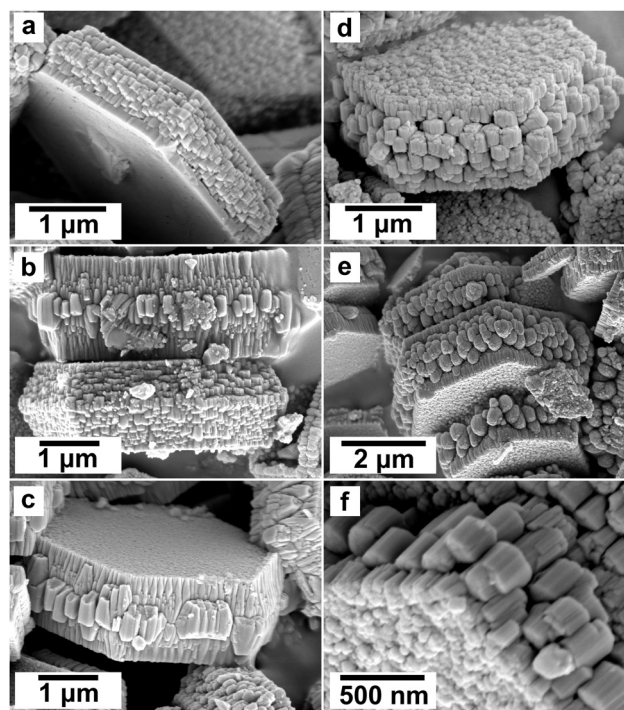


Fig. 8 SEM micrographs of bastnäsite replacing Nd-fluocerite at 90 °C (a, b and c) and 200 °C (d, e and f).



units of similar size and shape (sometimes reaching 500 nm), with the same orientation.

Cerianite, on the contrary, crystallised as nanoaggregates of crystals covering the surface of fluocerite during the early stages of the reactions, growing in size (up to ~ 100 nm) and developing a cubic or bipyramidal morphologies at higher temperatures and after long reaction times (Fig. 9).

A noteworthy observation is the non-random orientation of REE-fluorocarbonates, indicative of epitaxial overgrowth. Specifically, the fluocerite–bastnäsite epitaxial relationship was defined by the parallel alignment of the crystallographic axes of the bastnäsite to those of the underlying fluocerite substrate (Fig. 10).

Additionally, in the Ce experiments, a substantial number of cerianite crystals were identified, oriented precisely along the (001) plane of the fluocerite substrate (Fig. 9) with a perfect matching of the (111) crystallographic plane of cerianite to the (001) plane of fluocerite. At the highest tested temp-

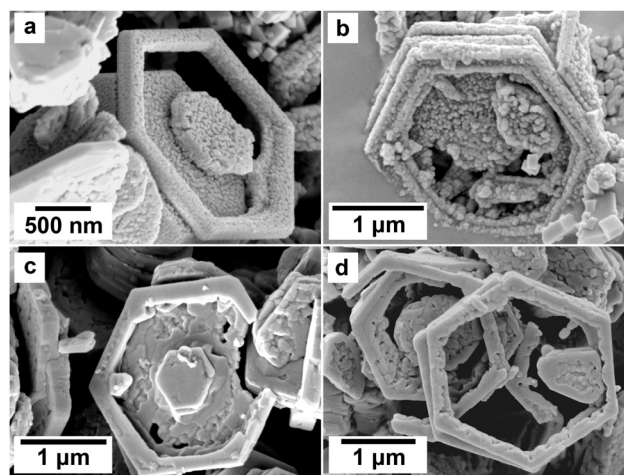


Fig. 11 SEM micrographs of skeletal Ce- (a and b) and La-fluocerite (c and d) crystals occurring at 200 °C after reaction times >48 h.

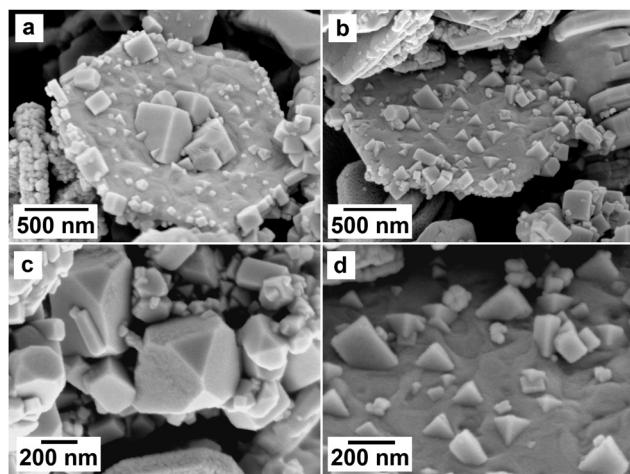


Fig. 9 SEM micrographs showing the crystallisation of cerianite on Ce-fluocerite. (a and b) Cerianite crystals growing on the (001) surface of Ce-fluocerite at 200 °C after 60 days. (c) Detailed view of cerianite nanoaggregates. (d) High-resolution view of a Ce-fluocerite crystal showing dissolution marks and cerianite nanocrystals growing on its (001) surface.

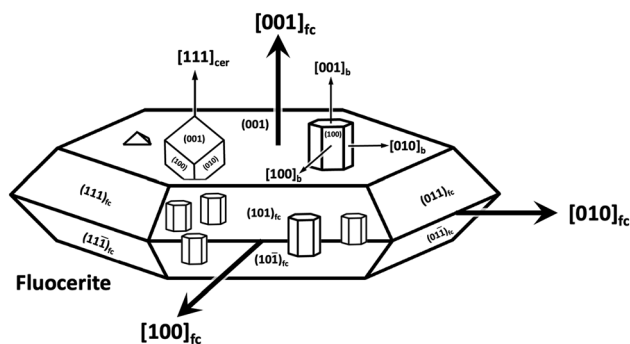


Fig. 10 Schematic diagram showing the epitaxial relationships between fluocerite, bastnäsite and cerianite crystals.

eratures and longest reaction times (*i.e.*, 200 °C after 60 days; see Fig. 11), intriguing dissolution features were observed in the original fluocerite crystals within both the La and Ce systems. These features manifested as dissolution progressing from the central area of crystal towards its edges (see Fig. 11b and d). In certain instances, this dissolution process advanced to the extent that skeletal hexagonal rings of the original fluocerite crystal were left behind.

Discussion

Our findings illustrate that REE ionic radii and temperature have a significant impact on the kinetics of the fluorite to fluocerite transformation reaction, as well as the crystalline morphology of the resulting solids. Additionally, the reaction of synthetic La-, Ce-, and Nd-bearing fluocerites with a CO₃-rich solution results in the formation of REE-fluorocarbonates and cerianite. The use of individual REE ions proves to be a key factor, profoundly influencing the crystallization pathways, polymorph selection, and the thermodynamically stable end product upon completion of the reactions.

Fluorite to fluocerite transformation

Fluocerite resulted to be the only product of the fluorite–fluocerite transformation reaction irrespectively of the REE ion present in the aqueous solution and of temperatures. The XRD and SEM data support a solvent-mediated surface precipitation and subsequent mineral replacement reaction that is dependent on temperature and the type of REE. The rate of transformation is directly proportional to the temperature and the atomic mass of the REE. In order to calculate the reaction rates (*k*) (Table 2), data obtained from fluorite to La, Ce and Nd fluocerite reaction curves (Fig. 1) were fitted to a JMAK (Johnson–Mehl–Avrami–Kolmogorov) particle nucleation model, which is based on the Avrami equation.



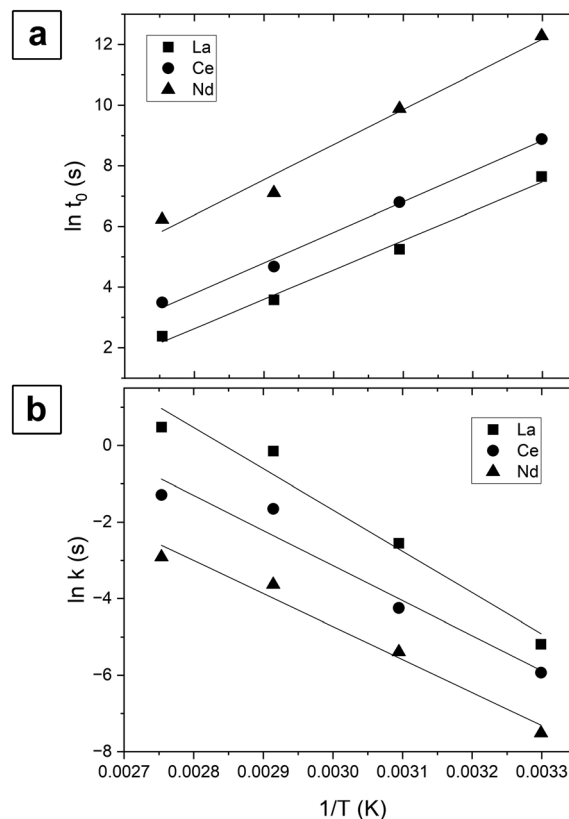
Table 2 Values of crystallisation rates (k), induction times (t_0) and Avrami exponents (n) for the fluorite to La-, Ce- and Nd-fluocerite transformation reactions, and activation energies of nucleation and crystallisation for La-, Ce- and Nd-fluocerite

T (°C)	La			Ce			Nd		
	n	k ($\times 10^{-3} \text{ s}^{-1}$)	t_0 (s)	n	k ($\times 10^{-3} \text{ s}^{-1}$)	t_0 (s)	n	k ($\times 10^{-3} \text{ s}^{-1}$)	t_0 (s)
30	1.31	0.006	2100	1.08	0.003	7200	2.91	0.001	216 000
50	0.88	0.078	192	1.36	0.014	900	1.62	0.005	19 800
70	0.76	0.864	36	0.95	0.193	108	1.30	0.027	1230
90	0.95	1.617	10.8	0.78	0.277	33	1.52	0.054	510
$E_{a(\text{cryst.})}$ kJ mol $^{-1}$	90(12)			76(12)			72(8)		
$E_{a(\text{nucl.})}$ kJ mol $^{-1}$	81(6)			83(5)			96(10)		

While the JMAK model is primarily associated with solid-state transformations and phase transitions in materials such as metals and alloys, it has also been adapted and applied to describe mineral-water interaction reactions, particularly in the context of mineral dissolution, precipitation, and transformation in aqueous environments (e.g.,^{53–56}). Each reaction curve obtained in this study fit to this model with a good R^2 (>90%). The data plots of $-\ln \ln(1 - \alpha)$ against $\ln t$ shows that reaction rates display close to parallel lines with similar n values (Fig. SI-4†). This indicates that the main crystallisation mechanism remains the same regardless of the different REE used or the temperature.^{46,47} The activation energies of nucleation and crystallisation for La, Ce and Nd fluocerites (Table 2), were calculated by plotting, respectively, the \ln of the induction times (t_0) and reactions rate (k) against the inverse of the temperature, as showed in Fig. 12.

The determined n values from the fluorite–fluocerite transformation reactions, using the JMAK model, are presented in Table 2. The n constant in the Avrami equation, denoted as $n = d + 1$, has been found to be related to the crystal growth dimensionality.^{55,57} For instance, a mineral undergoing spherulitic growth, occurring in three dimensions, would typically exhibit $n = 4$. In our experiments, a direct proportionality can be observed between the n values and the dimensionality of fluocerite growth for each REE, with lighter REE elements favouring linear growth (e.g., La-fluocerite with n values between 0.95 and 1.31; average of 0.98) and heavier elements displaying two-dimensional growth tendencies (e.g., Nd-fluocerite with n values between 1.52 and 2.91, average values of 1.84). It needs to be highlighted that in the hexagonal system, the two directions [100] and [010] within the basal plane (perpendicular to [001]) are equivalent in terms of crystallographic symmetry. Therefore, the growth of this basal plane should be considered unidimensional.

Data analyses reveal that La had the fastest reaction kinetics, followed very closely from Ce, while Nd showed the longest ones. These difference in kinetics can be explained by the progressive increasing of the ionic potential along the lanthanide series. REE³⁺ ionic radii decrease with increasing atomic number (a phenomenon known as the lanthanide contraction⁵⁸). This decrease in REE³⁺ ionic radii is translated into an increase in the ionic potentials, which are calculated by dividing the ion's valence by its ionic radius. In solution, all

**Fig. 12** Arrhenius plots for fluocerite nucleation (a) and crystallization (b).

ions are solvated, but they must remove their solvation shells to integrate into a crystalline structure. The ionic potential of the REE³⁺ cation determines its ability to retain water molecules in its solvation shell.^{59–61} Ions with higher ionic potentials require more energy to dehydrate before incorporating into mineral structures. Consequently, higher ionic potentials are translated into higher temperatures or longer times required for the initiation (nucleation) and completion of the crystallization reactions.^{59,62} The ionic potentials calculated for La³⁺, Ce³⁺, and Nd³⁺ are 2.60, 2.70, and 2.77 Å⁻¹ respectively.⁶³ La³⁺ would require the least energy for dehydration, followed closely by Ce³⁺, while Nd³⁺ would require the highest energy. This aligns with the longer kinetics observed in fluorite–fluocerite transformation reactions for heavier REE com-



pared to lighter ones. The activation energies for nucleation exhibit a rising trend, correlating to the REE atomic numbers, ranging from 81 to 96 kJ mol⁻¹. Conversely, the activation energies for crystallisation are similar between Ce and Nd (76 and 72 kJ mol⁻¹ respectively) and slightly higher for La (90 kJ mol⁻¹) (Table 2). These small discrepancies can be attributed to minor factors influencing the fluorite to fluorocerite replacement reactions, such as slight variations in the initial fluorite powder's particle size distribution or the presence of Ca impurities in the structure of fluorocerite affecting its solubility.

Fluocerite – REE-fluorocarbonates transformation

The carbonation of synthetic La-, Ce-, and Nd-fluocerite led to the formation of various phases, influenced by the specific REE present in each experiment. Bastnäsitate was predominant in La and Nd systems, while cerianite was the main phase forming in the Ce system (with minor amounts of bastnäsitate). The newly formed REE-fluorocarbonates nucleated on the surfaces of the host grains and continued to grow in size eventually forming a crust on the fluorocerite crystals. The final result consisted of a solution-mediated pseudomorphic replacement reaction in which the original morphology and dimensions of the fluorocerite grains remained constant, but most the host mineral was replaced by the newly formed phases.

A significant feature observed in the carbonation experiments is the presence of oriented overgrowths of the newly formed phases (bastnäsitate and cerianite) on the surface of fluorocerite. The epitaxial overgrowth observed between bastnäsitate and fluorocerite is not unexpected, given their shared crystal system (hexagonal) and similarities in space group. Fluocerite-Ce crystallises in the trigonal system with space group *P* $\bar{3}c1$, with lattice parameters of $a = b = 7.131$ Å and $c = 7.286$ Å.⁶⁴ Bastnäsitate belongs to the hexagonal crystal system with space group *P* $\bar{6}c2$, characterized by lattice parameters $a = 7.118(1)$ Å and $c = 9.762(1)$ Å.⁶⁵ These structural similarities suggest a favourable fit for epitaxial growth (Fig. 13 and 14).

The misfit values (Table 3) have been calculated using the expression:

$$mf(\%) = \frac{t_{[uvw]_{sub}} - t_{[uvw]_{over}}}{t_{[uvw]_{over}}} \times 100$$

and quantify the degree of lattice mismatch between the epitaxial overgrowth and the substrate,⁶⁶ further support this assertion. For most epitaxial directions, the misfit values are lower than -1.5%, falling well below the generally accepted limits (15–20%) required for epitaxial overgrowth to occur.⁶⁷ Their negative values indicate that the unit cell of bastnäsitate is very slightly contracted along the different directions of epitaxy of fluorocerite. The only exception arises in the [010] direction of fluorocerite, where the misfit with bastnäsitate (-25.60%) exceeds the typical threshold. This large value can be easily explained by the structural arrangement of bastnäsitate, with alternating CO₃ and Ce-F layers following [100] (Fig. 14).⁶⁵ The absence of CO₃ groups in fluorocerite results in the *c*-axes of this mineral being approximately 2 Å shorter compared to bastnäsitate.³¹

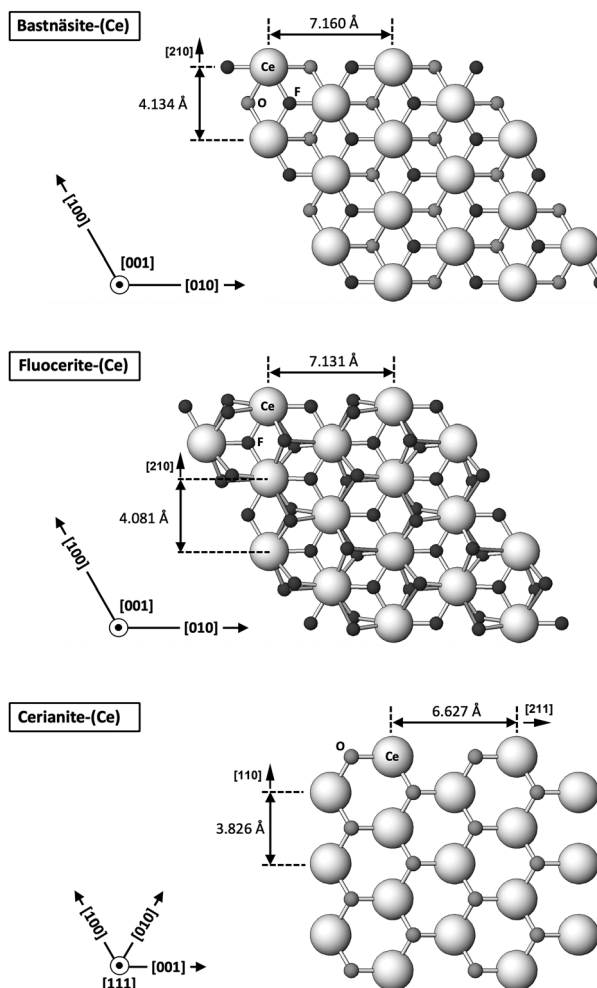


Fig. 13 Projection of the crystal structures of a (001) slice of Ce-bastnäsitate, a (001) slice of Ce-fluocerite and a (111) slice of cerianite showing the main directions of epitaxy and the repeating periods.

The other epitaxial overgrowth was only observed in the Ce experiments: when Ce-bastnäsitate undergoes complete breakdown as a consequence of the Ce³⁺ to Ce⁴⁺ oxidation process,⁶⁸ cerianite can grow epitaxially on the non-replaced remnants of fluorocerite crystals, specifically on their (001) surfaces. Although cerianite is a cubic mineral (space group *Fm* $\bar{3}m$; $a = 5.42(1)$ Å), the epitaxial overgrowth is feasible by the similar repeating periods of the (111) plane of cerianite and the (001) plane of fluorocerite (Table 3 and Fig. 13) with calculated misfit values <+8%, positive values that indicate that the unit cell of cerianite is slightly expanded along the different directions of epitaxy of fluorocerite. With time, the Ce³⁺ to Ce⁴⁺ oxidation would be translated into the full transformation of Ce-fluocerite into cerianite. The smaller molar volume of cerianite (32.21 cm³ mol⁻¹ (ref. 69)) compared to fluorocerite (23.86 cm³ mol⁻¹ (ref. 69)) generates porosity on the (001) face of the latter during the replacement, often resulting in skeletal crystals (Fig. 11).

All these epitaxial overgrowths between bastnäsitate, cerianite and fluorocerite have in common the close similarity in arrange-



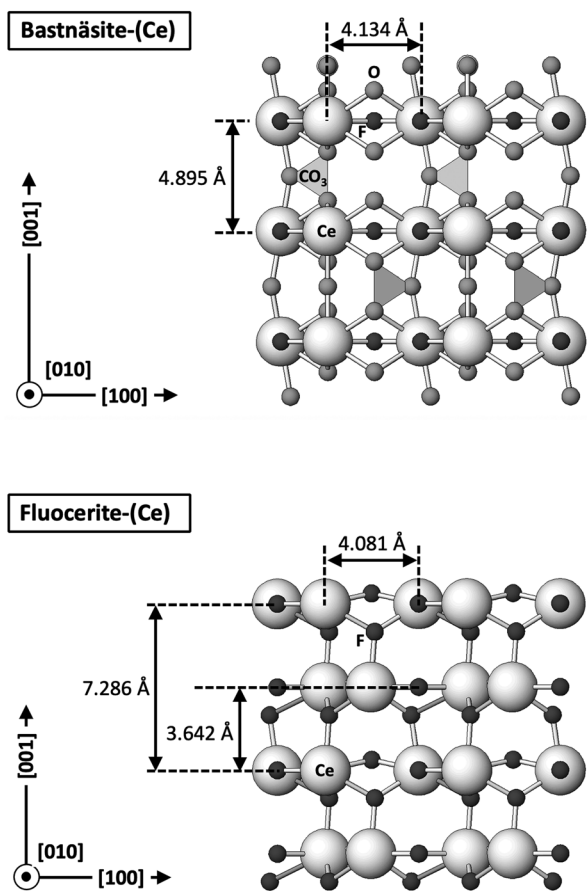


Fig. 14 Projection of the crystal structures of the (010) slices of Ce-bastnäsäsite and Ce-fluocerite showing the main directions of epitaxy and the repeating periods.

Table 3 Calculated misfits for alignment directions in fluocerite–bastnäsäsite and fluocerite–cerianite epitaxies

Minerals	Directions of alignment	Misfit (%)
Fluocerite–Bastnasite	$(001)_{fc}[010]_{fc} (001)_{bast}[010]_{bast}$	−0.42%
	$(001)_{fc}[210]_{fc} (001)_{bast}[210]_{bast}$	−1.28%
	$(010)_{fc}[100]_{fc} (010)_{bast}[100]_{bast}$	−1.28%
	$(010)_{fc}[001]_{fc} (010)_{bast}[001]_{bast}$	−25.60%
Fluocerite–Cerianite	$(001)_{fc}[210]_{fc} (111)_{cer}[110]_{cer}$	+6.66%
	$(001)_{fc}[010]_{fc} (111)_{cer}[211]_{cer}$	+7.61%

ment of the Ce–O and Ce–F groups of these three minerals (Fig. 13) and their similar bond distances (Ce–O distance of 2.34 Å in cerianite, Ce–F distances of 2.39 Å and 2.42 Å in bastnäsäsite and fluocerite, respectively)^{64,65,70} further supporting these structural matches. However, it is important to consider that the minerals obtained in our study were synthesized with targeted chemistries, using single REE. Although oriented overgrowth has been described in natural samples,⁶⁷ these three minerals often consist of complex solid solutions comprising multiple REE alongside other cations such as Ca and Sr.^{26,71,72} This chemical variability has the potential to influence the matching of unit cells and subsequently impact the

kinetics of mineral replacement processes, as well as the resulting solid morphology.

No experiment resulted in the complete carbonation of the original fluocerite phase at any temperature, irrespective of the REE used. To better understand this partial carbonation reaction, it is important to evaluate the structural relationship between the host mineral and the newly formed REE-fluorocarbonates, taking two main factors into account: (i) The higher molar volumes of the REE-fluorocarbonates (e.g., 42.52 cm³ mol^{−1} for La-Bastnäsäsite; 45.59 cm³ mol^{−1} for Nd-kozoite⁶⁹) compared to fluocerite (e.g., 32.99 cm³ mol^{−1} for Fluocerite-La⁶⁹) promote partial equilibrium situations (e.g.,^{68,73,74}). They occur when reactive host grains of fluocerite are covered by a crust of the precipitating REE-fluorocarbonate phases, which isolate the reactive solid from the aqueous solution, slowing down the carbonation reaction. (ii) Also, the bastnäsäsite overgrowth is aligned with the orientation of the fluocerite due to the lattice matching between the two minerals, with very low misfit values (Table 3⁶⁵). This epitaxial overgrowth of the REE-fluorocarbonates crystals relative to the host fluocerite surfaces results in a highly uniform coating, further enhancing the partial equilibrium situation. As a comparison, it is worth noting that while the molar volume of fluocerite (e.g., La-fluocerite 32.99 cm³ mol^{−1} (ref. 69)) is larger than that of fluorite (24.55 cm³ mol^{−1} (ref. 69)), the fluorite–fluocerite transformation experiments did not result in a clearly defined state of partial equilibrium. This can be attributed to the lack of oriented overgrowth of fluocerite on the host fluorite crystals, hindering the complete coating of the host fluorite and allowing a full replacement reaction. However, it cannot be ruled out that in some experiments some partial equilibrium could have occurred at the very late stages of the reaction. Supporting this hypothesis is the observation of very minor quantities of fluorite persisting for an extended period at the end of some of the experiments (Tables SI-4, SI-5 and SI-6[†]).

Carbonation experiments in the La and Nd systems showed very similar trends with small variations. For example, only the La experiment at 50 °C showed a full replacement of the host fluocerite, which could be explained by the small particle size of the host. Also, in the Nd experiments the formation of bastnäsäsite took place *via* kozoite (NdCO₃OH), an orthorhombic precursor of bastnäsäsite that belongs to the ancylite mineral group.⁷⁵ Its occurrence solely in the Nd system can be attributed to the longer lifetime of Nd-kozoite compared to La-kozoite in crystallisation from solution experiments and in mineral replacement reactions (e.g.,^{48,59,76}).

The Ce system deserves a separate discussion. Unlike the La and Nd systems, the Ce experiments did not achieve a state of partial equilibrium. To account for this behaviour, we propose a multi-step crystallization process: fluocerite → bastnäsäsite → cerianite. This way, the process in the Ce system would involve three main steps: (i) dissolution of host crystals in the CO₃^{2−} bearing solution with the release of Ce³⁺ and F[−], leading to the crystallisation of bastnäsäsite-Ce; (ii) redox oxidation from Ce³⁺ to Ce⁴⁺, which happens since Ce⁴⁺ has a stable electronic configuration of a nearest noble gas and is more stable than Ce⁰



and Ce^{3+} ,^{77,78} and (iii) subsequent decarbonation of bastnäsite-Ce and decrease of pH, followed by cerianite crystallization. The works of Szucs *et al.* (2023)⁶⁸ and Janos *et al.* (2017),⁷⁹ who synthesized cerianite *via* homogeneous precipitation, validate this hypothesis since they both reported Ce carbonates as intermediate phases. In our experiments, Ce fluorocarbonates are metastable and the decarbonation process during the transformation from bastnäsite-Ce to cerianite is translated into an increase of the porosity since the smaller molar volume of the latter ($32.21 \text{ cm}^3 \text{ mol}^{-1}$ (ref. 69)) compared to the former ($23.86 \text{ cm}^3 \text{ mol}^{-1}$ (ref. 69)). Cerianite is the most thermodynamically stable phase in this system, as its solubility ($\log K_{\text{cer}} = -59.3 \pm 0.3$ in 0.01 M NaClO_4 ⁸⁰) is lower compared to bastnäsite⁸¹ or fluocerite.⁸²

In La and Ce carbonation experiments conducted at $200 \text{ }^\circ\text{C}$ for 60 days, SEM analyses also unveiled intriguing morphologies characterized by skeletal crystals (Fig. 11). Remarkably, the (001) faces of fluocerite exhibited complete dissolution in the central region, leaving behind only the hexagonal rings of the original crystal. We suggest that dissolution may begin at the center of the fluocerite hexagonal crystals, spreading outward. Three main factors could drive this process: (i) enhanced isolation of the host fluocerite crystal rims from the fluid by a thicker bastnäsite crust during carbonation, (ii) preferential dissolution of the central area of thin fluocerite crystals, and (iii) porosity development during the coupled dissolution of Ce-bastnäsite and crystallisation of cerianite. Further in-depth experimental investigation into the dissolution of fluocerite would be needed to fully comprehend the specific mechanisms of formation of these skeletal crystals.

Implications

Our experiments demonstrate that fluocerite can act as a transient precursor to bastnäsite and cerianite. The primary mechanism for the formation of these minerals involves a dissolution-recrystallization process mediated by the hydrothermal solution, resulting in a pseudomorphic mineral replacement. The ionic potential of the three tested REE plays a crucial role in this process, influencing the kinetics and stability of the mineral phases. Additionally, while the aqueous solution facilitates these mineralogical transformations, the structural similarity between bastnäsite and fluocerite promotes partial equilibrium situations. This highlights the intricate interplay between ionic potential, solution chemistry, and structural factors in driving the observed mineral replacement reactions.

Our understanding of the crystallisation mechanisms of fluocerite and its role as a precursor phase to bastnäsite and cerianite allows comprehending the geological origin of critical minerals that crucial for rare earth element extraction. Our experimental results align with research on natural fluocerite, which has been found in REE ore deposits and associated with REE bearing minerals (*e.g.*,^{23,24,26,29}) and has been suggested to be a precursor phase in the formation of fluorocarbonates and cerianite.^{12,16,23,24,26,29} This work also provides new

insights into the formation of REE-fluorocarbonates and cerianite in REE hydrothermal systems, such as those found in carbonate deposits, of which an example is Bayan Obo REE deposit in China, the largest known in the world.^{3,4,10-12}

Although the association between F^- and REE is well established, there are gaps in the understanding of its significance for the genesis REE deposits.⁸³ REE deposits generally display an association between REE and Fe mineralizations and F-bearing minerals, suggesting the presence of F-rich mineralizing fluids.^{83,84} Hydrogeochemical modelling of high-temperature experimental data has revealed that Cl^- , SO_4^{2-} and $\text{HCO}_3^-/\text{CO}_3^{2-}$ ions play a main role as ligands for the transport of the REE at hydrothermal conditions^{33,59,85} as well as OH^- in near-neutral and alkaline conditions.⁵⁹ Also, F^- has long been considered one of the primary ligands facilitating the solubility and hydrothermal transport of REE.^{32,33,85} However, recent studies suggest that F^- functions as a depositional ligand, promoting the effective formation of monazite, xenotime, and bastnäsite in REE-bearing solutions containing phosphate and/or carbonate.^{32,33,85}

In our carbonation experiments we were able to synthesise bastnäsite at low hydrothermal conditions ($50\text{--}90 \text{ }^\circ\text{C}$). This result was unexpected because the non-fluorobastnäsite (hydroxylbastnäsite, REECO_3OH) is known to form at hydrothermal conditions, usually above $150\text{--}200 \text{ }^\circ\text{C}$.^{48,59,68,76,86} According to Hsu (1992)⁸⁷ and Wakita and Kinoshita (1979)⁸⁸ the lower limit of the La- and Ce-hydroxylbastnäsite is at temperatures comprised between 100 and $115 \text{ }^\circ\text{C}$ at 1 atm whereas La- and Ce-fluorobastnäsite resulted to be stable at higher temperature compared to their corresponding hydroxylbastnäsites.⁸⁷ We hypothesize that the incorporation of F^- into the hydration shell of REE^{3+} could lower the energy required for desolvation of the REE^{3+} ion, allowing the nucleation of fluorobastnäsite under low hydrothermal conditions. This REE-F association would play a similar role than the $\text{Mg}^{2+}\text{-SO}_4^{2-}$ ion in calcite: It is known that the structural incorporation of Mg^{2+} into growing calcite crystals requires high temperatures.⁸⁹⁻⁹¹ However, experimental work has found that SO_4^{2-} facilitates the desolvation of the Mg^{2+} ion during the growth of calcite in Mg-rich fluids.⁹²⁻⁹⁴ These studies, also supported by molecular modelling showing that substitution of the MgSO_4 ion into the calcite surface is energetically more favoured than pure Mg^{2+} , demonstrated that the $\text{Mg}^{2+}\text{-SO}_4^{2-}$ ion pairing effect enhances the incorporation of Mg^{2+} into calcite, especially at ambient conditions. Our findings indicate that the low-temperature formation of fluorobastnäsite could be facilitated by the formation of $\text{REE}^{3+}\text{-F}^-$ ion pairs, akin to a similar effect observed with Li^+ ions, where the energy barrier for Li^+ desolvation is decreased in the presence of F^- .⁹⁵

Although our experiments involve single REE ions, the results provide valuable insights into the mechanisms of bastnäsite formation in natural fluids. Bastnäsite is naturally enriched in light REE, as heavy REE are too small to form this mineral and typically enter as impurities along with light REE.⁶⁵ This suggests that bastnäsite would be the final thermodynamically stable phase in nature. While the for-



mation of metastable kozoite (a known precursor of bastnäsité^{59,76}) cannot be entirely ruled out, it would be very scarce compared to bastnäsité. The presence of kozoite-(Nd) in experiments not involving fluocerite is quite common across a wide range of temperatures, from ambient to over 200 °C, both in crystallisation from solution experiments⁷⁶ and in mineral replacement reactions.^{48,59,96} Also, natural fluocerite and bastnäsité containing some Ce are susceptible to oxidation, leading to the crystallization of cerianite that would replace the host solid at the nanoscale, accumulating as a discrete phase.⁹⁶

Our findings suggest that the formation of bastnäsité from fluocerite at low hydrothermal conditions could offer valuable insights for laboratory and industrial synthesis of these and other REE crystalline compounds. These minerals are valued for their optical properties^{97–100} and other industrial applications, including in oxygen sensors, scintillators, lasers, telecommunications, and photonic and biophotonic devices.^{35,99} The methodology employed in our study is simple, environmentally friendly, highly reproducible, and cost-effective for producing fluocerite and bastnäsité and could serve as a foundation for refining and innovating targeted synthesis methods for REE-bearing compounds.

Conclusions

This study presents groundbreaking insights into the crystallisation pathways, mechanisms and kinetics of fluocerite formation from fluorite and its subsequent transformation into bastnäsité and cerianite. It highlights the significant influence of REE mass, temperature, and structural compatibility between all these crystalline phases, as well as the interplay between the molar volume differences of the reacting minerals and the epitaxial overgrowth of bastnäsité on fluocerite, which can be translated into partial equilibrium situations during carbonation processes. Additionally, this work offers insights into the geological origins of these REE-bearing minerals, emphasizing the role of F⁻ ions in facilitating the low-temperature synthesis of REE-fluorocarbonates, which may also impact laboratory and industrial production of REE fluorides used in technological applications.

Author contributions

Luca Terribili: experimental work, data analysis and interpretation, manuscript writing – original draft, review and editing; Remi Rateau: data analysis; Melanie Maddin: data analysis; Juan Diego Rodriguez-Blanco: conceptualization, data analysis, writing (review and editing), supervision, validation, funding acquisition.

Conflicts of interest

The authors declare no competing financial interest.

Acknowledgements

This publication has emanated from research conducted with the financial support of Science Foundation Ireland, Geological Survey of Ireland and the Environmental Protection Agency under the SFI Frontiers for the Future Programme 19/FFP/6771 “SEparating Critical metals ThrOugh mineRal crystallization (SEleCTOR)” (rodrigid-SFI-17/RC-PhD/3481). We are extremely grateful for the help of Colin Reid, Paul Guyett and Leona O'Connor from the iCrag Lab at TCD, and Dr Robbie Goodhue in Unit 7, Trinity Technology and Enterprise Centre (TTEC).

References

- 1 K. Smith Stegen, *Energy Policy*, 2015, **79**, 1–8.
- 2 J. M. Klinger, *Extr. Ind. Soc.*, 2015, **2**, 572–580.
- 3 V. Balaram, *Geosci. Front.*, 2019, **10**, 1285–1303.
- 4 N. Dushyantha, N. Batapola, I. M. S. K. Ilankoon, S. Rohitha, R. Premasiri, B. Abeysinghe, N. Ratnayake and K. Dissanayake, *Ore Geol. Rev.*, 2020, **122**, 103521.
- 5 D. Gielen and M. Lyons, *Critical materials for the energy transition: Rare earth elements*, International Renewable Energy Agency, Abu Dhabi, 2022.
- 6 K. M. Goodenough, F. Wall and D. Merriman, *Nat. Resour. Res.*, 2017, **27**, 201–216.
- 7 B. Zhou, Z. Li and C. Chen, *Minerals*, 2017, **7**(11), 203.
- 8 S. Carrara, S. Bobba, D. Blagoeva, P. Alves Dias, A. Cavalli, K. Georgitzikis, M. Grohol, A. Itul, T. Kuzov, C. Latunussa, L. Lyons, G. Malano, T. Maury, A. Prior Arce, J. Somers, T. Telsnig, C. Veeh, D. Wittmer, C. Black, D. Pennington and M. Christou, *Publications Office of the European Union*, Luxembourg, 2023.
- 9 J. Lucas, P. Lucas, T. Le Mercier, A. Rollat and W. Davenport, *Rare Earths – Science, Technology, Production and Use*, Elsevier, Amsterdam, 2015.
- 10 Z. Weng, S. M. Jowitt, G. M. Mudd and N. Haque, *Applied Earth Science, Trans. Inst. Min. Metall., Sect. B*, 2013, **122**(2), 83–96.
- 11 M. R. Payne, A. P. Gysi and N. C. Hurtig, *Chem. Geol.*, 2023, **617**, 121256.
- 12 W. Zhang, W. T. Chen and A. E. Williams-Jones, *Contrib. Mineral. Petrol.*, 2023, **178**, 34.
- 13 Z. Weng, S. M. Jowitt, G. M. Mudd and N. Haque, *Econ. Geol.*, 2015, **110**, 1925–1952.
- 14 W. Zhang, W. T. Chen, T. P. Mernagh and L. Zhou, *Miner. Deposita*, 2021, **57**(6), 1–19.
- 15 M. Smith, J. Kynicky, C. Xu, W. Song, J. Spratt, T. Jeffries, M. Brtnicky, A. Kopriva and D. Cangelosi, *Lithos*, 2018, **308–309**, 65–82.
- 16 A. C. Strzelecki, A. Migdisov, H. Boukhalfa, K. Sauer, K. G. McIntosh, R. P. Currier, A. E. Williams-Jones and X. Guo, *Nat. Geosci.*, 2022, **15**, 327–333.
- 17 A. E. Williams-Jones, I. M. Samson and G. R. Olivo, *Econ. Geol.*, 2000, **95**(2), 327–341.



- 18 A. P. Gysi and A. E. Williams-Jones, *Geochim. Cosmochim. Acta*, 2013, **122**, 324–352.
- 19 A. P. Gysi, A. E. Williams-Jones and P. Collins, *Econ. Geol.*, 2016, **111**(5), 1241–1276.
- 20 C. D. Beard, K. M. Goodenough, A. M. Borst, F. Wall, P. R. Siegfried, E. A. Deady, C. Pohl, W. Hutchison, A. A. Finch, B. F. Walter, H. A. Elliott and K. Brauch, *Econ. Geol.*, 2022, **118**(1), 32.
- 21 A. K. Patel, B. Mishra, D. Upadhyay and K. L. Pruseth, *Econ. Geol.*, 2022, **117**, 683–702.
- 22 P. L. Verplanck, G. L. Farmer, R. M. Kettler, H. A. Lowers, C. A. Johnson, A. E. Koenig and M. J. Blessington, *Ore Geol. Rev.*, 2022, **146**, 104953.
- 23 S. I. Lahti and V. I. Suominen, *Bull. Geol. Soc. Finl.*, 1988, **60**, 45–53.
- 24 D. Holtstam and U. B. Andersson, *Can. Mineral.*, 2007, **45**, 1073–1114.
- 25 A. E. Williams-Jones and S. A. Wood, *Geochim. Cosmochim. Acta*, 1992, **56**(2), 725–738.
- 26 M. T. Styles and B. R. Young, *Mineral. Mag.*, 1983, **47**, 41–46.
- 27 E. W. M. Heinrich and E. B. Gross, *Am. Mineral.*, 1960, **45**, 455–459.
- 28 R. M. Perhac and E. W. Heinrich, *Econ. Geol.*, 1964, **59**(2), 226–239.
- 29 G. J. Beukes, H. D. Bruijn and W. A. Westhuizen, *S. Afr. J. Geol.*, 1991, **94**, 313–320.
- 30 A. E. Williams-Jones, A. A. Migdisov and I. M. Samson, *Elements*, 2012, **8**, 355–360.
- 31 M. M. Muller, H. J. Kleebe, S. Lauterbach and G. Zito, *Z. Kristallogr.*, 2011, **226**, 467–475.
- 32 A. A. Migdisov and A. E. Williams-Jones, *Miner. Deposita*, 2014, **49**(8), 987–997.
- 33 A. A. Migdisov, A. E. Williams-Jones, J. Brugger and F. A. Caporuscio, *Chem. Geol.*, 2016, **439**, 13–42.
- 34 B. Maximov and H. Schulz, *Acta Crystallogr.*, 1985, **41**(2), 88–91.
- 35 W. A. Crichton, P. Bouvier, B. Winkler and A. Grzechnik, *Dalton Trans.*, 2010, **39**, 4302–4311.
- 36 J. Berndt and S. Klemme, *Nat. Commun.*, 2022, **13**, 2892.
- 37 G. M. Yaxley, M. Anenburg, S. Tappe, S. Decree and T. Guzmics, *Annu. Rev. Earth Planet. Sci.*, 2022, **50**, 261–293.
- 38 J. D. Rodriguez-Blanco, S. Shaw and L. G. Benning, *Mineral. Mag.*, 2008, **72**, 283–286.
- 39 D. J. Tobler, J. D. Rodriguez-Blanco, H. O. Sørensen, S. L. S. Stipp and K. Dideriksen, *Cryst. Growth Des.*, 2016, **16**(8), 4500–4508.
- 40 L. Terribili, R. Rateau, A. M. Szucs, M. Maddin and J. D. Rodriguez-Blanco, *Cryst. Growth Des.*, 2024, **24**(2), 632–645.
- 41 A. A. Coelho, *Topas Academic v4.1*, Coelho Software, Brisbane, Australia, 2007.
- 42 W. S. Rasband, *ImageJ*, U. S. National Institutes of Health, Bethesda, Maryland, USA, 1997–2018.
- 43 M. D. Abramoff, P. J. Magalhaes and S. J. Ram, *Biophotonics Int.*, 2004, **11**(7), 36–42.
- 44 C. A. Schneider, W. S. Rasband and K. W. Eliceiri, *Nat. Methods*, 2012, **9**, 671–675.
- 45 M. Avrami, *J. Chem. Phys.*, 1939, **7**, 1103.
- 46 A. Putnis, *Introduction to Mineral Sciences*, Cambridge University Press, Cambridge, 1992.
- 47 F. Xia, J. Brugger, G. Chen, Y. Ngothai, B. O'Neill, A. Putnis and A. Pring, *Geochim. Cosmochim. Acta*, 2009, **73**(7), 1945–1969.
- 48 A. M. Szucs, M. Maddin, D. Brien, P. C. Guyett and J. D. Rodriguez-Blanco, *Global Challenges*, 2022, 2200085.
- 49 S. Shaw, S. E. Pepper, N. D. Bryan and F. R. Livens, *Am. Mineral.*, 2005, **90**(11–12), 1852–1860.
- 50 L. E. Davidson, S. Shaw and L. G. Benning, *Am. Mineral.*, 2008, **93**(8–9), 1326–1337.
- 51 J. D. Rodriguez-Blanco, S. Shaw and L. G. Benning, *Nanoscale*, 2011, **3**, 265–271.
- 52 A. C. Lasaga, *Kinetic Theory in the Earth Sciences*, Princeton University Press, Princeton, New Jersey, 1998.
- 53 N. Yee, S. Shaw, L. G. Benning and T. H. Nguyen, *Am. Mineral.*, 2006, **91**(1), 92–96.
- 54 K. M. Peterson, P. J. Heaney and J. E. Post, *Chem. Geol.*, 2016, **444**, 27–36.
- 55 G. Montanari, J. D. Rodriguez-Blanco, N. Bovet, S. L. S. Stipp and D. J. Tobler, *Cryst. Growth Des.*, 2017, **17**, 5269–5275.
- 56 M. Ossorio, T. M. Stawski, J. D. Rodriguez-Blanco, M. Sleutel, J. M. Garcia-Ruiz, L. G. Benning and A. E. S. Van Driessche, *Minerals*, 2017, **7**, 140.
- 57 L. Granasy, T. Pusztai, G. Tegze, J. A. Warren and J. F. Douglas, *Phys. Rev. E: Stat., Nonlinear, Soft Matter Phys.*, 2005, **72**, 011605.
- 58 A. Jordens, Y. P. Cheng and K. E. Waters, *Miner. Eng.*, 2013, **41**, 97–114.
- 59 A. M. Szucs, A. Stavropoulou, C. O'Donnell, S. Davis and J. D. Rodriguez-Blanco, *Cryst. Growth Des.*, 2021, **21**(1), 512–527.
- 60 L. B. Railsback, *Geology*, 2003, **31**(9), 737–740.
- 61 G. H. Cartledge, *J. Am. Chem. Soc.*, 1928, **50**(11), 2855–2863.
- 62 G. H. Cartledge, *J. Am. Chem. Soc.*, 1928, **50**(11), 2863–2872.
- 63 R. D. Shannon, *Acta Crystallogr., Sect. A: Cryst. Phys., Diffraction, Theor. Gen. Crystallogr.*, 1976, **32**, 751–767.
- 64 A. K. Cheetham, B. E. F. Fender, H. Fuess and A. F. Wright, *Acta Crystallogr., Sect. A: Cryst. Phys., Diffraction, Theor. Gen. Crystallogr.*, 1976, **32**, 94–97.
- 65 Y. Ni, J. M. Hughes and A. N. Mariano, *Am. Mineral.*, 1993, **78**(3–4), 415–418.
- 66 J. H. Van Der Merwe, *Crit. Rev. Solid State Mater. Sci.*, 1978, **7**, 209–231.
- 67 A. G. Walton, in *Nucleation*, ed. A. C. Zettlemoyer, Dekker, New York, 1969, pp. 225–307.
- 68 A. M. Szucs, M. Maddin, D. Brien, R. Rateau and J. D. Rodriguez-Blanco, *RSC Adv.*, 2023, **13**, 6919–6935.
- 69 A. V. Chichagov, D. A. Varlamov, R. A. Dilanyan, T. N. Dokina, N. A. Drozhzhina, O. L. Samokhvalova and T. V. Ushakovskaya, *Crystallogr. Rep.*, 2001, **46**, 876–879.
- 70 S. Vyas, R. W. Grimes, D. H. Gay, A. L. Rohl and J. Chem. Soc., *Faraday Trans.*, 1998, **94**, 427–434.



- 71 A. N. Zaitsev, A. R. Chakhmouradian, O. I. Siidra, J. Spratt, C. T. Williams, C. J. Stanley, S. V. Petrov, S. N. Britvin and E. A. Polyakova, *Mineral. Mag.*, 2011, **75**(6), 2813–2822.
- 72 Y. Singh, G. Nagendra Babu, R. Viswanathan, M. Sai Baba, A. K. Rai and P. S. Parihar, *Central Eur. J. Geosci.*, 2014, **6**(4), 492–505.
- 73 H. C. Helgeson, *Geochim. Cosmochim. Acta*, 1968, **32**(8), 853–877.
- 74 J. D. Rodríguez, A. Jiménez, M. Prieto, L. Torre and S. García-Granda, *Am. Mineral.*, 2008, **93**, 928–939.
- 75 R. Miyawaki, S. Matsubara, K. Yokoyama, K. Takeuchi, Y. Terada and I. Nakai, *Am. Mineral.*, 2000, **85**(7–8), 1076–1081.
- 76 B. Vallina, J. D. Rodriguez-Blanco, A. P. Brown, J. A. Blanco and L. G. Benning, *Nanoscale*, 2015, **7**, 12166–12179.
- 77 D. G. Brookins, Yttrium and the Rare Earth Elements (REE), in *Eh-pH diagrams for geochemistry*, ed. D. G. Brookins, Springer Berlin Heidelberg, Berlin, Heidelberg, 1988.
- 78 B. Bilal, E. Müller and E. Z. Naturforsch, A: *Phys. Sci.*, 1992, **47**(9), 974–984.
- 79 P. Janos, J. Henych, J. Pfeifer, N. Zemanova, V. Pilarova, D. Milde, T. Opletal, J. Tolasz, M. Maly and V. Stengl, *Environ. Sci.: Nano*, 2017, **4**, 1283–1293.
- 80 T. V. Plakhova, A. Y. Romanchuk, S. N. Yakunin, T. Dumas, S. Demir, S. Wang, S. G. Minasian, D. K. Shuh, T. Tyliczszak, A. A. Shiryayev, A. V. Egorov, V. K. Ivanov and S. N. Kalmykov, *J. Phys. Chem.*, 2016, **120**(39), 22615–22626.
- 81 M. Voigt, J. D. Rodriguez-Blanco, B. Vallina, L. G. Benning and E. H. Oelkers, *Chem. Geol.*, 2016, **430**, 7077.
- 82 T. Mioduski, C. Guminski, D. Zeng and J. Phys, *Chem. Ref. Data*, 2015, **44**(1), 013102–013101.
- 83 Y. Xing, B. Etschmann, W. Liu, Y. Mei, Y. Shvarov, D. Testemale, A. Tomkins and J. Brugger, *Chem. Geol.*, 2018, **504**, 158–176.
- 84 S. Salvi and A. E. Williams-Jones, *Geochim. Cosmochim. Acta*, 1996, **60**, 1917–1932.
- 85 E. P. Perry and A. P. Gysi, *Geofluids*, 2018, **5382480**, 1–19.
- 86 B. Vallina, J. D. Rodriguez-Blanco, J. A. Blanco and L. G. Benning, *Mineral. Mag.*, 2014, **78**, 1391–1397.
- 87 L. C. Hsu, *Mineral. Petrol.*, 1992, **47**, 87–101.
- 88 H. Wakita and S. Kinoshita, *Bull. Chem. Soc. Jpn.*, 1979, **52**, 428–432.
- 89 W. Mejri, A. Korchef, M. Tlili and M. Ben Amor, *Desalin. Water Treat.*, 2014, **52**(25–27), 4863–4870.
- 90 G. D. Saldi, G. Jordan, J. Schott and E. H. Oelkers, *Geochim. Cosmochim. Acta*, 2009, **73**(19), 5646–5657.
- 91 D. Di Tommaso and N. H. de Leeuw, *Phys. Chem. Chem. Phys.*, 2010, **12**(4), 894–901.
- 92 M. R. Nielsen, K. K. Sand, J. D. Rodriguez-Blanco, N. Bovet, J. Generosi, K. N. Dalby and S. L. S. Stipp, *Cryst. Growth Des.*, 2016, **16**, 6199–6207.
- 93 H. Sakuma, M. P. Andersson, K. Bechgaard, S. L. S. Stipp and J. Phys, *Chem. C*, 2014, **118**(6), 3078–3087.
- 94 D. Kralj, J. Kontrec, L. Brecevic, G. Falini and V. Nothing-Laslo, *Chem. – Eur. J.*, 2004, **10**, 1647–1656.
- 95 J. Zheng, Y. Wang, J. Wang, H. Yuan, Y. Liu, T. Liu, J. Luo, J. Nai and X. Tao, *Appl. Mater. Interfaces*, 2022, **14**(43), 48762–48769.
- 96 M. Maddin, R. Rateau, A. M. Szucs, L. Terribili, B. Hoare, P. C. Guyett and J. D. Rodriguez-Blanco, *Global Challenges*, 2024, 2400074.
- 97 G. Hass, J. B. Ramsey and R. Thun, *J. Opt. Soc. Am.*, 1959, **49**(2), 116–120.
- 98 M. M. Lezhnina, T. Jüstel, H. Kätker, D. U. Wiechert and U. H. Kynast, *Adv. Funct. Mater.*, 2006, **16**(7), 935–942.
- 99 V. Bartunek, V. Jakes, V. Kral and J. Rak, *J. Fluorine Chem.*, 2012, **135**, 358–361.
- 100 R. K. Sharma, A. V. Mudring and P. Ghosh, *J. Lumin.*, 2017, **189**, 44–63.

